PHASE I REPORT

SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS IN SOUTHWESTERN UTAH

Earth Sciences Associates

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Prepared for

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IN

SOUTHWESTERN UTAH

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PHASE I REPORT SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS

in

SOUTHWESTERN UTAH

I. INTRODUCTION

A. General

An investigation of geologic, tectonic, and seismic conditions in the vicinity of eight Soil Conservation Service (SCS) flood control dams located in southwestern Utah (see Figure 1) was conducted during late 1981 and early 1982 by Earth Sciences Associates. The conditions of the embankments and their foundations was also investigated by a drilling, test pitting, sampling, and testing program. Results of the first phase of this investigation (as described under Purpose and Scope below) are presented in this report. Detailed descriptions of the field investigation, laboratory investigations, regional geology and tectonics, and seismicity of the region are presented in the appendices contained in a separate volume.

B. Purpose and Scope

The two major purposes of the overall investigation are:

- o To determine the safety of these dams with respect to potential for surface fault rupture in the foundation areas (and consequent offset of the embankments); and
- o To evaluate the ability of the embankments to resist strong seismic shaking likely to occur in the region, along with determination of the characteristics of shaking that is probable at each dam site.

This "Phase I" report is a mid-investigation report being written after the completion of the field investigation phases of the investigation, but before the embankment stability analyses have been performed. It primarily reports on geologic/tectonic/seismic conditions in southwestern Utah and in the specific dam site areas, but also describes embankment conditions found to exist during the drilling and test pit excavation portions of the field investigation. The specific tasks reported on herein are listed in the ESA proposal for this project dated September, 1981 under "Phase I - Seismic Hazard Study" as Tasks a through m. A portion of the laboratory testing program (which is part of Phase II (Tasks IIa and b)), has been performed and is reported on in Appendix D.

End products of the overall study will include an evaluation of the safety of the dams with respect to faulting and seismic shaking, the development of recommendations for remedial measures should hazards be found to exist, and the development of recommendations for courses of action for the SCS to take both under existing conditions and following an earthquake (if any of the dams is determined to have significant safety problems). These matters will be addressed in the Phase II report to follow.

C. Background Information

The dams investigated during this investigation are relatively low earthfill embankments. They were constructed for the purpose of temporarily retaining flood waters that flow down the drainages of this semi-arid environment during infrequent periods of heavy rainfall ("thunder showers"). They were constructed by the SCS in order to protect downstream areas from flooding that had periodically occurred before construction of these dams. Since construction of the dams, rapid increases in population in the St. George and Cedar City areas has changed the downstream land use markedly. Areas that were once dominantly agricultural or rural are becoming more urbanized, and houses have been constructed in areas just downstream of some of the dams (Green Lakes 2 and 3, and Ivins, in particular).

As a result of these rapid changes in downstream land use, the consequences of failure of some of these structures has increased markedly. Consequently, even though these reservoirs are infrequently filled, the SCS has contracted to have the safety of these dams evaluated. This was done in order that remedial measures can be developed in the event that hazards are determined to exist, and so that there would be a basis for decisions about the dams' relative usefulness in mitigating versus creating potential hazards to downstream residents and property.

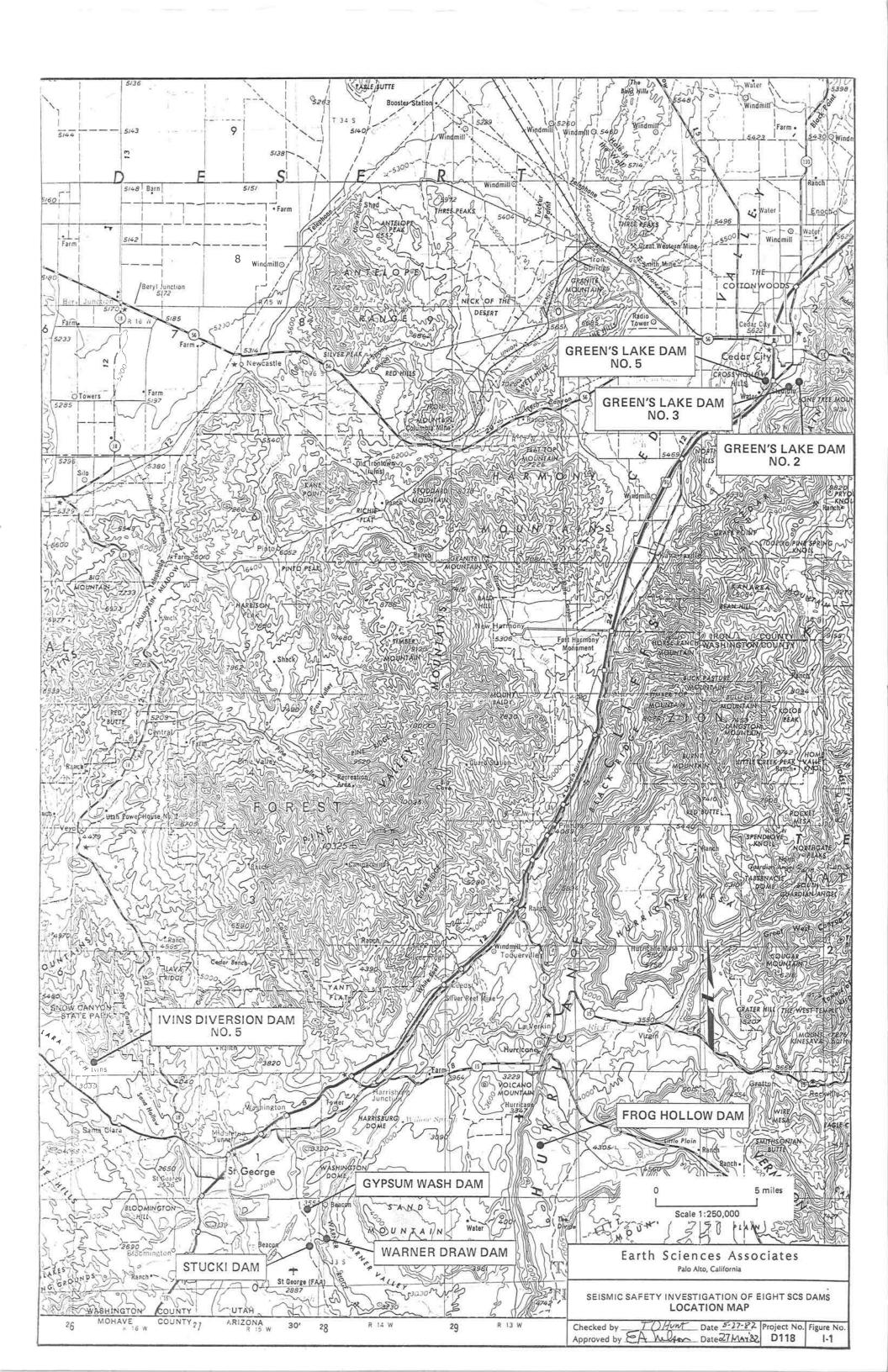
D. Performance

Work on this project was performed for the SCS West Technical Center in Portland, Oregon. C. Edward Stearns is the Project Coordinator for the SCS, and Joan K. Johnson is the Contract Administrator. Other SCS personnel who have been directly involved with this project are Don Wallin, Bob Nelson, Bob Rasely, and Claud Scoles.

ESA personnel who have worked on this project include: Douglas M. Yadon, who logged the drill holes and supervised the initial field operations; Richard Morris, who logged the test pits and performed the field density tests; Robert H. Wright, who performed the photogeologic studies, and developed the strip map of the Hurricane fault zone and the site specific photogeologic mapping; T. Dwight Hunt, who performed the field geologic mapping, logged the trenches, and developed the site specific geologic maps and trench logs; Michael L. Traubenik, who analyzed the existing SCS data and made engineering evaluations of the embankments; Julio E. Valera, who was Project Manager and coordinated the engineering and budget control portions of the project; and Eugene A. Nelson, who was Principal in Charge and coordinated the geologic portions of the work. The report was written by Wright, Hunt, Traubenik, Valera and Nelson, and typed by Kathleen McCracken. Tom Camara and Dave O'Shea performed the drafting.

Working as a consultant to Lindvall Richter and Associates (LRA), who are subcontractors to ESA on this project, Dr. Walter J. Arabasz performed the

evaluations of seismic and tectonic conditions in the region, and wrote the extensive discussions on these subjects that appear in Appendices E and F. Richard J. Proctor of LRA coordinated the seismic studies for LRA.



II. CONCLUSIONS AND RECOMMENDATIONS

Preliminary conclusions and recommendations pertaining to the subject matter of this initial report are presented below.

- 1. Trenches excavated near the south corner of Gypsum Wash Dam expose offsets in young-appearing sedimentary deposits that have been estimated to be of Holocene age by ESA's consultant on soil stratigraphy, Dr. Roy J. Shlemon (see Chapter VIII). This finding differs with previously available information regarding recency of faulting in southwestern Utah. Nevertheless, ESA believes that these near-surface offsets may be of consequence to the dam. The potential for damage to the embankments depends upon the recurrence interval between episodes of surface faulting and the amount of displacement possible during a single event. These factors, and the consequent potential for damage to the embankments, will be evaluated during Phase II of this investigation.
- 2. No evidence of faulting was found in trenches excavated at the Green's Lake 2 and 3 Dams, but it is obvious from the photogeologic studies that these dams are within a zone containing several strands of the Hurricane fault (see Figures IV-2 and VIII-5). Charcoal samples from the deposits exposed in these trenches have been submitted for age dating, and should provide an indication of the age of unfaulted surficial deposits at these sites. The potential for damage to the embankments of Green's Lake Dams No. 2 and 3 will be evaluated during Phase II of this investigation.
- 3. Relatively young (late Pleistocene) faulting was exposed in a trench excavated along the eastern margin of Green's Lake No. 5 Reservoir, and faults of similar age probably exist along the eastern front of the basalt near the main dam (see Figure VIII-5). These faults do not pass beneath the main dam, however, which is located west of the western edge of the fault-bounded north-south valley (see Figure VIII-10).
- 4. There is no potential for surface fault offset at the other four dam sites (Frog Hollow, Warner Draw, Stucki, and Ivins) (see detailed descriptions of geologic

conditions at the dam sites in Chapter VIII of this report). Strong ground shaking is likely to occur at all of these sites (see Conclusion 7), and the seismic stability analyses performed during Phase II of this investigation will include consideration of the characteristics of shaking likely to occur at the dam sites during their useful life.

- 5. A drill hole located on the crest of Frog Hollow Dam encountered very loose soil deposits in one portion of the foundation. Further drilling to either side of the first hole did not encounter similar conditions. SCS records indicate that the loose materials encountered may represent improperly placed backfill of a trench containing a pipe removed from the smaller original dam that was buried beneath the raised embankment. The SCS is evaluating this situation outside the scope of this contract.
- 6. The regional tectonic/seismic environment of southwestern Utah is complex, and has been comprehensively described by Dr. Walter J. Arabasz in Appendices E and F, which are summarized in Chapters IV, V, VI, and VII. The Cedar City-St. George area lies along a tectonically active intraplate boundary. While the level of seismicity is not as high as that along the Pacific/North American plate boundary, the historical record contains more than 20 earthquakes in the Magnitude 5 to $6\frac{1}{2}$ + range within 200 km of the study area (see Table F-1). The recurrence interval for larger events is longer than the available record (see Table F-3), but a Magnitude 7 to $7\frac{1}{2}$ event is thought to be the maximum credible earthquake for the region with a recurrence of 200 to 700 years for all of southwestern Utah.
- 7. Magnitudes of both the maximum credible and maximum probable earthquake which could occur on the major fault zones located in close proximity to each of the dam sites are tabulated in Table II-1. Also presented in this table are the recurrence intervals for these two earthquake levels. Assuming that they can occur in close proximity to each dam site, it can be seen from this table that the magnitude of the maximum credible earthquake ranges from 7 to $7\frac{1}{2}$ with a recurrence interval of 1,000 to 10,000 years. The magnitude of the maximum probable earthquake for all dam sites is a 6 with a recurrence interval ranging from 200 to 300 years.

Table II-1

Estimated Magnitude and Recurrence Interval of Maximum Credible and Maximum Probable

Earthquakes in the Vicinity of Dam Sites

	Olemant Beatle	Distance	M-4-1 T14	Maximum	Credible	Maximum	
<u>Dam</u>	Closest Fault to Dam Site	Distance to Fault (miles)	Total Fault Length (miles)	Magnitude	Recurrence Interval (yrs)	<u>Magnitude</u>	Recurence Interval (yrs)
Green's Lake No. 2	Hurricane	0	160	7.5	1,000-10,000	6.0	200-300
Green's Lake No. 3	Hurricane	0	160	7.5	1,000-10,000	6.0	200-300
Green's Lake	Hurricane	. 0	160	7.5	1,000-10,000	6.0	200-300
No. 5							,
Gypsum Wash	Washington Hurricane	0 10	40 160	7.0 7.5	2,000-10,000 1,000-10,000	6.0 6.0	200-300 200-300
Warner Draw	Washington Hurricane	1 9	40 160	7.0 7.5	2,000-10,000 1,000-10,000	6.0 6.0	200-300 200-300
Stucki	Washington Hurricane	0.5 10	40 160	7.0 7.5	2,000-10,000 1,000-10,000	6.0 6.0	200-300 200-300
Frog Hollow	Hurricane	2	160	7.5	1,000-10,000	6.0	200-300
Ivin's	Grand Wash	5	100	7.5	2,000-10,000	6.0	200-300

For purposes of this study, the maximum probable earthquake represents an event which has a reasonable probability of occurring during the life of these facilities. Although the magnitude of the earthquake selected for design should be based on the level of risk which the owner of the facility is willing to accept for each specific dam site, it is our judgment, and that of our consultant Dr. Arabasz, that a Magnitude 6 event represents a reasonable earthquake which could occur during the life of these facilities and one which each dam should be able to withstand without catastrophic consequences.

III. OPERATIONAL HISTORY AND PHYSICAL CHARACTERISTICS OF THE DAMS

The eight dams investigated during this study were constructed by the SCS in order to protect downstream areas from flood waters and to trap sediment. The three dams in the Cedar City area were constructed in 1958. The five dams near St. George were constructed during the period of 1974-78 (the existing Frog Hollow Dam represents a raise of an existing dam that was first constructed in 1956).

The intended purpose of the dams is to retain the infrequent large flows of water that result from thundershowers in this arid area, and then to release the water at diminished flow rates over the next few days. Consequently, the reservoirs behind the dams are dry except after heavy rainfalls. Available information indicates that the dams and reservoirs have performed in this manner since construction with one exception. The original trash rack on the outlet of Green's Lake No. 3 became plugged by debris in 1967, which caused water to remain in the reservoirs for 3 months. The standing water resulted in settlement of portions of the reservoir area and embankment (see Figure VIII-2). Table III-1 follows this page and presents information compiled from SCS records on performance of the dams since construction, along with observations on present conditions as noted during the field investigation.

The dams are relatively low earthfill embankments ranging in height from 17 to 60 feet. Information from the SCS files and ESA's current field investigation on the physical characteristics of the embankments and their foundations is presented in Table III-2. A detailed discussion of the nature and condition of each dam is presented in Appendix G.

Table III-1

History of Performance and Present Condition of Dams

Dam	Performance 1	Conditions ²
Green's Lake No. 2	Satisfactory. Some minor slumping of embankment occurred in 1980.	Some differential settlement along axis of embankment, and some cracks which run parallel to the embankment centerline were observed. Cracks were also noted in the foundation in vicinity of embankment.
Green's Lake No. 3	Operational problems caused by subsidence of foundation soils when water remained in reservoir for 3 mo. Extensive subsidence near east end of dam which extended under embankment. Erosion and piping caused cracks to widen. Some block rotation of embankment occurred as a result of subsidence. Cracks repaired by grouting with a soilslurry mixture.	Cracks along upstream face and transverse to embankment crest were observed. Settlement of dam crest also noticeable.
Green's Lake No. 5	Satisfactory.	Some minor cracking observed.
Gypsum Wash	Satisfactory.	Some buldges noted along downstream slope of embankment. No evidence of settlement or excessive cracking along crest of embankment.
Warner Draw	Satisfactory.	Good.
Stucki	Satisfactory	Good.
Frog Hollow	Extensive transverse crack- ing of embankment since 1978 raise. Longitudinal crack- ing along upstream face also noted.	Longitudinal crack along upstream face of embankment observed. Zone of weak soil encountered in foundation in one borehole, accompanied by an anomalous loss of drilling water.
Ivins Diversion No. 5	Satisfactory.	Some minor cracks and erosion of embankment noted.
Notes:		

- (1) Performance of dams based on available SCS reports and correspondence.
- (2) Condition of dams noted during field investigation phase of this investigation.

TABLE III-2 SUMMARY OF EMBANKMENT AND FOUNDATION CHARACTERISTICS

							EMBA	NKMENT MAT	<u>ERIALS</u>	FOUND	ATION MATER	IALS
DAM	DATE OF CONSTRUCTION	TYPE	MAX. HEIGHT (FEET)	LENGTH (FEET)	SLOPE (H:V)	DOWNSTREAM SLOPE (H:V)	CLASSIFICATION	RELATIVE 1 COMPACTION - (%)	SPT BLOW COUNTS (BLOWS/FOOT)	CLASSIFICATION	RELATIVE 2 COMPACTION - (%)	SPT BLOW COUNTS (BLOWS/FOOT)
Green's Lake No. 2	1958	Zoned	20	1315	3:1	2:1·	Zone II/Core ML,CL	81 to 111/ -	25 to 100	Alluvial and Colluvial Deposits Gravelly with Increasing Depth	82 to 96	12 to 82 (Typically 20 to 40)
							Zone I/Shells SM,ML with Gravel and Cobbles	105/ 85 - 90	_			
Green's Lake No. 3	1958	Zoned	17	2030	3:1	2:1	Zone II/Core ML,CL	93 to 102/ -	20 to 52	Alluvial and Colluvial Deposits Gravelly with Increasing Depth	87 to 93	9 to >100 (Typically 15 to 30)
_							Zone I/Shells SM, ML with Gravel and Cobbles	98 to 106/ 79 to 95	-			
Green's Lake No. 5	1958	Homogenous	22	235	3:1	2:1	ML,CL	95 to 107/ 70 to 82	17 to 63	Alluvial and Colluvial Deposits Generally Gravelly	84 to 91	48 to >100
Gypsum Wash	1974- 1975	Zoned	30	3128	3:1	2:1	Zone I/Core SC,SM,ML	95 to 105/ —	23 to >100 (Typically 30 to 40)	Thin Deposits of SC,SM,ML over Siltstone and Shale	76 to 106	25 to >100
							Zone III/Shells GM,SM with Cobbles	95 to 104/ 85 to 88	-			
Warner Draw	1974	Zoned	60	1300	3:1	2:1	Zone I/Core SC,SM	95 to 107/ —	15 to 74 (Typically 40 to 60)	Sandstone, Siltstone and Shale, Left Abutment - SP,SM,SC	96 to 106	Left Abutment— 16 to >100 (Data from SCS Borings)
							Zone III/Shells SM,SC with some Gravel and Cobbles	95 to 101/ 92 to 104	50 to 60			oco borniga,
Stucki	1974	Zoned	30 (Above Ground Surface) 45 (Total Fill)	1400	3:1	2:1	Zone I/Core SC,SM	95 to 103/ —	20 to 84 (Typically 30 to 50)	Alluvial and Colluvial Deposits. Left Abutment - Silty Sandstone	92 to 97	25 to >100 (Typically 30 to 50)
			(Total Fill)				Zone III/Shells SM with Gravel and Cobbles	96 to 105/ 93 to 106	40 to 50			
Frog Hollow	1956 Raised in 1978	Zoned	48	1900	3:1	2:1	Zone I/Core SC,SM, CL,ML with Gravel and Cobbles	94 to 128/ 95 to 100	26 to 48 (Typically 30)	Basalt Flows and Alluvial and Colluvial Deposits Near Toes of Embankment	81 to 89	13 to >100 (Typically 30 to 60)
							Zone II/Shell GM in CL	-/ 79 to 92	-			
							Original (1956) Embankment GM,SM	-	-			
Ivins Diversion No. 5	1977	Homogenous	3 20	5300	3:1	2:1	SM,ML with some Gravel	95 to 101/ 91 to 105	30 to 90	SM,ML Over Siltstone	82 to 95	14 to 80 (Typically 40 to 50)

NOTES:

- 1) First range of relative compactions from construction records as determined by either the Standard Proctor test procedure or ASTM D678-70 method A or C. (See Appendix). Second range of relative compactions from tests performed in test pits excavated as part of this investigation. (See Appendices B, C, and D).
- 2) Ranges of relative compactions from tests performed in test pits excavated as part of this investigation. Test pits were dug at either the upstream or downstream toe of embankment. Tests were performed in the soils present beneath outer shells of the embankments. (See Appendices B,C, and D).

IV. SUMMARY OF REGIONAL GEOLOGY AND TECTONIC FRAMEWORK

This chapter summarizes regional geologic and tectonic conditions in the St. George and Cedar City areas of southwestern Utah as they relate to the performance of the eight dams that are the subject of this report. A detailed discussion of regional geology and tectonics, including an account of the events that led to the development of present-day geologic structure in this area, appears in Appendix E.

Although the details of geologic and tectonic conditions in this area are complex, the controlling factors related to the safety of the dams are: 1) the faults present within the vicinity of the dam sites, 2) the potential for ground rupture on these faults at the dam sites, and 3) the size of earthquake that is likely to occur in the vicinity of the dam sites during their useful life (along with the characteristics of the shaking that will be generated at the base of the embankments).

A brief description of the tectonic framework of the region surrounding the dams follows, along with descriptions of the fault zones that are of concern to the dams. Chapter V presents a discussion of what is known about rates of fault movements in this region, which relates to the potential for ground rupture at the various dam sites. Chapters VI and VII discuss "Historic Seismicity and Earthquake Recurrence", and "Maximum Credible and Maximum Expectable Earthquakes", respectively.

A. Tectonic Framework

The SCS Utah dams study area lies within a transitional intraplate boundary zone between the Basin and Range and Colorado Plateau physiographic/geologic provinces. This boundary zone is coincident with a segment of the Intermountain Seismic Belt, a major zone of seismicity within western North America between the Basin and Range province and the Middle Rocky Mountains-Colorado Plateau (see Appendices E and F and Figures E-1 and F-1).

Regardless of specific models proposed to explain the tectonics of the region, it is clear that there is a major, though gradual, change across the boundary zone

from generally thin weak crust and lithosphere on the west, to thicker, more stable crust and lithosphere on the east. In addition, stress orientation changes across this boundary from extensional WNW-ESE along the eastern basin and range, to compressional WNW-ESE within the Colorado Plateau.

Although the boundary zone between the Basin and Range and Colorado Plateau tectonic provinces is transitional, it is roughly delineated by major tectonic features of the region. The eastern boundary of the northern Basin and Range tectonic province is represented by (from southwest to northeast) the Grand Wash-Cedar Pocket Canyon-Gunlock-Veyo fault system, the north boundary of Pine Valley Mountains, the Hurricane-Parawon-Paragonah/Monocline-fault system, and the Wasatch fault zone (see Figure IV-1).

The three SCS dams in the Cedar City area (the "Cedar City dams") are essentially within the northernmost portion of the Hurricane fault zone (see Figures IV-2 and VII-11). In this area, a complex fault-monocline structure forms a "bridge" between the northernmost segment of the Hurricane fault and the Parawon-Paragonah fault, which is located about 35 km northeast of Cedar City.

The five SCS dams being investigated in this study in the Hurricane-Frog Hollow, St. George, and Ivins areas (the "St. George area dams") are within the northern part of the Grand Canyon subprovince of the Colorado Plateau, which is characterized by NNE-trending, west-down normal faults. The Washington fault, which follows a north-south trend through the vicinity of Gypsum, Warner Draw and Stucki Dams (see Figures IV-1 and VIII-12), is one of these normal faults located about 10 miles west of the Hurricane fault. Descriptions of the major fault zones in the study region follow.

B. Hurricane Fault Zone

The Hurricane fault zone is a major tectonic feature in northwestern Arizona and southwestern Utah. The fault zone has been mapped northward from the vicinity of Peach Springs, Arizona (Wilson and others, 1969), to Kanarraville, Utah, where it turns northeastward and continues to the vicinity of Cedar City, Utah, over a total distance of approximately 256 km (Hintze, 1980). Throughout

much of its length, particularly north of the Grand Canyon, the Hurricane fault zone is marked by a prominent west-facing topographic escarpment called the Hurricane Cliffs.

The Arizona portion of the Hurricane fault zone and that portion in Utah south of Kanarraville, lie within the Grand Canyon subprovince of the Colorado Plateau tectonic province (see Appendix E, Figure E-3). Within this subprovince the Hurricane fault is one of several major west-down normal faults. These faults bound the series of northeastward - tilted fault blocks which form the Grand Staircase transition from the eastern Basin and Range province to the higher-standing Colorado Plateau to the east.

Northeast of Kanarraville, the Hurricane fault zone is coincident with the boundary between the Basin and Range province and the High Plateaus subprovince of the Colorado Plateau Province (at least as far northward as Cedar City). This boundary becomes progressively more complex north of Cedar City, where it involves horsts, grabens, and ramp structures with scissors-like displacement (Best and Hamblin, 1978). The Hurricane fault zone cannot be clearly extended northeast of Cedar City to connect with the Parawon-Paragonah fault (see Appendix E, Figures E-9 and E-10, and Section E 3.2.)

The age of initiation of movement on the Hurricane fault zone is not certain, but its formation may have been related to relative down-dropping of the eastern Basin and Range province as a result of collapse of regional upwarping during late Cenozoic time. Dating of ash-flow tuffs of regional extent suggests that structural differentiation of the provinces began some time after 29 m.y. ago, with vertical movement along faults including the northeasternmost segment of the Hurricane fault that bounded the west side of the Colorado Plateau province.

This faulting began about 26-18 m.y. ago in late Oligocene - early Miocene time (Rowley and others, 1978, 1979; Best and Hamblin, 1978). Anderson and Mehnert (1979) suggest that major movement on the Hurricane fault zone occurred after Miocene time. Total displacement across the fault zone seems to be at least 2,000 m, with current displacements rated on the order of 300-500 m/m.y. (0.03-0.05 cm/yr) (see Appendix E, Section E3.1).

The Hurricane fault zone in Utah is considered to be both seismically active and geologically active in the sense that strands of the fault displace Quaternary geologic units. The fault zone is located within the intermountain seismic belt (see Appendices E and F), a major zone of intraplate seismicity within western North America. Two historic earthquakes thought to be in the range of 5 to $5\frac{1}{2}$ magnitude have occurred on the Hurricane fault zone in the Cedar City area (see Appendix F, Section F.1.2., and Figure F.3), and several earthquake swarms have also occurred in the vicinity of the fault zone in the Cedar City area during historic time (see Appendix F, Figure F.4)

Quaternary movement along the Hurricane fault zone in Utah is well documented (Anderson and Miller, 1979). At several locations along the zone, Pleistocene basalt units and latest Pleistocene alluvial and alluvial fan deposits are displaced.

As part of this study, new, high quality, 1:24,000 scale low-sun angle (both AM and PM), black and white stereoscopic aerial photography was flown in October, 1981 along the mapped location of the Hurricane fault zone in Utah. The photography was analyzed for lineations suggestive of tectonic origin, and possibly related geomorphic and geologic features. The results of this analysis is presented in map form on Figure IV-2. In summary, this map shows that the Hurricane fault zone in Utah is generally represented by a single, usually continuous trace, along the base of the Hurricane cliffs escarpment.

In two areas, south of Ash Creek Reservoir and south of Cedar City, the generally steep and narrow Hurricane cliff escarpment is complicated by large-scale, complex, westward directed landsliding. These landslides have created a wider, less steep escarpment, with a base that bulges out over the projected fault trend. At these locations, the Hurricane Fault zone projects across the landslide as a wide zone of discontinuous traces which offset the landslide masses.

There are several locations where strands of the Hurricane fault zone appear to offset late Pleistocene alluvial fan deposits. Such relations exist just north of the Utah-Arizona border, southwest of the Wart, at several locations between the Hurricane airport and Taylor Creek, northeast of Pintura, and at Shurtz Creek, south of Cedar City.

unde a

The drainage basin of Shurtz Creek between the Hurricane fault zone and the crest of Cedar Mountain to the east is unique among west-draining basins along the Hurricane cliffs. Landslides at higher elevations on Cedar Mountain have shed significant amounts of debris westward. This debris has ponded behind high-standing, elongate NE-SW aligned bedrock ridges just east of the fault zone, creating large areas of upland alluvial fan deposits. Only locally, at Shurtz Creek and an unnamed creek just to the north, is there evidence that this alluvial material spilled through narrow gaps in the bedrock ridges, across the Hurricane fault zone, and out into the valley to the west. Displacements along strands of the Hurricane fault zone probably played a role in the ponding of alluvial deposits upslope to the east as indicated by prominent fault scarps in the alluvial fans at locations where they cross the escarpment. A more subdued and probably older scarp in alluvial fan deposits occurs upslope to the east of the main escarpment.

The 20 m high fault scarp at the mouth of Shurtz Creek is the most prominent and youngest appearing scarp along the entire Hurricane fault zone in Utah. The age of the displaced alluvial materials is pre-Holocene and possibly more than 50,000 years b.p., and the age of displacement is thought to be latest Pleistocene (see Appendix F, Section F.2.1.2.).

Several apparent graben structures were identified along the Hurricane fault zone during this study. Just north of Toquerville, immediately north of the mouth of Shurtz Creek, and in the landslide area south of Cedar City, there are apparent small-scale graben structures within the main Hurricane fault zone. A short distance north of the Utah-Arizona border, lineations and the distribution of alluvial units suggest that a zone of Quaternary faulting extends several miles west from the main Hurricane fault zone, forming an apparent horst and graben structure.

On a larger scale, a zone of complex horst and graben structures at least 5 km wide extends northwest from the main Hurricane fault zone in the area northeast of Anderson Junction. Prominent fault scarps occur in both late Pleistocene alluvium and in Pleistocene basalt flows NNE-trending structures associated with the Virgin Anticline project to intersect the Hurricane trend in this area.

Another large apparent graben structure occurs southwest of Cedar City. Fault traces exposed in trenches excavated during this investigation in this area cut Pleistocene alluvium and overlying basalt, and may cut late Pleistocene colluvium along the east side of the Cross Hollow Hills (see Chapter VIII, and Figure VIII-1). These discontinuous fault traces and associated air photo lineaments were mapped to the southwest along the east side of North Hills to a projected intersection with the main Hurricane fault zone near Murie Creek. This relationship suggests that the alluvial valley(s) southwest of Cedar City between the Hurricane Cliffs and the Cross Hollow and North Hills may be a late Pleistocene graben, here referred to as the Shurtz Creek graben. This structure is similar to structures located northeast of Cedar City along the southern part of the Paragonah fault (see Figure E-11).

C. Washington Fault Zone

The Washington fault is one of the major tectonic features in northwestern Arizona and southwestern Utah (see Figure IV-1). The fault has been mapped northward from just south of Wolf Hole, Arizona to just north of Washington, Utah, a total distance of approximately 68 km (Dobbin, 1939; Cook, 1960; Cook and Hardman, 1967; Hamblin, 1970a,b; Best and Hamblin, 1978; Wilson and others, 1969). Along much of its length in Utah, the Washington fault lies just west of, and parallel to prominent west-facing topographic escarpments and high ground, including the southwest end of Washington Dome, Warner Ridge, and the west side of Beehive Dome.

The Washington fault lies within the Grand Canyon subprovince of the Colorado Plateau Tectonic Province (see Appendix E, Figure E-3). The Washington fault is one of several major west-down normal faults within this subprovince. These faults bound the series of northeastward-tilted fault blocks which form the Grand Staircase transition from the Basin and Range province to the higher-standing Colorado Plateau to the east.

The age of initiation of movement on the Washington fault is not certain, but its formation may have been related to relative down-dropping of the eastern Basin and Range province as a result of collapse of regional upwarping during Late Cenozoic time. Structural differentiation of the provinces began some time after 29 m.y., with vertical movement along faults including the Washington fault.

Displacement on the Washington fault increases southward. Vertical displacement has been estimated to range from several hundred feet (a few hundred meters) near the Virgin River, where the fault breaches the northeast-trending, pre-Quaternary Virgin anticline (Washington Dome, in part), to 2,500 feet (760 m) at the Arizona state line (Dobbin, 1939; Cook and Hardman, 1967).

The Washington fault in Utah is considered to be both seismically active and geologically active (this report), in the sense that strands of the fault displace Quaternary geologic units. The fault is located within the intermountain seismic belt (see Appendices E and F), a major zone of intraplate seismicity within western North America. One historic earthquake thought to be in the range of 4 to $4\frac{1}{2}$ magnitude occurred in 1891 in the St. George area (see Appendix F, Section F.1.2, and Figure F3), and may have been associated with the Washington fault.

Prior to the present investigation, Late Quaternary movement along the Washington fault in Utah had not been documented. Consequently, the Washington fault is not shown on the Quaternary Fault Map of Utah (Anderson and Miller, 1979), and is shown as concealed (dotted) beneath Quaternary deposits on the Geologic Map of Utah (Hintze, 1980). As part of this study, however, detailed field mapping and exploration were conducted in the immediate vicinity of the SCS dams in the St. George area. The results of these investigations are discussed in Chapter VIII C, D and E of this report and shown in Figure VIII-2 (scale 1:24,000). In summary, the results of the detailed field mapping and exploration in the Gypsum Wash-Warner Draw-Stucki area indicate that the Washington fault, which passes between Warner Draw and Stucki Dams and through Gypsum Wash Dam, displays probable early Holocene, normal, east-side-up displacement.

In addition to field mapping and exploration, new, high quality, 1:24,000 scale low-sun angle (both AM and PM), black and white stereoscopic aerial photography

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was flown in October, 1981 along the mapped location of the Washington fault in Utah. The photography was analyzed for lineations suggestive of tectonic origin, and possibly related, geomorphic and geologic features. South of the Stucki dam area, the Washington fault appears on the aerial photography as a prominent, apparently single linear. This linear sharply juxtaposes a bedrock pediment along the base of Warner Ridge against dissected, Quaternary older alluvial fan (Qof on Figure VI-2) bajadas along a subtle west-facing scarp.

From just southeast of Stucki dam northward toward Washington, the fault appears as a number of discontinuous bedrock lineations, tonal lineations in Quaternary alluvium (Qa) and alluvial fan deposits (Qaf), and west-facing scarps juxtaposing Quaternary older alluvial fan deposits (Qof) and/or alluvium (Qa; see Figure VI-1). North of Gypsum dam, the Washington fault is not readily identifiable on the aerial photography.

D. Grand Wash Fault Zone

The Grand Wash fault zone, also designated the Cedar Pocket Canyon-Gunlock fault (Figure IV-1 and Figure E-13), and the Gunlock-Veyo fault (Figure E-9), is a major tectonic feature in northwestern Arizona and southwestern Utah (Cook, 1960; Dobbin, 1939; Anderson and Mehnert, 1979; Best and Hamblin, 1978). The fault zone has been mapped as extending from the vicinity of Grapevine Wash in northwestern Arizona (Wilson and others, 1969), northward to the vicinity of Gunlock, Utah, a total distance of approximately 159 km.

The Grand Wash fault forms the boundary between the Basin and Range and Colorado Plateau provinces in northwestern Arizona. North of the Utah Border, the boundary is less well defined, but is interpreted (Anderson and Mehnert, 1979) to follow the Cedar Pocket Canyon-Gunlock-Veyo fault zone (the northern continuation of the Grand Wash fault). The boundary then apparently steps eastward along the north side of the Pine Valley Mountains to join the Hurricane-Parowan-Paragonah/Monocline-fault system near Cedar City. The structural block between the west-down normal Grand Wash fault zone and the Hurricane fault has been uplifted at least 64 m/m.y. in Late Cenozoic time (Hamblin and others, 1981).

The age of initiation of movement on the Grand Wash fault is not certain, but its formation may have been related to relative down-dropping of the eastern Basin and Range province as a result of collapse of regional upwarping during Late Cenozoic time. Structural differentiation of the provinces began sometime after 29 m.y., with vertical movement along faults including the Grand Wash.

The Grand Wash fault zone in Utah is considered to be seismically and probably geologically active (Anderson and Miller, 1979). The fault zone is located within the intermountain seismic belt (see Appendices E and F), a major zone of intraplate seismicity within western North America.

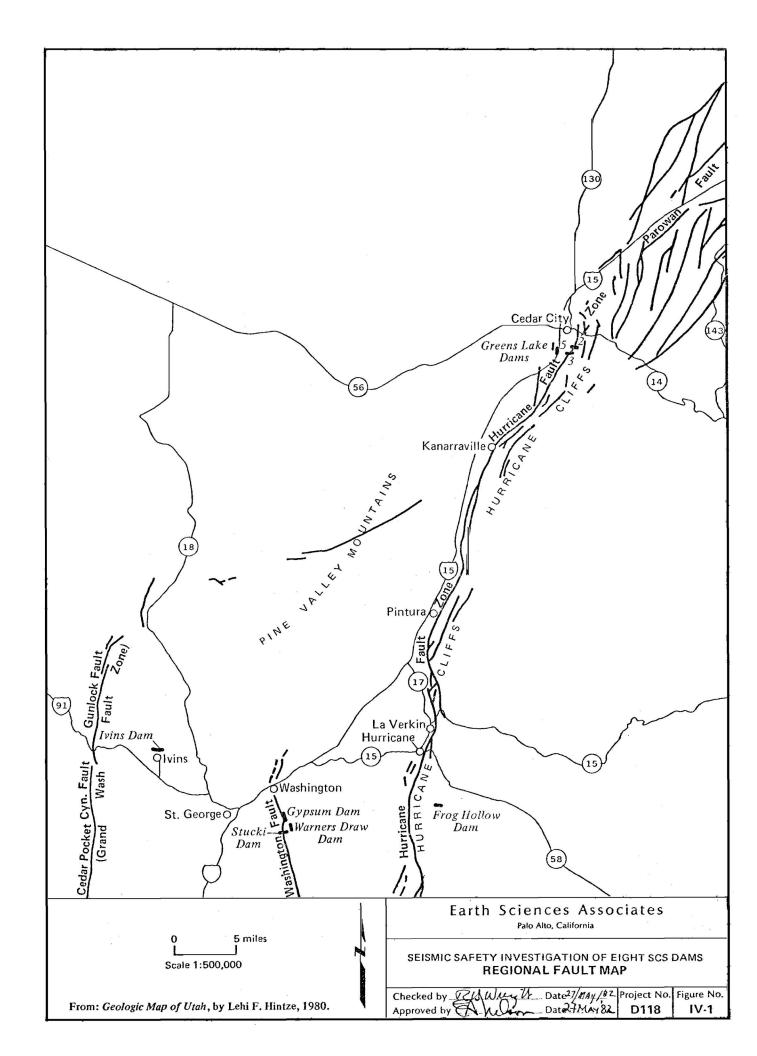
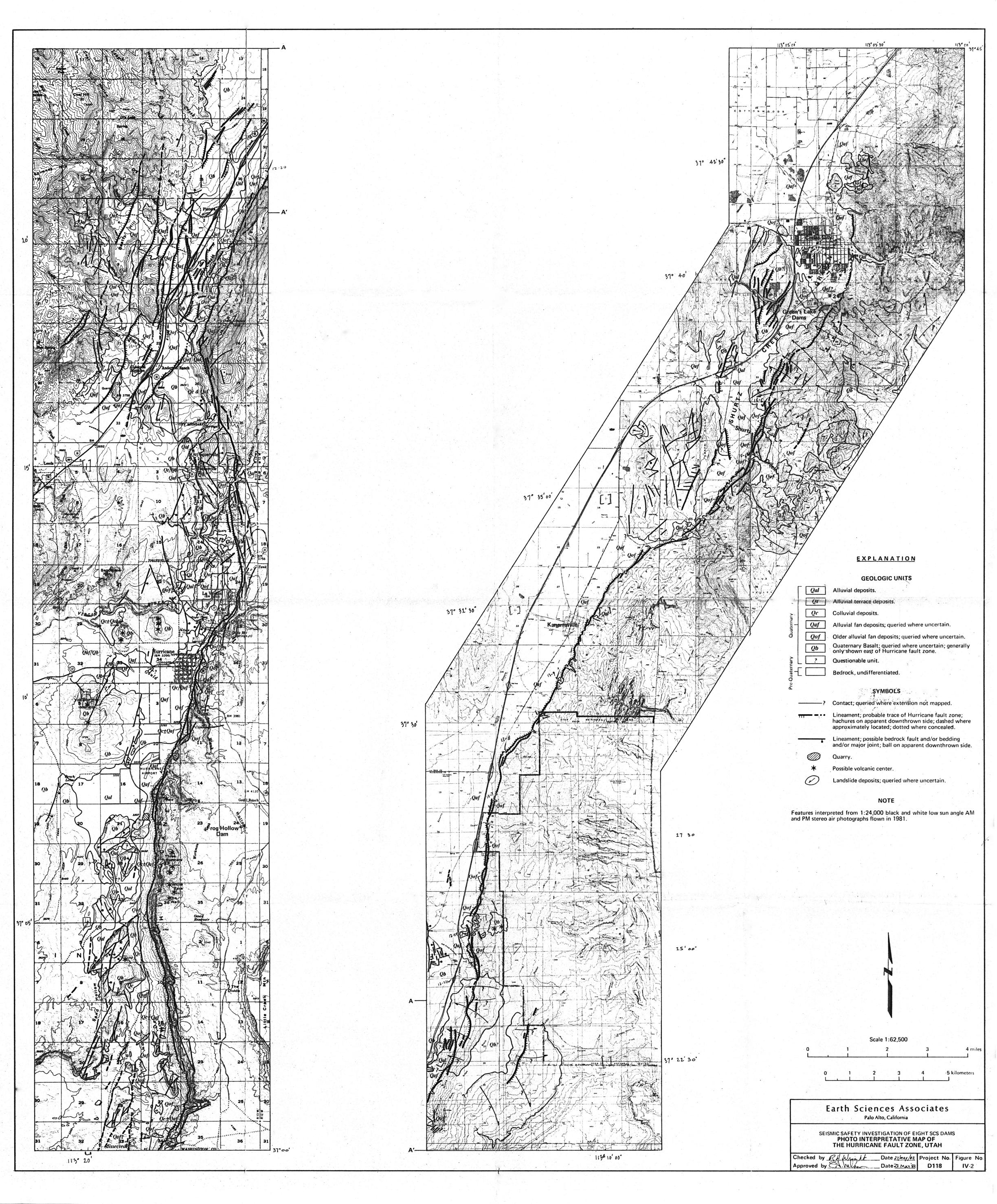


Figure IV-2 PHOTO INTERPRETIVE MAP OF THE HURRICANE FAULT ZONE, UTAH



V. FAULT DISPLACEMENT RATES AND RECURRENCE OF SURFACE FAULTING

A. Summary

Although both the Colorado Plateau and the Basin and Range provinces are presently being uplifted, their rates of uplift are not the same. The Colorado Plateau is rising faster than the area to the west, and this differential movement is being accommodated by displacements on the predominantly north-south faults that mark the boundary zone between these two provinces.

Fault displacement rates for the faults in southwestern Utah indicate average slip rates of fractions of a millimeter to a few millimeters per year (see data compiled by Doser and Smith, 1982). Based on these slip rates, very long recurrence intervals (of the order of thousands of years) for surface rupture events at any one location on these faults are probable.

B. Fault Displacement Rates

Systematic age dating of late Cenozoic basalts displaced by major faults, such as the Hurricane and Grand Wash faults, provides some of the best estimates of fault-displacement rates in southwestern Utah and northwestern Arizona (Hamblin and others, 1981; Anderson, 1980; Anderson and Mehnert, 1979). Hamblin and others (1981) have carefully studied the stratigraphic offset of dated basalt flows in drainage systems along the Basin and Range-western Colorado Plateau border and conclude the following:

"The Grand Wash area in the Basin and Range is rising about 26 m/m.y. and reflects the regional relative uplift rates between the southwestern Colorado Plateau and adjacent Basin and Range province. The block between the Grand Wash and Hurricane faults is rising at an additional minimum rate of 64 m/m.y. (a total of 90 m/m.y.), and the block east of the Hurricane fault is rising at an additional rate of 300 m/m.y. (a total of at least 390 m/m.y.)" (Hamblin and others, 1981, p. 298).

Anderson (1980, p. 535) cites displacement rates for the Hurricane fault of 470 m/m.y. near Pintura, 40 km south-southwest of Cedar City, and 400 m/m.y. in the North Hills area, about 15 km south-southwest of Cedar City.

A summary of fault-displacement data compiled by Doser and Smith (1982) as part of their study of seismic moment rates in southwest Utah supports the interpretation of fault-displacement rates of the order of a millimeter or less per year (1 mm/yr = 1,000 m/m.y.) An anomalously high displacement rate of several millimeters per year at Braffits Creek is now judged to be related to a tensional collapse structure, apparently deforming aseismically (Anderson, 1980). It should be noted that many of the displacement or slip rates listed by Doser and Smith (1982) involve significant uncertainty because of the approximate age estimated for a measured offset. Data for the Shurtz Creek fault scarp are a case in point. Nevertheless, the observations indicate the correct order of magnitude, and they basically agree with the more precise rate estimates of less than a millimeter per year derived from displacements of accurately dated basalt flows.

Elsewhere in the Utah region, the most reliable information on rates of fault displacement have come from recent studies of the Wasatch fault. Fission-track dating of apatites within the uplifted Wasatch Mountain block imply a long-term late Cenozoic slip rate of 0.4 mm/yr, but detailed trenching studies indicate Holocene slip rates of about 1-2 mm/yr and late Pleistocene slip rates as high as 4 mm/yr along certain segments of the Wasatch fault (see Swan and others, 1980).

C. Recurrence Intervals

Extraordinarily long recurrence intervals (of the order of thousands of years) for the repetition of surface faulting at the same site are now well recognized for faults in the Basin and Range province (e.g., Wallace, 1981). The Wasatch fault, which has been perhaps the most active locus of recurrent surface faulting along the eastern Great Basin, displays average recurrence intervals for surface faulting of the order of several hundred years to three thousand years at the same site—as determined from detailed trenching studies across four different fault segments (Swan and others, 1980; Hanson and others, 1981). The average net vertical tectonic displacement for surface faulting earthquakes on the Wasatch fault is about 1-4 m per event (Hanson, 1981), and the recurrence interval for such scarpforming earthquakes is 50-430 years for the entire 370 km-long fault zone (Swan and others, 1980).

Recurrence rates of surface faulting on the Wasatch fault probably represent an upper limit for any active faulting in the Utah region. In western Utah, some prominent Holocene fault scarps have been estimated to reflect recurrence intervals for surface faulting greater than 10,000 years (see Wallace, 1981). In the Cedar City-St. George area, the northern Hurricane fault appears to have the greatest displacement rate associated with tectonic earthquakes, but the rate of recurrence of surface faulting along it is unknown. Arguments of Bucknam and others (1980) and Anderson (1980) suggest a surface faulting recurrence interval greater than 10,000 years for the Hurricane fault. Additionally, the Hurricane fault is apparently moving at a grossly similar slip rate to other faults in southwestern Utah, and recurrence intervals for surface faulting at the same site of at least a few thousand years are indicated for these other faults at all studied locations.

VI. SUMMARY OF HISTORIC SEISMICITY AND EARTHQUAKE RECURRENCE

A comprehensive discussion of the historic seismicity of southwest Utah by Dr. Walter Arabasz, Jr., is presented in Appendix F of this report. The major findings related to historic seismicity and earthquake recurrence are summarized in this chapter.

A. Historic Seismicity

Historic seismicity in the region surrounding the dam sites (as defined below) was documented for the period of record from 1850 to 1981 by reviewing the earthquake data files of the University of Utah.

The University of Utah master catalog for the Utah region comprises three parts (Arabasz and others, 1979, 1980): (1) the 1850 - June 1962 historical catalog; (2) a catalog for the period July 1962 - September 1974, consisting of systematically revised, instrumental earthquake locations and magnitudes; and (3) a catalog of seismicity since October 1974 based upon data from an extensive network of telemetered seismic stations.

The earthquake data for the regions of interest have been considered in several different ways:

- 1) All earthquakes of 3.0 or Magnitude greater within 200 km of the center of the study area (lat. 37° 25′N, long. 113° 15′W) have been listed and plotted. The listing of these earthquakes is provided in Appendix G.
- 2) All earthquakes of 3.0 or greater Magnitude within 150 km radial distance of four specific points, corresponding to the locations of the various dam sites have been listed. The geographic coordinates of the four points are as follows:

Cedar City:	37 ⁰	39.00'N	113 ⁰	04.30'W
Frog Hollow:	37 ⁰	07.00'N	113 ⁰	15.30'W
St. George:	37^{O}	03.30'N	113°	29.00'W
Ivins:	37 ⁰	10.30'N	113 ⁰	40.00'W

Listings for the Cedar City and St. George sites are included in Appendix H.

3. A list of all earthquakes within a rectangular area encompassing St. George and Cedar City (lat. 36.75°N - 38.00°N, long. 112.00°W - 114.25°W), appearing in the University of Utah earthquake catalogue, has been compiled and the epicenters plotted. This earthquake listing is contained in Appendix I. A plot of the instrumental earthquake epicenters for the period from July 1962 - December, 1981, is presented in Figure VI-1. In this illustration, instrumental seismicity is plotted with different symbols to distinguish earthquake epicenters located during the period July 1, 1962 - September 30, 1974 (open circles) from more accurately located epicenters (closed circles) corresponding to the period October 1, 1974 - December 31, 1981.

Twenty-two earthquakes of estimated Richter magnitude 5 or greater have occurred within 200 km of the Cedar City-St. George study area between the period 1850 - 1981. These are tabulated in Table VI-1. Nine of these have had an instrumentally determined magnitude of $5\frac{1}{2}$ or greater. The largest within the 200 km radius region of interest occurred in 1901 near Richfield, Utah (182 km NE of the center of the study area) and had a maximum Modified Mercalli intensity of 8 to 9. This event did not produce any observed surface faulting, and Arabasz and Smith (1979) estimate that it had a local Magnitude (M_L) of $6\frac{1}{2}$ +. A map showing the epicenters of the largest historical earthquakes known from the Utah region during the period of 1850 - 1978 is also presented as Figure F-3 of Appendix F.

The largest historical earthquakes within the immediate vicinity of Cedar City and St. George are the following:

- (1) An earthquake of April 20, 1891 ($I_0 = 6$, $M_L \sim 5$), which shook Washington County and caused minor damage in St. George (Williams and Tapper, 1953).
- (2) An earthquake of November 17, 1902 ($I_0 = 8$, $M_L \sim 6$) near Pine Valley, Utah, about 30 km north of St. George. Many buildings in St. George

Table VI-1. Largest Earthquakes Within 200 km of Study Area,
1850 through 1981

Date (GMT)	Lat. (ON)	Long. (OW)	<u>I</u> o	Magnitude (M _L)	Location	Distance (km)
1887 Dec 05	37.1	112.5	7	$(5\frac{1}{2})$	Kanab, Ut.	77
1891 Apr 20	37.1	113.6	6	(5)	Washington Co., Ut. (St. George)	34
1901 Nov 14	38.8	112.1	8-9	$(6\frac{1}{2}+)$	Richfield, Ut.	182
1902 Nov 17	37.4	113.5	8	(6)	Pine Valley, Ut.	24
1902 Dec 05	37.4	113.5	6	(5)	Pine Valley, Ut.	24
1908 Apr 15	38.4	113.0	6	(5)	Milford, Ut.	110
1910 Jan 10	38.7	112.1	6-7	$(5-5\frac{1}{2})$	Elsinore, Ut.	171
1910 Jan 12	38.7	112.1	6	(5)	Elsinore, Ut.	171
1912 Aug 18	36.5	111.5	6-7	$(5\frac{1}{2})$	Williams, Ariz.	186
1921 Sep 29	38.7	112.1	8	(6)	Elsinore, Ut.	171
1921 Sep 30	38.7	112.1	7	$(5\frac{1}{2})$	Elsinore, Ut.	171
1921 Oct 01	38.7	112.1	8	(6)	Elsinore, Ut.	171
1933 Jan 20	37.8	112.8	6	(5)	Parowan, Ut.	60
1934 Apr 15	38.0	115.0	-	5.0	SE Nevada	167
1942 Aug 30	37.7	113.1	6	(5)	Cedar City, Ut.	34
1942 Sep 26	37.7	113.1	6	(5)	Cedar City, Ut.	34
1945 Nov 18	38.8	112.0	6	(5)	Glenwood, Ut.	186
1952 May 24	36.1	114.7	6	5.0	ArizNev. border	195
1959 Feb 27	38.0	112.5	6	(5)	Panguitch, Ut.	93
1959 Jul 21	37.0	112.5	6	$5\frac{1}{2}$ -5-3/4	Kanab, Ut.	81
1966 Aug 16	37.5	114.2	6	5.6	NevUt. border	80
1967 Oct 04	38.5	112.2	7	5.2	Marysvale, Ut.	158

 $^{^1}Summarized$ from compilation of regional seismicity (M \geq 3.0) within 200-km radial distance of study area. Table includes earthquakes of estimated Richter magnitude 5 or greater. I = maximum Modified Mercalli intensity. Magnitudes in parentheses estimated from I $_{0}$ (M=1+2/3 I $_{0}$). Distance measured from a point half-way between Cedar City and St. George.

sustained damage and almost every chimney reportedly went down in Santa Clara (Williams and Tapper, 1953).

- (3) An aftershock of the 1902 Pine Valley earthquake that occurred on December 5, 1902 ($I_0 = 6$, $M_L \sim 5$).
- (4) An earthquake of January 20, 1933 ($I_0 = 6$, $M_L \sim 5$) near Parowan, 30 km northeast of Cedar City. Minor damage occurred in Parowan and Minersville (Williams and Tapper, 1953).
- (5) Two earthquakes, apparently as part of a swarm sequence, near Cedar City on August 30, 1942, and September 26, 1942 (each with $I_0 = 6$, $M_L \sim 5$). Minor damage occurred in Cedar City during both of these shocks (Williams and Tapper, 1953).

B. Earthquake Recurrence

Rates of earthquake recurrence in any seismically active region can be estimated using both seismological and geological data. If a catalog of historical and/or instrumental seismicity is available, then relationships of earthquake frequency versus magnitude can be used to estimate the average recurrence interval or inter-event time for earthquakes of a specified magnitude. Recurrence estimates can also be made from seismic moment rates determined from available geologic data on Quaternary faulting.

Each of these approaches has been recently applied to seismically active areas in Utah, and available information on earthquake recurrence in southwestern Utah is summarized here. A detailed discussion of this subject is presented in Section F.3 of Appendix F.

Estimates of the frequency of occurrence for a range of earthquake sizes were made for the southwestern Utah region, which encompasses an area of 56,160 km². In making these estimates, several relationships derived from the historical and instrumental seismicity were used (Smith and Arabasz, 1979). The three different cases considered are listed at the bottom of Table VI-2, which also

Table VI-2 - Estimated Frequency of Occurrence of Damaging Earthquakes (in years)

Earthquake	Recurren	ice for Entir	e SW Utah	Recurrence Within 50 km of any Site			
Size	Case I	Case II	Case III	<u>Case I</u>	Case II	Case III	
$M_{T} \ge 7.5$	313	241	660	2,240	1,720	4,720	
Б							
$M_{L} \ge 7.0$	148	125	282	1,060	894	2,020	
1.2							
$M_{L} \ge 6.5$	70	65	120	500	465	858	
D							
$M_{L} \ge 6.0$	33	34	51	236	243	365	
r-							
$M_{L} \ge 5.5$	16	18	22	114	129	157	
Г			_				
$M_{L} \ge 5.0$	7.4	9.5	9.3	53	68	66	
T	••-	2.5	•••				

Case II: Historical data, 1850-1978:
$$\log N_{e} = 2.38 - 0.65 \text{ M}_{L}$$

Case II: Historical data, 1850-1978: $R = 1/[-F(x)]$

where

$$F(x) = \exp\{-[-\exp[-(x-3.50)/1.14]\}, -\infty < x < + \infty$$

Case III: Instrumental data, 1962-1978: $\log N_{e} = 2.73 - 0.74 \text{ M}_{L}$

 $M_L = Local Richter magnitude.$

 N_c = Annual number of events equal to or greater than a given magnitude.

F(x) =Probability that the largest earthquake in a year will have intensity less than or equal to x.

R = Return period.

presents rates of recurrence for a range of earthquake sizes. These values were determined assuming that the seismicity of the southwest Utah study area is uniformly distributed over the area. This is a reasonable assumption for earthquakes up to Magnitudes 6 to $6\frac{1}{2}$.

A fundamental assumption regarding the estimation of earthquake recurrence is that the data used to calculate the required parameters accurately represent the long-term seismicity of a region. Ideally, the data should represent a period of time long enough to include at least one repeat interval of the largest earthquake. Consequently, while rates of earthquake recurrence for magnitudes up to about 6- $6\frac{1}{2}$ in southwest Utah can be estimated with a high degree of confidence on the basis of historical and instrumental seismicity, there is clearly considerable uncertainty in extrapolating recurrence intervals for earthquakes of larger size.

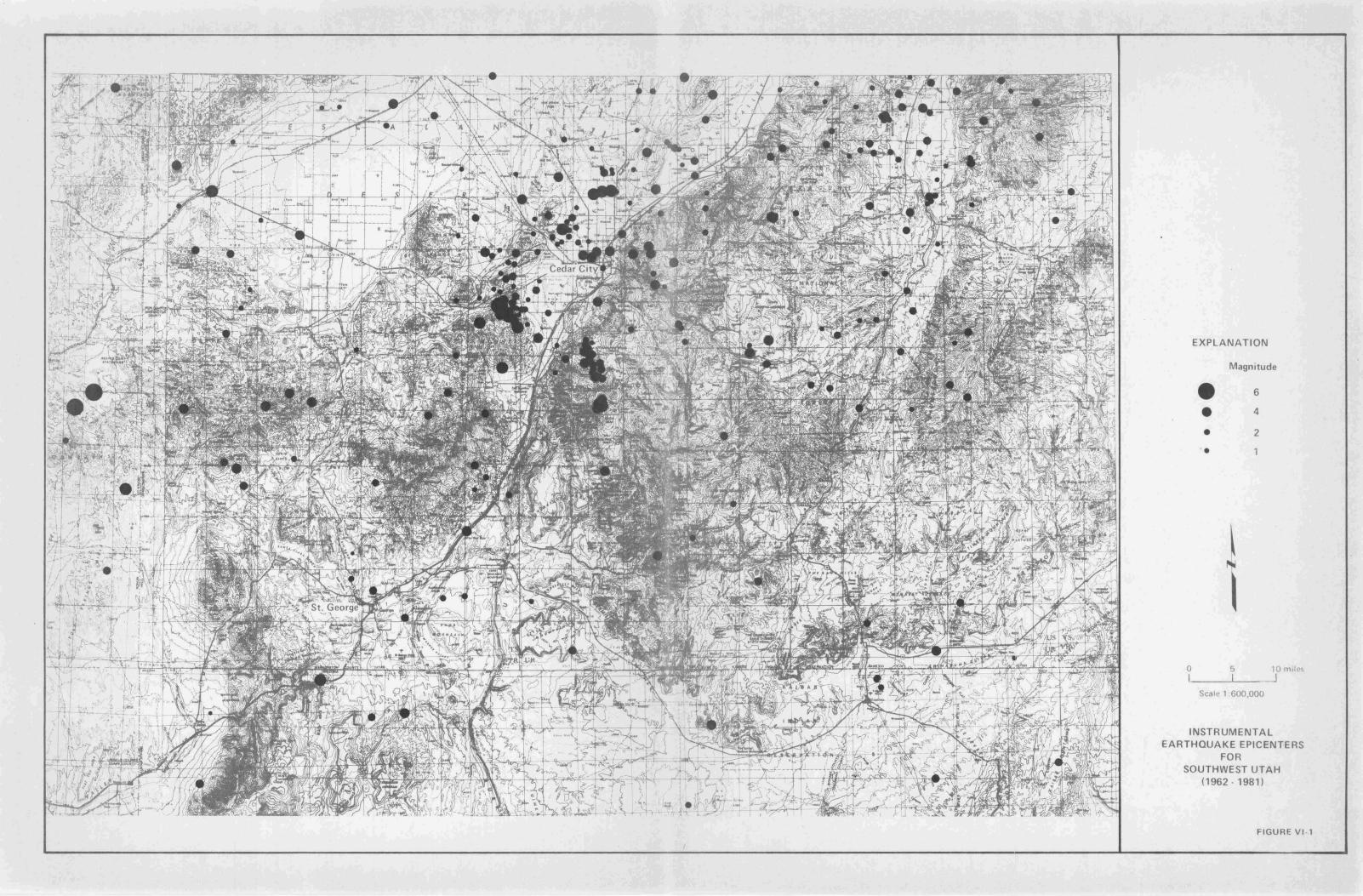
Since the seismicity of the region has been assumed to be uniformly distributed, it is possible to obtain an estimate of the frequency of occurrence of earthquakes within a smaller area. If we consider an area within 50 km of any of the eight dam sites (which amounts to 7,854 km²), the recurrence interval within this smaller area is inversely proportional to the ratio of the two areas. Thus, as can be seen from Table VI-1, the recurrence interval for an earthquake of any given size occurring within 50 km of any of the eight dam sites is about 7 times longer than the recurrence interval within the entire southwest Utah region.

The frequency of occurrence of moderate-to-large earthquakes can also be estimated from geologic data by relating geologically determined slip rates on individual faults to seismic moment rates (Anderson, 1979; Molnar, 1979). Such an approach complements analyses based on the historic record of seismicity, which is generally too short for valid evaluation of the long-term seismic activity over hundreds or thousands of years. Doser and Smith (1982; see also Doser, 1980) have applied the moment rate method to various regions in Utah, including the southwest Utah region (Region III, Figure F-12a) defined by Smith and Arabasz (1979) for recurrence analysis of seismicity. They first determined a moment-magnitude scale for Utah based on the spectral analysis of 19 earthquakes in the Utah region in the magnitude (M_L) range 3.7 to 6.6. Moment rates were calculated from a compilation of geologic information on slip rates of Quaternary faulting

in southwest Utah. Earthquake recurrence rates were then calculated using the slopes of the relationships established by Smith and Arabasz (1979) and an appropriate coefficient defining the moment-magnitude relationship.

For southwest Utah, the frequency of earthquake occurrence based on geologically determined moment rates is essentially in agreement with that calculated from historical earthquake data. The estimated recurrence interval for an earthquake of $7.0 \le M_{T} \le 7.5$ somewhere in the southwest Utah region is between 200 and 600 years (Doser and Smith, 1982), where the range results from an uncertainty in the moment-magnitude scale for Utah. For the same magnitude range and area, extrapolated seismicity (Smith and Arabasz, 1979) would predict a recurrence interval between 200 and 500 years. The discrepancy between such an expectation of a scarp-forming earthquake somewhere in southwestern Utah every few to several hundred years and the apparent absence of Holocene fault scarps in southwest Utah may in large part be the result of the extremely active erosional/depositional environment. Anderson (1980) and Bucknam and others (1980) believe that there has been no radical change in the rate of occurrence of large earthquakes during Holocene time compared to the late Quaternary record in southwest Utah. Two possible implications of the lack of evidence of surface faulting are: (1) the recurrence interval of surface faulting on individual faults in southwest Utah may be very long, and (2) there may be different frequencymagnitude relations for earthquakes smaller than and larger than, respectively, about magnitude 6-6½ in southwest Utah.

Based on consideration of the historical seismicity and the geology of the region, it is our opinion that the recurrence interval for an earthquake of Magnitude 7 to $7\frac{1}{2}$, in close proximity to any of the dam sites, would range from 1,000 to 10,000 years. For earthquakes of Magnitude 6 to $6\frac{1}{2}$ or smaller, the values given in Table VI-2 for an area within a 50 km radius of any of the dam sites are appropriate.



VII. MAGNITUDE OF MAXIMUM CREDIBLE AND MAXIMUM PROBABLE EARTHQUAKES

The largest historical earthquake in southwest Utah (which produced no observed surface faulting) had an estimated magnitude of $6\frac{1}{2} \leq M_L \leq 7$. The largest historical earthquake in the entire Intermountain seismic belt had a measured magnitude (M) of 7.1--resulting in a normal fault scarp with a maximum surface displacement of 6.7 m.

The selection of a maximum size earthquake for any area conventionally is based on consideration of: the structural and geological features of a region; the historical experience of faulting within the same (or a similar) tectonic province; and empirical relations between observations of fault-rupture length, fault displacement, and earthquake magnitude (e.g., Slemmons, 1977).

The maximum size earthquake for the entire Utah region is generally considered as a "Magnitude $7\frac{1}{2}$ ". The U.S. Geological Survey (1976) postulated the occurrence of earthquakes as large as magnitude (unspecified) 7.5 on the Wasatch fault. In the latter study, mention is made of the magnitude (M) 7.1 Hebgen Lake, Montana, earthquake of 1959 and the magnitude (M) 7.6 Pleasant Valley, Nevada, earthquake of 1915 as plausible upper limits for the "maximum credible" earthquake on the Wasatch fault. Here M refers to a "revised magnitude" determined by the Seismological Laboratory in Pasadena that is distinct from, but essentially equivalent to, a surface-wave magnitude $M_{\rm S}$ (see Geller and Kanamori, 1977, regarding differences between various magnitudes computed and published by Gutenberg and Richter).

Swan and others (1980) relate measurements of pre-historic fault displacements on the Wasatch fault to Slemmons' (1977) curve for normal-slip faults to assert that "surface faulting events associated with earthquakes in the magnitude range $6\frac{1}{2}$ to $7\frac{1}{2}$ have occurred repeatedly along (segments of the Wasatch fault)" (Swan and others, 1980, p. 1458).

Bucknam and others (1980) review data on historic surface faulting in the Basin and Range province. Their tabulation shows that, excluding the M_L = 8.0 Owens Valley, California, earthquake of 1872 (max. displacement = 6.4 m), the

largest event (coincident with the largest maximum displacement of 5.6 m) is the $\rm M_L$ = 7.8 Pleasant Valley, Nevada, earthquake of 1915. Bucknam and others (1980) also note that all seven historic earthquakes in the Great Basin larger than $\rm M_L$ = 6.3 have produced surface ruptures. Without clear explanation, they assume a maximum magnitude (unspecified) of 7.6 for the Wasatch fault zone, and they discuss the likelihood of infrequent large earthquakes of "M = 7.0 to 7.6" in southwest Utah (Bucknam and others, 1980, p. 306-308). M in their usage is simply "magnitude". The value $\rm M_L$ = 7.8 for the Pleasant Valley earthquake is ascribed by Bucknam and others (1980) to a NOAA data file summary, in which some conversion was apparently made from the published value of M = 7.6 for that earthquake (e.g., Richter, 1958).

In the Cedar City-St. George area, Anderson (1980) states that "the 20 m high scarp in alluvium along the trace of the Hurricane fault at Shurtz Creek clearly implies a displacement history comparable to other faults in the Great Basin for which a credible magnitude of $7\frac{1}{2}$ would readily be assigned". Anderson also sees no evidence that the Shurtz Creek scarp is a composite scarp (although evidence for multiple movements admittedly would be difficult to recognize in the coarse bouldery alluvium), and he believes that it may represent a single displacement event (oral communication to geologists participating in a field inspection of SCS dam sites in the Cedar City-St. George area, January 1982).

ESA considers it likely that the 20 M high Shurtz Creek scarp is the result of multiple movements. A 20 M high single displacement scarp on the Hurricane fault is not a credible event within the context of the historic record. The largest historic vertical displacement on a normal fault is 14.34 M on the Yakutat Bay fault in Alaska in 1899 (Bonilla, 1970). This displacement was the result of a 8.5-8.6 M earthquake, a magnitude significantly greater than that assigned to the Hurricane fault. Evidence of multiple events in the Shurtz Creek environment would not be expected to be well preserved. Robert E. Wallace of the U.S. Geological Survey estimates that a one meter scarp in the Shurtz Creek environment would not last more than 500-1,000 years, and a two meter scarp not more than two to three thousand years (R. E. Wallace, USGS, oral communication, 5/82).

A maximum earthquake size of Magnitude $7\frac{1}{2}$ is a reasonably conservative value to assign to the study area. The precise number--within differences of

a few tenths of a magnitude unit--is insignificant, in our opinion, in view of: (1) inconsistent specification of magnitude scales, (2) general problems with magnitude, in general, as a reliable measure of earthquake source properties, and (3) the range of uncertainty involved in relating magnitude to a predicted value of displacement, rupture length, or any ground-motion parameter. Adoption of a maximum earthquake magnitude of $7\frac{1}{2}$, which in practice will likely be measured as a surface-wave magnitude will be a larger magnitude than that for any historic earthquake which has occurred in the Intermountain Seismic Belt (ISB). This value will be consistent with the general consensus of geologists and seismologists familiar with earthquake risk in the ISB.

On the basis of the historical record of seismicity, the pattern and nature of known faults, the anticipated distribution of known faulting and the inferred temporal relationships of regional strain, an earthquake of magnitude 7.5 is concluded to be a conservative maximum credible event for the southwest Utah region.

Magnitudes of both the maximum credible and the maximum probable earthquakes which could occur on the nearby faults and which could produce significant ground shaking at the various dam sites have been estimated and are summarized in Table VII-1, together with estimates of their recurrence intervals. For purposes of this study, the maximum probable earthquake represents an event which has a reasonable probability of occurring during the life of these facilities. Although the magnitude of the earthquake selected for design should be based on the level of risk which the owner of the facility is willing to accept for each specific dam site, it is our judgment, and that of our consultant Dr. Arabasz, that a magnitude 6 event represents a reasonable earthquake which could occur during the life of these facilities and one which each dam should be able to withstand without catastrophic consequences.

Table VII-1

Estimated Magnitude and Recurrence Interval of Maximum Credible and Maximum Probable

Earthquakes in the Vicinity of Dam Sites

	Closest Fault	Distance to	Total Fault	Maximum Credible		Maximum Probable	
Dam	to Dam Site	Fault (miles)	Length (miles)	<u>Magnitude</u>	Recurrence Interval (yrs)	Magnitude	Recurence Interval (yrs)
Green's Lake No. 2	Hurricane	0	160	7.5	1,000-10,000	6.0	200-300
Green's Lake No. 3	Hurricane	. 0	160	7.5	1,000-10,000	6.0	200-300
Green's Lake	Hurricane	0	160	7.5	1,000-10,000	6.0	200-300
No. 5							
Gypsum Wash	Washington Hurricane	0 10	40 160	7.0 7.5	2,000-10,000 1,000-10,000	6.0 6.0	200-300 200-300
Warner Draw	Washington Hurricane	1 9	40 160	7.0 7.5	2,000-10,000 1,000-10,000	6.0 6.0	200-300 200-300
Stucki	Washington Hurricane	0.5 10	40 160	7.0 7.5	2,000-10,000 1,000-10,000	6.0 6.0	200-300 200-300
Frog Hollow	Hurricane	2	160	7.5	1,000-10,000	6.0	200-300
Ivin's	Grand Wash	5	. 99	7.5	2,000-10,000	6.0	200-300

VIII. GEOLOGIC CONDITIONS IN THE VICINITY OF THE DAM SITES

The investigation of the geological environment of each specific dam site began with photo-geologic evaluation and identification of lineaments which could represent active fault traces, followed by field reconnaissance mapping, with emphasis on the nature of the air photo lineaments and certain geologic features mapped by previous investigators. Sites determined by the field reconnaissance to be located on or near suspected active faults were investigated by exploratory trenching. A discussion of geologic conditions determined to exist at each of the dam sites is presented in the order listed below:

- A. Green's Lake No. 2 and No. 3 (GL-2 and GL-3)
- B. Green's Lake No. 5 (GL-5)
- C. Gypsum Wash
- D. Warner Draw
- E. Stucki
- F. Frog Hollow
- G. Ivins Diversion No. 5

Of the eight dams investigated, five are located on faults, or within fault zones, and these sites are discussed first.

A. Green's Lake No. 2 and No. 3

1. General Geology

Green's Lake Dam Embankments No. 2 and No. 3 (GL-2 and GL-3) are located at the toe of a major N-S-trending escarpment which rises eastward to the Kolob Terrace of the Colorado Plateau province. The escarpment or hillfront is considered to be the locus of the main trend of the Hurricane fault. South of Cedar City, the escarpment is characterized by moderate to very steep slopes which are underlain by complex landslides of varying ages, and by areas of tectonically deformed bedrock. Figure VIII-10 defines bedrock and landslide areas within the escarpment, and shows a prominent strand of the Hurricane fault which cuts both bedrock and landslide zones south of GL-3. This strand apparently extends beneath young alluvial fan deposits to the north.

The hillfront escarpment in this area has been dissected by numerous drainage channels. Periodic torrential runoff has deposited coarse rubble and boulders of sandstone, limestone, and basalt up to five or more feet in size in these ravines. The coarse material deposited in the re-entrants grades finer and less chaotic to the west across the front of the escarpment, grading to well stratified alluvial fan deposits which form an apron along the hillfront and underlie the GL-2 and 3 dam site areas. The fan deposits consist of silt, sand, and gravel. Rock types present in the gravel clasts include limestone, siltstone, sandstone, and gypsiferous shale derived from bedrock sources to the east. Gravel clasts are typically angular to flat and up to 2 or 3 inches in size, although cobbles and boulders are common within the fans near the toe of the hillfront to the east. The alluvial fan sediments are commonly medium dense to dense silt and sand mixtures with lesser coarser lenses of sand and gravel. As determined by the trench excavations, the fan deposits will stand vertically without support for an indefinite time. Site geology in the vicinity of Green's Lake No. 2 and No. 3 is shown in Figures VIII-1 and VIII-2, respectively.

2. Faults

Previous geologic reconnaissance mapping of the Green's Lake dam sites by R. L. Bridges and D. H. Griswold (data compiled from SCS files) during the late 1950's indicated fault-like features trending generally south-southwest to north-northeast near the toe of the hills and underlying portions of the embankments. The locations of these suspected faults, as interpreted from the available SCS literature, are plotted in Figure VIII-3 (see also Figures VIII-1 and VIII-2), and are assigned letters A, B, C, D and E.

Air photo studies during the current investigation confirmed the existence of several lineaments at or near the previously mapped faults on the alluvial fan deposits. However, subsequent field reconnaissance revealed extensive man-made alterations to the original ground surface both within and outside of the main dam embankments. These modifications include cutting and filling operations that were apparently carried out during dam construction or subsequent repair operations associated with basin maintenance and spreader dike construction. Because surficial geologic evidence for faulting or fault-like features on the fan deposits were not observed at many of the locations previously mapped by Bridges and/or Griswold, it is assumed much of the surficial reworking occurred after their

studies. Most of the air photo lineaments currently identified on the alluvial fans apparently result from anomalous surficial features not related to subsurface geologic conditions.

Earth Sciences Associates' current field reconnaissance located two prominent scarp-like features at the approximate fault locations previously mapped by Bridges and/or Griswold. These scarps also coincide with lineaments identified on air photos, and are shown in Figure VIII-3 as Scarps I and II. Scarp I trends N35E and is located at the west edge of the present tree cover. Several air photo lineaments are located in this same area and it is near previously mapped fault traces "A" and "B". The relief of scarp I is approximately 5 to 6 feet, and the surface dips at about 250 to the west. Trench GL-2a was excavated across Scarp I and encountered locally stratified alluvial fan and chaotic boulder fanglomerate deposits, indicating the extreme agradation-degradation processes which have occurred here. Trench exposures revealed no indication of displaced or faulted stratigraphy at the scarp location, but a sharp discontinuity exists to the west of the topographic scarp at Station 0+98. At this station, carbonate-cemented gravels and a silt lens are juxtaposed across a vertical contact with loose, coarser gravels (see log of Trench GL-2a and Photograph VIII-1). This discontinuity is overlain by 6 to 7 feet of younger, unbroken fan deposits. Trench GL-2d was excavated 40 feet north of GL-2a across the N5E-striking trend of the discontinuity in order to determine the extent of the feature, but no similar discontinuities or other disruptions of stratigraphy were exposed. To the north, Trench GL-2c extends across the projection of the break found in Trench GL-2a, but no disruptions of the stratified, fine-to-coarse-grained alluvial deposits were observed in this exposure. Although it is possible the discontinuity at Station 0+98 in Trench GL-2a resulted from faulting, in view of the additional unbroken trench exposures along the projection, we believe that the feature most likely represents a buried stream channel margin.

North-northeast of the GL-2 dam site and parallel to the GL-1 dam embankment, Scarp II (Figure VIII-3) trends approximately N20E. It is generally irregularly defined, but typically has 2 to 4 feet of relief with the west side down. Trench GL-2b intersected Scarp II (which conceivably coincides with fault "A" or "C" as mapped by Bridges), and exposed unbroken fine-grained alluvial fan deposits.

Trench GL-2c was excavated across the southern projection of Scarp II, across the previously-mapped fault trace "A", as well as across the northward projection of the break found in Trench GL-2a. As noted above, no evidence of faulting was detected in the exposed stratified alluvial fan deposits. Scarp-like features I and II are apparently either the result of earlier excavations or natural erosional features, or are a combination of both.

Surficial features similar to Scarps I and II (which projected through GL-2) were not observed during reconnaissance mapping near GL-3. The prominent air photo lineament located in the hills immediately south of Site No. 3 (Figure VIII-10), which is likely a major strand of the Hurricane fault, and previously-mapped suspected fault "D" intersect the embankment, however. Trench GL-3a was positioned on a zone of ground apparently undisturbed by construction activity across the northward projection of the prominent air photo lineament and suspected fault "E". Trench GL-3c intersects mapped fault D, which trends N20E from approximately Station 11+00 of the GL-3 dam embankment (as plotted from SCS data). Both trenches exposed well stratified, fine-grained alluvial fan deposits which displayed no evidence of fault disturbance.

Trench GL-3b was located on ground where minor surficial alteration had occurred during the past, but extended across the prominent Hurricane fault lineament where it traverses directly along the toe of the hillfront immediately west of GL-3's right abutment. Extensively sheared and faulted bedrock was exposed at this location, although the rocks are clearly overlain by an undisturbed, younger sequence of fine-to-coarse-grained alluvial fan deposits (see trench log).

Materials in the SCS files show that an area within the GL-3 reservoir subsided substantially when water remained in the reservoir because of a clogged outflow trash rack during 1967. The area of subsidence extended beneath a portion of the embankment, and is shown in Figure VIII-2. Because GL-3 dam is located within the Hurricane Fault Zone, it was suggested that the subsidence was in some way related to faulting.

Other than the settlement of this zone, no additional evidence was found during the current investigation to suggest Holocene faulting at this site. Nearby zones where suspicious fault-like features were identified and subsequently

trenched revealed no indication of fault disturbance or subsidence deformation in the alluvial deposits (Trenches GL-3a, 3b, 3c). Furthermore, it is doubtful that the subsidence occurred as a result of differential settlement of underlying weak, sheared bedrock materials within a fault zone. Such deformation would likely have been present in Trench GL-3b where sheared bedrock conditions were exposed, but were found to be overlain by flat-lying, undeformed alluvial sediments. Both the intact bedrock materials and the weak, sheared bedrock zones are considerably harder and denser than the overlying alluvial deposits exposed in the site area.

In summary, all known evidence suggests that the subsidence which occurred here is the result of solution of gypsiferous materials within the alluvial fan materials, or collapse of porous mud flow deposits with associated settlement of near-surface soil deposits.

Because of the proximity of the Hurricane fault zone and the associated lineament or lineaments which project into the alluvial fan deposits from the flanking hills, the positions of the GL-2 and GL-3 trenches (and GL-3 subsidence zone) may very well overlie a strand or strands of the Hurricane fault. However, if such fault traces exist, they are buried by sequences of younger, unfaulted alluvial fan deposits which are at least as thick as the depth of the exploratory trenches. Figure VIII-5 is a sketch of inferred geologic conditions in the Green's Lake dams vicinity, and illustrates suspected fault-alluvial fan relationships.

A determination of the age of the undisturbed fan deposits would indicate the minimum amount of time an underlying fault, if present, has been inactive. Although attempts to assign ages to the alluvial sediments were hampered because of the general lack of soil development, the exposed sediments overlying the discontinuity in Trench GL-2a contained a weakly developed soil with cambic and locally weak argillic horizons. Underlying these deposits (adjacent to the discontinuity of probable erosional origin) near the base of the exposure, an incipient buried paleosol is present within channel fill material. Estimated ages of these two horizons are approximately 3,500 to 5,000 years and 5,000 to 10,000 years B.P. (before present), respectively (Dr. R. J. Shlemon, written communication), indicating that no fault activity has occurred at this location for at least 5,000 years. Detrital charcoal samples collected from specific stratigraphic horizons within trenches GL-2d, GL-3a, and GL-3b have been submitted for Carbon 14 age-dating, but results are not expected until early June, 1982.

B. Green's Lake No. 5

1. General Geology

The Green's Lake No. 5 site (GL-5) is a flood basin formed by the main retarding dam at Cross Hollow, the north and south dikes, and the natural topography. Quaternary-age basalt forms the west, south and southeast margins of the basin, and underlies the relatively high Cross Hollow Hills area west of the site. In the GL-5 vicinity, the basalt is underlain by older alluvial deposits which comprise the GL-5 basin floor, and, as determined by exploratory Trench GL-5a, the right abutment of the north dike and the northeast margin of the basin. These deposits were penetrated by exploratory borings GL-5-1 and 2 at the main dam during the current study, and are exposed beneath the basalt approximately two miles southwest of the GL-5 site in road cuts.

The basalt is characterized by rough, prominent outcrops where steep scarps exist, notably along the Cross Hollow drainage west of the main dam, and along the northeast-southwest-trending hillfront adjacent to the GL-5 basin. Where exposed, the basalt is very hard and prominently fractured, typically forming crude blocks several feet in size. At the top of the east-southeast-facing hillscarp which flanks the main dam site, the exposed basalt exhibits a gentle, east-dipping downward warp, conceivably the result of either gravatational slump-type movement along this upper margin, or a tectonically induced flexure which may be associated with adjacent faulting.

The older, underlying alluvial deposits consist of dense, stratified sandy silt, sand, and gravel lenses which consist of sandstone, limestone, siltstone, shale and basalt clasts. As determined in Trench GL-5a, most gravel clasts are angular or flat in shape, with re-worked or rounded clasts comprising an estimated 5 to 10% of the volume. The alluvium ranges from medium dense to very dense, and stands vertically in unsupported excavations.

At the base of the hillslopes and below the more prominent scarps on the east side of the Cross Hollow Hills, talus and colluvial deposits have accumulated in poorly defined zones which extend outward over the alluvium. The colluvial

deposits exposed in Trenches GL-5b and GL-5c consist of silt to boulder-sized material. They are commonly a gravelly silt containing erratic basalt cobbles and blocks with a prominent, overlying hardpan zone of leached calcium carbonate. Where uncemented by the carbonate, the colluvium is typically loose to medium dense. Site geology in the vicinity of Green's Lake No. 5 is shown in Figure VIII-4.

2. Faults

Figure VIII-10 illustrates the locations of the GL-5 embankments, the prominent topographic scarps which trend northeast-southwest and form the margins of the GL-5 basin, and the numerous air photo lineaments present in the Cross Hollow Hills to the west. Approximately two miles southwest of the site and along the projected trend of the scarps and lineaments, road cut excavations expose both the basalt and underlying older alluvium. At this location, both preand post-basalt age shear planes are present as illustrated by photographs VIII-2 and VIII-3 (both north-facing views). Photograph VIII-2 defines a steeply east-dipping shear within the older alluvium which is clearly truncated by the overlying basalt. Photograph VIII-3 shows another east-dipping shear which offsets both the alluvium and younger overlying basalt.

At the northeast margin of the GL-5 basin, exploratory Trench GL-5a is located along a northeast-southwest scarp where basalt outcrops are absent. The trench exposures indicate this scarp is underlain by apparent down-to-the-west displacements within alluvium which is similar in character to the alluvium underlying basalt at the road cut to the southwest. Several high angle and apparent dip-slip offsets which displace well stratified sand and gravel and a well developed overlying carbonate horizon are present. The age of the latest offsetting movement (based on the development of carbonate within the trench exposures) is estimated to be latest Pleistocene, or at least 10,000 B.P. (before present) (Dr. R. J. Shlemon). Detrital charcoal sampled from strata cut by one of the high angle shears (see log of Trench GL-5a) has been submitted for Carbon 14 dating, and results from this analysis are expected in June, 1982.

The prominent, southeast-facing scarp underlain by basalt at the west margin of the basin is locally mantled by colluvial deposits which were exposed by

exploratory Trench GL-5b. A nearly vertical discontinuity was observed at Station 0+62 (see trench log) which juxtaposes a distinctive carbonate hardpan (Unit 2a) on the east side with a stratigraphically older colluvial gravelly sand (Unit 3) on the west. The apparent east-side-down displacement of Unit 2a coincides with the overall sense of displacement along the west margin of the GL-5 basin. Overlying the discontinuity, a well developed and undisturbed carbonate layer (Unit 2) is present, which is expected to be similar in age to the offset carbonate horizon exposed (Unit 5) in Trench GL-5a to the east (R. J. Shlemon).

Southwest of Trench GL-5b and along the trend of the same scarp (but southwest of the main retarding dam at the Cross Hollow hills drainage), colluvial deposits of unknown depth are exposed by Trench GL-5c below a relatively smooth, consistent slope. The location of this trench is plotted in Figure VIII-10. At this locale, a continuous, well-developed leached carbonate horizon is displaced approximately 1 foot on a near-vertical plane in an apparent east-side-down sense (Photograph VIII-4). This feature may be associated with faulting, or alternatively the result of gravity induced down-slope movement occurring entirely within the colluvium.

All the offsets observed along the east margin of the Cross Hollow Hills are apparently normal (or dip-slip) displacements. The schematic cross section, Figure VIII-5, illustrates the likely geologic configurations in the GL-5 basin area.

C. Gypsum Wash

1. General Geology

Gypsum Wash dam is located on a gently westward-sloping surface west of Warner Ridge. The eroded core of a north-south trending anticline lies between the dam and Warner Ridge. The dam site is on the eroded (or planed-off) west-dipping flank, and Warner Ridge forms the prominent east-dipping flank. These relationships are shown on the schematic cross section presented in Figure VIII-7.

Bedrock underlying the site is mapped as Triassic-age Moenkopi Formation, consisting primarily of thinly bedded gypsiferous shale and siltstone with minor

interbeds of sandstone and limestone. These rocks locally display tight, contorted warps and randomly oriented gypsum or calcite-lined shear planes, but exposures in the vicinity indicate that bedding is typically coherent for extended distances.

Between the dam site and Warner Ridge and along the eroded anticline core, the topography is characterized by low, rounded hog-back hills with sharp "bad-lands" relief, which has resulted from erosion of the weak shale bedrock. Thin, older alluvial fan deposits overlie minor portions of the bedrock terrain. Both the older fans and the bedrock are well dissected by erosion.

In the vicinity of the dam embankment and the diversion berm to the south, exposures of the bedrock and old fan deposits (which are common to the east of the site) are buried beneath younger alluvial fan deposits. These younger fans thicken westward toward the center of the valley.

The mapped trend of the Washington fault strikes approximately north-south through this transition area between exposed and buried bedrock, and locally forms a sharp demarcation between bedrock and alluvial materials at nearby locations.

Exploratory trenches excavated along the trend of the Washington fault encountered both the older fan deposits exposed at the surface to the east, and the younger, overlying fan deposits. The younger deposits are typically well stratified and consist of loose-to-medium dense silt, fine-to-coarse-grained sand, and minor fine gravel. The sand is characterized by medium- to coarse-grained, weakly stratified lenses with a "salt and pepper" contrast. Loose silt and fine gravel zones are common, and form conspicuous, short, discontinuous lenses with sharp contacts. Soil profiles are virtually absent in these sediments.

The older alluvial fans are typified by relatively denser and coarser-grained materials. They consist of well stratified silt, sand, and gravel mixtures of varying percentages with internally chaotic gravelly beds or lenses common locally. The coarser lenses contain abundant angular and flat clasts of siltstone and gypsiferous shale. As with the younger fan deposits, no soil profiles or paleosols were observed in the older fans, since they either have been removed by erosion or were never developed.

2. Faults

The trace of the Washington fault has been documented along the east side of the valley in which Gypsum Wash Dam is located. The fault zone extends northward beyond the town of Washington, and southward into Arizona for a total distance of approximately 42 miles. The main trace of the fault is clearly exposed approximately two miles south of the site. In this area, a prominent scarp with sharp relief separates dissected Triassic-age bedrock on the east from Quaternaryage alluvial fan deposits on the west (see Figure VIII-11). Between the well-defined scarp and the dam site to the north, the fault traverses the drainage which extends westward from Warner Draw. No clear exposure of faulting is present where the mapped location of the Washington fault crosses the primary Warner Draw drainage ravine, but a prominent, high-angle fault cuts the Moenkopi bedrock somewhat east of the mapped trace of the fault. This offset probably represents a major splay of the fault, if not the actual main trace. At this location, the faulted bedrock is capped by unbroken, deeply dissected alluvial fan deposits which may be of similar age to the older fans exposed at the dam site to the north. Photograph VIII-6 illustrates the prominent shear and capping alluvial deposits at this location.

The trace of the Washington fault near the dam site was apparently located by pre-construction exploration, which involved excavation of test pits at the embankment site. These pits exposed ruptured bedrock and fan materials (Test Pits 118A, 118B, 120B, SCS data) which resulted in re-location of the dam further east to the present position. Subsequent studies during this investigation indicate that a wider zone of sheared earth materials than the pre-construction excavations revealed is present, and that the zone extends further eastward beneath portions of the dam.

Exploratory trenches excavated during the current study were located along lineaments or anomalous topographic features observed on air photos and during field reconnaissance mapping. Figure VIII-6 shows geologic conditions at the dam site along with the locations of the trenches. Trench G-1 was positioned across a north-northwest trending photolineament which coincides with the abrupt western end of the low, dissected hills located immediately south of the south end (or bend) of the dam embankment. Pre-construction test pit excavations to the north

exposed discontinuities with trends approximately parallel to this lineament. Exposures in Trench G-1 reveal dense, approximately horizontally stratified fan deposits which are cut by numerous high angle shears (see log). The shears trend north-northwest with displacements commonly ranging from a few inches to several feet. Both down-to-the-west and down-to-the-east relative displacement are present, resulting in a Horst and Graben structure. However, most offsets are down-to-the-west and coincide with the normal displacement associated with the Washington fault. At the west end of trench G-1, a thickening wedge of less consolidated, younger alluvial fan deposits (Units 1 and 9) overlie the older material. These younger deposits do not appear to be involved in the shearing which disrupts the underlying older fan deposits.

Trench G-4, located at the toe of the dam between the pre-construction test pits and Trench G-1, revealed offsets in older fan materials similar to those exposed in G-1. At the G-4 location, well stratified younger fan deposits are thicker and more clearly exposed. The uppermost deposits identified as Unit 8 on the trench log are clearly unbroken. Unit 1, underlying Unit 8 and overlying the older fan deposits, consists of a generally massive, loose zone of sandy silt. This zone is apparently involved in the offsetting shears. It conforms to the irregular surface (which is disrupted by offsets) of the older underlying fan, and is present as in-filling material in fissures and pull-apart structures along certain shear planes (Stations 0+26, 0+30, 0+39). The character of the Unit 1 material more closely resembles the younger of the fan deposits. The age of Unit 1 is estimated to be in the 5,000 to 10,000 year old range (Dr. R. J. Shlemon, written communication).

South of Trench G-1, Trenches G-2 and G-3 were sited at the western end of an exposure of gypsiferous shale which forms small, hog-back type ridges. These trenches exposed a sharp, nearly vertical shear plane which juxtaposes gypsiferous shale bedrock on the east with dense, older alluvial fan on the west. The shear plane trends N-S through both the G-2 and G-3 trench exposures, and extends upward to within 2 feet of the ground surface (see logs of Trenches G-2 and and G-3, and Photograph VIII-5). Young alluvial fan materials overlie and are in sharp contact with the older fan deposits, and are clearly offset by 2 inches on the shear plane in a down-to-the-west sense (the amount of bedrock and older fan offset is at least 4 feet). These offsets are located at Sta. 0+25 in trench G-2 and at Sta. 0+15

in Trench G-3 (see logs). The displacement of the young fan material most likely results from either direct tectonic fault displacement along the shear plane (indicating recurrent movements), or differential settlement across the shear plane by dissolution of gypsiferous material, or differential compaction across the shear plane (conceivably as a result of strong seismic shaking). Information on the depths and nature of alluvial deposits underlying the trench west of the shear plane could lend weight to one of the above possibilities. In any case, it is assumed the bedrock offsets exposed in G-2 and G-3 define a major trace of the Washington fault at this site.

Between Trenches G-3 and G-1, Trench G-X was excavated and inspected during a field meeting at the site on January 26, 1982. This trench was not logged, since it was excavated in an attempt to resolve to those present, the nature of the offset relationships between the younger and older alluvial units as defined in Trenches G-2 and G-3. Trench G-X revealed a single zone of rupturing within older alluvial deposits. An overlying sequence of loose, young alluvium a few inches in thickness was found to be unbroken.

The relationship of the Gypsum Wash dam embankment and the main trace of the Washington fault or fault zone is shown on Figure VIII-6. At the southeast end of the dam, the fault location is based on the northward projection of the fault as exposed in Trenches G-2 and G-3, the observed location of the bedrock/alluvium contact just south of the dam embankment, and bedrock exposures within the basin just to the north. Bedrock is at least 10 feet below the surface just to the west, where Trenches G-1 and G-4 expose the sheared alluvial deposits. Within the trenches, these shears increase in frequency eastward (see trench logs), supporting the concept that fault activity has been concentrated along the projection of the locally exposed bedrock/alluvium contact, as illustrated on Figure VIII-6.

Near the central and northern portion of the dam embankment, areas where the low, rounded hills are present generally coincide with areas underlain by bedrock. Here the available evidence indicates the Washington fault is located along the western edge or toe of the hills as shown in Figure VIII-6, separating bedrock from the essentially flat-lying alluvial fan deposits to the west.

The exact location where the trend of the Washington fault trace or zone intersects the dam alignment between its location at the southeast end of the dam and the generally well-defined hillfront scarp just northwest of the dam is speculative because of the lack of bedrock exposures and locally altered topography near the dam. An available SCS construction drawing ("Cracks Exposed in Bedrock Foundation and Cutoff Excavation, Gypsum Site", 5/74) indicates a series of northwest-southeast trending cracks exposed on the surface and within (?) the embankment cutoff trench between Stations 28+30 and 30+25, which may represent the fault location. The geologic logs of the foundation excavation and cutoff excavation ("Geologic Profile - Cutoff Trench and Foundation Excavation, Gypsum Site", 4/74; SCS data) only illustrate lithologic zones and do not indicate the presence of cracks or shears at these or other locations along the dam alignment.

Figure VIII-7 is a schematic sketch of a simplified south-facing cross section of geologic conditions at the Gypsum Wash site. The shear relationships exposed in the trenches during the field investigation are summarized by the fault offsets illustrated at the circled letter "A", "B", and "C". "A" represents the northward projection of the shear located in Trenches G-2 and G-3 through the low hills just south of the southernmost corner of the dam. It is assumed the contact of bedrock and old alluvium existing at this location is a fault contact, although no excavations were made here because no potentially datable soils are present.

"B" represents the conditions exposed in Trenches G-2 and G-3 where the prominent shear plane offsets bedrock, older and younger alluvial deposits as described above. The estimated age of the young fan deposits in these trenches is no older than 1,000 to 1,500 years (Dr. R. J. Shlemon, written communication).

"C" illustrates the relationships in Trench G-4, where alluvial Unit 1 with an estimated age of 5,000 to 10,000 years is involved in shear offsets. The youngest, uppermost sediments at this site (Unit 8) are estimated to be of late Holocene age, and are unbroken. The older alluvial fan deposits exposed in all of the Gypsum Wash trenches are estimated to be late Pleistocene, or between 10,000 and 25,000 years old (Shlemon). Although no materials were found in the trenches which could be positively dated, conditions exposed indicate that fault-related displacements of the alluvial materials underlying the dam are likely to have occurred during Holocene time.

D. Warner Draw

1. General Geology

In the dam site locality, Warner Ridge strikes approximately north-south and forms the western margin of Warner Valley. The ridge is underlain by the east flank of an anticline, the axial zone and western flank of which are now eroded away. Runoff from Warner Valley has eroded a low gap in the ridge forming Warner Draw, a narrow westward-draining channel where Warner Draw Dam is sited. Most of the foundation and right abutment of the dam is underlain by well-bedded Triassic-age rocks of the Moenkopi and Chinle Formations, which also underlie Warner Ridge and dip eastward approximately 20 degrees from horizontal. These rocks consist mainly of weak shale with prominent interbeds of harder siltstone and fine-grained sandstone, and individual beds can commonly be traced for many hundreds of feet where exposed. The character of these rocks is described further in materials contained in the SCS files.

The left abutment and a portion of the southerly foundation area is underlain by old alluvial sand deposits, which are at least partially eolian. The alluvial deposits range from well sorted (poorly graded) silty sand to sandy silt, and are typically massive with a few minor stratified pockets of coarse sand and gravel. Leached carbonate commonly cements the alluvial sediments, forming dense, stable deposits which have been locally incised by erosion. According to previous records produced during dam construction, the alluvium extends to considerable depths, filling an old stream channel meander located just south of the present drainage channels (SCS data).

2. Faults

The prominent gap at Warner Draw suggests the presence of a weakness or discontinuity in the Warner Ridge bedrock at this location. Nearby, prominent shears cut the conspicuous promintory located about 3,000 feet southwest of the site, where the point of the ridge has been displaced downward to the west (see Figure VIII-12). The possibility that a weak zone at Warner Draw resulting from another similar fault splay was considered, since the Washington fault is located only $\frac{1}{2}$ mile to the west.

Several air photo lineaments which extend close to or in the direction of the Warner Draw gap were identified during this investigation. Their orientation was anomalous to the general north-south trend of the Washington fault, however, and field examination of bedrock outcrops and ravine channels cut in alluvium where the air photo lineaments were plotted revealed no evidence that the lineaments were fault related. Reconnaissance of the Warner Draw drainage indicates that the bedrock is essentially coherent near the dam, with a progressive increase in fracturing and distortion westward to the Washington fault zone. (The major shear displayed in Photograph VIII-5 and described in Section VIII-C.2 is located at the eastern margin of the fault zone.) However, a small bedrock fault exists near the Warner Draw channel bottom which intersects the downstream toe of the dam (Figure VIII-8, fault (A)). The fault has an apparent thrust offset of three feet and is oriented N20E, dipping 30 southeast. Photographs VIII-7 and 8 illustrate this feature which is a continuous, consistent tight plane traceable throughout the bedrock exposure. An offset of similar orientation is locally exposed in bedrock higher in the stratigraphic section, approximately 1,000 feet to the south. During dam construction, another bedrock fault (Figure VIII-8, fault (B)) was mapped at Station 15+70 of the embankment alignment in the foundation cut-off trench (SCS data). Information as to the strike and amount of displacement is not available, but it is conceivable this shear parallels the above described offsets.

Inspection of the adjacent alluvial deposits along the projection of the dam site faults revealed no signs of disturbance. Exploratory trenches WD-1 and WD-2 were excavated across the projected strike of the fault shown in Photographs VIII-7 and 8, and across the speculated trend of the fault mapped in the foundation cut-off trench. The trench exposures revealed dense-to-very-dense, fine-grained silty sand deposits which display no evidence of faulting or other disturbances. The faults present in the adjacent bedrock are considered to be ancient features which most likely resulted from compressional stresses associated with folding of the Warner Ridge anticline.

E. Stucki

1. General Geology

Stucki dam extends for 1,400 feet across a broad, north-draining Valley which parallels the Washington fault and Warner Ridge. These features are located approximately 2,000 feet and 3,000 feet to the east, respectively (see Figure VIII-11).

Alluvial fan deposits extend westward from the ridge front and obscurely merge with valley fill and stream channel alluvium along the lower east side of the basin (see Figure VIII-9). The fan and alluvial deposits underlie all but the extreme western end of the dam embankment. The west (left) abutment of Stucki dam and the southward-extending basin margin is formed by a continuous, east-facing slope which is eroded into older, very dense sandy alluvium. The flat upper surface of the old alluvium west of the left abutment slope is underlain by a well developed calcrete or hardpan layer, which results in a prominent ledge at the top of the east-facing slope. At the toe of this erosional escarpment, the alluvium within the drainage basin is in disconformable contact with the older, dense alluvium. A description of the lithology of the earth materials present at the site is presented by Deming and Bridges 1971 construction report in the SCS files.

2. Faults

The Washington fault is prominently exposed east of the site, where alluvial fan deposits on the west are juxtaposed with Triassic rocks on the east. The current investigation disclosed no traceable surficial evidence of westward-branching splays extending toward the embankment, but the prominent left abutment and basin margin (which forms a continuous, north-northwest south-southeast escarpment) was regarded as a suspicious feature in view of the similar configuration of the nearby Washington fault.

Field reconnaissance along the escarpment just south of the embankment established the continuity of horizontal beds at and just below the calcrete. These beds can be traced several tens of feet westward within a prominent erosional

gully present here. The central and lower portions of the escarpment were further investigated by exploratory trenches SD-1 and SD-2. These trenches exposed very dense sand and silt in thick, poorly defined flat-lying beds, overlain locally by colluvial deposits (see log of Trench SD-1). At the toe of the slope, Trench SD-1 exposed the dense old alluvium at an abrupt, near-vertical contact with loose, Rubbly material derived from the older alluvium extends younger alluvium. eastward from the contact for a few feet below the overlying younger deposits as illustrated on the trench log. Shear planes, dragged structures or other features indicative of faulting were not exposed. Trench SD-2, excavated 100 feet to the north at the toe of the slope and along the trend of the vertical contact, exposed an eastward-thickening wedge of younger, looser, valley alluvium, forming a smooth, shallowly-east-dipping contact with the underlying dense old alluvium. The above conditions indicate the vertical contact configuration of the alluvial units in Trench SD-1 represent an old wash or ravine bank which was subsequently in-filled by the younger alluvial deposits. On the basis of the exposures afforded by the erosional gully below the calcrete and the exploratory trenches, the left abutment escarpment is considered to be a natural erosional feature and not the result of faulting.

Additional geologic investigation concerning the remainder of the dam site was performed by the SCS during construction operations when a cut-off trench was excavated into the alluvial deposits prior to placement of the embankment. Faults or shear planes were not reported during construction, and none are illustrated on the log of the cut-off trench (SCS data). Considering the extent of the site exploration, it is reasonable to assume the foundation and abutments of Stucki dam are free of active fault traces.

F. Frog Hollow

1. General Geology

Frog Hollow dam site is located approximately 2 miles east of the conspicuous north-south trending Hurricane fault scarp. Quaternary-age basalt flows (which are common in the vicinity of the fault) extend eastward from vents near the scarp and underlie the dam site. North and east of the site, alluvium of varying

thickness overlies basalt flows which have been mapped as Stage 3, or oldest of the flows in the vicinity (SCS construction reports). South and west of the dam, a younger basalt flow (Stage 4) forms an obvious contact at the western margin of the alluvium that underlies the drainage basin upstream from the dam (see Figure VIII-12).

Geologic conditions exposed in Workman Wash (the Frog Hollow dam site channel) before and during placement of embankment materials were investigated by the SCS and are described in pre-construction geologic reports available in SCS files. These studies indicate relatively pervious zones at the upper and lower margins of the basalt flows, and a southerly-dipping alluvium/basalt contact underlying the present embankment. Basalt exposures currently observed near the dam indicate only hard, fresh materials, particularly in the channel bottom where abrasion from stream flow has scoured the rock. Earlier reports indicated "extremely weathered" basalt apparently in or along the channel, but only fresh rock was observed at the available downstream exposures during the current study.

Basalt exposed in the channel downstream from the dam is fractured along two major joint set orientations, both of which are nearly vertical. One set strikes north-south to N15W, with joint spacing ranging from 2 or 3 feet to 10 feet, but typically about 5 to 6 feet (Photograph VIII-9). The second joint set cuts the basalt at approximately right angles to the first mentioned set, striking N40 to 60E with spacing from about 3 to 10 feet, averaging 6 feet or greater (Photograph VIII-10). Near the rock surface, many of these joints are open several inches or more as a result of erosion and scouring, but appear to be generally tight at depth. Surface exposures in the channel indicate these joint sets are poorly developed, generally extending for a few tens of feet where they die out or are truncated by other fractures. Near-horizontal flow structures, including fractures and cooling vesicules, are exposed throughout the channel, and result in more prominent and continuous zones of discontinuity than the joint sets.

Alluvium overlies bedrock upstream from the site. At least 8 feet of alluvial cover is currently exposed where deep gullying exists at the margin of present storage basin. Preconstruction excavations and borings in the upstream area indicate the alluvium has a maximum thickness of several tens of feet in the areas explored.

2. Faults

Both the review of previous geologic mapping and reports together with the current air photo analysis and field reconnaissance, indicates the Frog Hollow site is free from active faulting. The Triassic sedimentary rocks which underlie the basalt in this area may be faulted, but no evidence of such faults, if they exist, are found in the younger, overlying Quaternary basalt or alluvium.

G. Ivins Diversion No. 5

1. General Geology

The Ivins Diversion No. 5 embankment is sited on an alluvial fan deposit which extends south-southeastward from the prominent erosional scarp at the south end of the Red Mountains. The steep escarpment is cut in nearly horizontal beds of gypsiferous shale, siltstone and sandstone of the Triassic-age Kayenta Formation, above which Navajo formation sandstone forms vertical, prominent cliffs. At the west end of the diversion dams and just north of the town of Ivins, an old block slide is present which rests on the lower portion of the escarpment (see Figure VIII-13). Large, integral masses of both the Navajo and Kayenta formations are exposed in this slide deposit. The degree of erosional dissection and amount of colluvial debris which masks portions of the block slide suggest this feature is most likely pre-Holocene in age.

2. Faults

Several prominent lineaments are present in the Navajo sandstone which caps the southern end of the Red Mountains (see Figure VIII-13). Field reconnaissance revealed no evidence to indicate these lineaments extend below the Navajo sandstone into the underlying weaker shales. Prominent, vertical fractures cut the Navajo sandstone at the cliff-top exposure and presumably account for the lineaments observed.

Additional reconnaissance of an air photo lineament in Quaternary deposits east of the site failed to disclose any evidence of disturbed materials. Based on the conditions observed, the Ivins dam site is considered to be free of faulting.

Major faults which are present in the region include the Grand Wash fault located approximately 5 miles west of the site, and the Washington and Hurricane faults, located 10 miles and 21 miles east, respectively. Each of these faults has experienced normal, down-to-the-west displacements.



Photo VIII-1 View of south wall and floor of Trench GL-2a, Station 0+98. Coarse, rubble zone with locally loose gravel to left of sharp discontinuity; lighter-toned carbonate cemented gravels and underlying silt lens (on floor) to right. Most observers agreed this feature represents an old erosional channel margin. Overlying sediments at red flagged nail are continuous.



Photo VIII-2 View approximately north of road cut at south end of the Cross Hollow Hills. Basalt mapped as Quaternary in age disconformably overlies older, stratified alluvial deposits which have been offset. Sharp, continuous basalt/alluvium contact indicates rupturing occurred prior to basalt flow.

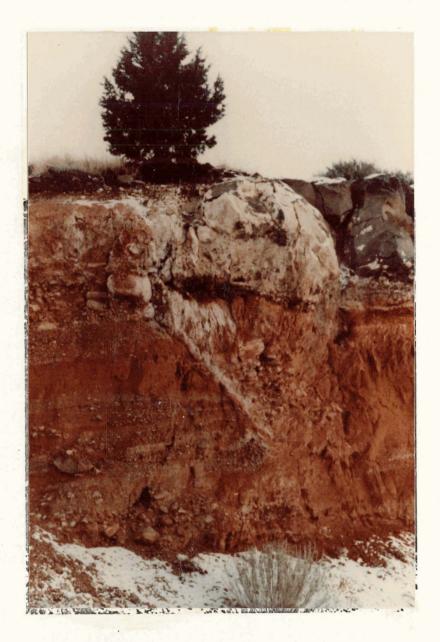


Photo VIII-3 View of the same road cut just east of Photo VII-2. At this location, both the alluvium and younger overlying basalt are displaced by shearing, indicating a post-basalt flow rupture. Assuming the basalts of Photos VII-2 and VII-3 are the same age, recurring displacements are indicated.

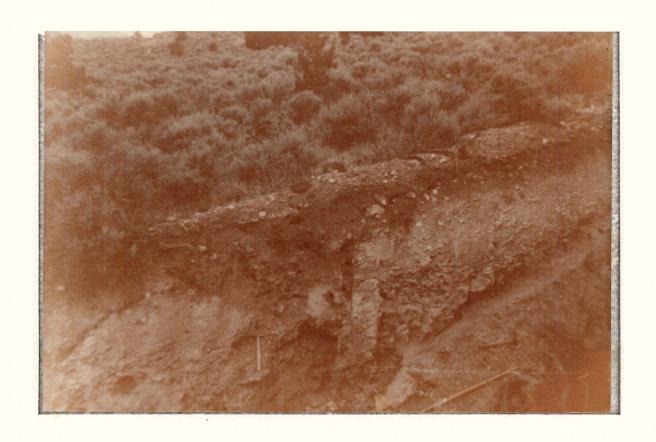


Photo VIII-4 View of south wall of Trench GL-5c. Colluvium derived from Cross Hollow Hills basalt to right mantles slope. Prominent carbonate (hard pan) horizon 1 or 2 feet below surface is disrupted locally along near-vertical planes. Trench GL-5c is located along southern projection of prominent scarp and an offset exposed in Trench GL-5b.

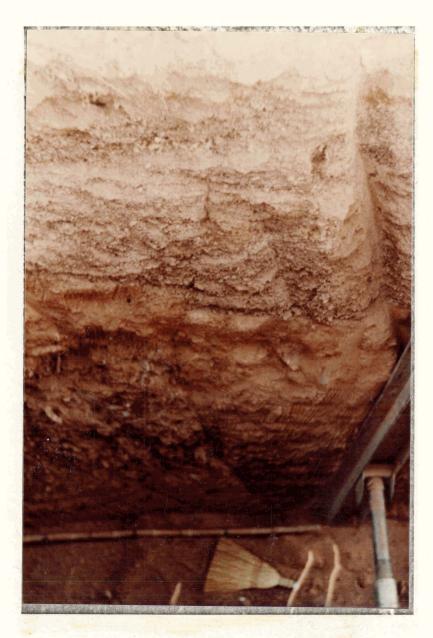


Photo VIII-5 View of south wall of Trench G-2. Shear plane extends from bottom of wall just left of shore upwards to 1 inch from top of photo. Gypsiferous shale bedrock at lower left, bedrock debris and older alluvial fan at lower right across shear. Young alluvial fan overlying is pale gray and is offset 2 inches at base.



Photo VIII-6 View of north wall of ravine which extends westward from Warner Draw. Deeply eroded vertical channel in disrupted shale and siltstone of Moenkopi Formation is prominent shear zone associated with Washington Fault. Undisturbed alluvial fan deposits unconformably overlie Moenkopi Formation at contact approximately 1.75 inches below top of photo. Pine Valley Mountains north of St. George in background.

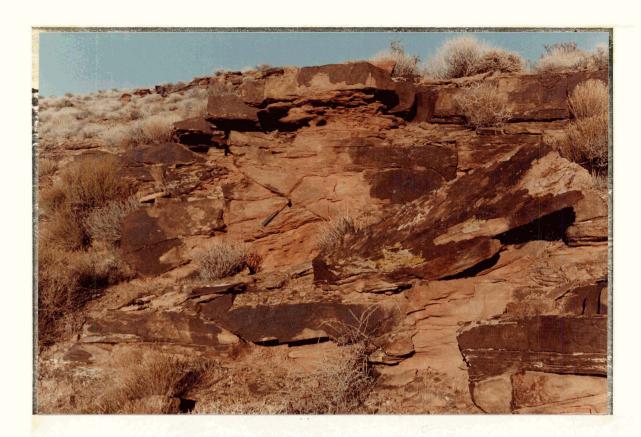


Photo VIII-7 View northeast of Chinle Formation outcrop on southeast side of ravine approximately 100 feet downstream from toe of Warner Draw dam. Rock hammer inserted in bedrock fault plane with apparent 3 feet of thrust displacement. Shear plane is a tight, continuous plane oriented N20E dipping 30% SE. No evidence of disturbance was found within younger materials explored along the projection of this fault.

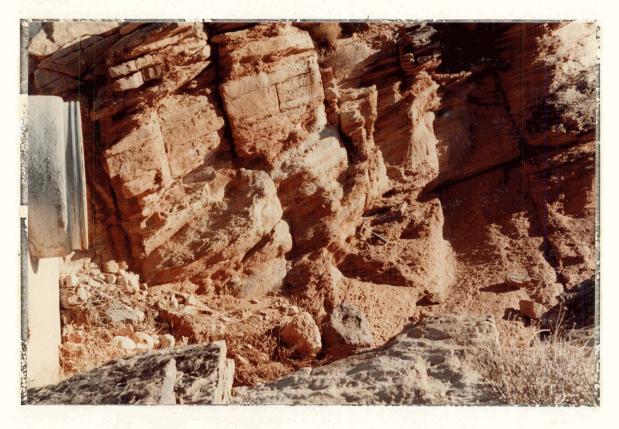


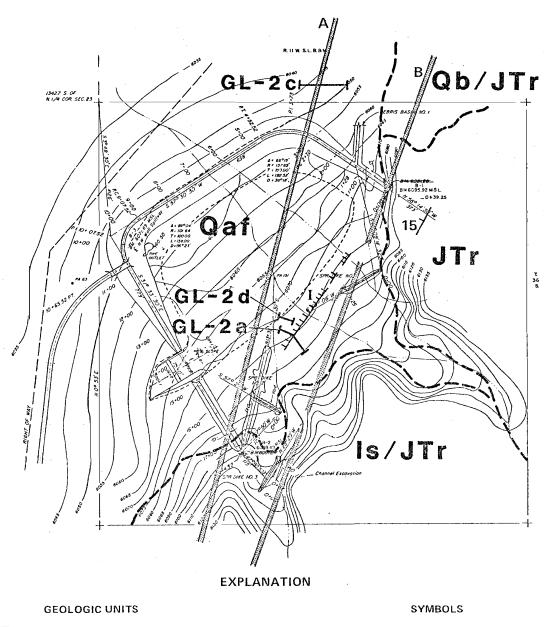
Photo VIII-8 View southeast of same fault plane as in Photo VIII-7 where it intersects outlet pipe at toe of Warner Draw dam. Plane is located just above head of rock hammer and appears to coincide with bedding planes at this configuration.



Photo VIII-9 View north of basalt (mapped as Stage 3 in SCS data) approximately 500 feet downstream from toe of Frog Hollow dam. Prominent joints oriented N5W are well defined at the surface but appear to tighten with depth. Abrasion-polishing of basalt in lower portion of channel obvious.



Photo VIII-10 View east of basalt downstream from Photo VIII-9. A second well defined joint set oriented N55E is exposed here. Rock hammer located where erosion has cut laterally into the channel wall along joint plane. Approximately horizontal flow banding and associated fractures are apparent here and form the most continuous structures within the basalt.



Alluvial fan deposits: silt, sand, and gravel; locally rubbly. Basalt: poorly defined zones of limited areal extent. Landslide deposits:poorly defined accumulations of Ωb and JTr. Jurassic and Triassic rocks, undifferentiated: includes sandstone, siltstone, gypsiferous shale, and limestone.

150

300 feet

Geologic contact; approximately located.

Topographic scarp located during ESA field reconnaissance; hachures on down-dip side.

Suspected fault traces located from SCS data; refer to Section VI. A-2 for discussion.

____15 Strike and dip of bedding.

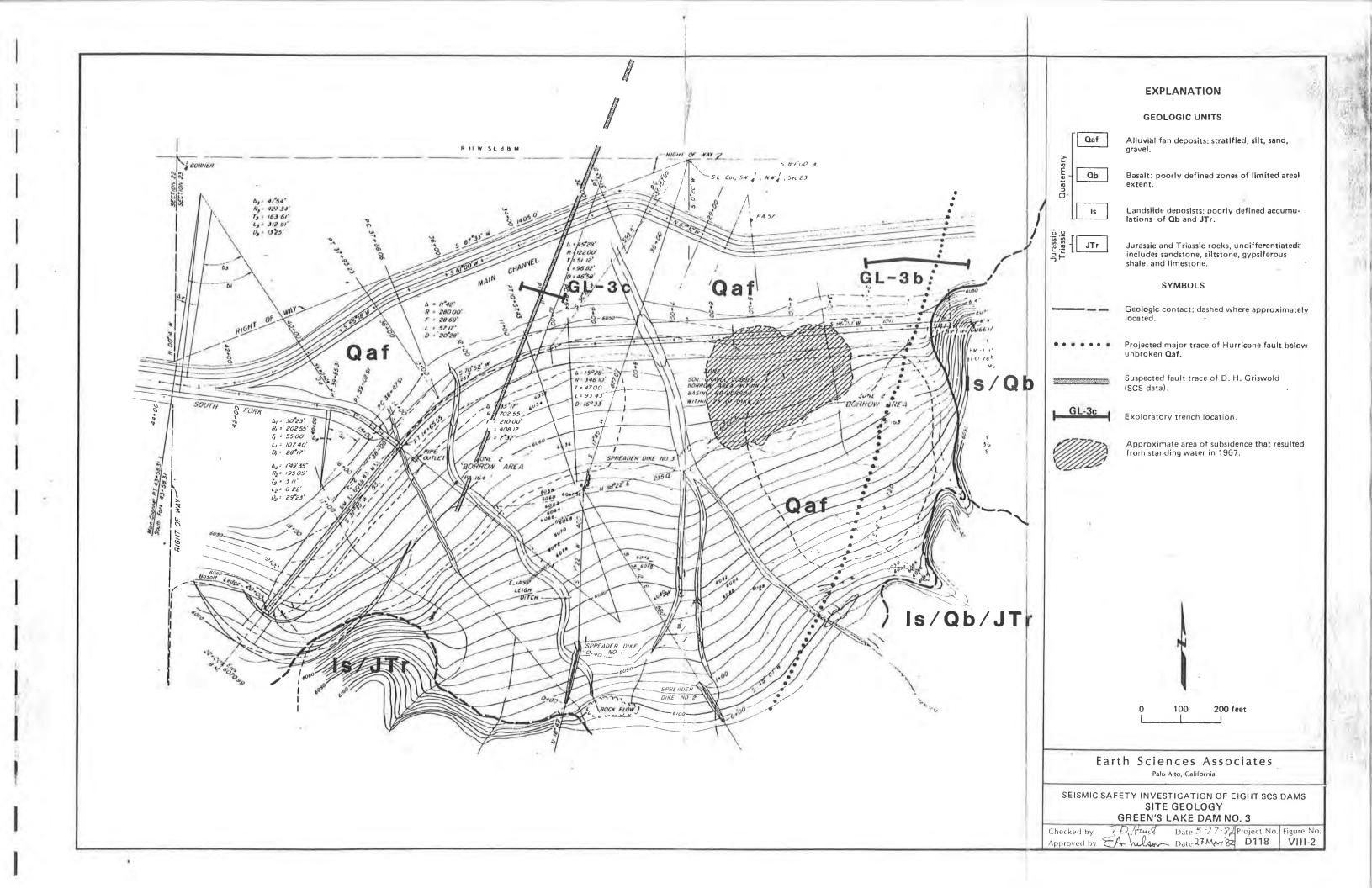
GL-2a Exploratory trench location.

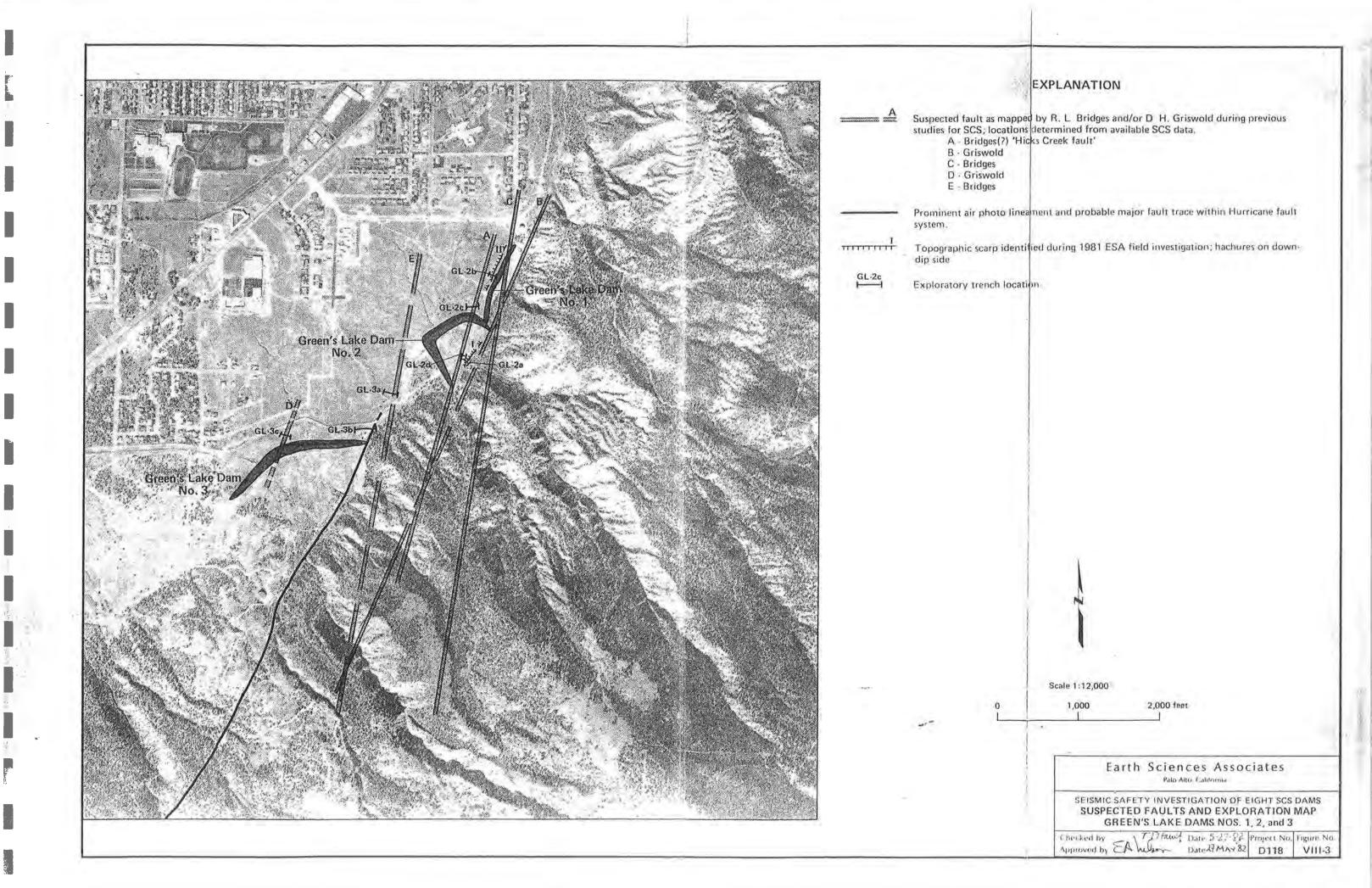
Earth Sciences Associates

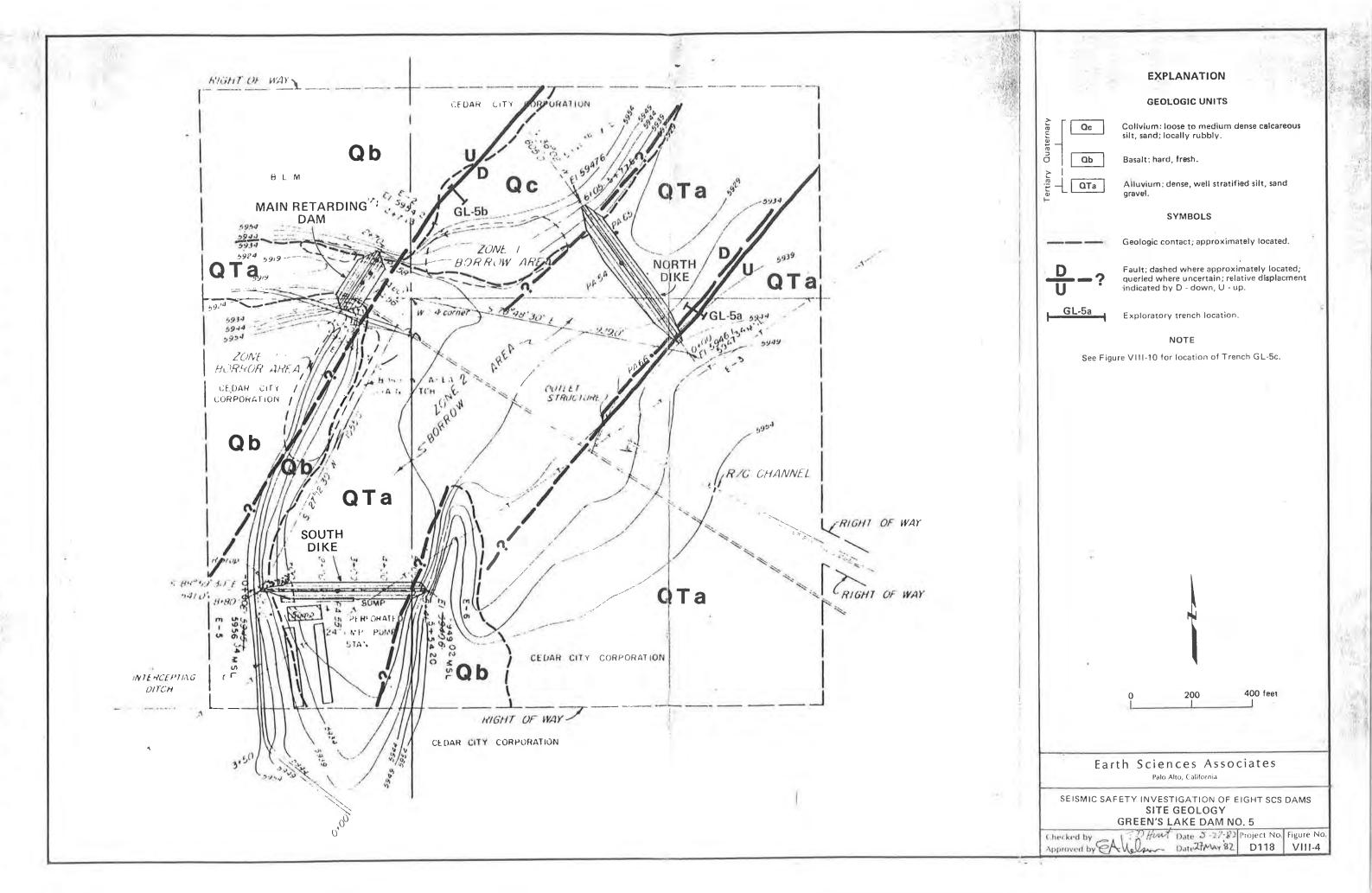
Palo Alto, California

SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
SITE GEOLOGY
GREEN'S LAKE DAM NO. 2

Checked by Date 5-27-82 Project No. Figure No. Approved by A. Nelson Date 27 May 82 D118 VIII-1

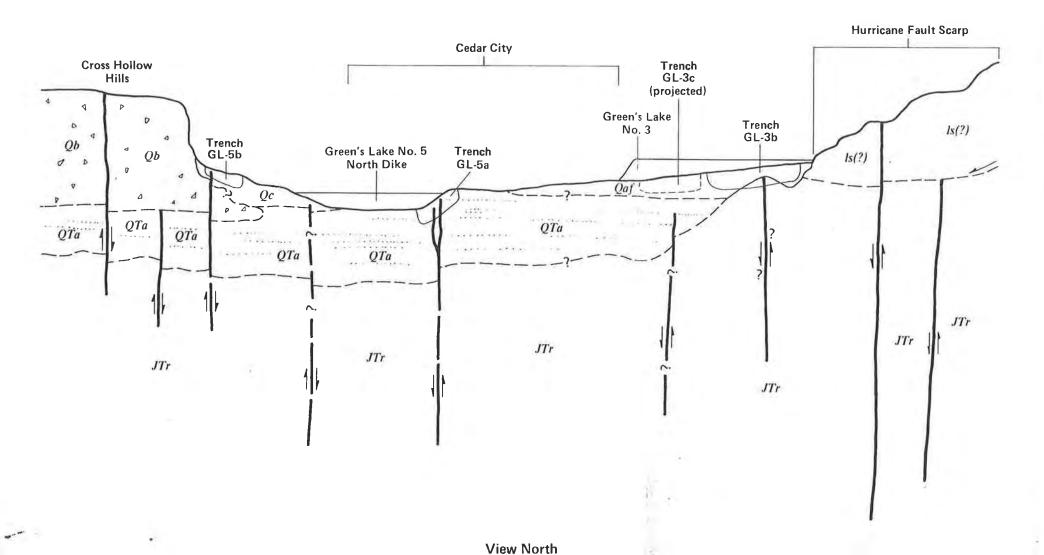






West

East



No scale

GEOLOGIC UNITS

Qaf | Alluvial fan.

Qb Flow basalt.

QTa Alluvium.

Is Landslide deposits.

Jurassic and Triassic rocks (includes local intrusions of basalt).

SYMBOLS

Geologic contact: approximately located; queried where uncertain.

Normal fault within Hurricane fault zone: arrows indicate relative displacement; queried where displacement or existance speculative.

NOTES

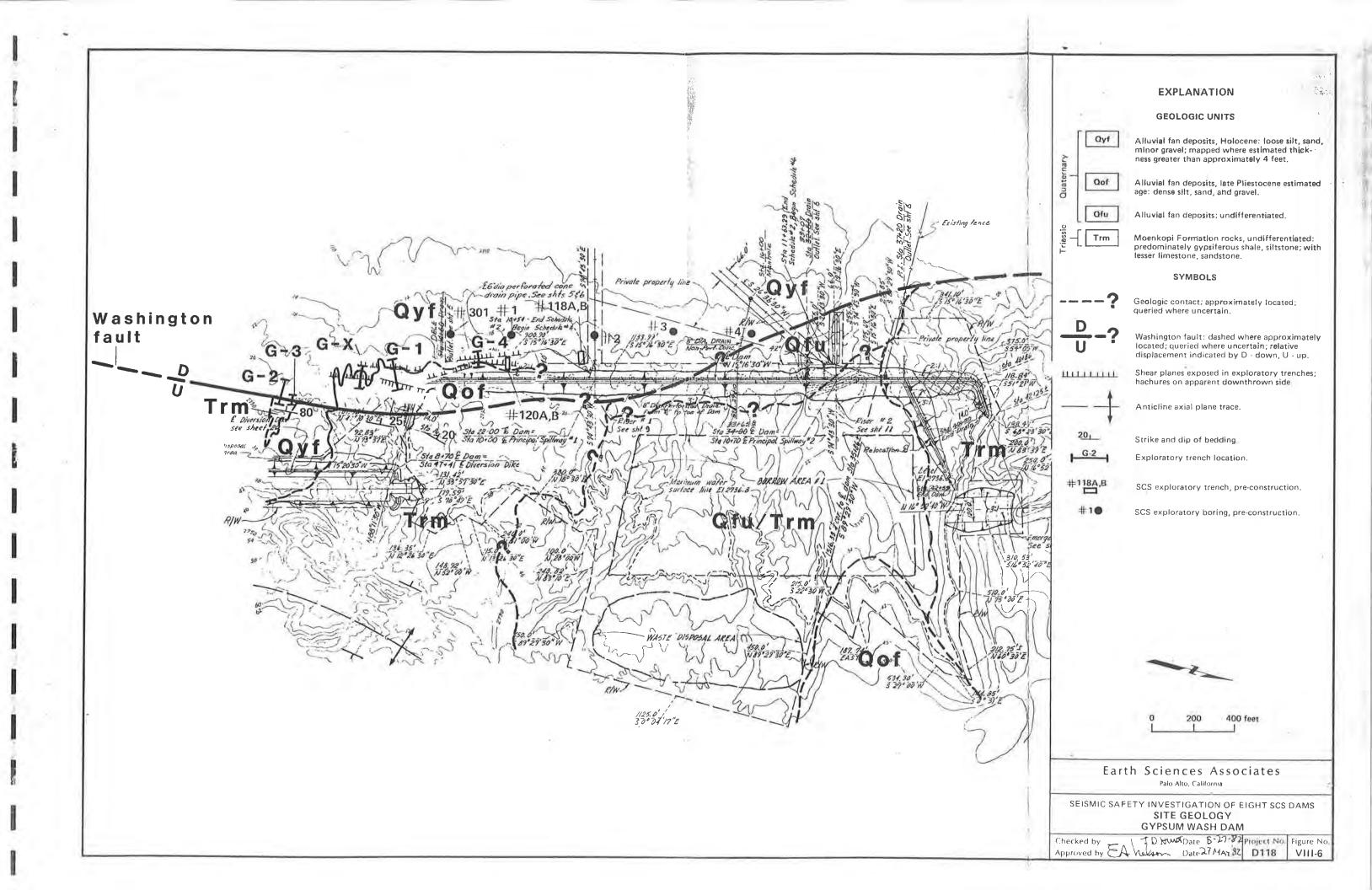
Section is a diagrammatic sketch summarizing the geologic conditions determined during the current investigation; locations of faults and geologic contacts will vary from actual field conditions.

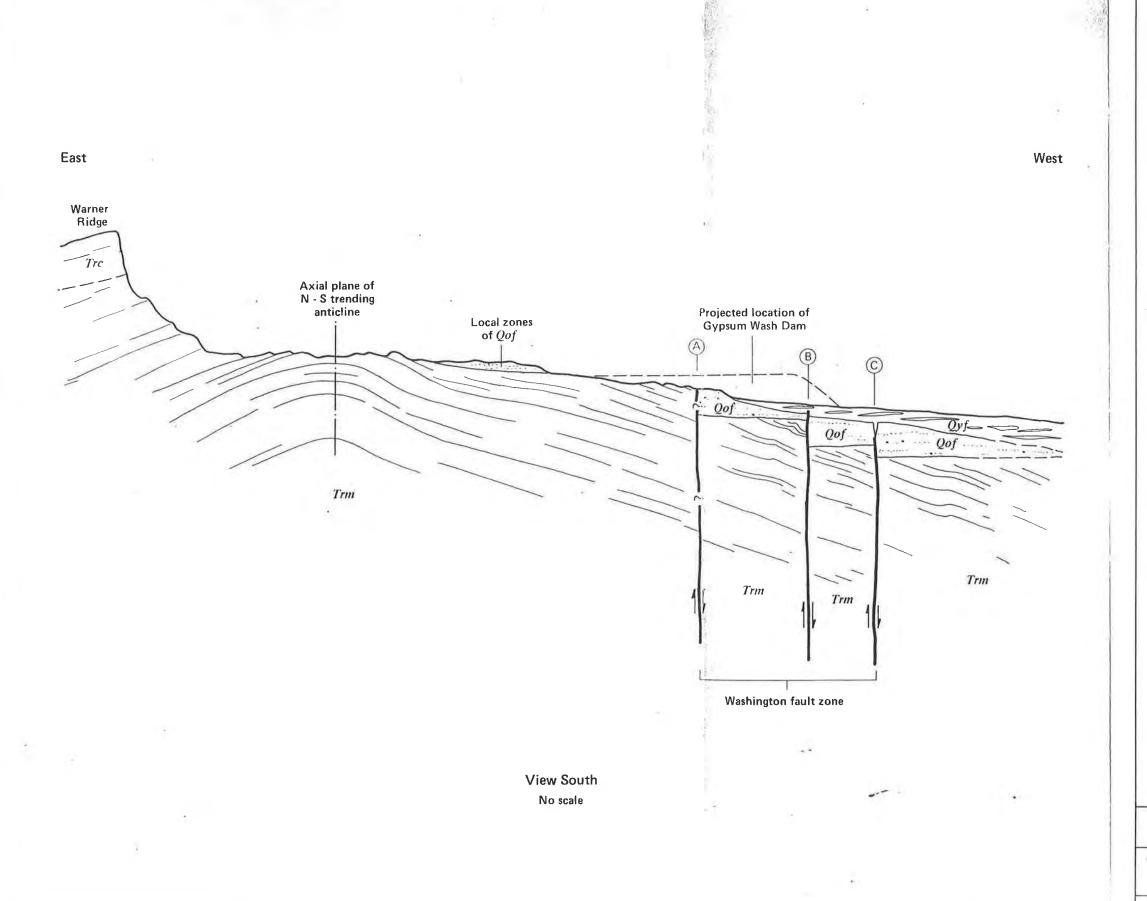
Earth Sciences Associates

Palo Alto, California

SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
SCHEMATIC GEOLOGIC SECTION
CEDAR CITY VICINITY

Checked by A Project No. Figure No. Approved by EA whom Date 27 PM 82 D118 VIII-5





GEOLOGIC UNITS



NOTES

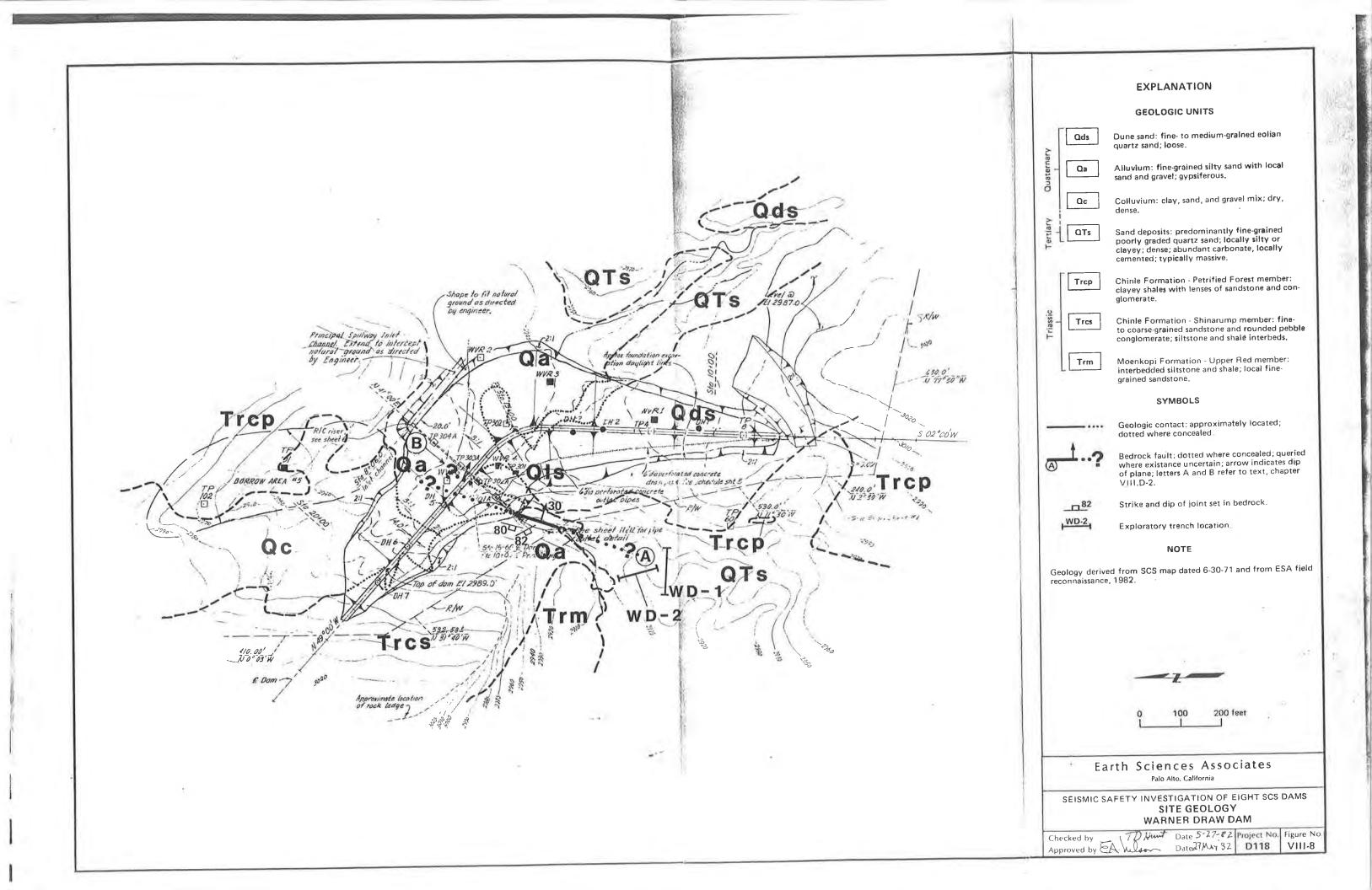
- A Structural relationship apparent in low hills located at south end of embankment and north of trenches G-2 and G-3; major fault trace juxtaposes gypsiferous shale on east with old alluvial fan on west; no surficial scarp or younger overlying sediments to aid in dating of last displacement.
- B Structural relationship exposed in trenches G-2 and G-3, bedrock offset down to west with older alluvial fan juxtaposed on west side; overlying young alluvial fan broken by shear at contact with older fan; shear dies out upward in young fan.
- © Structural relationship exposed in trench G-4; shears present in older alluvial fan deposits. Younger fan deposits overlie older fan, and are involved in shears as in-filling material in pull-away structures; no through-going shears into youngest alluvial lenses near surface.

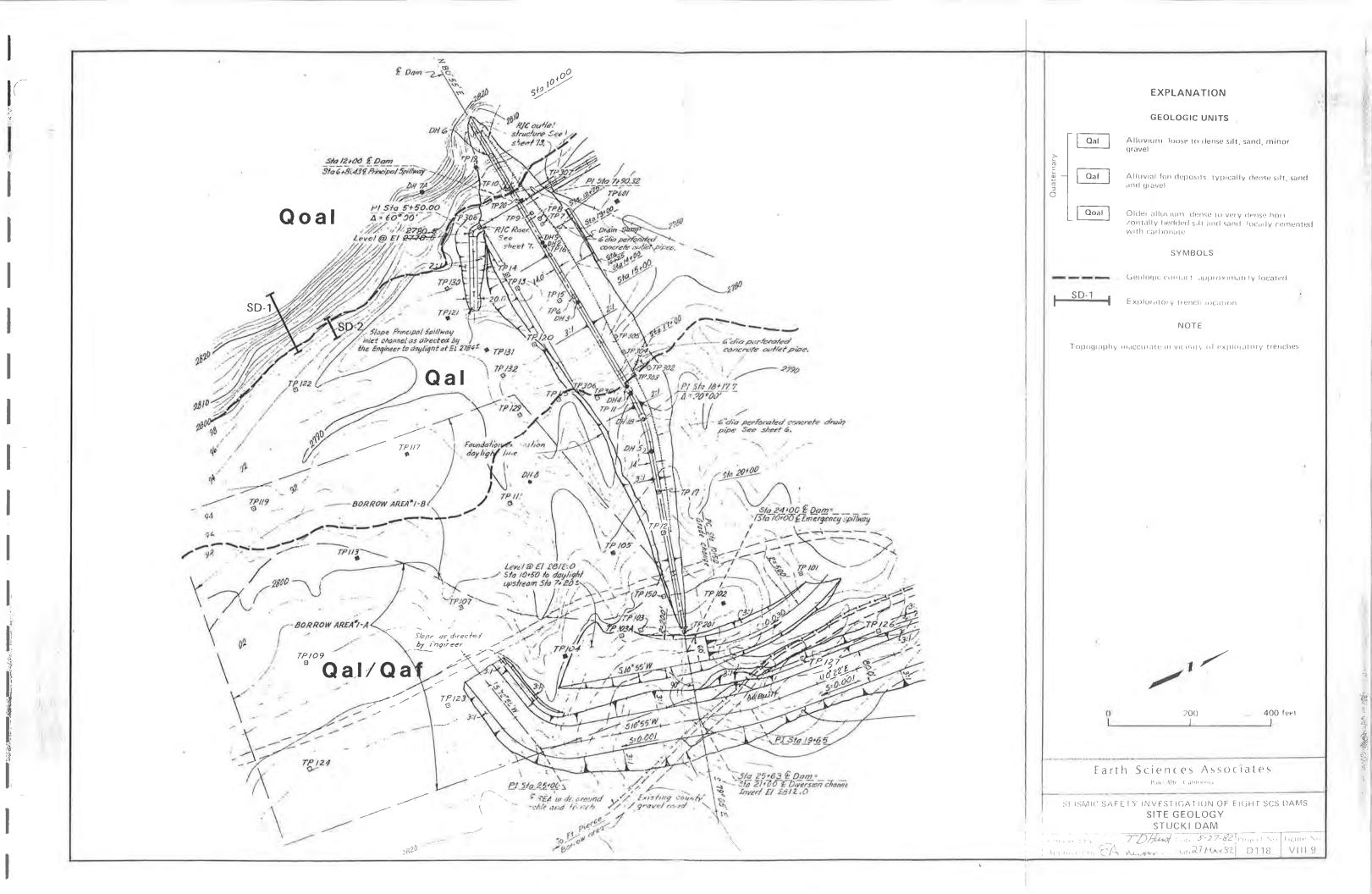
Earth Sciences Associates

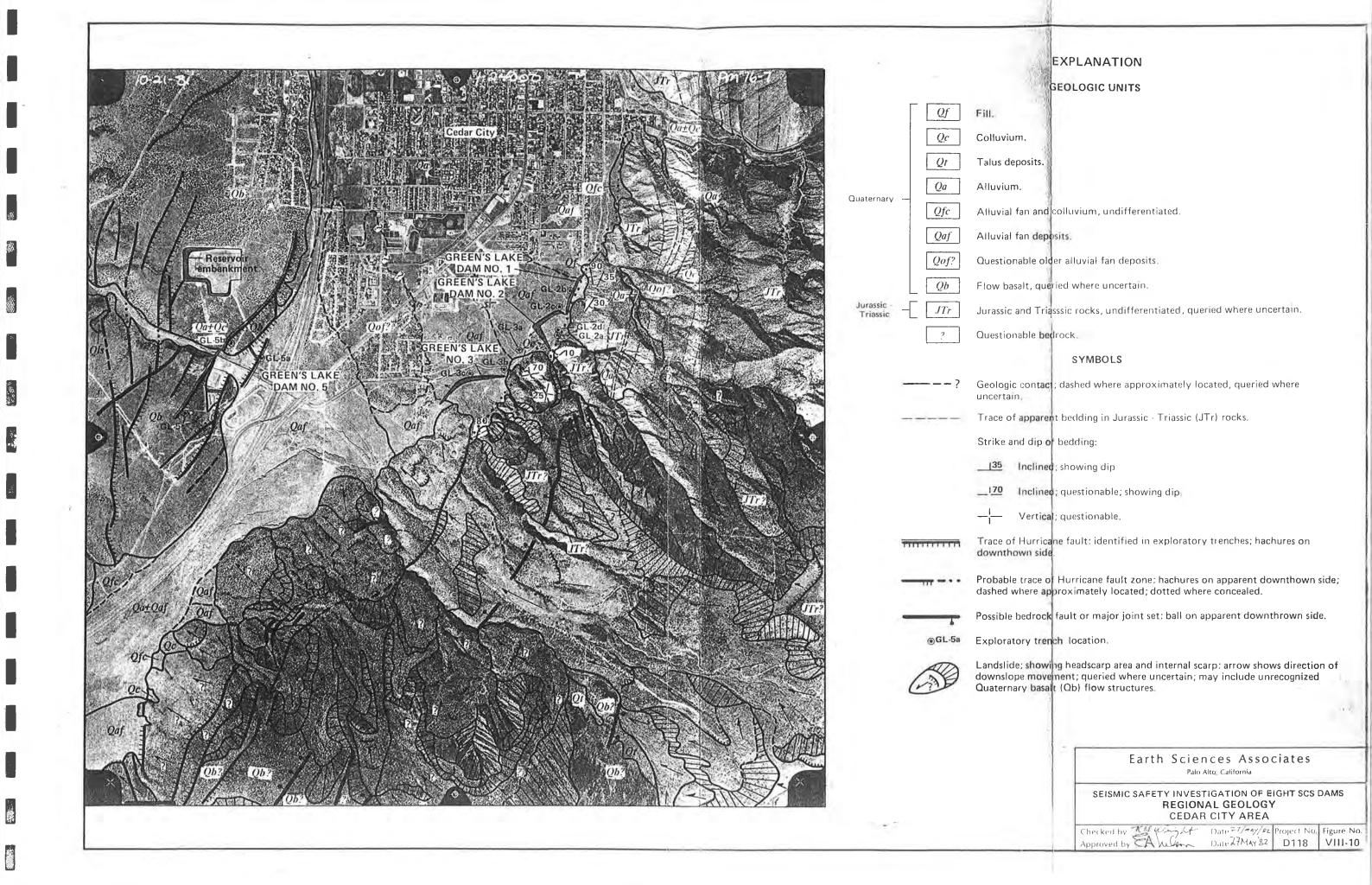
Palo Alto, California

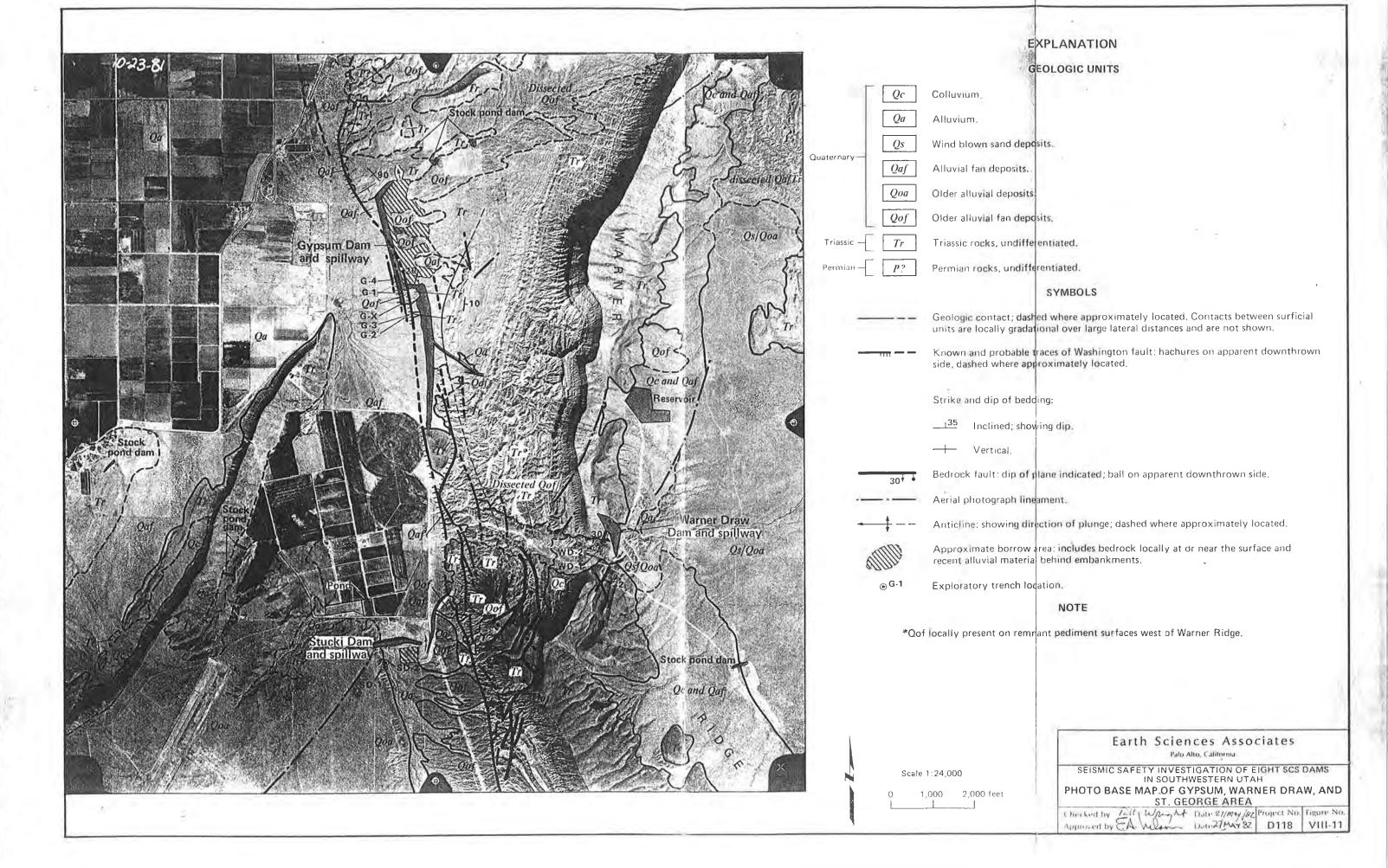
SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
SCHEMATIC GEOLOGIC SECTION
GYPSUM WASH DAM SITE VICINITY

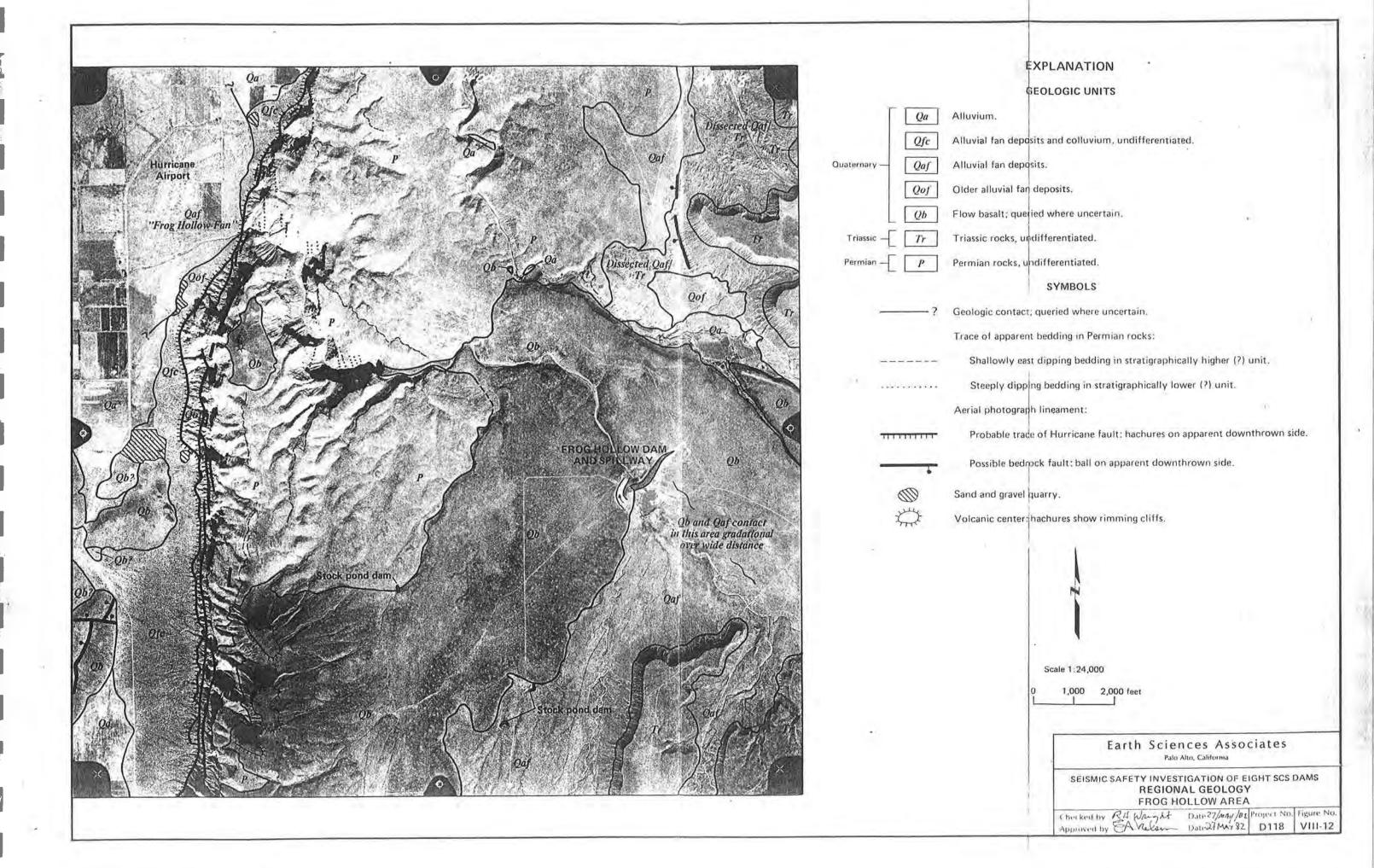
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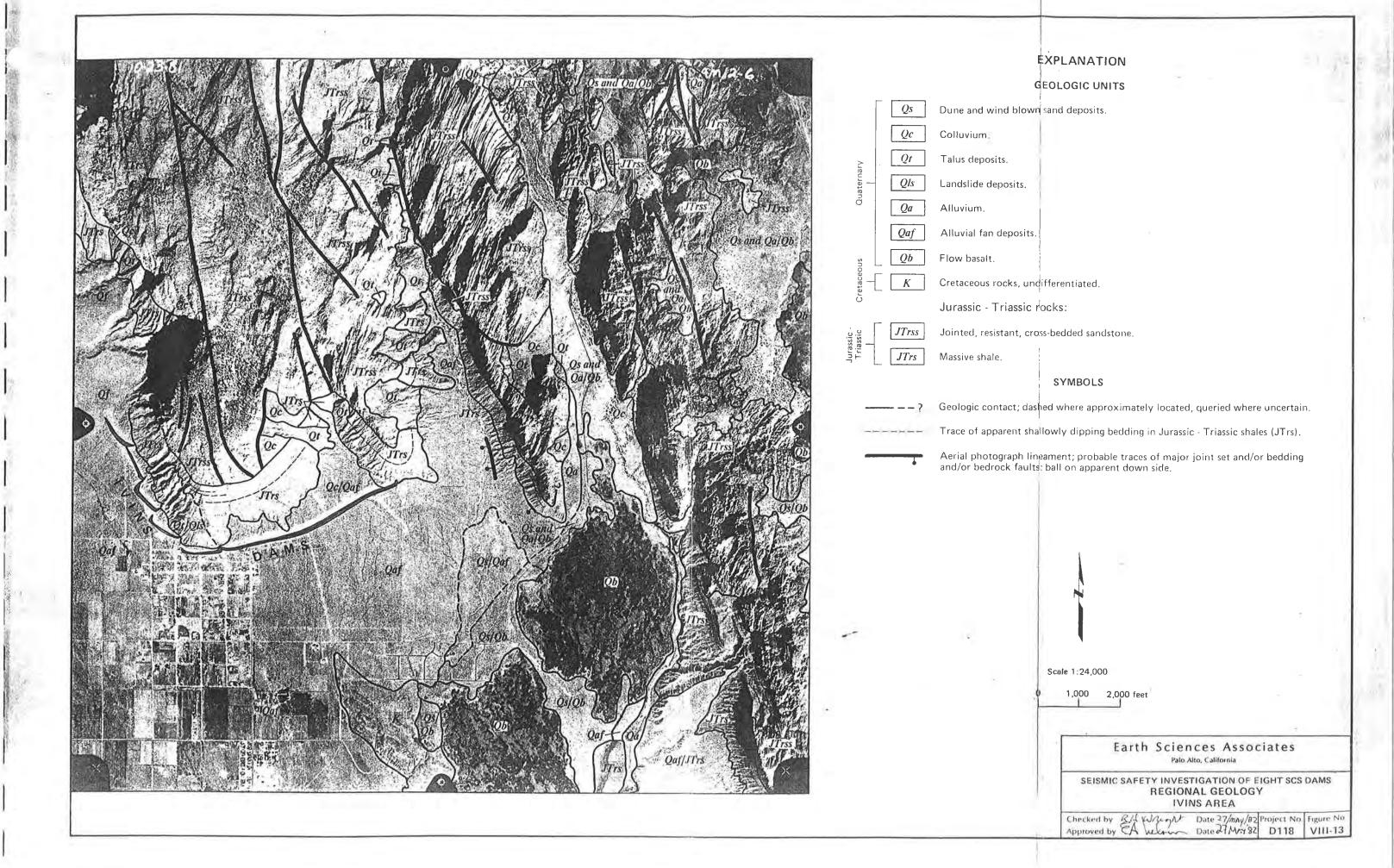












APPENDICES
TO
PHASE I REPORT
SEISMIC SAFETY INVESTIGATION
OF EIGHT SCS DAMS
IN SOUTHWESTERN UTAH

APPENDICES

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APPENDIX A

TRENCHING

Appendix A TRENCHING

The exploratory trenching program began in the Cedar City vicinity on December 10, 1981. Reconnaissance geologic mapping at and in the vicinity of Green's Lake Dams 1, 2, 3 and 5 prior to selection of trench sites was conducted during the four following days, and occurred intermittently at later dates as the project required. Exploratory trenches GL-2a, GL-2b, GL-2c and GL-2d investigated subsurface conditions near Green's Lake dams nos. 1 and 2; trenches GL-3a, GL-3b and GL-3c were located near dam no. 3. Trenches GL-5a and GL-5b were excavated along the margins of the Green's Lake No. 5 debris basin located west of the above sites. Trench GL-5c consisted of an existing bulldozer cut along a mapped lineament, and this exposure was included in the study.

Reconnaissance mapping in the St. George vicinity included the Gypsum Wash, Warner Draw, Stucki, Frog Hollow, and Ivins sites. Mapping commenced on January 6, 1982 in the Gypsum Wash, Warner Draw, and Stucki dam area. Prior to trenching operations, a field tour of the proposed sites was conducted on January 9 for the Bureau of Land Management in order to obtain an archaeological clearance permit.

Four exploratory trenches (G-1, 2, 3, and 4) were excavated at the Gypsum Wash site, followed by two trenches each at Warner Draw and Stucki dam; these excavations occurred between January 13 and January 20, 1982.

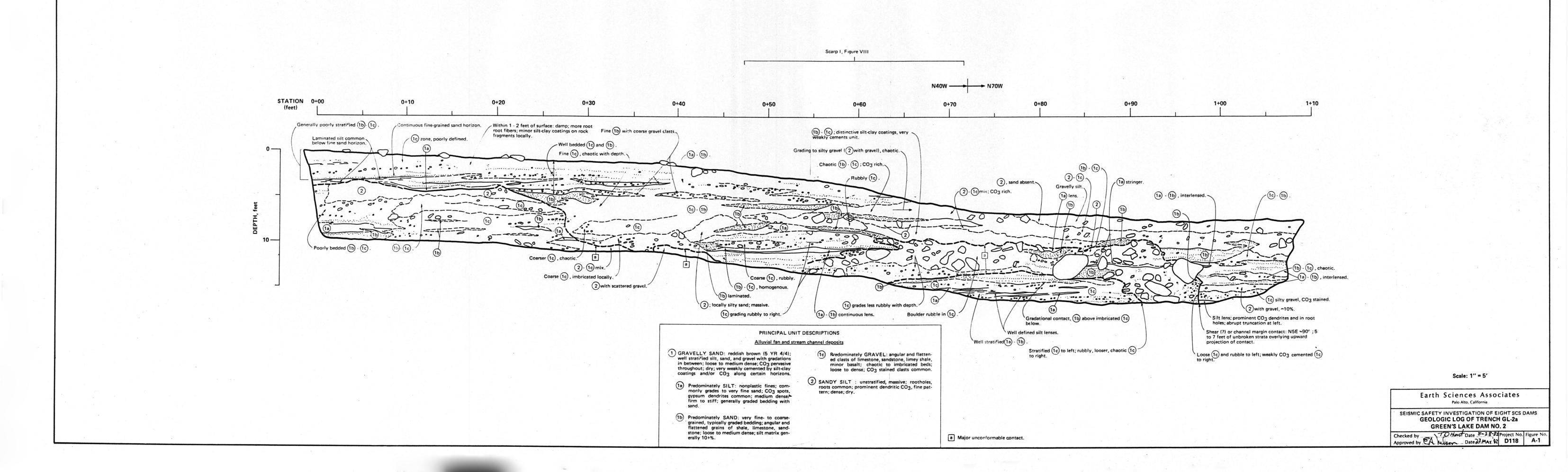
Trench exposures in the Cedar City and St. George areas which revealed faulting or suspicious fault-related features were examined by Dr. R. J. Shlemon, consulting soil stratigrapher, during the week of January 25, 1982. A field tour of these sites attended by SCS personnel, Dr. Shlemon, R. E. Anderson of the U.S. Geological Survey, Denver, Walter Arabasz of the University of Utah, in addition to ESA geologists was conducted on January 26 and 27, 1982. Following the field meeting, subsequent work was planned at the Green's Lake No. 3 site. Reconnaissance mapping of the Frog Hollow and Ivins sites and the additional study at Green's Lake No. 3 was completed by February 12, 1982.

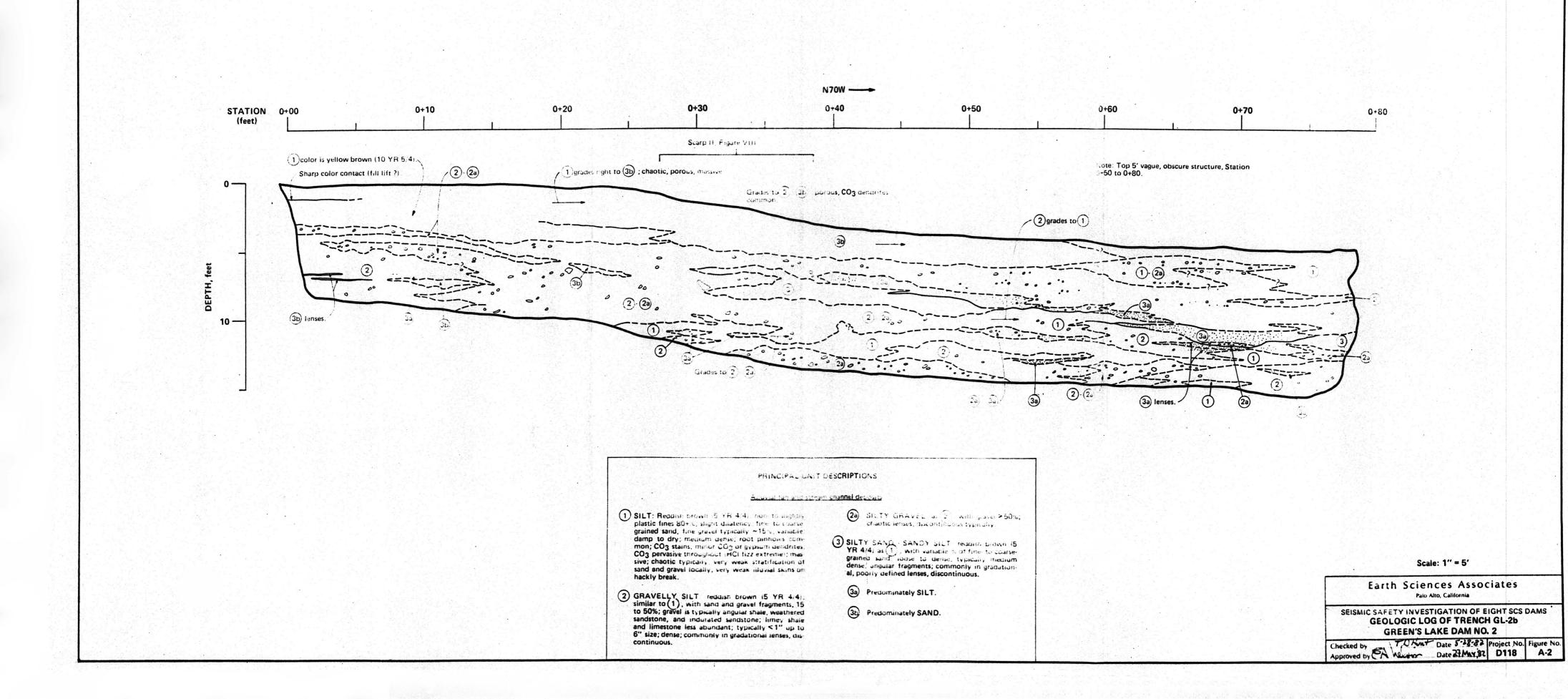
The field investigation was interrupted by bad weather throughout the term of the project.

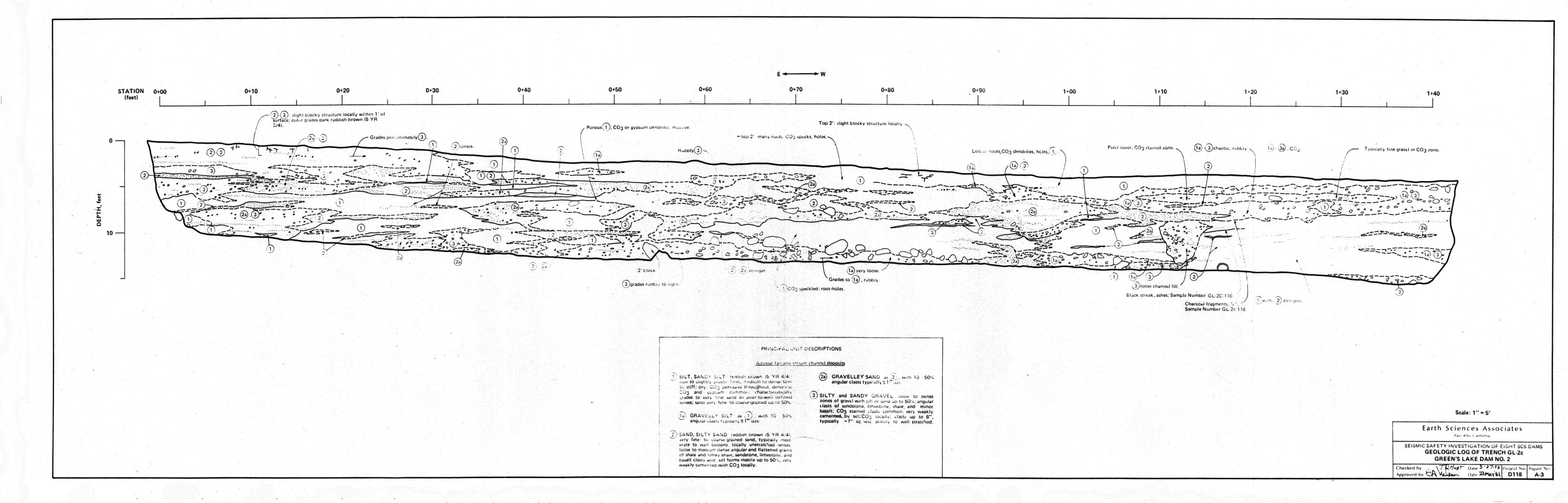
Exploratory trenches were logged in detail at a scale of 1 inch equals 5 feet, and reached depths of 14 feet below the ground surface. The trenches were supported by portable hydraulic shoring and locally by timber shoring where support was needed for extensive time periods. Barbed wire fencing provided protection for the excavations when left unattended. All trenches were backfilled by February 9, 1982.

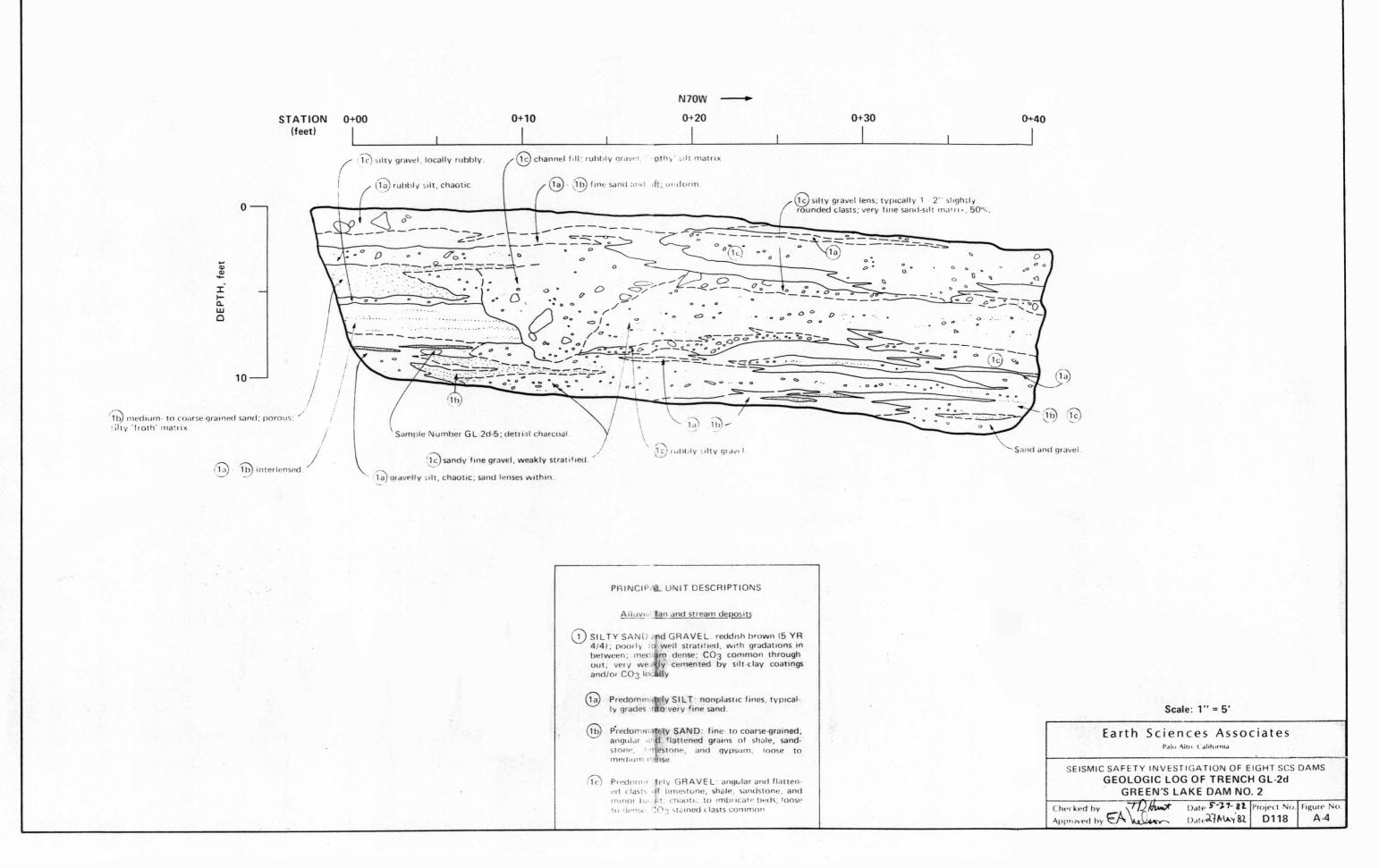
In the Cedar City area trenches, charcoal samples were obtained from key stratigraphic horizons for Carbon-14 age-dating purposes. Samples of charcoal submited for age-dating were recovered from the following sites: Trench GL-2d, Sta. 0+05; Trench GL-3a, Sta. 0+57, Sta. 1+54; Trench GL-3b, Sta. 0+42, Sta. 1+27; Trench GL-5a, Sta. 0+05, Sta. 0+10.

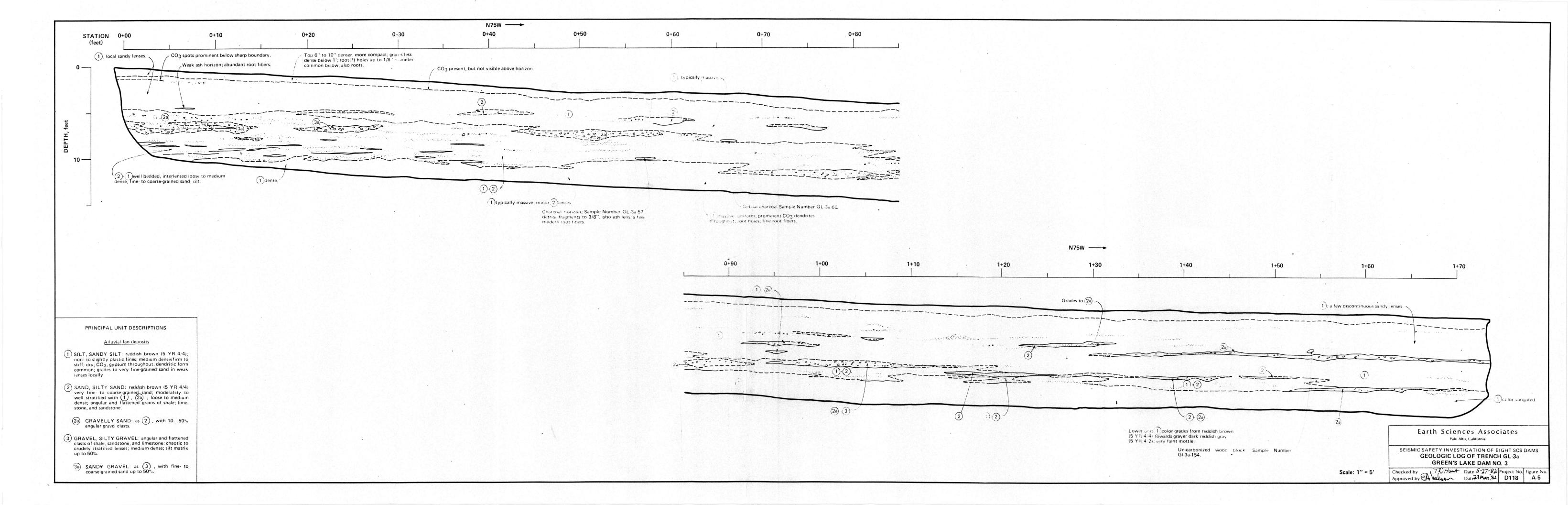
Field reconnaissance mapping, trench logging, and charcoal sample recovery was performed by T. D. Hunt of Earth Sciences Associates. Backhoe service was provided by George Ziegler and Son of Cedar City, Utah. Dr. R. J. Shlemon provided geomorphic and soil stratigraphy analysis for geologic age determinations. The charcoal samples are presently at Teledyne Isotopes (New Jersey), and Geochron, Krueger Enterprises (Cambridge, Massachusetts) laboratories for agedating analyses. The results of these tests are expected in May, 1982.

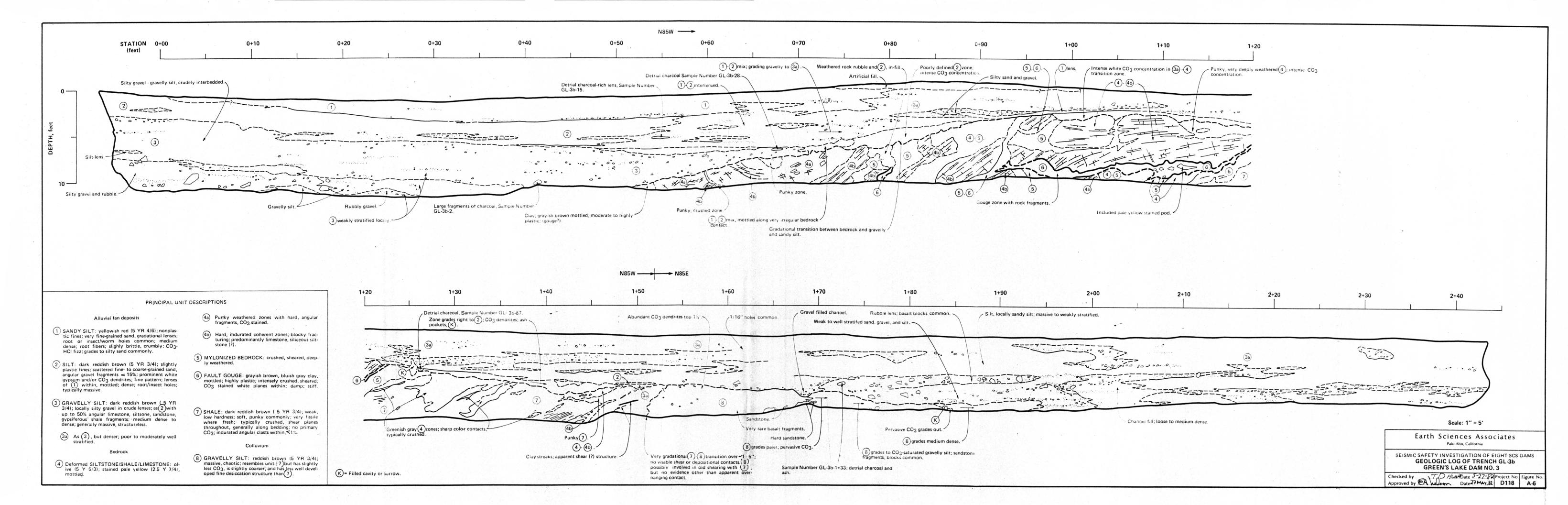


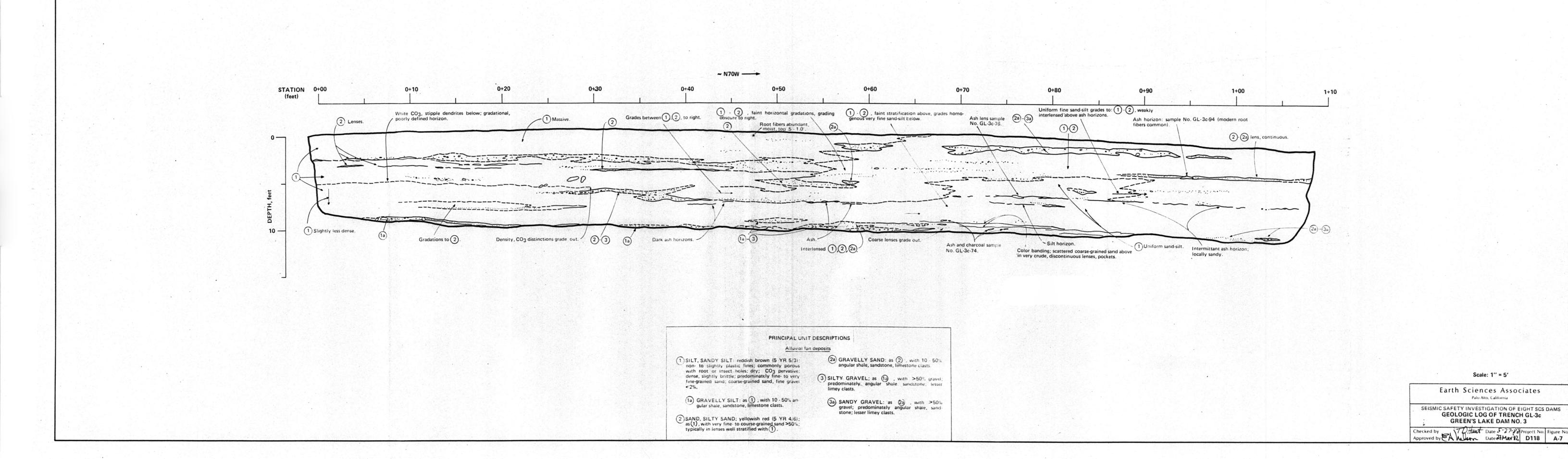


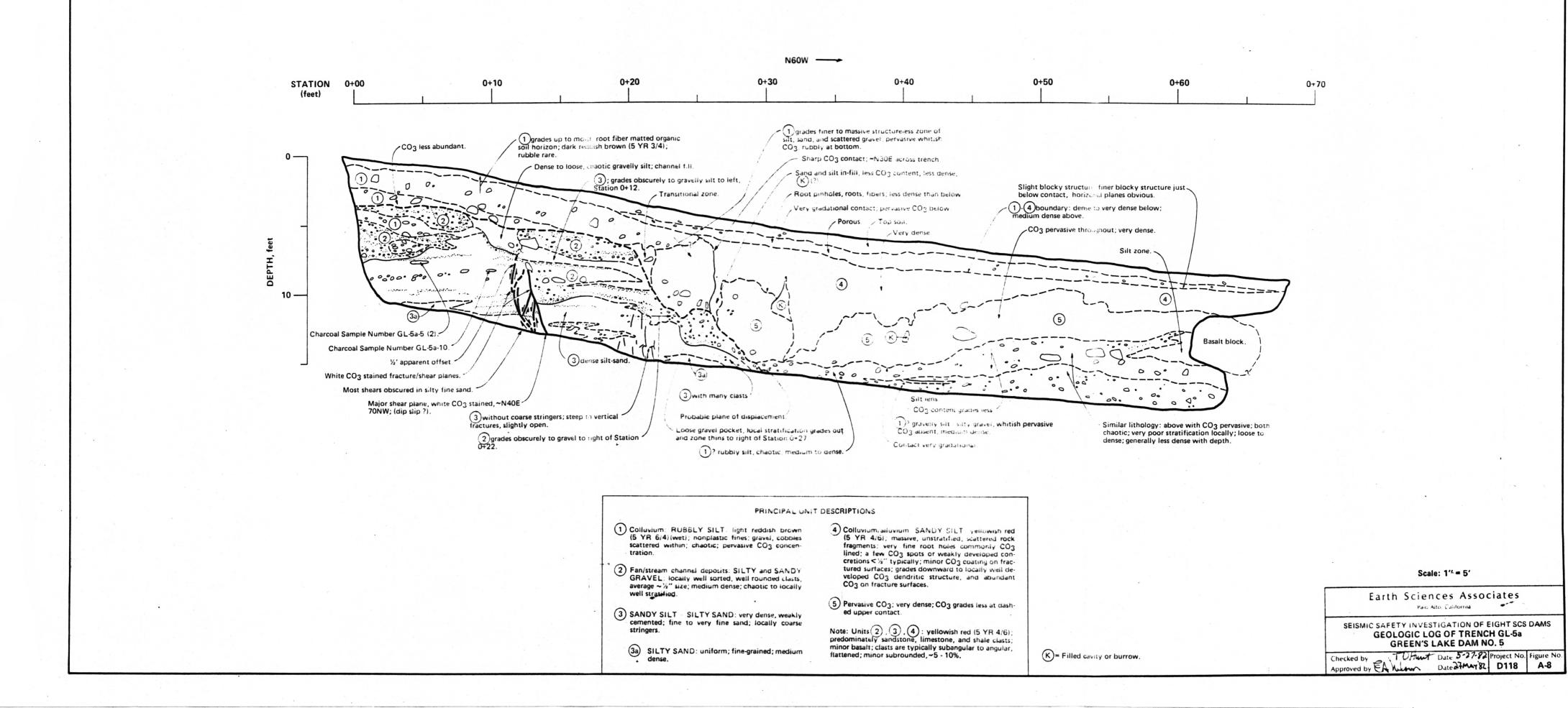


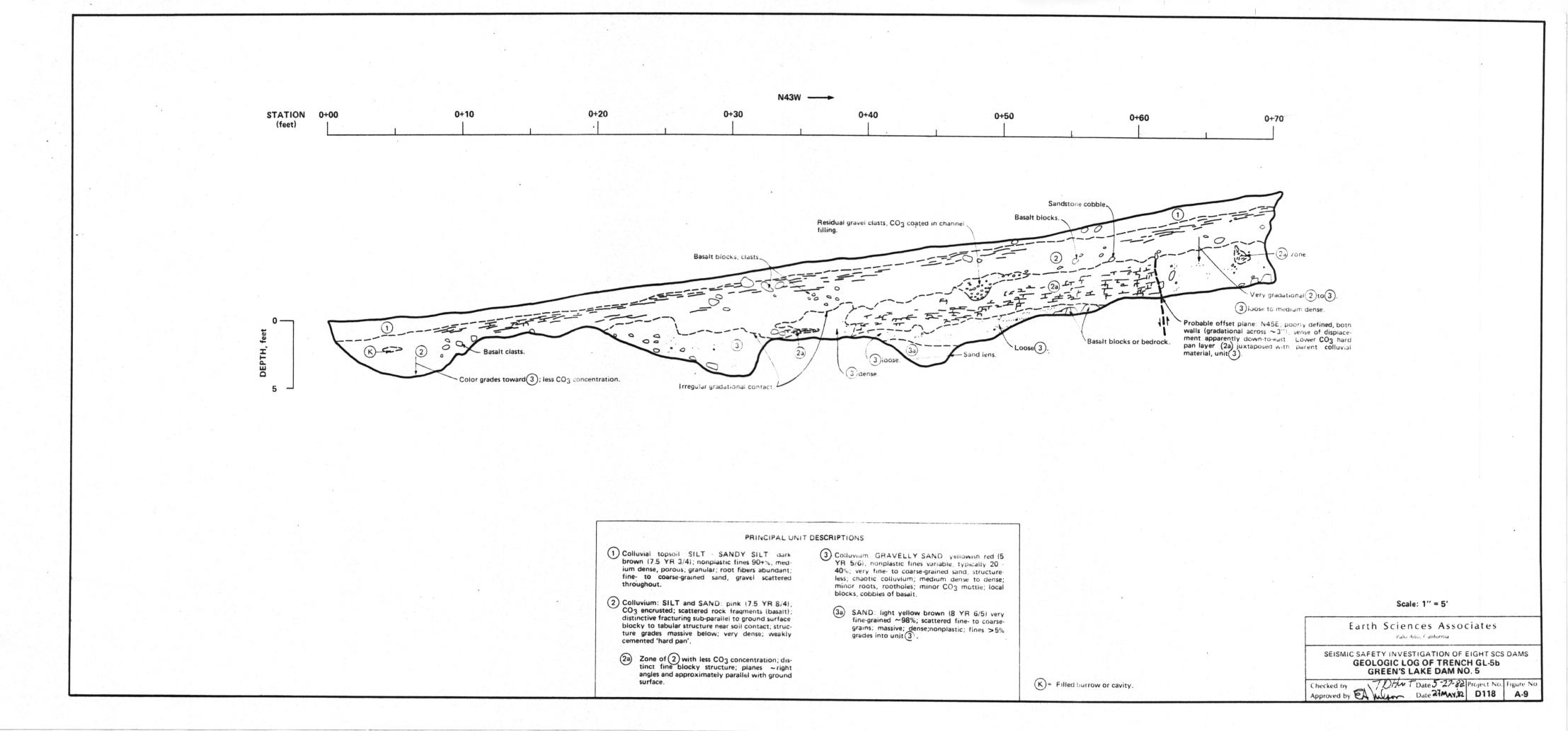


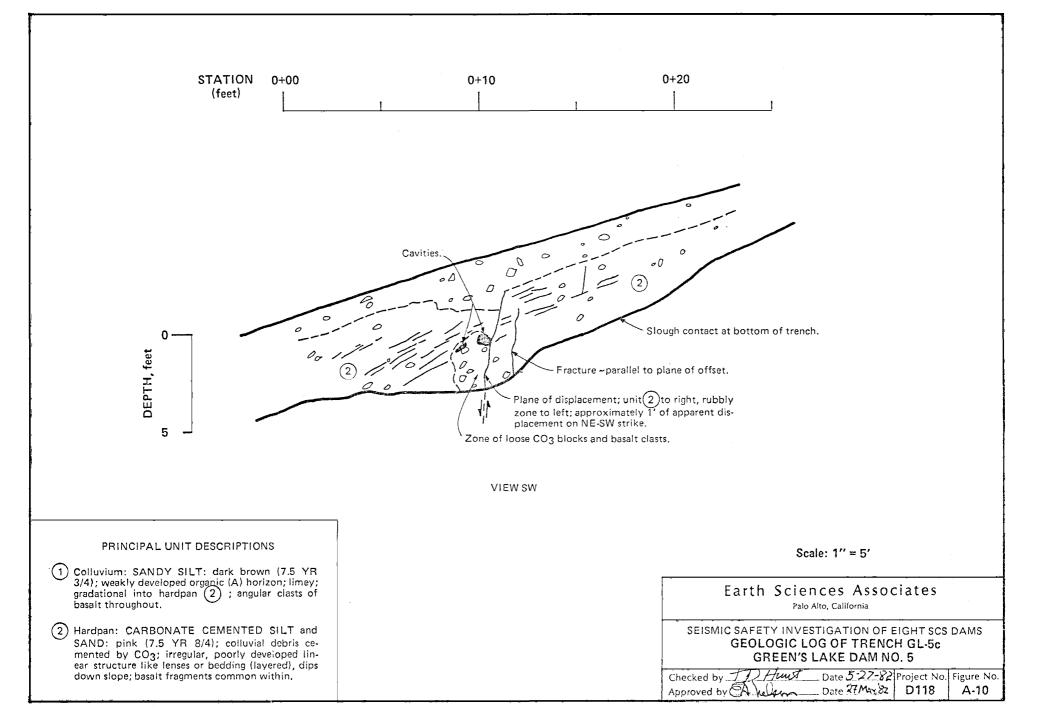


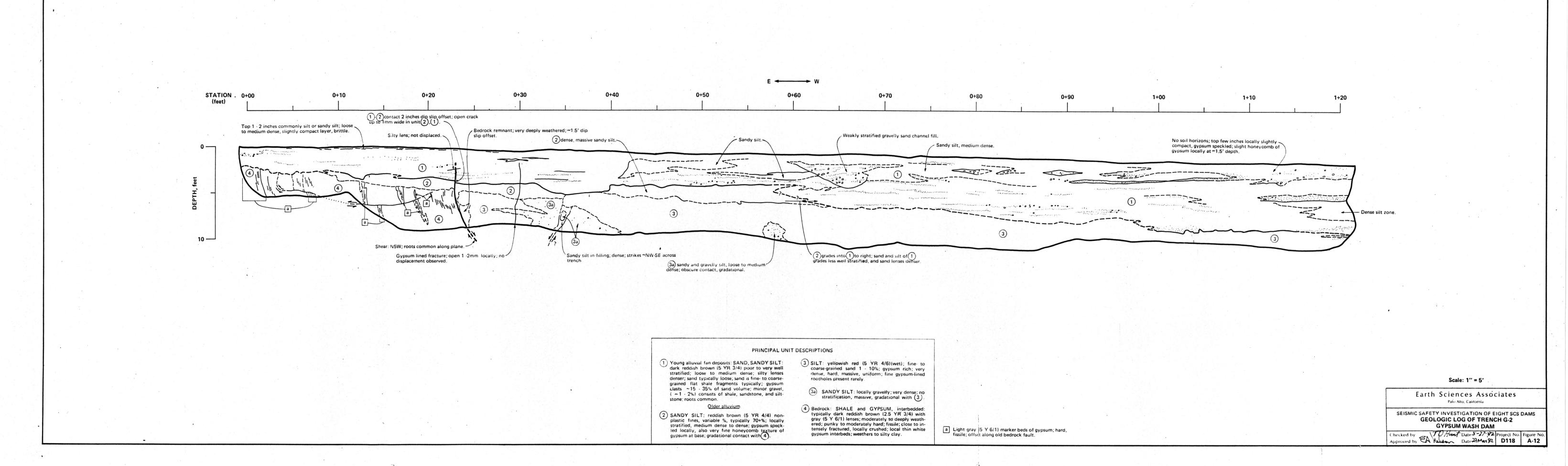


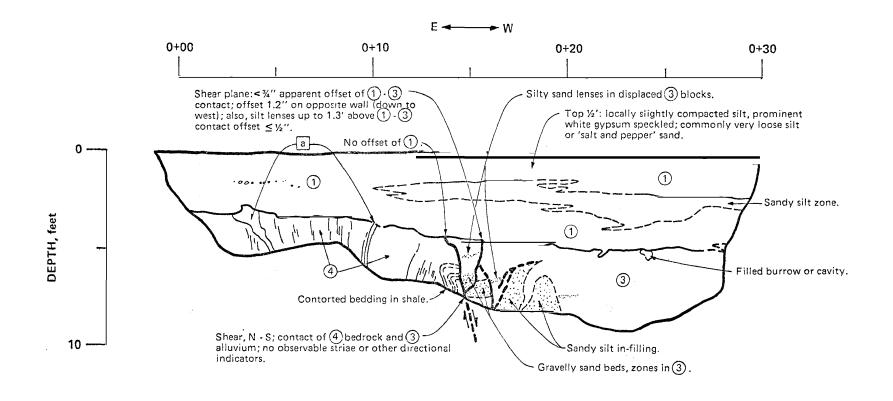












Note: Units described on trench G-2,

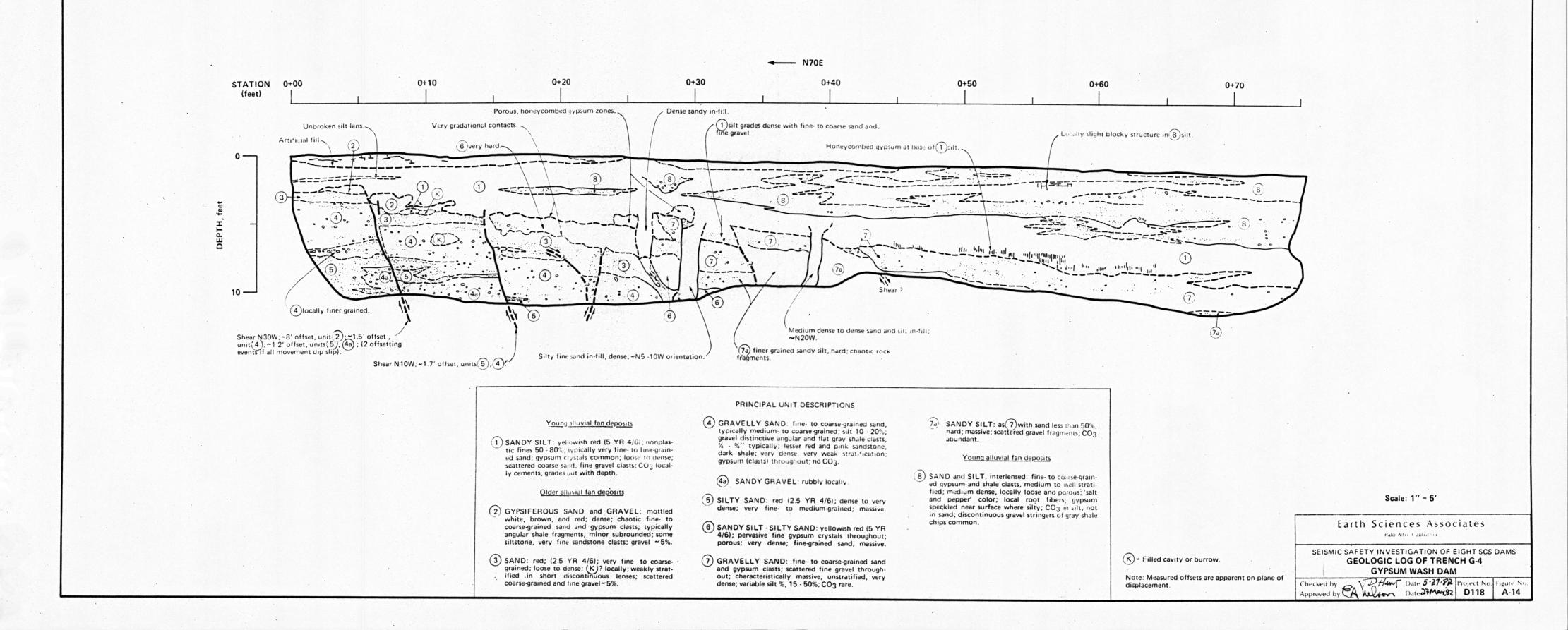
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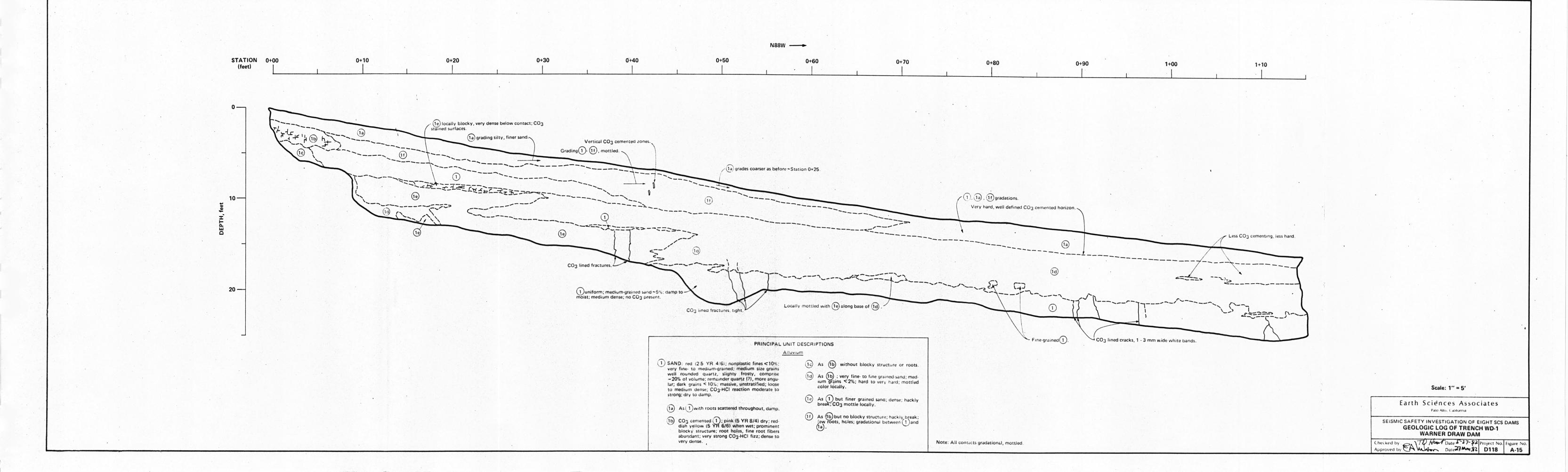
Earth Sciences Associates

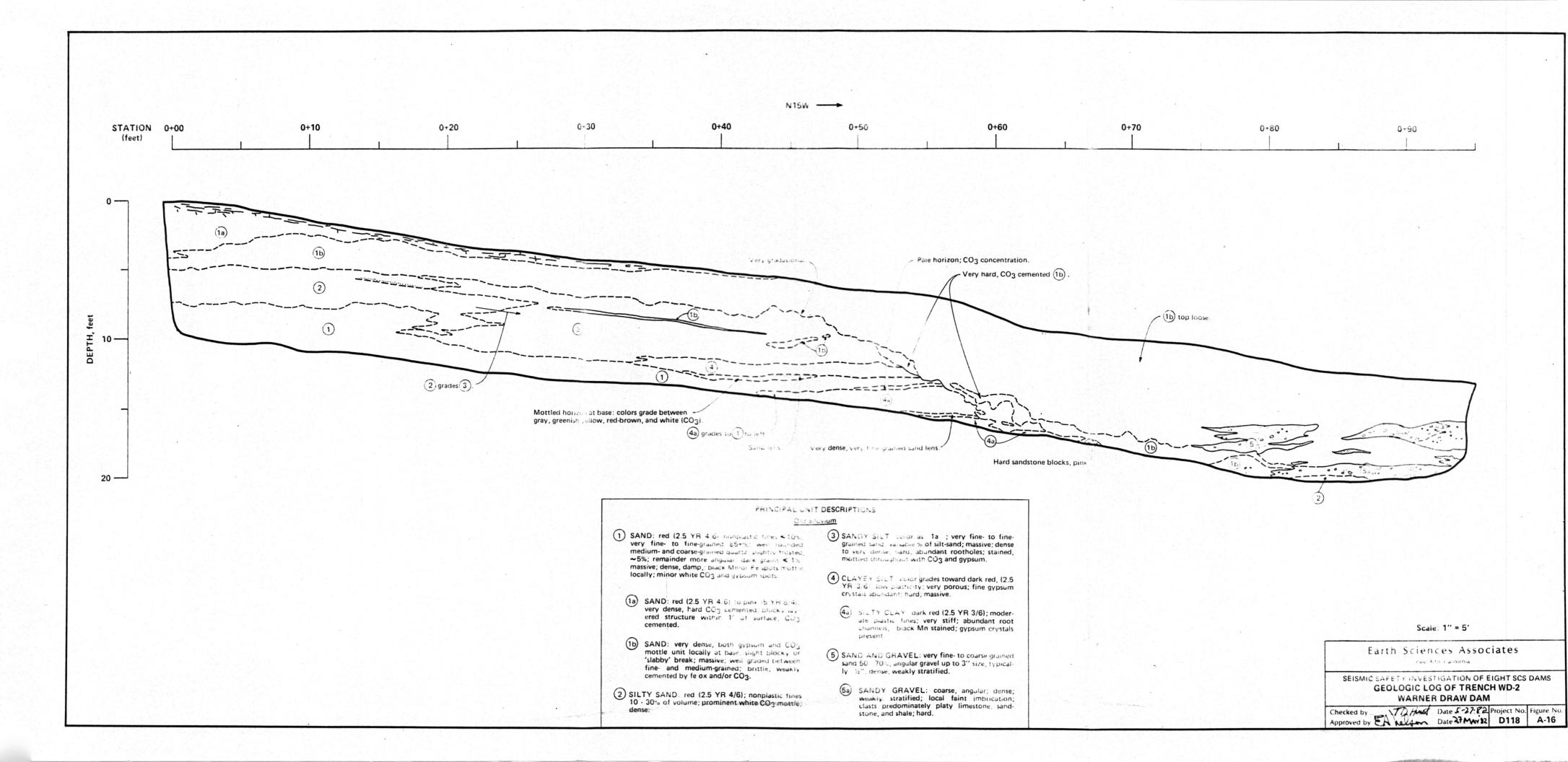
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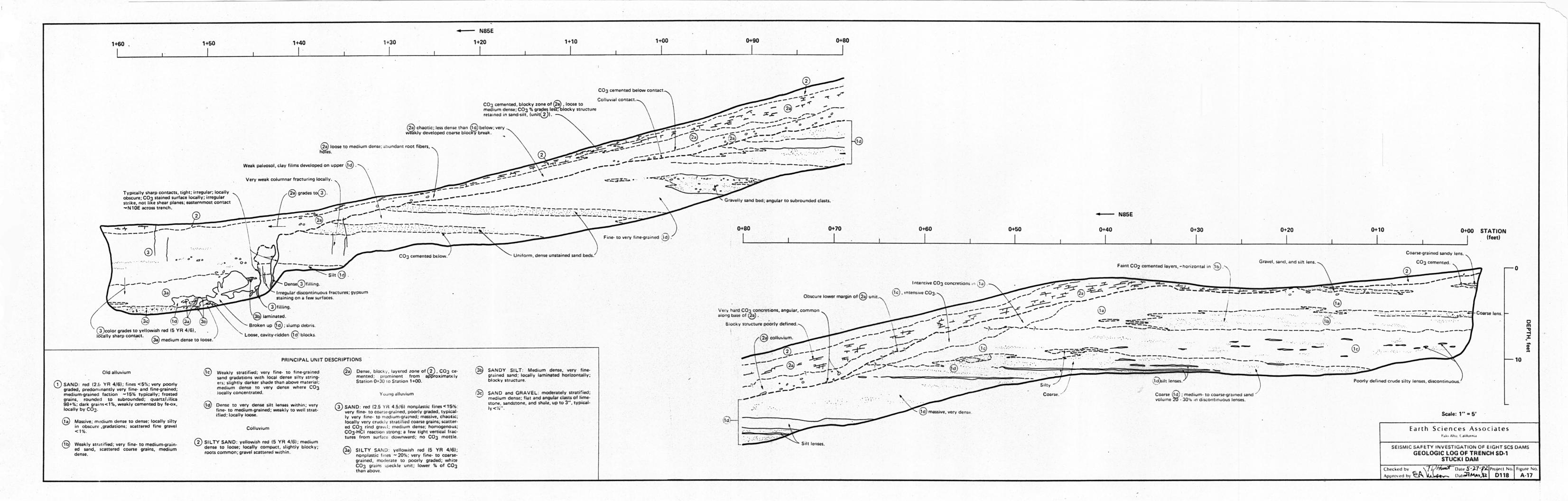
SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
GEOLOGIC LOG OF TRENCH G-3
GYPSUM WASH DAM

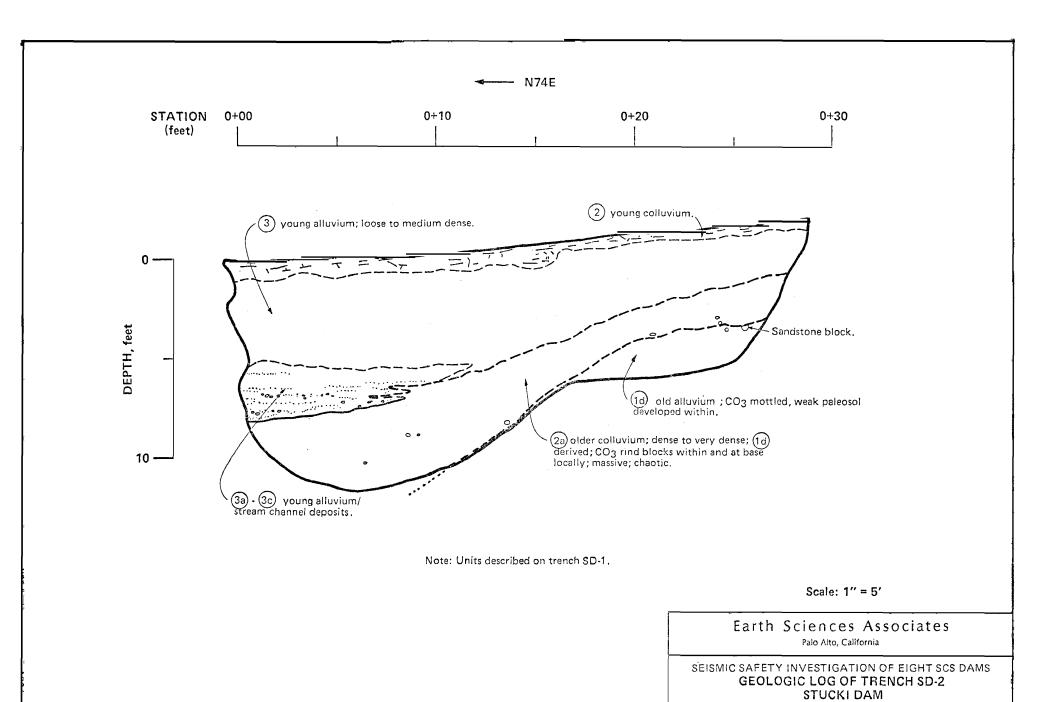
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APPENDIX B

DRILLING AND SAMPLING

Appendix B DRILLING AND SAMPLING

The geotechnical field investigations for the seismic evaluation of eight SCS dams in the Cedar City-St. George, Utah area were conducted from September 28 through October 20, 1981. The sampling and drilling phase of the investigations was performed by Douglas Yadon (Engineering Geologist) of Earth Sciences Associates. Drilling operations were subcontracted to Pitcher Drilling Company of East Palo Alto, California.

The objectives of the drilling and sampling phase of the investigations were:

- 1. To provide information on the materials comprising the reservoir embankments and their underlying foundation.
- 2. To perform Standard Penetration Tests (SPT).
- 3. To obtain disturbed and undisturbed samples for laboratory testing.

Twenty-six exploratory borings were drilled in the embankment and foundation of all eight dams totaling 1096.7 lineal feet. All borings penetrating embankment materials were drilled along the dam centerline from the crest and then into the underlying foundation soil or rock. Most of the foundation borings were located adjacent to the upstream toe of the embankment. The locations of the borings are summarized in Table B-1 and are shown in Figures B-1 through B-8.

A Failing 1500 rotary drill rig was used to advance all borings. Holes were drilled using the direct rotary method with water circulation. In most cases drilling fluid additives were unnecessary due to relatively low fluid losses. Bentonite drilling mud and/or salt-gel were added in some instances to retard fluid loss and help flush heavy gravel cuttings. Each boring was started with a 6-inch flight auger and a short section of surface casing was set and sealed with bentonite.

All of the borings were sampled as continuously as possible with a 3-inch Pitcher barrel sampler and a standard penetration split spoon sampler. The usual

procedure employed was to cut a $2\frac{1}{2}$ foot relatively undisturbed sample with the Pitcher barrel, conduct a standard penetration drive test (SPT) through the next $1\frac{1}{2}$ feet, clean out the hole with a 4-7/8-inch tricone rock bit to the bottom of the SPT, and then repeat the procedure to total depth. If significant debris remained in the hole after the Pitcher barrel sampling, the hole was cleaned with the rock-bit before conducting the SPT. Where coarse gravel and cobbles prevented undisturbed sampling or meaningful drive testing, the rock bit was used to penetrate the coarse interval. Table B-2 summarizes the drilling and sampling performed.

Details of the design and operation of the Pitcher barrel sampler are described in the accompanying manufacturer's literature. Shelby tubes recovered were processed as follows:

- o Excess drilling fluid and slough drained from top.
- Sample recovery measured.
- o Ends of sample logged.
- o Samples weighed.
- o Spacer placed on top of sample and wax-sealed.
- o Plastic caps placed on ends of tubes.
- o Caps sealed with tape and waxed.
- o Sample identification marked on tube.
- o Samples boxed and stored.

In spite of the presence of some gravel in most of the materials sampled, the overall average Pitcher barrel sample recovery for each dam site was good to very good as summarized in Table B-2.

Standard penetration tests were conducted by driving the thick-walled standard split spoon sampler with a 140-pound slide hammer falling freely through a distance of 30 inches. The number of hammer blows required to drive the sampler through three successive 6-inch intervals were recorded. The number of blows to penetrate the last 12 inches is referred to as the standard penetration resistance, N_{STD}. In some instances the sampler met refusal before being driven the full 18 inches due to the presence of gravel, cobbles, very hard soil or rock. In these instances a minimum of 50 blows was applied to the last 6-inch increment or

fraction thereof. Samples from the split spoon were logged and placed in labeled heavy-weight zip-lock plastic bags for storage.

Details of materials encountered and data on the samples recovered are recorded on the Drilling and Sampling Logs accompanying this appendix. Results of SPT tests are shown in Figures B-9 through B-21 of this appendix.

In general, the field exploration program proceeded smoothly and, with the exception of Frog Hollow Dam, conditions encountered in the field were about as anticipated from review of SCS files. Data on Frog Hollow dam available to ESA prior to field exploration were incomplete, particularly in regard to the as-built configuration of the embankment, foundation (and old embankment) and location of the old and new service outlets. This situation resulted in siting boring FH-1 on the dam crest very close to or directly on the alignment of the old outlet pipe. Severe drilling fluid loss occurred in this hole below a depth of approximately 50 feet and either a void or very soft, unrecoverable material was present from a depth of 50.5 to 54.8 feet. Two additional borings were authorized by SCS and were drilled 6.5 feet on either side of FH-1. No fluid loss or voids were encountered in these holes. This condition is currently under review by SCS and it is ESA's understanding that some additional exploration is planned.

Table B-1

Locations of Exploratory Boreholes

Dam	Boring	Station	€ Offset (ft)	Approx. Elevation (ft)
	GI A 1			6080
Green's Lake	GL2-1	7+00		6070
No. 2	GL2-2	5+75	64 u/s	6053
Green's Lake	GL3-1	11+20		6067
No. 3	GL3-2	9+45	50 u/s	6053
	GL3-2A	9+30	50 u/s	6053
Green's Lake	GL5-1	1+29		5940
No. 5	GL5-2	1+73	82 u/s	5926
Warner Draw	WD-1	15+33		2989
Warler Draw	WD-2	17+49		2989
	WD-3	15+45	210 u/s	2940
Stucki	STK-1	14+06	201 100	2814
Diucki	STK-2	17+85		2814
	STK-3	12+79	97 d/s	2782
Gypsum Wash	GW-1	33+65		2740
Gypsum wasn	GW-1A	33+55		2740
	GW-1A GW-2	20+01	68 u/s	2725
	GW-2 GW-3	26+83		2740
The authorize	DII 1	11+26		$119\frac{1}{1}$
Frog Hollow	FH-1 FH-2	10+52	106 u/s	801
		10+32 11+19	100 u/s	1191
	FH-3			1191
	FH-4	11+33		119
Ivins	IV-1	12+48		3189
Diversion	IV-2	20+00		3189
No. 5	IV-3	36+11		3189
	IV-4	19+62	43 u/s	3179
	IV-5	20+02	42 d/s	3175

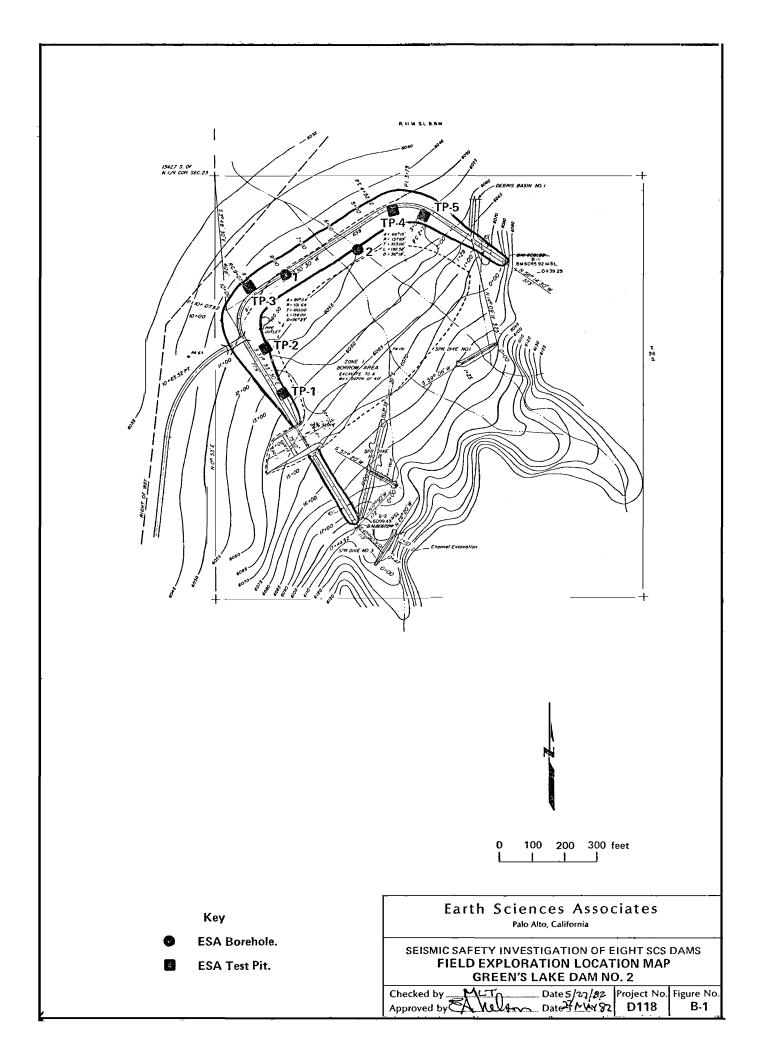
 $^{^{1}}$ arbitrary datum.

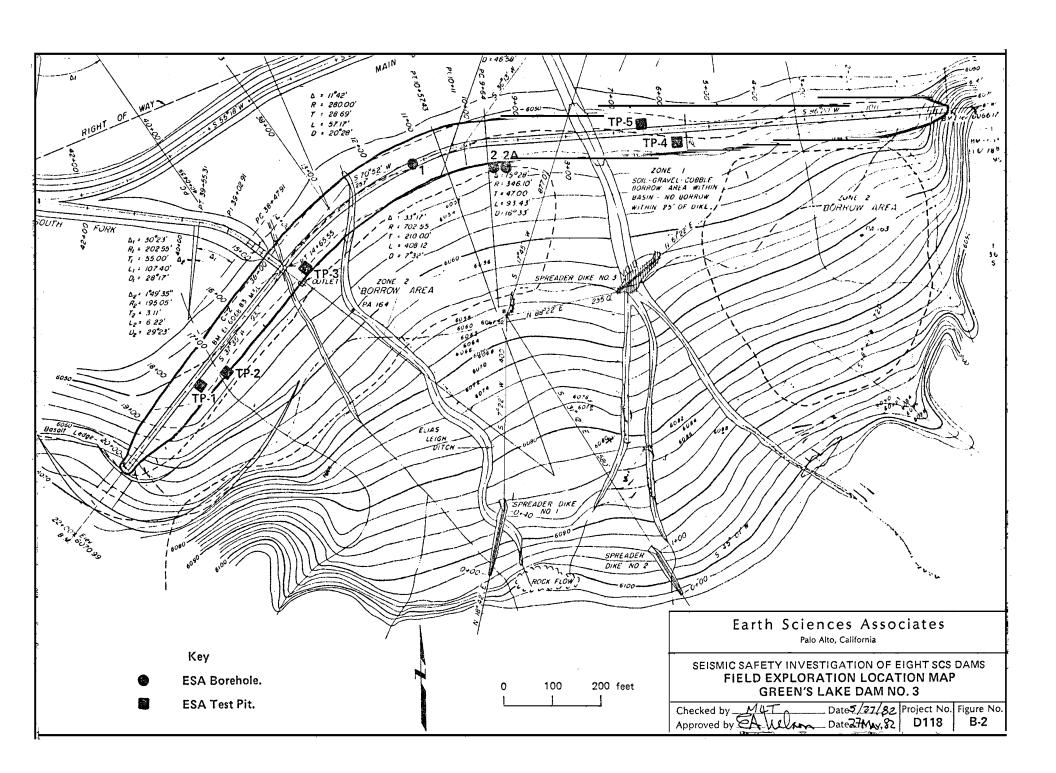
Table B-2
Summary of Drilling and Sampling

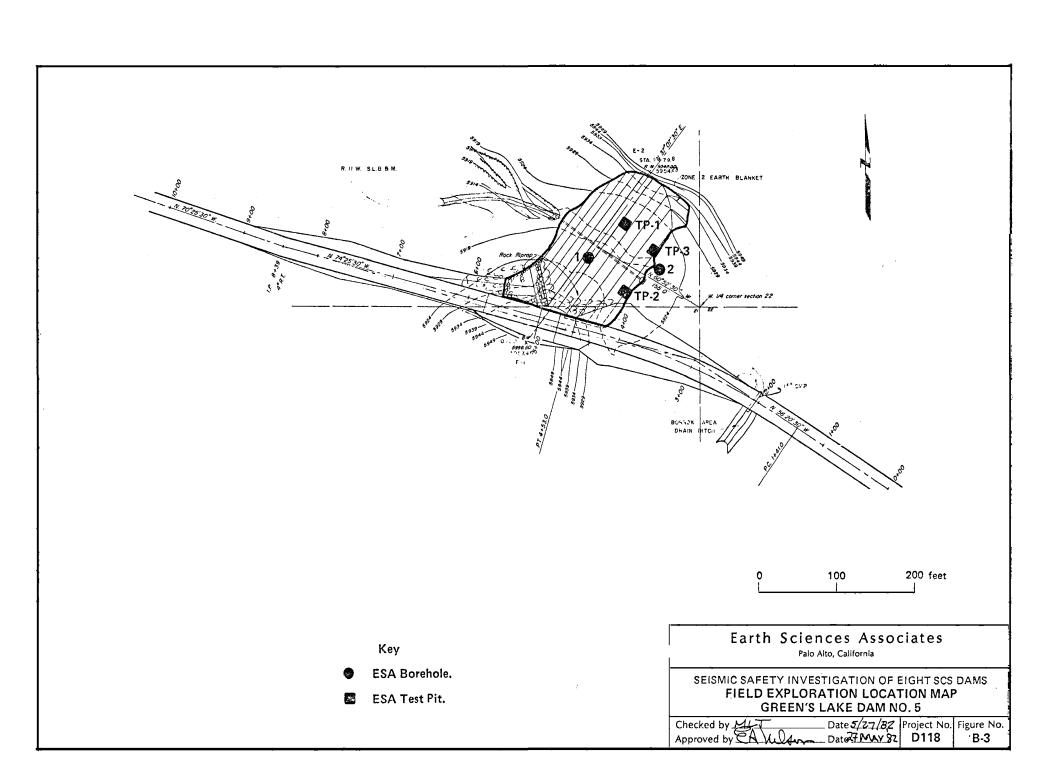
<u>Dam</u>	# of Borings	Total Footage	# PB Samples	# SPTs	AD/RD ¹ Footage	Average Recovery Percent (Includes Embankment and Foundation)
Green's Lake No. 2	2	90.0	22	24	8.9	78
Green's Lake No. 3	3	98.3	23	23	16.2	75
Green's Lake No. 5	2	85.8	18	20	24.7	77
Warner Draw	3	160.9	38	40	12.2	95
Stucki	3	189.0	42	42	31.0	89
Gypsum Wash	4	115.7	26	26	21.8	92
Frog Hollow	4	224.0	31	31	104.9	90
Ivins Diversion No. 5	<u>5</u>	<u>133.0</u>	<u>26</u>	41	8.0	78
TOTALS	26	1,096.7	226	247	227.7	

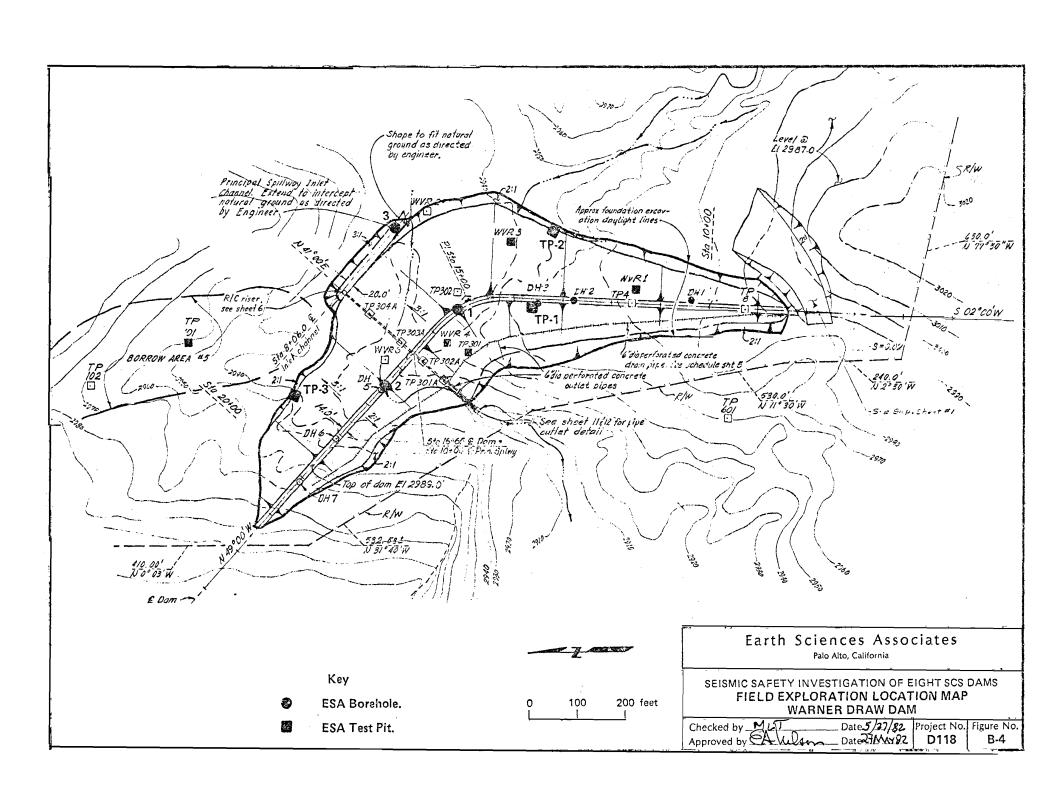
RD - Rotary drilling.

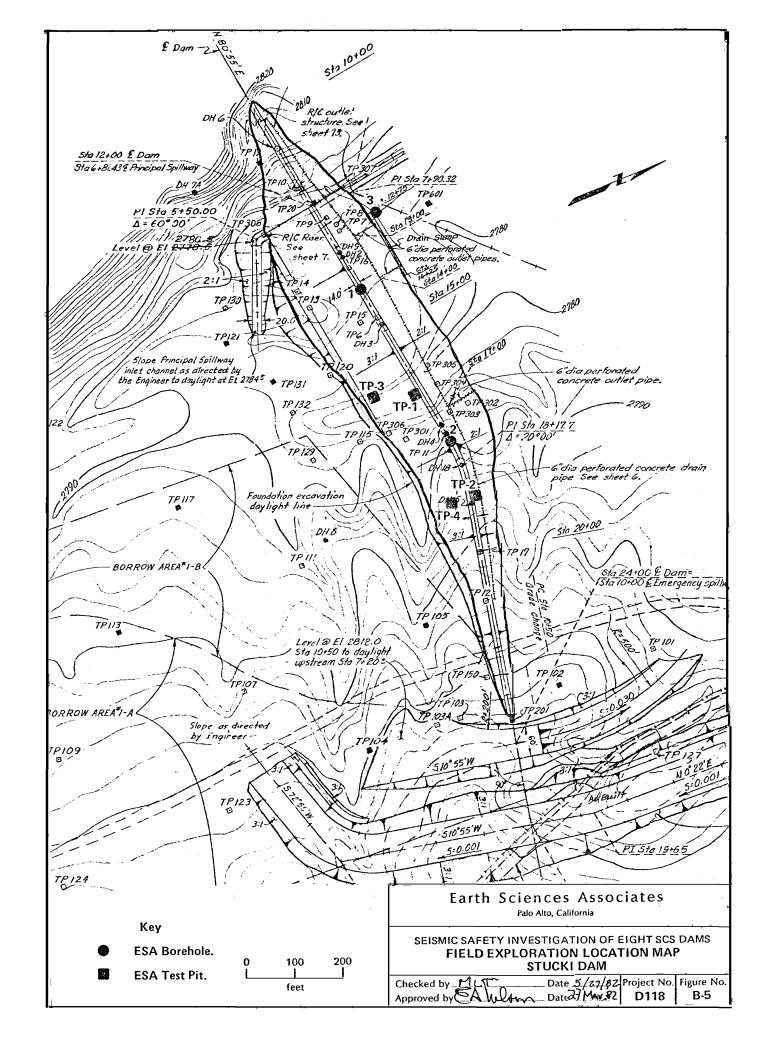
¹AD - Auger drilling.

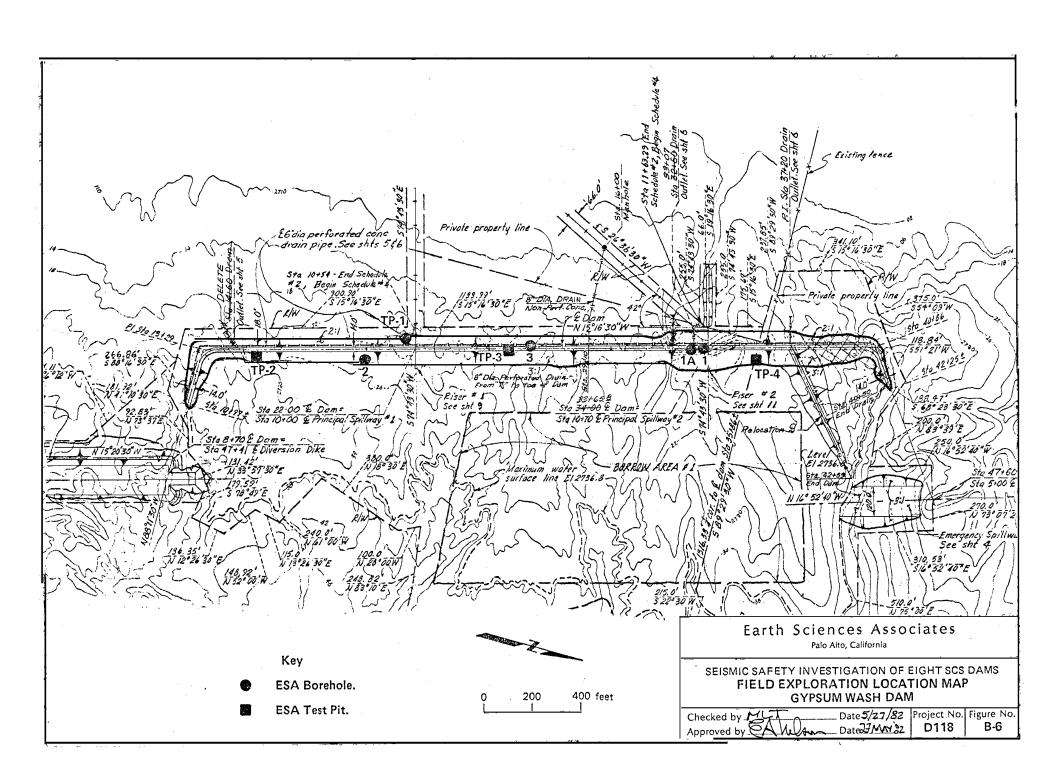


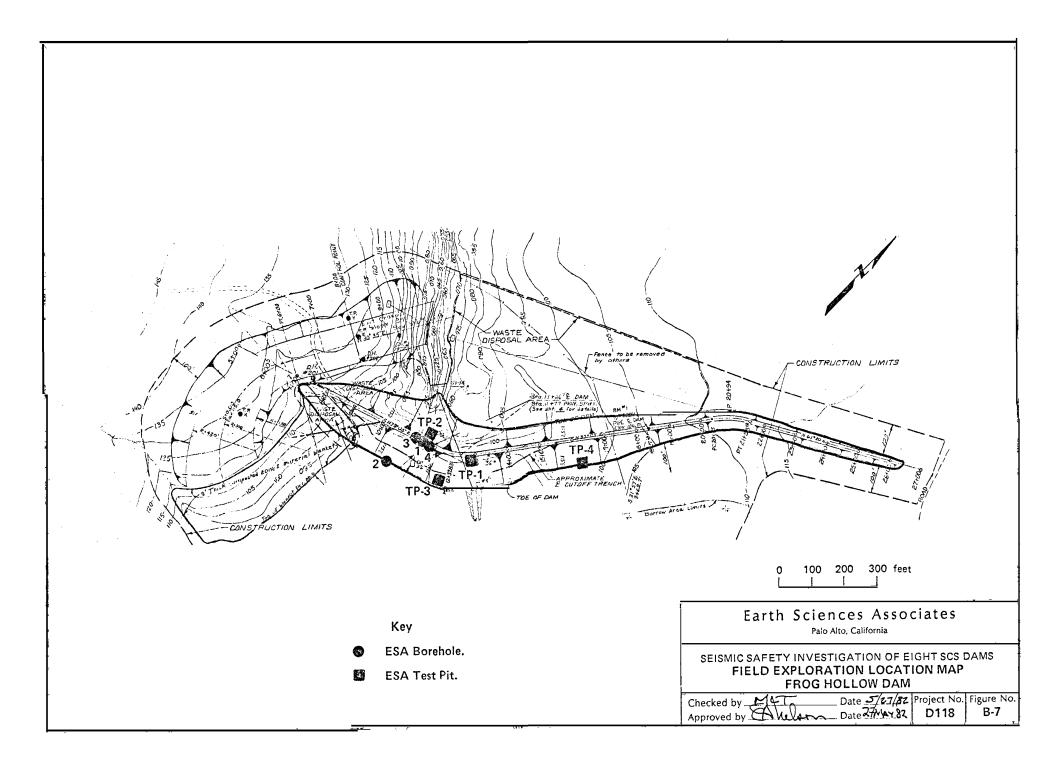


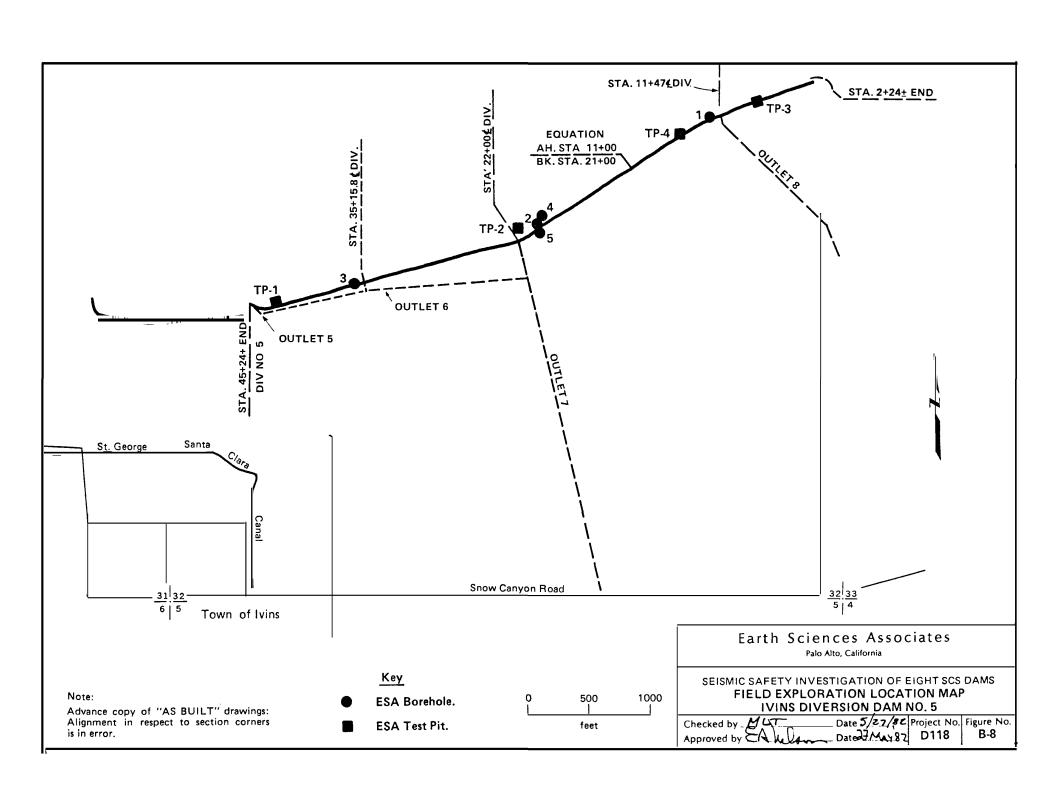


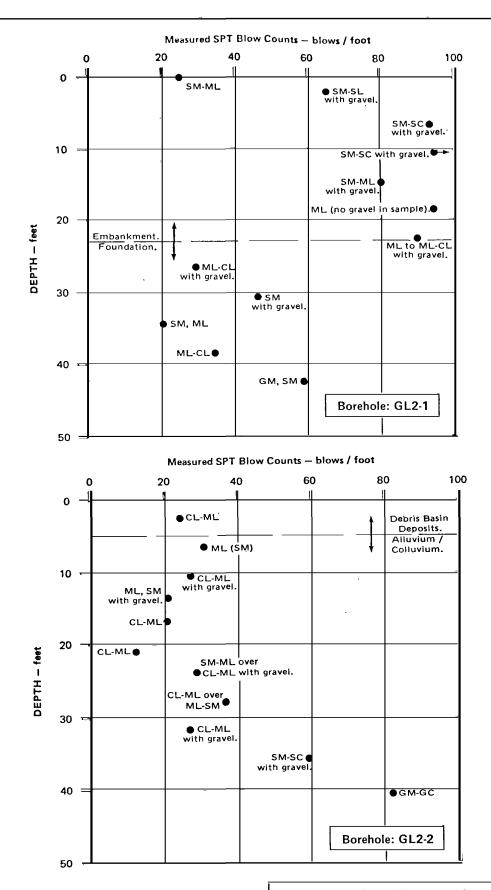










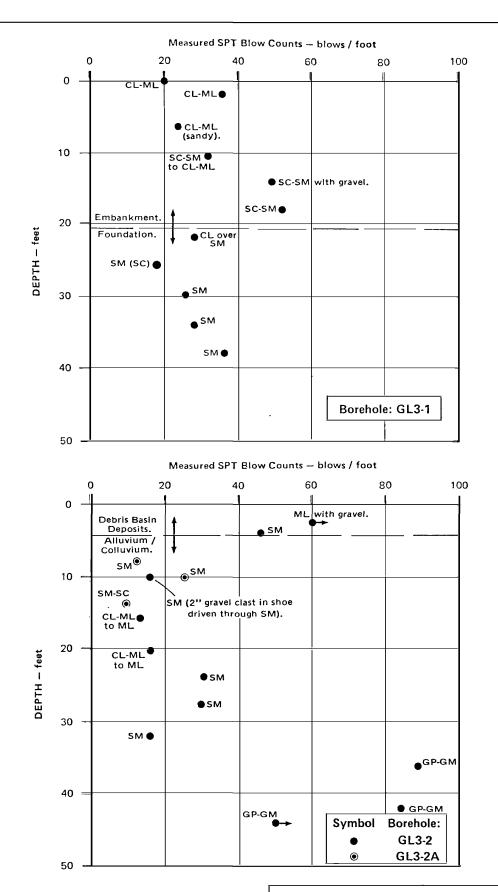


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SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS STANDARD PENETRATION TEST RESULTS GREEN'S LAKE DAM NO. 2

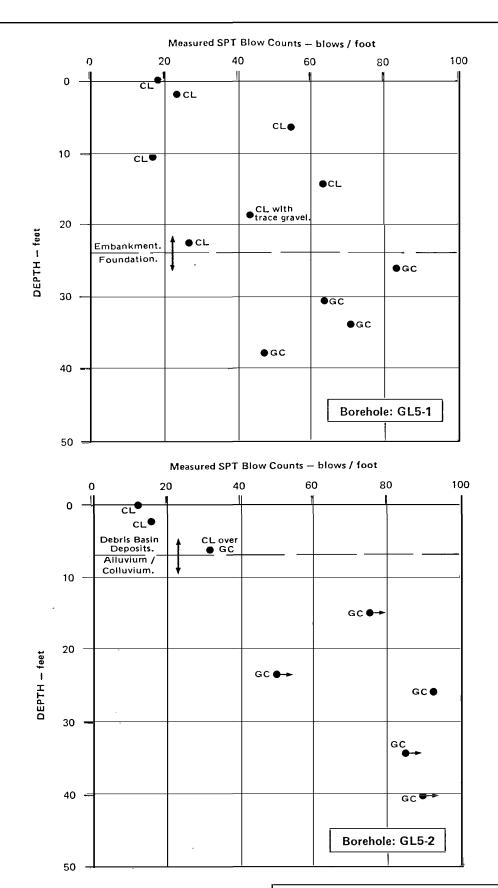
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SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
STANDARD PENETRATION TEST RESULTS
GREEN'S LAKE DAM NO. 3

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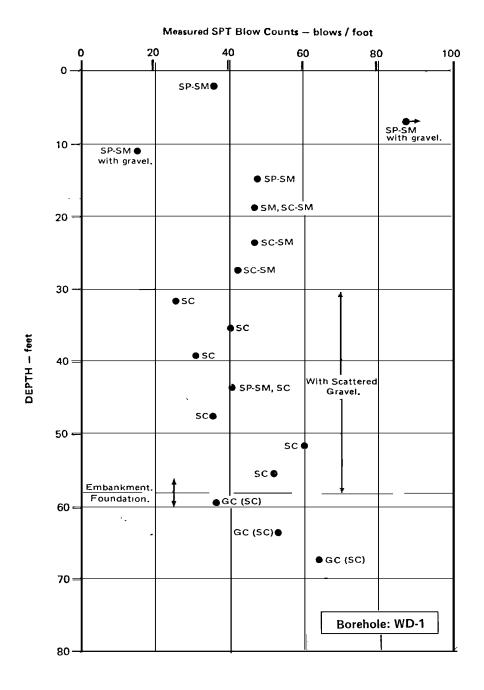


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SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
STANDARD PENETRATION TEST RESULTS
GREEN'S LAKE DAM NO. 5

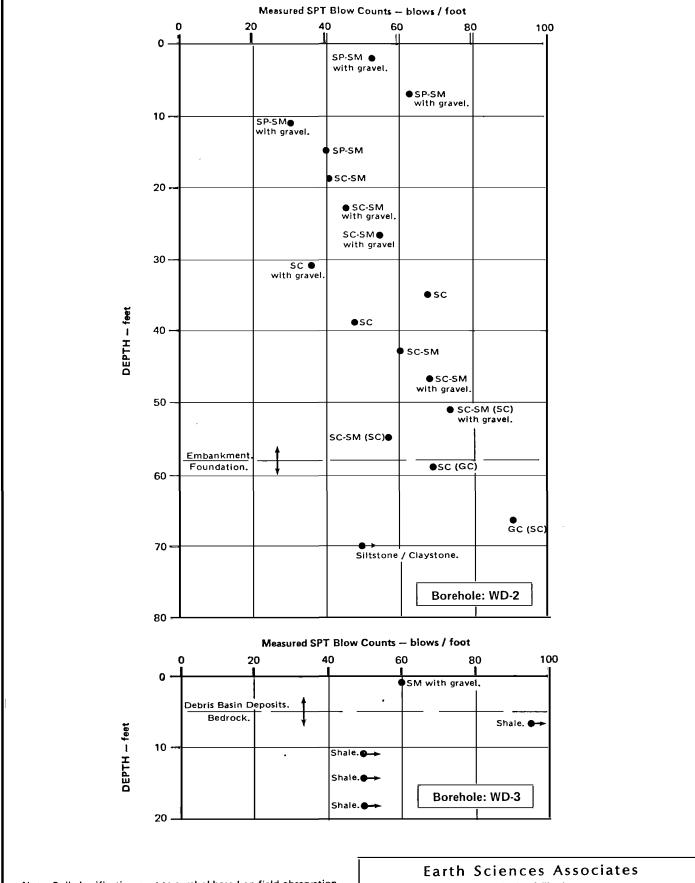
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SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
STANDARD PENETRATION TEST RESULTS
WARNER DRAW DAM

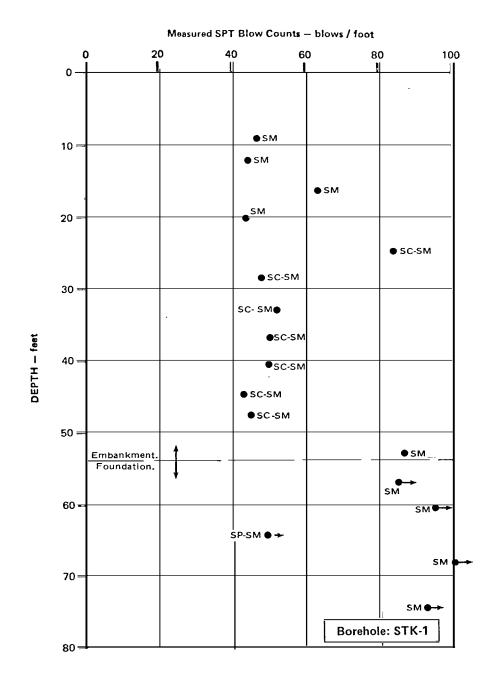
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SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
STANDARD PENETRATION TEST RESULTS
WARNER DRAW DAM

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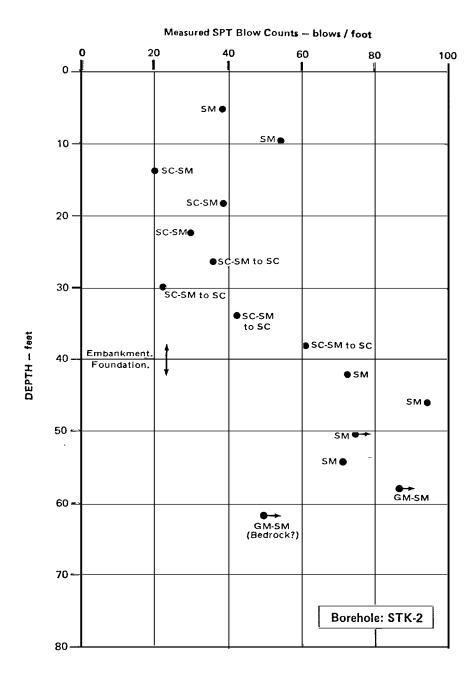


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SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
STANDARD PENETRATION TEST RESULTS
STUCKI DAM

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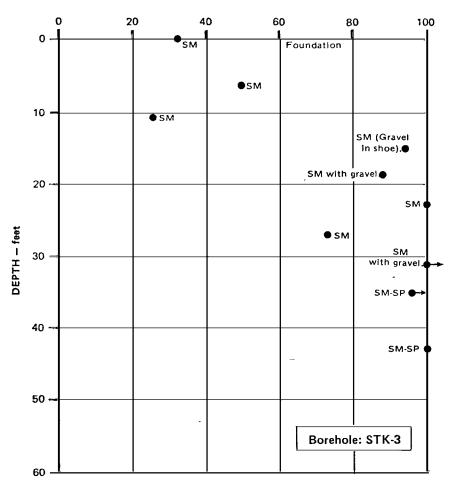
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SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
STANDARD PENETRATION TEST RESULTS
STUCKI DAM

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Measured SPT Blow Counts - blows / foot

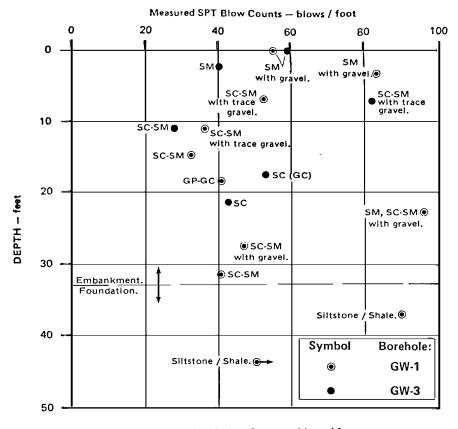


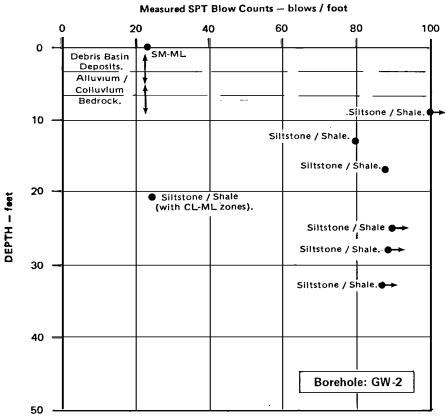
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SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
STANDARD PENETRATION TEST RESULTS
STUCKI DAM

Checked by MUT Date 5/27/82 Project No. Figure No. Approved by Date 27 May 82 D118 B-16





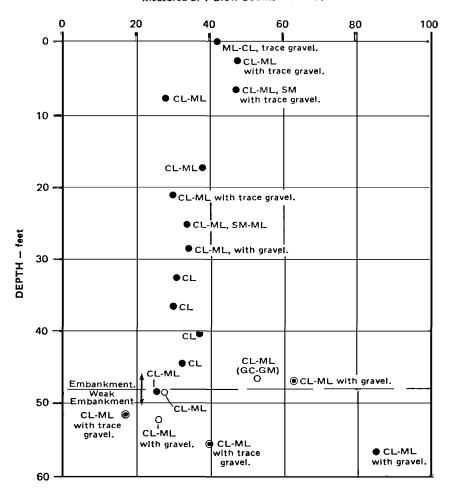
Earth Sciences Associates

Palo Alto, California

SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS STANDARD PENETRATION TEST RESULTS GYPSUM WASH DAM

Checked by MUT Date 5/27/82 Project No. Figure No. Approved by Date 3/40/87 D118 B-17

Measured SPT Blow Counts - blows / foot



Key

Symbol	Borehole:
•	FH-1
Ο	FH-3
•	FH-4

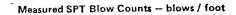
Note: Soil classification next to symbol based on field observation.

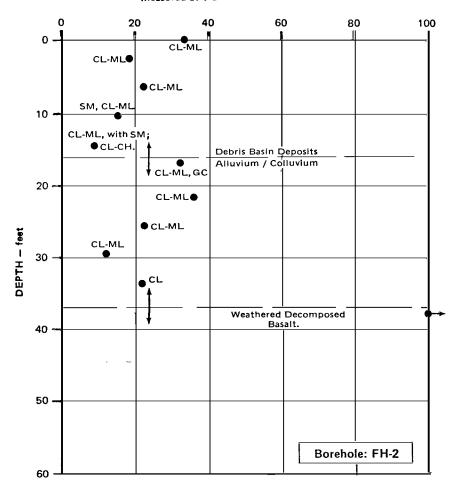
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Palo Alto, California

SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
STANDARD PENETRATION TEST RESULTS
FROG HOLLOW DAM

Checked by M4T	Date 5/	27/82	Project No.	Figure No.
Approved by August	Date 27	Mm 92	Ď118	B-18
Approved by The ROSE	Dutce			



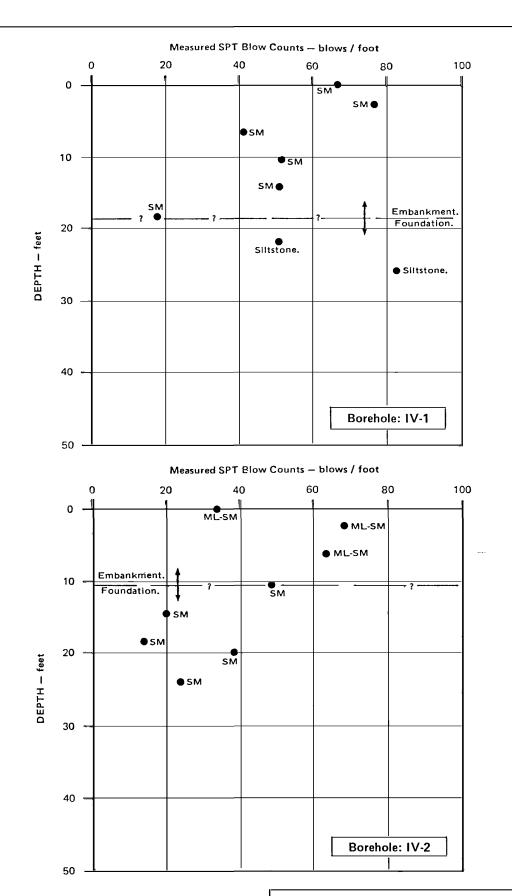


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SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS STANDARD PENETRATION TEST RESULTS FROG HOLLOW DAM

Checked by MUT	_ Date 5/27/8 2	Project No.	Figure No.
Approved by A	_ Date 5/27/82 _ Date 7 May 92	D118	B-19

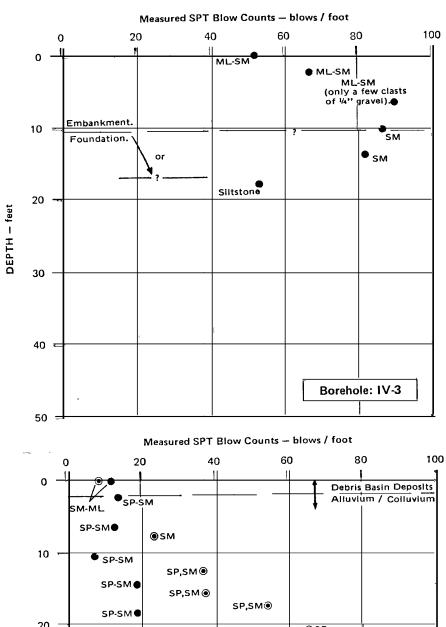


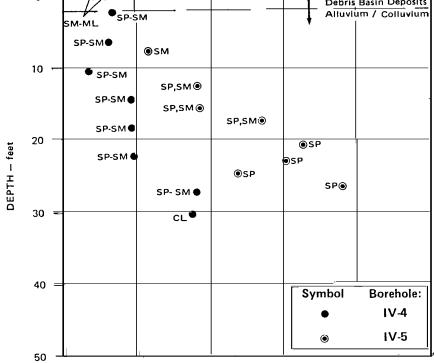
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SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
STANDARD PENETRATION TEST RESULTS
IVINS DIVERSION DAM NO. 5

Checked by M4T	Date 5/27/82	Project No.	Figure No.
Approved by	Date 27 MAY 8	D118	B-20





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SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS STANDARD PENETRATION TEST RESULTS IVINS DIVERSION DAM NO. 5

Checked by Date 5/21/82 Project No. Figure No. Approved by Date Date 7/4/82 D118 B-21

EARTH SCIENCES ASSOCIATES

DRILLING AND SAMPLING LOG

`		DRILLING AND SAM		_		
ROJECT_	<u>118,50</u>	S Dams - Wah DATE DRILLE	D 10/2-3/	<u> 181</u>	HOLE NO. 662-/	
CATION	E atc	cest, Dam #2 @ N sta. 7+00	GROUN	D SURF	ACE ELEV. ~60 10 (79%)	./
RILLING CO	ONTRAC	CTOR Pitcher Dilling Co. LOGGED BY	<u> Drill</u>	DEPTH T	ro ground water <u>aotengua</u> fall <i>140 l</i> b. 30	ye,
PE OF RIG	<u>railing</u>	1500 HOLE DIAMETER 47/8" HA	MMER WEIGH	HT AND	FALL 1 childs cont	
IRFACE CO	NDITIO	INS dirt embantment	w	EATHER.	scottend couds; cool;	
		T		 	1	11
DEPTH C	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS	36
		· · · · · · · · · · · · · · · · · · ·	<u> </u>	<u> </u>		7
0.0	SM-	EMBANEMENT FILL		SPT-1	start drilling \$5:00	
4	ML	0.0-5.0 SILTY SAND to SANDY	8-1	1.2/	Proces fathard spirt spoon 4/0.5 9/0.5 14/0.5	z
<u>‡</u>		5/LT: 1 77 1/15	<u>: </u>		170.5 170.5 -70.9	
z.0 ±	[Vary 40-60%, quick dilatancy,	<u> </u>	RD :	<u> </u>	
‡		Sand stap qualed cul 5-10 plastic. Sand stap qualed cul 5-10 to mito. L. qualed of to u, f. qualed.		SPT-Z	15/0.5 50/0.5 -gravel	را
Ī		6. 9101 ne d, 30-55% fu, f. quaised	<i>B</i> -2			
7		trace f. sondst., limest. gravel; strong rac. to Hel; dense; dry.			10/2 endshift \$ 5:45	
" a İ		30-2 C 1/1 H amy GM. SINGIT		RO	10/3 startshifta 7:30	
40 =	,	clast in shoe of split span; specks, small soft nodules of	40	PB-1	Smph w/3" Aitcher Barrel	
<u> </u>]	specks, small soft nodules of Caco-	Ĺ	' ~ '	large (~3") cobble in end	1
Ŧ-	Tem	5.0-11.5 SILTY & CLAYEY SAND:	8-3	[, , ,]	of tube - sample loose, disturbed - hot weighed	
· ^ ‡	50	; 17 nes +	·	1/18	ţ -	
6.0 		10-30%, vary non-to low plastic;		7.5	₽	
‡		sand 55-85% pad figning sand, about 15-20% of sand	;	SPT-3:	ļ.,,	
Ŧ		Is Mc. grained locally grown	B-4	1.25	33/05 43/0.5 50/0.5	19
1		varies 15-15-10, most 14-34", +	<u> </u>	1.25°	f Chammering gravely	
8.0 =		vavies 5-15%, most 1/4-74, 50m2 161/2, pAd sandst and -	- 41	PB-Z:	Chammening gavel) RO to clean past some coarse gravels	
Ŧ		mod hard clasts much dissim-	£ ``.	70-6		
Ŧ		mod hard clasts much dissem- inated CaCO- w/ from reactoticle dense to very dente moist to wet (may be due to diffl fluid).	F _ ,	-	end of tube bent over	1
Ŧ		(Mai be due to diffil fluid).	5-/	2.01	some slough on top of	
10.0		7	F	2.50	+ we = 516.130z.	
Ī]		SPT-4	Wf = 17/6.202.	
+		1 −	13-5	0.7/0.7	45/0.550/0.2	9
₽		11.5-14.4 SANDY SILT:	E	RD	(10 gravel in spoon-min)	'
12.0 +	ML	; 60-70% non-plasticting;	-4z	<u> </u>	FROTO IN TO CHESA HOLE	
ŧ		mun a sand than above unit as	f	PB-3	of large cobble-use	
-		Iscattend grains; Ass to no gravel: 1	-	} -	30 psi down-pressure	1
‡	13	compactiess encertain; moist tower (may be due to drift floid).	: <i>S</i> -2	1.6,	ensoftwhe bout - 3"	
14-0	10	grades to -	<u> </u>	24	quarel slowsh (2) an top we = 516 1302.	Ì
‡-		14.4-18.81 SANDY SILT & SILTY		SPT-5.	wf=1516840z.	
+	₩ .	SANO:	B-6	1-9/10-	30/05 50/05 Residen	ہے
‡	SM- ML	1			cobbe or gravel	-
16.0 +	100	sand poorly to skip graded to 4	43		NO to 16' to dean hole	
†		140-70% locally, NS-10% 15m-	'5	PB-4	end of the dented	
• 🛨		c. grained trace to few of fim. gravel, some black baralt, some	F / 3	4	We= 516. 1202.	
ŧ		mugh dissem. CACO; dense to	5-3	2.54	Wf=19166/202	
18.0		IV. dence - moist (dill fluid?)	F	7.5	 -	
‡	\ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \	v. dents - moist (drill fluid?) -	8-7		E	
+	MI	1/8.8-22.8 SANOY SILT 7, SILT: 1		SPT-6:	35/0.5 45/0.5 57/0.45 (no grove) in sample)	19
‡	1016	mod. reldish brown (10R,416); 25-909; non-plastic fires, goick dilation tough ness; 20-1596 phd. v.ff. sand	B-7A	1.3/1.45	I cutter / of 7	1
20.0 f		tough ness; 20-1595 pmd. v.tt. sand t		1073	SHEET / OF Z]

ROF

DEPTH	CL	ASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS	Ī
Z0.0	<u> </u>				. MODE	KEMAKKS	4
	ŧ	ML	1 1000 011 -C. 3001 G 90 3011 (C) 7 - ~ //3	; =	PB-5	I end of tube hadly bent	١
•	‡		no grove, cacos as soft, white pod			= We=516.12/202.	
	Ŧ		From 16 to 12", some distan. Caco.	3;= <i>S-4</i>	2.28,	= wf=not weighed	
22.0-	‡		clightly moist.	[[2.5		
· · · · · · · · · · · · · · · · · · ·	<u> </u>		ALLUVIUM		3PT-7	-	١
<u> </u>	-	ML	22.8-30.5' SANDY to ERAVELLY CAY-SILT: MOTHED W/	F B-8	0.8/1.0	35/0.5 55/0.5 Rayal	
, ,	=	CL	; finer 55-	. [= on gravel clast -ro to clean out hole	
24.0	= _		55%, vary from v. low-low plast	5-45			١
	=	6C	75-35 % mod well graded sand for fi-c. grained commonly 10-20%	%F	176-6	end of tube dented	١
	-	(5)	Fm. gravel, clasts 14-1 "Sonds	ا ہے ۔ ا		we = 516.1174 oz.	١
26.0	_		stat, limest, clasts mod. deeply	'E 5-5	2.30	wf=1916.90z.	١
	<u> </u>	ļ <u> </u> ļ	weath, subang-subrad-mod- strong reac. to Hel, some nodu	Æ l	12.5		١
‡	_	M-	of catos; may be some gypsom,	Ě	SPT-8		١
		1	dense; moist.	E 8-9	1.1/5	19/0.5 14/0.5 15/0.5	١
28.0-		w/ grovel	24.0-26.5 may grade to >50 gravel in this introd (chec	% <u>F</u>	1.5		1
Ī	=	17000	gravel in this inflival (chec 5-5); offermise as for first	<u> </u>	273-7	- 	
	<u>-</u>		of unit	‡	'~ '	most of samp addipped from tube; Amajning	ᆀ
‡	=			B-10	0.3	E sample Ant cheirland	
30.0-	_			+ " "	1/2 =	- due to disturbance	۱
1			30.5-34.5 SILTY SAND W/	‡ ——			۱
<u> </u>	_	1 7 1	GRAVEL:	· ‡	SPT-9-	21/0.5 23/0.5 23/0,5	.
32-0 -	-	w/ pmie/	30-40 % low-non-plastic fines;	‡ <i>B-1</i> /	0.9/	2" gravel clast inshoe	
32-0-	 -	[""]	anaded to stip-graded w/most	47		- ,	ľ
<u>‡</u>	<u>:</u>		F grained; 5-15% F-in group!	I	PE-8	end of the dented	
7	:		mist sandst or apprificacles' state out distinctive gray color. strong Hel tear. to dissen, and pro	<i>i</i> 主	1 7	w= 516.113/4 0Z.	
% 0-	<u> </u>		module cacar gypson as white-	\frac{1}{2} \simeq \frac{1}{2}	1.69	wf=16/6.121/202	
<u> </u>	<u>:</u>		gray subory Mads - becomes soft,	Ŧ 1	2.5		ᆡ
	-	SM.	pasty in water, no sear, to Heli-dense to u. dence; slightly morst.	1	587-10		
-	•	ML	34.5-38.5 SANDY SICT 40 SICTY	‡ <i>B-1</i> Z	1/1/	8/2-01	
<i>₹</i> .0-‡	-		SAND:	-	11.5	8/0.5 9/0.5 11/05	ŀ
‡			Fines 40-6040, U-low-non-plastic;	1 48	P8-9-	end tob at.	.
	<u>-</u>		35-55% pred. f. grained sand, w/ some 111c. sand; 5-10%	4	' -	of sample slipping out	
• ‡	•		Scatted and f. quivel of sandit	,] 5-7	1.45	-lough on top	
æ.o -}	_		PIRCHORE; Cacos and ayp. as in party moist.	<u>-</u>	25	- We = 516 12/4 12.	
Ē	•	MI	38.5-41.0(?) SANOY CLAY-CILT:	<u></u>		mf = 13/6/0/202.	
=	-	<i>/</i> /	similar to time-grained portroll	<u> </u>	597-11=	_	- 1
,]		I I	of 22.8-30.5'above; varuing it apparent in finer analysis w/some	8-13	1.3/	12/0.5 14/0.5 20/0.5	
40.0			dombr. amanic-nich lauers: Cacos =	- 10			
3			as small pods and mothing iscomments		P3-10=	end of the basty bent	
3		l · . I	Slightly moist.	ㅋ 1		We = 5/6. 12/4 02.	
73 7	_	kw, l	41.0(2)-44.0 INTERBEDGED SUTY	5-8	7-22	wf=17/6 63/402.	
42.0		kW	SAND and SILTY, SANDY GRAVEL:	=	2.5	- 7702.	
E			30.5-34.5 but w/quave 20-609	<u> </u>	\$07-17		
」」当			30.5-34 5 but w/qrave/20-609 qpacks one limestone, some sandst; tense-u derse(s); slightly moist;	3-14	0.9/	16/0.5 22/0.5 31/as	ŀ
44.0=	_	<u> </u>	B.H. @ 44.0	디 ' ㅣ	0.9/=	SHEET 2 OF Z	
					7	ruff, days, best siles w	1

DRILLING AND SAMPLING LOG

PROJECT <u>0/18 SCS Dams - Utah</u> DATE DRILLED 10/3-4/81 HOLE NO. 622-2

LOCATION <u>Up of Ham the Dam #2, NSta. 5+75;64 U/S</u> GROUND SURFACE ELEV. N6053(topo.)

DRILLING CONTRACTOR <u>Pitcher Drilling Co.</u> LOGGED BY <u>DMY</u> DEPTH TO GROUND WATER <u>not encountered</u>

TYPE OF RIG <u>Failing 1500</u>HOLE DIAMETER <u>47/8</u> HAMMER WEIGHT AND FALL <u>140 16s</u>; 30

SURFACE CONDITIONS <u>brushy flat</u> WEATHER <u>ceathered</u> <u>clouds</u>; cool

DEDTU	CL ASS	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS	N
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE			flows Ft.
0.0	CL- ML	BASIN DEPOSITS 0.0-5.0 SANDY CLAYEY SILT:	50	57-1	Pushed shelby tube (3") Sample loose in tube	
		dark eddish orn. (5483/4) trines	5-1	2.55	Un = 516. 12/4 02. Uf= 1616. 15/2 02.	
2.0		mod quick dilatany, low- no toughness; sand 35-45 %,		2.5 SPT-1	Dove standard splitspoon	
		mod well graded v-f-m., anly scattered & grains; quevel to	8-/	0.7	10/05 11/05 13/05	
4.0		basalt, s. Hist.; strong reac. to Hel; some gyp- as white-gray	51	PB-/	Sample w/3" litelar Barrel tube end bent, small wid a bottom, sample 100 se in tube	2
6.0	ML (SM)	Dasalt, s.H.t.; strong Rac. to HCl; some gyp. as white-gray now ker to 116 "; dense; dry to slightly damp ALLUVIUM 5.0-9.0 SANDY SILT to	5-2	2.49/	- wf=1816.640z.	
_		silty SANO dt. Rd. bm(5183/4); similar to above except fines are 40-60%, slightly	<i>B</i> -2	SPT-Z	16/0.5 15/0.5 16/0.5	. 2/
8.0	-	test plastic; may be introduced	52	1.0/ /1.5 PB-Z	<u>.</u>	
	?	9.0-12.5 SANOP CLAYEP SILT: similar to 0.0-5.0, but may have slightly less frond	5-3	1.61	tube v. badly bent on end; some slouph on top we = 516.151202.	
10.0		in 5-12 5 as above hutuit	-	25 5PT-3	- wf = 14/b. 23/402.	•
12.0	ML (w/	946 as 1/16 " nodules common:	<i>B-3</i>	11.5	6/0.5 13/0.5 14/0.5	
	SM- ML	12.5-17.0 INTERBEODED SANDY SILT, SILTY SAND, GRAVELLY - SANDY SILT: FINES AR All fow-	53	PB-3 0.0/1.0	-RD widing bit, then will tricone bit to push 19-cobble into sidewall -PB-3, run on gobble-	<u>/</u>
14.0	- ML,	non-plastic, goick dilatanoli	B-4	SPT-4	end bady bont no	
	ML	parall common - aun present:	-		<u>-</u> .	
16.0		med dense moist to wet may be due to drill fluid: 12.5-13.9 quiels to 1/2" w/	- B-5	1.43	end of tube badly deak sample not weighed	
	CC- ML	15% of total 12-74" of	P.	SPT-5	9/ 9/ -/	
18.0		sitty sand and sardy sitt	B-6	' ' +	= 90.5 9/0.5 12/0.5	2/
20.0		17.0-33.2' SANOY to GRAVELLY CLAYEY SILT: relish bin(54R4/3);	5-4	0.36,	saturated, disturbed SHEET _ OF 3	1

-						- -
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS	
20.0	ML		5-4 (con'4)	PB-5: (con'f)	, , , , , , , , , , , , , , , , , , ,	
22.0 -		fines vary 250-65%, locally may be 45-50%, low plosticity, mod-goick dilatancy, low-no tousiness;	<i>B-</i> 7	SPT-6:	5/0.5 6/0.5 6/0.5	12
22.0		graded, fc. grained, probly,	56	PB-6	10/3 end shiff @ 5:30 10/4 startshift@ 7:45 186 on gavels/cobbles	
24-0 -	ML (w/ grave)	more f-m. than c.; quall varies 5-20%, pred. f-m. clasts from 14-1", some c clasts, small		RD	Ro to za' to pass quavels/cobbles	
		some sitet.; gareially strong reac.	8-8	SDT-7	20/25-15/25 14/25	29
26.0	ML	nodules and mottles common med dence dence shift v. stiff : slightly moist.	57	PB-7	end of tube badly tom,	
20.0	SM- MC (W/)	22.5-24.5 MOR grove/4 to 20-25% W/most clasts 2</td <td>5-5</td> <td>1-18,</td> <td>may be some slough' of top = 516. 15/202.</td> <td></td>	5-5	1-18,	may be some slough' of top = 516. 15/202.	
29.0 -	90(R)	locally may grade to sanditaries - diphty less plostic - 26.0-27.8' similar to 22.5-		RD	RD to clean hole to 20'	
	M1- 5M	24.5' 28.4-29.1' quades to soundy silt-sifty sound con up to	B-10	SPT-8	11/0.5 16/0.5 21/0.5	37
30.0	(w/ (m/ growe)	Fines slightly less plustic	58	_	end o.k. good sample	
		29.1-33.2 gravel much common to 10-15-00; cruse approx.	5-6	1-60/	$w_e = 516.15\% \text{ oz.}$	
32.0		honzontal straffication; thin strugers, motified zones of Cacos, ayp. fairly common	B-11		10/0.5 12/0.5 15/0.5	27
		33.2-46.0 INTERESTORSO CLAYEY	59	15	- 10/0.5 12/0.5 15/0.5 CSPOON confained a stoff c.sand, f.gravel slough)	
340-	6M- 6C 40	SICTY GRAVEL and CLAYEY SICTY - SAND dr. relish bourn (548314); -	<u>:</u> :	0.0/1.0	no recovery	
	SM-	tines Vary 20-40%, low-v. low plasticity, mod. guick dilatancy, low toughness; sand vanes from	<u>:</u>	RO	RD in growelly unity - Dushing quiels/cobble - From above?	
36.0		15-20% in quivelly sections to so-60% in sonly sections, -	13-12	SPT-10	- 20/0.5 24/0.5 25/0.5	က်က
		5-15% in sandy sections, 50-60-	<u> </u>	1.5	collection of spoon	
39.0		1/2"-3" with most 1/2-1" clasts are pred sandst, limestone out - some busait; cacoz throughout uf	<u>:</u>	1 7	AD thro qually section	
		dense -u. dense (?); slightly mout.		PB-10	tube bert on end, small void send of sample	
40.0	- 6M-	35.5-27.0 crode, approx. honcontal bidding o.1-0.2 thick apparent in variations	-	507-11	wf=1016.140z.	
	50 (SM-	of grain size in sand and	B-13	0.3/0.8	32/0.5 50/0.3 (Rfusal) (spoon w/o.7'c.sand, f. govel slough)	<i>8</i> 2+
42.0	- (sc)	40.3-46.0 grades predominally of less sandy	-	I '`` - ∓	Ro thru gravelly rection	
44.0		43.5 Rovered 2/2" Ad linethre	61 B-14	PB-11 =	only Acovered 21/2"cobble SHEET Z OF 3	
	<u> </u>		<u> </u>	<u> </u>		

PROJECT DILA, SCS Dams - Utah DATE DRILLED 10/3-4/8/ HOLE NO. GLZ-Z

DEPTH	CLA	\$5.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS
44.0		6M-	33.7-46.0' INTERESCUED CLAYEY SICTY SAND: (CON'T)		SP7-12º	RO to clean co6465/ gravel from hole
			SANU : (CON 1)	B-15	0.8/	17/0.5 25/0.5 44/0.5
46.0-			B. H. @ 46.0'			Terminated hole @46' w/sufficient data
-			-	-		Backfilled hole w/ mud and cultings
	-		_	-		
-	-		-			
-	_		<u>-</u>			
-	-		<u>-</u>			
			<u>-</u>	-		
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-	-		<u> </u>	<u>-</u>		
	ŧ			-		SHEET 3 OF 3

DRILLING AND SAMPLING LOG

PROJECT DIPS, SCS Dams-Utah DATE DRILLED 10/1-2/81 HOLE NO. 623-1

LOCATION CHEST of Dam #3 @ N Sta. 11+20 on & GROUND SURFACE ELEV. 16067 (topo.)

DRILLING CONTRACTOR Pitcher Drilling Co. LOGGED BY DMY DEPTH TO GROUND WATER not encountered type of RIG Failing 1500 HOLE DIAMETER 478" HAMMER WEIGHT AND FALL 140 16. 30"

SURFACE CONDITIONS dirt embankment WEATHER Statked clouds, sprinkles occ.,

	CONDITIO	NS OFF EMPONICMENT	V	VEATHER:	COOL
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS
0.0	CL-	EMBANKMENT FILL	- .	SPT-1:	Once standard split spoon
	ML	0.0-5.0 CLAY-SILT: dark . frayish brown (10 112 4/z); fines	<u> </u>	1.0/1.5	5/25 9/0.5 1/0.5 2
z.o -	_	125-920/2 /our south's mode out		AD :	Auger culs "Flight auger
	-	dilatancy, low toughness; sand- 10-15%, poorly quaded, most fv.f. quined; trace fine qual, scattered throughout; strong	<i>B-</i> 2	1.2/1.5	16/25 17/0.5 19/0.5 3
		しょりゅく プルグダム/ デカめつりのんんます こっぱなる	† I		set surface casing to la
4.0	-	dence to cense; dry to vislightly moist winc. depth.	19 5-1	PB-1	sample of 3 Pitcher Barrel Part of Sample Slipped
-	<u> </u>	5.0-7.0 <u>SANOY CLAY-SICT</u> : as for 0.0-5.0 but w/15-25%		7.34	from tube o surface We = 516.1102.
ø.o -	- ML	fv. f. sand . w/some Hacks of black carbinactout material; w/ yel. small pods (1/15) of Cacos.	8-3	2.5	wf = 1816. 814 oz.
	; ;	small pods (167) of Cacos. 7.0-12.5 CLAYET SILTY SAND		SPT-3	
80	SC- SM	to SANDY CLAYEY SILT: as a love	2-4	1.3/	15/0.5 11/0.5 12/0.5 2
	to cl-	pred. tV-t. grained, trace of m. grained sand, scattered throughout,	‡ 2 0	PB-Z	Part of sample slipped from tube
	- ML	grains.	‡ []	2.07	Ue= 516.10/402.
10.0	<u>-</u> :	-	(B-5)	1.5	Wf=1716.150z.
-	:		± = B-6	SPT-4=	10/0.5 14/0.5 18/0.5 3
12.0	: -	_	+21	1.3/ 1.5	<u>†</u>
-	- SC-	17.5-18.0' SICTY CLAYEY SAND: Redish brn. (SYRAVA); Fine: 33-	B-7		chatter 0.5' into PBrun
14.0	$\cdot \mid \mathcal{W} \mid \cdot $	40%, low-non-platic; sand 30- 60%, poorly graded, pred. f-v.f-		0.85	(sample disturbed, not weight d)
74.0	- gavel	gravel generally 10-15-6, locally	‡ P. 0	SPT-5	17/25 77/2/
	-	cobbles of basait, sandst-clasts generally 1/2-1/2 strong. Hell leac. May be some gyp; med.	8-8	11.5	12/0.5 22/0.5 27/0.5 4
K.0	-	dense (1); slightly moist.	72	PB-4-	end of tube bent We = 516.93402.
			5-3	2.25	Wf = 316. 12/4 oz.
18.0		18.0-21.0 CLAYEY SILTY SAND	<u>+</u>	7.5	
‡ -1	SC- SM	104842): 40-25% 10W-non-plantic	F R-9	1.5	23/0.5 30/0.5 22/0.5 5
20.0		fines, 50-60% pad. poorly graded five to sand, some mine sand, there figures; caron approprietassing	235-4	11.5 PB-5	SHEET 1 OF 2

,] .				, DEMARKS
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS
20.0		18.0-21.0' CLAYEY SICTY SAND: (con4)	1	28-5	end tube badly bent
-	SM	red. dence to dence; slightly moist.	5-4	(con'X)	1
-	- CL	21.0-22.7 'SILTY CLAY: dk.	(con4.)	7.46	wf = 1816.840z.
22.0 -		488136 6m (5 YR 3/3); 85-90% -	 	<u> </u>	<u>,</u>
	∄		8-10	SPT-7:	10/0.5 14/0.5 14/0.5
-	≤M	fv.f. giarnof sand; no growli	8-11	1.3/.5-	790.3 170.3 170.3
	(54)		1 24	PB-6	
24-0	+ (30)	stiff; slightly moist.	[1, 2, 9, -	
	<u> </u>	22.7-39.5 SILTY SAND: d.k.	+		We = 516. 11/2 02-
-	†	Brown (7.5 PR 4/2) w/ 200 wh that	5-5	1.46	U4 = 13/6. 15 3/402.
]	Twhen drugging man 20-00 90	‡	12.5	
76.0	-	v. low-non-plastic; mod. guick - dilatancy, low-no toughness; sand	[SFT-8:	
<u>:</u>	፟	160-90 70, poorly groded V.FT.	B-12	0.9/	10/25 11/0.5 7/0.5
:	<u>†</u> 1	grained, scattered cm. plains; trace wisely scattered fc. gravel	-	<u> </u>	‡
z8.0 -	‡	I Strong Rac to HC/ Throughout; 44F-	ZS	P8-7_	
20.0	‡	dence; u-slightly moist.	Į		Uc = 516. 11/4 02.
-	‡	26-8-27.0 lens of coorser	<u> </u>	1.64	Wf = 1316.143/4 02.
•	‡	sould of some f. gravel; 942- sificul (?) stif, sondit.	፤	2.5	
30.0	<u> </u>		B-13	SPT-9:	<u></u>
:	‡ sm	12.5-30.4' contains thin lenses or hels where times grade low plastic to ce-ML	<u> </u>	0.9/	865 11/25 151 -
	Ŧ	grade low plastic to ca-ML	8-14	17.5	8/0.5 11/0.5 15/0.5 10/1 end shift@5:15, 10/2 start shift@7:35
32.0 <u>-</u>	<u>‡</u> │	30.4-31.5' contains chancool -	726	PB-8_	10/2 start shift @ 7:35
	‡	Frags	∄		
-	<u> </u>		‡ 5-7	1.57	we = 5/6. 11/2 02.
	‡		‡	2.5	Uf = 13/6.14/4 oz. (some slough on top)
34.0 -	-	34-0' scotled a said fonce!		SPT-10	
	-	34.0 scatterd c. said, f.govet at +id of PE-8	Ī B-15		11/0.5 13/0.5 15/0.5
	‡		‡~′′	0.8/	
7/ 0	<u>ŧ</u> .	_	27	PB-9_	end tube slightly bent
<i>3</i> 6.0 -	‡		‡		Un= 516. 11/2 02.
-	<u> </u>		I 5-8	1.73/	Wf = 1416.113/4 02.
	‡		Ŧ	2.5	(Some slough on top)
<i>38</i> .0 -	‡	- 341. 34 claridas al same sin	‡	SPT-11	-
•	<u> </u>	34.1-34.5' grades w/ some sm.	B-16	377-77	14/0.5 15/0.5 22/0.5
-	-	B.H. @ 39.5'	‡ ~ ~	1.6/1.5	1,000 0000 0000
40.0 -	E		‡		Terminak hole @ 8:45 a.m. @ 39/2 w/suffer
40.0 -	-	·	-		Edata
-	<u>‡</u>		‡		Backfilled hole w/mud and cuttings
	ŧ		‡		t and cuttings
42.0 -	<u> </u>		<u>+</u>	-	-
	‡		Į		E
•	-		<u>+</u>	.	<u> </u>
44.0	E I		+	1	SHEET Z OF Z

DRILLING AND SAMPLING LOG

PROJECT <u>Oll 8</u>, SCS Dams - Utah DATE DRILLED 10/2/81 HOLE NO. 5C3-2
LOCATION <u>Opstwom the Dam#3</u>, NSta. 9+45; 50'U/S GROUND SURFACE ELEV. N6053' Ctopo.)
DRILLING CONTRACTOR <u>Pitcher Drilling Co.</u> LOGGED BY <u>DMY</u> DEPTH TO GROUND WATER <u>not encountered</u>
TYPE OF RIG <u>Failing 1500</u> HOLE DIAMETER <u>418</u> HAMMER WEIGHT AND FALL <u>140 16</u>. 30"
SURFACE CONDITIONS <u>brushy Flat</u>
WEATHER <u>scattered</u> clouds, warm

		/	_			
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS	
0.0		BASIN DEPOSITS	28	57-1:	Push shelly take (3")	1
Z-0 -	X space	non-postic sand to stylo poorly must and to stylo poorly must sand to stylo party must suitely scattered, o-5%, 14-1/27, med dense (2); dry.	S-1 B-1	7.15 7.2 40 5PT-1 5.2/a6	240 psi to 2.0', 250 psi to 2.2' 1250 psi to 2.2' 1250 psi to 2.2' 1250 psi to 2.0' 1250 psi to 2	60+
4.0		3.0-4.0 grades slightly more	8-2	RD 5PT-2		
	SM W	4.5-B.O SILTY SAND: dk. 400V	B-3		15/0.5 24/0.5 21/0.5 Sample w/3"pitcher barnel	45
6.0	copple	fc. gravel, pred sandit, basalt;		0.01	sample washed during sampling or slipped from tube	
8.0		about disemminated cocos continuous some thin stringer or leases - grading to sandy clay-silt; med dense; slightly moist.	30	12.5 PB-2	Sample cots aproly- may be washing	2
10.0		8.0-10.5 recovered sov. baratte gravel clarts 1/2-2/2" in sixty to slightly clayey sond matrix	8-4	0.3	during sampling	
12.0		10.5-12.0 Recovered any a zumbly worked stilly sand		SPT-3:	9/0.5 6/0.5 9/0.5 (2" base If clast in short puched through, sitty sand in-situ-rock	15
	-	basatt cobble	-	PB-3 0.0/ 12.0	one z-7 "basalf cobble" in end of kube - mf	
14.0		, -	32	57-2	Push shelly tobe 2.0 uf 180 psi; no acovery; I dent in end of tibe	
16.0 -	CL-	16.0-24.5 SANDY SILT to SANDY CLAY-SILT : RESTIGHT		0.0, /2.0: SPT-4:	<u> </u>	
18.0	ML to ML	60000 (54R 4/3); Fines vary	B-5"	1.3/ (00)	Note: over dioue spon of	13
20.0		trace distely scattered f. gravel, most sondst. strong reactotheli- come gyp (3); stff-U.Stff; slightly moist.	5-2	PB-4 1.16/	We = 516.1502. We = 13 16. 6/402. SHEET OF _3	

			DATE DRILLE	· ·		HOLE NO. 623 2	-
DEPTH	CLAS	S.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS	
20.0		12- 12	16.0-245' <u>SANDY SILT TO SANDY</u> <u>CLAY-SILT</u> : (CON't)	(8-2,4.)	(con 4) SPT-5		
-20		to ML		8-6	1.25	9/0.5 6/0.5 10/25	10
<i>22.0</i> _				34	PB-5		
-				5-3	1.33	ως = 516.10 oz. ως = 1316 234 oz.	
24.0			24.5-34.0 <u>SKTY SAND</u> : Robish :	<i>B</i> -7	12.0 SPT-6	-	
<u>-</u>	5	SM	brn (5 x 4/3) w/ yel. orn. tint;	8-8	1.35	13/0.5 14/0.5 16/0.5	3
26.0 -			plastic, quick dilatancy, v. 1000 10	75	PB-6_	_emall dent in end of	
-			graded, mostly f-u.f. sand, w/ some mc. sand locally truce to 5-10% pred. f. gravel, mostly	5-4	1 1	we = 516.83/402. wf = 1716.21/4 oz.	
28.0-			saidst; occasional small specks		1.38	-	
			uned dense; slightly moist-	8-9	SPT-7	13/05 13/0.5 16/0.5	2
30.0			245-310 occasion small	B-10 36	1.5	end slightly dinged	
30.0 -			and Aravily mother concentrations of Cacos	5-5		We= 516.90Z.	
=			m. sand and some figure, pad. gray sandst.		1.80	wf=1516.113/40z.	
<i>32</i> .0 -			<u> </u>	B-11	SPT-8	<u>-</u>	
<u>-</u>			-	B-12	1/1.5	rough drilling-RD to	′
34.0			34.0-78.0 SILTY CLAY: raddish 6 rn. (5 yr 4/4); 85-90% low-mod.	37	RD PB-8	rough drilling-RD to _34 before PB-8	
: -	, -	<u></u>	plastic fines: 10-150/ poorly grided of the true of black or bonace gus material the strings of black or bonace gus most to maist.	B-13	1.5	end of bbe v. badly tom, bent	
25. 0 -			30.0 -44.30/C/ P 103/10/11 GRAVEL:	-	. 10	Ream hole to 25.5 % clean out grovel	
:		FM	molti-colored clasts in relaish brn. (54R413) matrix; 10-15-90 mon-plastes sitty fines; 25-30% mod. well		SPT-9	· •	
38,0 -			and tam anuel winost class	8-14	0.9	45/0.5 46/0.5 43/0.5 - Chammening in quavel)	3
<u>.</u>			12 de v. deeply weatherd, sub-		RD	rough dilling 38-39.5	
40.0			hd; rel fresh organic defins as mothers twing present; mod reac.	38	PB-9	smooths out a little @39.5 Hard chatter in PB-9	1
40.0			nd; rel. tiesh organic degris as notitets, twirs present; mod reac. to Hellin matrix: degree of compodiness uncertain; moisture content uncertain.	8-15	3.59	highly disturbed	
			ONGER I SOLVE	-	RO	<u>. </u>	
4z.0 -			<u></u> :	B-16	597-10	34/05 mb = 1.1	
-			<u>-</u> -	- "	0.7/1.0	34/0.5 50/0.5 Rhsal on 2/2" sandst elast	ક
44.0	‡			-	RO	SHEET Z OF 3	

PROJECT D118, SCS Dams - Utah DATE DRILLED 10/2/8/ HOLE NO. 663-2

-KOJECI -		, 00	CA NEWS STEET DATE DRILLE	00	,	HOLE NO	
DEPTH	CLAS	ss.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS	
				<u> </u>			
DEPTH 44.0 46.0	CL AS		FIELD DESCRIPTION B.H. © 44.3'	SAMPLE	MODE VIIII III III III III III III III III	REMARKS SPT-11 50/0.3 - refusal on gravel /cobble - No recovery Terminated hole & 44.3 w/sufficient information acquired Backfilled hole w/ outtings and mud Decide to drill auxillary to get samples and the try to try to get samples and additional spts in the 5.0-16.0 sitty sand intrual-see tog of boring \$6.3-2.A	50+
•	‡			<u>:</u>	‡	-	
	Ε		,	<u>:</u>	1	SHEET 3 OF 3	

DRILLING AND SAMPLING LOG

PROJECT D118, SCS Dams - Utah DATE DRILLED 10/2/8/ HOLE NO. 663-ZA
LOCATION Upstram toe of Dam#3 @ ~ Sta. 9430; 50' US GROUND SURFACE ELEV. ~6053 (topo)

DRILLING CONTRACTOR Pitcher Drilling Co. LOGGED BY DMY DEPTH TO GROUND WATER not excountered type of RIG Failing 1500 HOLE DIAMETER 478" HAMMER WEIGHT AND FALL 140 16. 30"

SURFACE CONDITIONS brushy flat WEATHER scattered clouds, warm

DEPTH	CLASS	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS
2.0		Note: see log of 663-2 for description of units from 0.0-8.0		20	Start drilling @ 3:35 P.M. drilling @ 3:35 RO to Rock Sampling interval Onills erratically from 0-7/2 w/ gradations
4.0	-		<u>:</u> <u>-</u>		On'lls eria titally from 0-T'z'w/ growntims From sandy to sitty; some chattering
6.0			<u>-</u>		
8.0	51	1 8.0-14.5 51LTY SAND: dk. gray brown (10484/2); fixes v. low to	8-/	SPT-1 1.2/	diviling smooths notably 8 +1/2-8 Drove standard splitspoon 4/1.0 8/0.5
10.00 -		non-plastic, 20-25%; 75-80% porty graded, pred. fu-? grained sand; trace scattered gravel; race to Hell is mod; mod dense; moist (may be due to drill fluid).	39 5-1 (13-2)	PB-1_	Sample w/s"Pitcher Barrel cycled pump on/off during PB-/ sampling run to a unit washing sample; end of tube
12.0		- and tes to do as to		SPT-Z	sample slipping out of - tobe, we = 516.12/202. 12/0.5 10/0.5 15/0.5
140	SM:		B-3	0./1.5	(B-2 saved from end of 700e) 6/0.5 4/0.5 5/0.5 Terminate hob @ 14/2' W/sufficient info.
K.O -	_		<u>:</u>		- Packfill w/ mub and cuttings
18.0			-		_
20.0			- -		SHEET_ / OF /

DRILLING AND SAMPLING LOG

PROJECT <u>DIPS</u>, <u>SCS Dams-Utah</u> DATE DRILLED <u>9/30/81</u> HOLE NO. <u>625-1</u>
LOCATION <u>Crest Dam #5 @ Nsta.H29 &; Codar City</u> GROUND SURFACE ELEV. <u>5940</u> (topo.)

DRILLING CONTRACTOR <u>Pitcher Drilling Co.</u> LOGGED BY <u>DMY</u> DEPTH TO GROUND WATER <u>not encounted</u>

TYPE OF RIG <u>Failing 1530</u> HOLE DIAMETER <u>478</u> HAMMER WEIGHT AND FALL <u>140/6.</u>; 30"

SURFACE CONDITIONS <u>dirt embankment</u> WEATHER <u>Scattered</u> clouds, cool

, o () A O E		NS OFF PASARMANT			- CATT (RA E110 05) COOT
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS
0.0	.]	EMBANKMENT FILL		SPT-1:	Diace standard split spoon
	CL	0.0-26.5 SILTY GLAY: dk. wolling	8-1	0.8/	7/0.5 8/0.5 19/0.5
		brown (5 PR 3/4) -> 95% mod.		AD	Auger w/5" flight auger
2.0	:	plastic fines, slow dilatancy, moditations of v.f.		SPT-Z	70 2.0
1		sand-no gravel: trace gyp: as small xystls, nodubs; mod. to strong- eac. to Hel throughout, w/small veins, strongers, of buff Calos; very stiff to hard; slightly moist.	B-2	10/15	8/0.5 10/0.5 13/0.5
- - היא		veins, storages, of buff calos;		RO	set surface assing to
4.0	- :	Very 2777 TO NOTO ; 21191111 WILL	/	PB-1	somble with from formed with the sell
-			5-1	2.75/	$\omega_f = 17/6. 130z.$
خ.o -				2.5	
				SPT-3	
1111			B-3		10/0.5 27/0.5 28/0.5
80		-	7	PB-2	<u>-</u>
	[<u> </u>			We = 5/6 4/4 02.
			5-2	2-1/2.5	wf = 16/6.9/40E.
10.0		-	_	72.5	
-]		SP1-4	
_			8-4	1.1/1.5	5/0.5 8/0.5 9/0.5
12.0	-		- 3	PB-3	-
-					most of sample slipped
-			B-5]	note to be down
14.0	<u>-</u>			2.5	1076
			-	SPT-5	Ream withicone to 14/2
-			B-6		14/0.5 23/0.5 44/0.5
16.0	-	=	4	PB-4:	We = 516.14 402.
1		<u>.</u>			We = 18 16. 15 3/4 02.
-			 5-3	7 77/	† ' I
18.0 -	-	18.0-26.5 grades w/trace		2.33/	sample loose at bottom of tube - disturbed
1		of widely Ecattened c. sond: to f. quotel sized basalt clasts, angular and hard		SPT-6	}
7		ciasts, anyonar and hard -	8-7		14/0.5 20/0.5 24/0.5
20.0				1.7/.5	SHEET / OF Z

PROJECT DITA - SCS Dams DATE DRILLED 9/30/81 _HOLE NO. 625-1 REMARKS SAMPLE MODE FIELD DESCRIPTION DEPTH CLASS. 0.0-26.5 SILTY CLAY: (CON't) \$5 PB-5 20.0 We = 516. 14/4 02. Wf = 17 16. 10/2 0Z. 5-4 2.5 27.0 -SPT-7 B-8 10/0.5 12/0.5 15/0.5 27 24.0 P13-6 We = 5/6. 14 4 0=. Wf=1716. 302. 5-5 2.14, 2.5 ALLOVIUM/COLLUNIUM 26.0 -26.5-47.0' CLAYEY SANDY SPAUSCE SPT-8. multi-colored closts in reddish om. 60 1.0/1.5 15/0.5 33/0.5-5% 3 83 (SVR 4/4) motrix - 10-20% low-mod plastic fines : 20-35% mod. well graded sand, most ang. to subang quains; 50-70% quael 280 -Fream to 29 w/thicons RDfrom 1/4-3" w/ most 1/2-1/2" smoother drilling @ 28/2 We = 516. 15 to az. limestone, sitst, sandst. and in most commonly bacath, all mod. to deeply weathered, weak to his ang- to sobrounded; degree of compactness concertain; in-situ moisture uncertain. PB-7 Wf = 15 16. 2 14 02. (endof cutoff) 2.0 ent end of tob; 30.0 63+ SPT-9± 15/0.5 48/0.5 B-10 RD FREAM to 32 w/fricone Note: locally unit may 32.0 be endely stratified all some thin layer of clayey sond to sandy clay as seen in sat-10 from 34.0-35.5 PB-8 + we = 5 16. 13 1/2 oz. : WF = 1216. Ooz 1.35/2 of bent end of 1/262 - disturbed sample 5-7 34.0 -SPT-10 0.95 = 27/0.5 37/0.5 34/0.5 71 B-11 RD = Ream hole to 36 w/frican 36.0 -PB-9 o.z/ = 21/2 "basalf clast in 12.0 Tube 16k taking wakrimix SPT-11 1/2 sacks mud B-12 38.0 - 1-3/1.5 = 20/0.5 25/0.5 22/0.5 47 8-13 RD bit caught by cabble in hale lonscruped fish & PB-10 Ream we trice to 40/2 0.6 Refusal & 42 40.0 -B-14 - Terminak hole Odz w/sufficient into. 42.0-B.H. @ 47.0' Back fill hole w/ mod & SHEET 2 OF 2 44.0

DRILLING AND SAMPLING LOG

PROJECT DIB, SCS Dams-Utah DATE DRILLED 9/30-10/1/81 HOLE NO. 665-2

LOCATION Upstram to Dam #2 @ ~ Sta. 1+73; 82 U/S GROUND SURFACE ELEV. ~5926 (tops.)

DRILLING CONTRACTOR Pitcher Drilling Co. LOGGED BY DMY DEPTH TO GROUND WATER not removinted type of RIG Failing 1500 HOLE DIAMETER 478" HAMMER WEIGHT AND FALL 140 16., 30"

SURFACE CONDITIONS Sediment behind dam WEATHER scattered clouds, cool

DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS
0.0	CL	CATCHMENT BASIN DEBRISCO) 0.0-7.0' SILTY CLAY: dk. Rd.	B-/		Onvestanlard split smon
2.0		brown (5483/4); 95-100% low- mod. plastic fines, slow dilat, mod. toughness; few old to truck of f. v.f. sand; no grouel; cacos		AD SPT-2	Auger w/s" flight auger to z
4.0	-	throughout; some partially to fully carbonized fine organis dolors: stiff; dry to slightly	8-2	1.91.5 RD	5/0.5 6/0.5 9/0.5 Ream hole w/47/21 thicope rock, bit; mix Simplify " " thor Barrel We = 5 10. 1402.
	- -	moist.	5-/	P13-1	We = 5 16. 14 02. Wf = 1516. 11/2 02. bottom, 0.3 of PB-1
6.0		ALLUVIUM /COLLUVIUM	8-3	2.5 SPT-3	slipped out of tube, saved in 8-3
8.0	60	5.0-43.8 <u>CLAYEY SANDY</u> BRAVEL: multi-colored clasts IN reddish orn. (54244) matrix; Fines vary, 10-30%, pred. low-	B-4	1.2/1.5 PB-Z	6/05 13/0.5 18/0.5 We = 516.13 402.
10.0	-	mod-plaitie, locally low-non- plastic - sand vanes 15-3090, mod. well graded, most ang- sob-ano; gravel vanes do-752	5-2	0.79/a.9 RD	Wf = 10 16. 5 1/4 02. Only w/tricone rock bit in quavelly section w/ some cobbles
12.0	-	sandst, limest. stst., hasalt, led volc. C2), ang-sub bunded, moddeeply weathered; reac. to the mod. the mughout: 20 cally unit is endely stratified and some lences quode to sandy or jourly clay and claypy sand; degree of compactines woncortain;	- B-5	PB-3	- Onilling smooths out at 12.2-12.4 (sample not weighed)
14.0	- - - -	moistine in-situ uncertain. 7.0-8.0 generally sandy Clay w/10-20% sand and 5-10% grouel		10.8 RD	Orill w/rock bit in gavelly section
16.0	- - (3-6	SPT-4 6.2/6.8 RD	25/0.5 50/0.3 Drill w/mck bit in quauelly section NIG'-hard rattling/ rig shaking
18.0	<u>. </u>	18.0-19.0 cmbe strutifica- tion from clayey sandy growel to gravelly sandy	14 B-7	PB-4	Drilling smooths out some a ~18 Campled highly distribut, not weighed

PROJECT 0118, SC	S Dams - Utah DATE DRILLE	0 9/30-10/11	19/ HOLE NO. 625-2	
DEPTH CLASS.	FIELD DESCRIPTION	S.AMPLE M	ODE REMARKS	
70.0 = GC	7-0-43.8' CLAYEY SANDY GRAVEG		nit) Podshift QS:20; starts 20 Ro to zi fo alean hou	the state of the s
22.0		5-3	2-5 = end tube badly bent = (some quavel slouph on top of disturbes sample -not weight 2.0 = 12 16.4% oz.) (1)
240		50	RD in gravel w/sor	on 50+
26.0	26.2-27.3 v. little to no binder; cardsf./baselt quavel	B-8 0.	Add z-3 16s. poly-dr to help Flush coth T-6 26.2 5/1.1 27/0.5 43/0.5 50/0.1	93+
zə.o=	. <u>-</u>	<u>-</u>	20 = 50/0.1 - ho 20 = 0.0/0.1 an gravelo = (0.3 slowed in spoor)	clast
30.0	-		Hole caved/shughed 25 /2 after taki 507-7 Mix 1 /2 sacks bents to help stabilize h Flush cuttings	70 19 vaite
32.0	_		Roto 32/2 while drill smooths somewhat to low plastic fines in B-6 - Cuttings	Ving more 1
34.0	34.5-35.3' v. little to no fines; ptd. bosalt, some sandst. qnowl, c. sand	5-4	77 = UP we = 516. 143/40.	02. 2. 11½" 8C1
38.0	36.5-38.0 quides less quouel	 	RD cast of sample is show	·
39.0	-		RO in gravelly section of the sectio	611
40.0		<u>-</u> - -	end of the cuito Bock - sample dist We = stately U4= 1716. 1.12/ Note: begin placed 25 - card board spacer Wax plug a top pt-9 - PB samples que 86.3 - 39/0.5 50/0.3 in	17 oz. 19 20. d 18 all 1900 n 874
42.0	43.0-43.8 31/2 "basaH cobble Request in PB-8	-	RD - KOM gavelly section graves more graves was	2114
44.0	B. H. @ 43. 8		OB-3 - (Sample not weight) 1.3/0. ST SHEET OF	_ !/
-		Terminare opn East-411 hole of	Mistan graffors	つなけ

DRILLING AND SAMPLING LOG

PROJECT <u>Olla-SCS Dams</u>, Otah Date Drilled 10/6/8/ HOLE NO. WD-/
LOCATION Warms Daw Dam, NSta. 15+23 ON & GROUND SURFACE ELEV. 12989 (topo)

DRILLING CONTRACTOR <u>Pitcher Drilling Co.</u> LOGGED BY <u>DMY</u> DEPTH TO GROUND WATER <u>Cot to the survey of type of RIGE Trilling Co.</u> HOLE DIAMETER 47/8 HAMMER WEIGHT AND FALL 40 16-, 30

SURFACE CONDITIONS 1854 of earth tabankment, cobbles WEATHER 2/200, warm

DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS	617
0.0	sf- SM	EMBANKMENT FILL 0.0-15.0 SAND to SILTYSAND		AD	Auger w/6" flight auger thro, surficial,	
Z.0 -	972 V	RESSISH Brown (2.5 YR A/4): 5-159	 		ganelly, cobbly makend to 25'-clean hole by hand for SPT	
-	sp. sm	1101 TOXIDE STAINED ATE. GIAINIS	# 8-1	SPT-1 -	Orive standard split Spoon 19/0.5 = 5/2.5 16/0.5 (scattend fn. 9nuel)	3
4.0		slow to mod. reac to Hel enditates some Caloz; gup as v. f. clear crystals; these carbonized black		RD :	set up tub, surface acycg to 3 mix I sack of bentonite, wz/6. zero ge/	,
-		med dance to dense; dry to design penetrated	+		sample w/3"PitcherBornel	/
6.0		by drilling fluid). 0.0-7.0 quavelly, couldy silfy 5010, w/20-70% c. 9002)	Ŧ '	7.42/	We = 5/6.15/202.	
8.0 -	-	small coebles commonly, 1-4 Row cobbles to 6"chosts pred. red sandst, 117/t color gtzite(?)	+ 1	597-Z 1.0/1.0 RO	35/0.5 53/0.5 (Absolon goul clost)	2
<u>-</u>	-	gtzite(?)" '		PB-2	one ding in end of to be void along sick of sompli	
10.0			<u>+</u> 5-2	2.40	we = 5/6.1502. wf = 20/6.3/202.	
1z.o –	SP- SM gavel	11.0-12.2 gades w/10-20% coarse sond and fire gasel one 1" clast in certar of SPT-3	B-3	SPT-3	5/0.5 6/0.5 9/0.5 (sathud gavel, one 1"	
<u>:</u> :	SP- SM	381-3	-		end of tube o.k. we = 516.150z.	
140 -			5-3	1.8,	wf = 17/6. 2/4 oz.	
16.0	sc- SM	15.0-30.0 CLAYEY SILTY SAND Rodish brn. (54RA/4); 25-30% low plastic fines, mod. quick.	- 1	SPT-4	17/25 15/25 33/25 Cface F. gavel)	9
-		low plastic fines, mod. guick dilatancy, low toughtessibs zing poorly gived, f V.f. att. sand: o-5 % scatters c.sond f. provided clasts /3-1/2", most /2-3/4", of	7	PB-4	end of this according	
18.0 -				?/z.5	Wa = 516.15/202. Wf = 2016.13/202.	
70.0	5M	leac. to HCl-some Calls : 948- as v.f. controis soil into potder to 18-14" of soil possibly of p. ; truce cartained back organic cours; dense; slightly moist.	B-5	SPT-5-	13/0,5 70/0.5 27/0.5 Cather f. groups SHEET OF	4

PROJECT ₄	<u> 0118-sc</u>	5 Dans, Wah DATE DRILLE	D 10/6/81	HOLE NO. WO-/
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE REMARKS
Z0.0 =		= -	B-5(0n4)	RD - soon growill slough
22-0	≸M H		5-5	PB-5 end of file, and in We = 5 16. 15/2 02. Z-62
24.0	† + + + +	23.5-25.0' slighty more fines	B-C	SPT-6 25/0.5 30/0.5 17/0.5 47
26-0		- - - - -	5-6	PB-6 = end danted, either great sample No.z' or have, that much tight slough 7.72 = on top of sample. UR = 516.153/402. 2.5 = Wf = 2216.7/2 oz.
28.0		27.5 qualing more plastic, fines to 35-40%; culting laminar of non-plastic silty sond & bottom of 15-6	8-7	SPT-7-27/0.5 21/0.5 21/0.5 42 1.3 (scattered f., deeply weath. 1.5 sandst. 9 muel)
30.0	SC w/	30.0 -54.0 CLAYEY SAND: - reddish brn. (54R4/4), less Hd tint, than plev; 35-45% low- mod. plastic fines; slow to mod.	<u>s-7</u>	$P8-7$ $W_{0} = 516.14^{3}/402.$ $W_{0} = 2116.1202.$
32.0	acic.	Sand to 50-65%, mod. poorly gaded, pied. fv.f., some M. e. giains: 0-10-15% gravel,	B-8	SPT-8 = 8/0.5 10/0.5 16/0.5 21 1.2/ (fw 3/8-1/2" gravels) 1.5
340		mod. reac. to Hel due to cacot; - some 940 a v.f. dear any stuly of buff, soff i req nodules to 18;	<u>5-8</u>	P8-8 = end of tube o.t. EUQ = 516. (4/4 oz. EVG = 1916. 14/2 oz. 2.23,
38-0		specks; contains thing clayed sand to sitty sound as described below; - med dence to dence very stiff med dence to an blow count);	F	SPT-9= 10/0.5 20/0.5 20/0.5 1.2/ (southerd c. sand, f. growl) 4
38.0 -		31.5-33.9 for % c. sand, -	<u> </u>	PS-9 = PNG of the CRATE of We = 516.14/2 oz.
40.0		35.5' coane sand to u10%. 35.5'-37.0' quader more c. sand and f. quarel 39.5' brick red-orange sand to siff sand (spsm), fv. f. quined, trace. c. sand	8-10	5PT-10 10/0.5 12/0.5 18/0.5 3
£.0 -	; 5C	41.5-41.8 clean spsand, t.	5-10	PB-10 end of the slightly dented We = 516-14/202. 2.21 WY = 19/6. 1102.
44.0	2 5PSN	23.5 sand to sity sand w/	B-//	2.5 SHEET Z OF 4

1	PROJECT_	0118-5	CSDams, Utah DATE DRILLE	0 10/8/8	P/HOLE NOWO-/	
	DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE REMARKS	
,	94.0	5C W	30.0-54.0 CLAYEY SAND: (CON4)	B-14)	(cont.) 15/0.5 17/0.5 24/0.5 1.3/1.5	4/
	46.0 -	9 rave			PB-11 end of tube badly flavor out-not capped, but wax-sealed	
	-			5-11	2.74 We = 516. 14/402 WF = 2216. 5/202	-
	48.0			B-12	SAT-12 1/0.5 16/0.5 20/0.5	36
ı	59-0				PB-12 end of tube at. We = 5 16.14/2 02	
,	_			5-12	2.26 = WG = 20 16. 102.	
y	52-0		<u>:</u>	8-13	SPT-13-16/0.5 32/05 28/0.5	60
יני מיני מל	~a n		ALLUUIUM/COLLUVIUM		18-13 end slightly but We = 516.1402.	
nkwen!	54.0	SC (6C)		5-13	2.49 Wf = 21/6.302.	
t'e mba	56.0		w/yel.bm to dk. bm clasts;10- 30 % low plostic fines; sand wanes 20-locally 50 %, wanty 20-30 %, both fv.f. and	B-14	SPT-14 176.5 186.5 34/25 0.9 Coone a sand, trace figure	- 52
. ma o	58.0-		well graded; grave I was zo-	B-15	PB-14 = 2/2-3 "cobb & in and a tible; sample distorted	`.l
	, d	(SC)	and sitst., 14-3" most 14-1";	R-16	z.zo up = not weighed	
	Ø-0		matax w/ some cacos, 949.) couldly stratified w/ layers ently stratified w/ layers ently ped. gravelly orsandy some clayer sands as for 30.0-	B-17	SPT-15 140.5 16/0.5 20/0.5 1-4/.5 (some quavels to 1/2")	36
	62.0		organic defens as small specks;		PB-15 end of the body bear we = 516. 1502.	
	4 		by samples, not comented; slightly moist. 540-58.0 predominantly	5-14	2.50 Wf = 2016. 7/202	
	64-0		SRO-67 C'DRE dominum Hy	B-18	SIT-16 14/0.5 14/0.5 34/0.5 1.3, (a)/250% sandst.grave	5.
	60.0		less clayey gravely sand = 163-64' cobr grading less Radish tent		PE-16 End of tube bady bent, sample loose in type	
	-		BEDROCK 67.5-71.5 SICTSTONE KLAYSTONE	5-15	2.35 WF= 18 16. 14 02	
<u> </u>	63.0		(see next sheet for description)	8-19	SPT-17 SHEET 3 OF 4	

PROJECT_	1118-50	S Dams, Utah DATE DRILLE	0 10/8/8	P/ HOLE NO. WO-/	
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE REMARKS	
68.0	1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	67.5-71.51 <u>SILTSTONE/CLAYSTONE</u> : (con/f) vary H. Omy(NT/E) fosh w/	B-19 (con't)	RO RO 685 while on	k) Ung
70-0	40 10 ST27 AS 12 ST27	vary 14. any (NT/E) fosh w/ brownish yellow (10 YR 6/6) westh. Staining, interferded HI. softer claystone and harder sittitione	- B-ZO	PR-17 - Sample disturbed, sav	
<u>-</u>	200	to f. graces sandstone; s/tst/ sandid. w/ blocky froctors, joints @ 10,40,50; cwde bedding @	8-2/	1/2 d we = not weight	d 76.5
77.0	-	deeply weatherd; clory		Terminated hole of d'into puck cul sufficient data	6-
j		Fractized; dayst. w/low hardness; styst./sandst. w/mod. hardness; clayst. weak to frable, sifst./sandst. weak to mod. strong.	- -	Backfilled hole w/ - cultimos/mud; pour bentonche-nck, slure unto top 3	æd
-		B.H. @ 7].5'		into top 3	ry
				 	
		<u>-</u>		 	
-					
<u>:</u>		-			
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-		_		<u> </u>	
<u> </u>		-			
		-	-		1
	# 	•	- - -	SHEET 4 OF 4	·

DRILLING AND SAMPLING LOG

PROJECT D118-5CS Doirs, Otah DATE DRILLED 10/7/8/ HOLE NO. WO-Z
LOCATION Warner Draw Dam, ~ Sta. 17+49 on & GROUND SURFACE ELEV. ~ 2989 (tops)
DRILLING CONTRACTOR Ficher Drilling Co. LOGGED BY DMY DEPTH TO GROUND WATER 10 PRICOUPLE &
TYPE OF RIG Failing 1500 HOLE DIAMETER 47/8" HAMMER WEIGHT AND FALL 140/6, 30"
SURFACE CONDITIONS CRSt of worth Embankment, cobbly WEATHER Clear, worm

SURFACE	CONDITIO	NS CRIST of earth embankment,	<u> </u>	VEATHER.	ZIFAR, WURM
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS 6kg
0.0	≤P- ≤M	EMBANKHENT FILL		AD :	Auger w/6" flight auger
2.0	- w/ 970パル かれた	0.0-15.7 SAND to SILTY SAND:			Auger w/6" flight auger to z' to clear goverly, cobbly surficial mantle clean hole by hand Down startard splitspus
	SA	redish brown (2.5 PROLE): 5-15-70 non-plastic times, guick dilatary, low-no town was . 80-90% poorly graded, pred fv.f. grained and;	B-1	SPT-1:	15/05-21/05 32/05-5
40		Rw dards to 1-2 " minor disper. Cacoz - 11/2-mod. Rac. to HCl;	-	20 -	Chammering graul orcabbio, shoe empty) setup must to b; casing to 3, RD in scatters grave!
-				PB-1	sample w/3"PitchrBarn
d. o -	<u>:</u> 	0.0-2.0 w/ 20-25% c., 9muel, some cobbles to	5-1	2.24	$w_{e} = 516.1402$ - $w_{f} = 1916.602$.
80			B-2	SPT-Z	236.5 296.5 34/0.5 6 (1/4" quad clast paar bottom of sample)
-		8.3'w/1/4" gravel clost, - grades slightly clotey		PB-Z	end of tabe slightly dinged
20-0			- S-Z	?/2.5	We = 516.140Z- Wf = 18168740Z. (Forgot to massue acovery, probably >21)
17.0	- -		_ <i>B</i> -3	1.4/5	15/0.5 15/0.5 15/0.5 3
-		<u>:</u>		PB-3	end of lube o.k. We = 516143402.
14.0-			<u>5-3</u>	2.25	wf=1918.10/202.
16.0	\$0-	15.7-29.0 CLAYEYSILTY SAND: Rddish brn. (5 YR4/4); 25-50%	- <i>B</i> -4	11/5	16/0.5 18/0.5 22/0.5 (Scatter & F. gnus) 00 Hom 0.4 of sample)
	SM	Redish brn. (5 YR4/4), 25-50% tow plastic tipes: 60-65% v. poorly quadro sand, pred. f. V.f. quained, subang-subrad, 10n-oxide stoined to jun-1/2		PB-4-	end of take o.k.; sample loose in take we = 516.1202.
18.0 -		I SCATICAL TO TO TOUR LIGHT A TILL TO	-5-4	2.2/	wf=notweighed
20.0		generally fow Rdc. To TTC, but who some mother, small notules GCDs: trace black Carbonized agains ! dease; slightly moist.	B-5 B-6	5PT-5 1.5/.5	14/0.5 19/0.5 22/0.5 4 SHEET_/_OF_4

PROJECT_	0118-	SCS Dams, Utah DATE DRILLE	D 10/7/	181	HOLE NO	- -
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS	
20.0	SC- ≤M	15.7-29.0 CLAYEY SILTY SANO: (CON4.) 19.0-19.7 grades to v. low to non-plostic fines	(Gon't.) - 5-5	SB (- S) Con 7.) PB-5 2.5.3	end of tibe dightly digged, Z"clast removed from bottomend UR = 516. 12/2 02. UF = 21/6. 7/4 02.	2
24.0	-	23.0-24.3 gnades slightly more plastic fines	- 8-7	11.5	16/0.5 23/0.5 23/0.5 (scatter of f. graves)	46
26.0		NZ5.0-Z6.0 grades w/ NIO-15% coane rand, f. govel as sitst., sandst. clasts	<u>-</u> ≤-6	PB-6-	end slightly dinged we = 516.1440z. wf = 2016.41/20z.	
28.0	-		8-3	587-7 14/ 11-5	140.5 26/0.5-29/0.5 (10-15% c. sand, f. grave) as sitst., sandst.) end of tobe o.k.	54
30.0	SC	29.0-43.0 CLAYEY SAND: reddish brown (54R4/4); 35- 45% 10w-mod. plastic fines; 45-55% mod. poorly graded sand, pred. fv.f. grained, but w/mc. grains present; 5:15%;	S-7-	234	we = 516.11 /2 02. Wf = 1916.143/4 02.	
32.0		w/m-c. quams present, 5-15% quall common as weath. sarding start to 12-1" some dissem. a some mothed cacos: trace questione small specks black corongered aganic debits dense to hard (from 6 houcount); slightly moist.		1.2/	8/0.5 14/0.5 21/0.5 (10-15% c.sand, f.gravel) end of tube o.k.	35
34.0		f. quavel	<i>S-</i> 8	2.60	we=516. 161/402.	
<i>₹</i> 6.0 -		35.0-35.3' clayey sitty sand	+	SPT-9 1-1/ 1.5	206.5 30/0.5 30/0.5 (10-150/0 c.sand, f.gmuel) end of tube U.slightly	රෙපි
38.0			- ≤-9	2.32	$w_{\ell} = 516.11/402.$ $w_{f} = 2016.57402.$	
40.0	SC	sq.0' w/15-20% non-plastic - fines at end of PB-9 39.0-39.3' non-plastic silty - sand	B-1/	191.5	19/0.5 19/0.5 28/0.5 end of hips a.t., but	47
42.0		AUEU CIVYU CONA.	5-10	7.42	sample 100se in tube. we = 516. 143140z. wf = not weighed	
44.0	sc- sm	43.0-50.0 CLAYEYSICTYSAND: v. similar to 15.7-77.0; dense- v. dense; slightly marst.	18-12	SPT-11: 12/1.5	27/0.5 33/0.5 27/0.5 SHEET Z OF 4	60

1-1100				TARBUME, OTATE DRILLE	<u> </u>		HOLE NO. WO Z	_
DEP	тн	CLA	ASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS	
44.	٥	, , , , , , , , , , , , , , , , , , , 	SC- SM		8-13-4.)	\$PT-11\\ \(\rangle B - 1 \)	end of the o.k. We = 616.151/2 02.	
46.	O -			<u>-</u>	5-11	2.44	wf= fire of to weigh	
48.	o -			· <u>:</u> -	<i>B-13</i>	1-3/	36/0.5 36/0.5 33/0.5 CSCOTHERS CHEPH Weath. Sandst. gravel) end of tube bent in one place	69
\$0.	0 _		<u>ح</u> ر-	50.0-59.0 CLAYEYSAND and - CLAYEY SILTY SAND: Inkulayening	5-12	2.64	- place We = 516. 14/4 oz. Wf=2116.13-14 oz.	
57.	0 _		M2 (22)	of units as described for 15.1- 29.0' and 29.0-47.0, but w/ more c. sand f. gruvel common;	B-14	2.5. SPT-13.	19/0.5 34/0.5-40/0.5- (9/0.41/4 to 20-300/0)	74
54.	0		 S(\f SM)	layers were wo 4-0.8 thick and presumably represent lifts in the compacted fill.	5-13	PB-13-	end of fibe heat over $\omega_{e} = 516.15$ oz. $\omega_{f} = 201611/2$ oz.	
56.	0 -		्रिस् ५ ८		- B-15	12.5. 587-14 1.4/.5	18/0.5-23/0.5-34/0.5 (eafferd and t. gmoel)	57
- 58.	0 -	# # # #	<u>5M</u> 50	ΑΙΔΟΝΙΟΜ/ΚΟΣΣΟΝΙΟΜ	-5-14	2.44	end of table diriged. no 1x2 "void @ botto'm we = 5/6.113/4 oz. wf = 20/6 zoz.	
6-1			(ec)	59.0-67.5 CLAYEY GRAVELLY - SAND-SANDY GRAVEL: OK. RESIGN DOWN (SYRJA), W)	B-16		23/0.5 36/0.5 33/0.5 (w/40-50% graval)	69
62.0)		 &c	than above; 10-30% low-med. plastic firs; 20-30% to asmich as 50% mod well quided sond, more fv.t. sand for sandier intervals; 40-70% quicels from	5-15	Pb-15	end of tube dented We = 516. 14 oz. Wf = 2216. 11 oz.	
64.			(2C)	c. gravel, sonall coopes, mostly sandst., stst.; truce of vel. Fresh organic debnis (twips), carbonisod spots, specks; some	_B-17	5PT-16=	(missed blow count, wasn't anomabusly) - 10w)	
65.6				caco, in matrix; truck 94p() as f. custals; compactors; uncertain, probably moderate, uncertainted; slightly moist. 59.0-62.0 ped. sondy	-	RD I	RO through somewhat more gravelly section ROSMOOTHS SOME Q 42/2	
6 3.0		1		67.5-70.8 CLAYTONE/SILTSTONE:	B-18 5-19	5PT-17- 1.5/5-	30/0545/05 46/05	9/

PROJECT D118-SCS Dame () tah DATE DRILLED 10/7/81

____HOLE NO. WD-Z_____

ROJECT_	1118-10	S DAMS () tah DATE DRILLE	:D 19/1/8/		HOLE NO. <u>WO-Z</u>
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS
70.0	SICTSTONE/	67.5-70.8 SICTSTONE/CLAYSTONE: (con't) dk.gray to ol. quay (544/1,544/2); claystone w/shaloy, crude fissility, has somewhat sizazo look, thin Cally veins, w/few harder sttst.inclusions, Phy CondIntency factored, but factores are tight; low hardness; weak; mod. weatlered.	5-16 B-20	1.5, 2.5 307-18	end of twhe cleared up we = 516.13/402. wf = 1416.4/202.
77.0	-	Phy Cond Intentity fractived but Fractives or tight from hardness; weathered. 70.5-70.8 grades to olive! harder sitts fore / shak B. H. @ 70.81		1/2 2 2	(dnumg in rock) Terminated hole e 70.8 Terminated hole extraction sufficient datacquiad Backfilled hole whenthey and mud
	- -				
	-				
1 1 1 1 1 1 1	- - -	-			
‡ ‡ ‡	- -	- -			- - - - - - -
	-				
	- -				
‡ ‡			-		SHEET <u>4</u> OF <u>4</u>

		DRILLING AND SAM	PLING LOC	3		
PROJECT 4	11/18-50	S Mans. Utah DATE DRILLE	D 10/8/8	/	HOLE NO. <u>UA-3</u>	_
LOCATION	Worner	Amustan, usta. 15+45; 200 Ul	S GROUN	ID SURF	ACE ELEV. ~ 2940 (400)	<u>.</u>)
ORHLING	CONTRAC	TOR Pitcher () million Co. LOGGED BY	(_0/47	DEPTH 1	TO GROUND WATER <u>パケアやアペンジ</u>	prig
TYPE OF RI	G Failing	STOO HOLE DIAMETER 47/8 THA	MMER WEIGI	HT AND	FALL 140 10., 50	-
SURFACE	CONDITIO	NS loose dois land by ts; br	W.	EAIHER.	CHAP, ODERM, WINGS	-
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS	N Blows
0.0	ML	DEBRIS THEM LET WITS (?):	B-/	5T-1:	Astad 3" shelby tube; end cased in, sample loca and disturbed we = not weight of the spoon	
2.0	SM	reddish om (2.5 YR 4/4); 25-50-6340 low-nun-plastro finis; 50-6340 poorty graded fv.f. sand; - 20-70% armel, 14-1/2, prod.	8-2	1.2/1.5	23/0.5 29/0.5 31/0.5	60
4.0		10-20% growel, /4-1/2, prod. 10-20% growel, /4-1/2, prod. 10-20% growel, /4-1/2, prod. 10-20% growelly pac. to Helipping. 20.00 dense of the prod. 20.10/2/1/1/2/ less growelly		RO	RA w/4 T/2 "throng bit the qually section	
4.0		BEDROCK SHALE to SHALEY 5.0-13.6 SKISTONE CLAYSTONE		PB-1	Sample w/3"Pitchr Barnel end of tube lastly bent saved disturbed sample	<u>/</u> -
6.0	CL- ML	gayish ad (5RA/Z) when teshit	8-3	1/2.0	tube	
		most weathered; weres from v. could by to fairly well developed shall fissility, so its eiter	8-4	SPT-2 1.0/1.4	33/0.5 45/0.5 50/0.4 (hammarry in rock)	95+
8.0	¥2₩.	a shale or claystown to - sitistone; hadding not appoint from, samples examined; fine		RO -	end of two bant	
10.0	#\s	hackly, any or booky structure whom the toward. Phy condition of the chief to the condition of the front of the course, deeply to moderately weathered.	_5-/	2.3	we = 6.11/b.	
12.0		5.0-6.0 very deeply weath, decomposes sock; semilar sosily to low plastic clay-silt	8-5	5PT-3 RD PB-3	Champioring in rock) ROTO Clean hole	50+
14.0		clay-silt	5-2	2.52	ent of the o.t. we = 6.1016. wf = 22.56/6.	
, 7, 0		-	B-6	SPT-4	50/0. = -0. =/0. = (haminoring in NOK)	50+
16.0		<u>:</u>		PB-4	RD to 16.0' where nick steems slightly softer end of tube o.k.	
19.0		MITO goodes hard enough that it com't to mirolded (i.e., low hardress to mod. hard) B.H. @ 18.6	S-3	7-38/	we = 6.04 16. cuf = 27.78 16	
20.0			8-7	\$ P7-5	Soloil; 0-1/0.1 Themme ring on nock) Terminated has a 19.6 13.6 into nock w/soffward SHEET OF	53+
<u> </u>	, !	r.	<u> </u>	1	data acquired Back Frish Not af Entry	J

DRILLING AND SAMPLING LOG

		DRILLING AND SAM	IPLING LO	,
PROJECT_	0118-5	SCS Dams, Wtoh DATE DRILL	ED 10/8-9	7/8/ HOLE NO. 57K-/
LOCATION	Stucki	Dam: NSta. 14+06 on &	GROU	IND SURFACE ELEV. 2814 (10,00.)
DRILLING	CONTRAC	CTOR <u>Pitcher (Drilling Co.</u> LOGGED B	Y 0MY	_DEPTH TO GROUND WATER Not PACOUN \$25
TYPE OF RI	IG <i>Harlim, I</i> -	300 HOLE DIAMETER 414 H	AMMER WEIG	GHT AND FALL, 140 10, 30
SURFACE	CONDITIO	ins exist of earth embankment	-, 9100000 \ 666/4	WEATHER 2/ear, aarm
Ī		<u> </u>	1 7	<u> </u>
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE REMARKS blows
			ļ	
0.0	6M-		‡	Auger w/6" Hight auger
	SM		‡	AD + casy aspense in
	<u> </u>	SAND- SANDY EXAUEL: as	∄	AD = easy avgening in growlly section to
2.0	-	for 40-245 but w/40-60%	ᅷ │	+ ~3/2-4
	!	hard sandst. quel and small cobbbs from 14-6 al most 1-1/2 or 2.	Į	
	<u> </u>	most 1-1/2 or 2".	Ī	
1, 1	†	20 2001 2000 00000	‡	= augers firms ond
4.0		4.0-24.5 SILTY SAND: -	Ŧ	set up mud tub; casing
] 3	SM	Rddish brm. (5 YR4/4); 20-25% non-plastic fires; mos. guisk dilatancy; no-vilow thughrass;	<u> </u>	RD IRO W/tricons not bit
	<u> </u>	dilatancy; no-vilous toughness;	‡	(47/87)
6.0	<u> </u>	75-80% poorly goods fine tov! fire sond, sutary-sub may inn- oxide stoiched step 23.; 023	<u> </u>	Sample w/ 3/Pitcher Barrel
	†	oxide stocked step 23. ; 0=10	.‡	PB-1 dent in end of tibe
-	-	c. sand fine gravel, 44-1" radice sendst. and glossy volc., shall	‡	ue = 6.0416
	<u> </u>	1 all Subject to Subdest Delocation	+ 2-/	2.47 wf = 20.48 16
0.0	-	common as v.f. ckor cristal and	<u>+</u>	1 2.5 +
	[ineq nodules 16-1/2" + truce block car bonized organics / Mino string; dense; slightly moist to moist.		SPT-1 = 22/0.5 25/0.5 25/0.5 47
	[dense; slightly moist to moist.	F B-/	1.2/ = 22/0.5 22/0.5 25/0.5 47
10.0		8.5-10.0' ander slightly more	‡ '	
	<u> </u>	grave/14	<u> </u>	PB-Z two dents @ end of tobe
-	-		I	₩e = 6.1716.
	[‡ 5-2	2.19 Uf= 19.02 /6.
12.0	-	122-127'sinder silfier w/	-	2.2
	‡	12.2-13.7 gades silfier w/- 30-40% v. lowplosticfines	1	SPT-Z 18/0.5 20/0.5 24/0.5 44
	 		‡ B-2	14 (trace F. gravel) 44
1,70	[1	end of the dented,
14.0		-	Ŧ	PB-3 = Sample loss &
	<u> </u>		<u>‡</u>	$\pm \omega_c = 6.0616.$
			5-3	2.73 Wf = 22.64 16.
16-0	<u> </u>	-	I	2.8
	‡	16.5-18.0' w/3" long conc. of	‡	32/ 30/
	-	TYP. ignoses from sitty sand to relictean sand to sandy,	 	SPT-3= 22/0.5 29/0.5 35/0.5 64
	<u> </u>	sift w/ dipth-tayering prob-	<i>₹ 13-3</i>	1.2/ 5 (locally Gyp. Comen Rd) 64
18.0 -	<u> </u>	ably reflects fill lifts -	 	tout of the ak
	‡		‡	PB-4 Us = 6.06 16.
	F		5-4	2.44, Uf= 21.01/b
20.0	Ŧ [l	‡ '	2.5 SHEET / OF 4

r	VI.O =	DATE DRILLE	.U <u>79/9</u>	1	HOLE NO	- T
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS	
20.0	SM	4.0-24.5' <u>SILTY SAND</u> : (con4.)	5-4 (con'4) B-4	SPT-4	16.5 166.5 28/0.5 (scattered f.g.raufl)	44
22.0	- - - -	<u>-</u> -		PB-5	end of the curled over void & bottom of the up = 6.0716.	
24.0	<u>. </u>	24.5-51.0 ZLAYEY STLTY SP1.03	5-5	2.40	wf = not weighed	
26.0 -	5C- SM	dk. 46 (2.5 4% 3/6); 35-45% 10W -	8-5	1.5	(maybe poording into) (maybe poording into) 91041 & low sample)	84
28.0	-	inn-oride stained ste. 5-5% of a stained steen or 10 more, intrology as in pay, orit; pervatives cacos; gyp. as f. cristals, small irreg nodoks · V. little carbonized org. debris; some w/coude planar	5-6	PB-6 2.50	end of fully dented, sample 1005e We = 6.081b. Wf = 20.0516.	
30.0	-	faint personally due to fill compaction; dense; moist. 27.0-32.5 quarter - lightly more plastic fires	B-6		11/0.5 20/0.5 20/0.5 (scathed f. 9 must)	48
	<u>.</u>	more plastic times 28.5' grades locally to lois% c. sand, f. grovel	<i>1</i> 8−7		sample slipped fromthe while handling esurface, toke not demograd	
32.0	<u>-</u>	Note: % fines us 45 sand warres somewhat thinguehout	B-8	2.5	11/0.5 ZZ/0.5 30/0.5	
34.0	- -	unit over a.1-1.01 commonly, presumably due to- variations in fill lifts, as well as within a single lift essentially all- material observed in field	B-9	1.3/1.5	(scattered figures)	5Z
35.0 <u>.</u>	-	material observation teld is sc-sm	5-7	PB-8 2.45	end of the o.k. We = 5.9716. Wf = 21.4316	
380	-		B-10	1.5	15/0.5 21/0.5 29/0.5	50
40.0	-		s-8	PB-9	end of take o.k we = 6.02/6. wf = 20.25/6	
			B-//	SIT-9-	18%.5 26%.5 24%.5	50
42.0	-		5-9	P8-10 2.49 2.5	end of fuhe o.k. we = 6.00/6. wf = 21.62/6. SHEET = OF 4	

Ī	PROJECT.	<u> 400 - 3</u> 	DATE DRILLE	<u> </u>	7 7 7	HOLE NO. 377	1
	DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS	
	44.0	<i>S</i> :	1 / ./\	(5,4.)	(con4)	11/2 - 1 - 1 - 2/]
	46.0	5 5 ≥ 2 5 ≥ 2 5 ≥ 2 5 2 2	(con4.) 44.5-46.0 contains several stringers to 12" of clean mod. plastre (c2) clay	B-12	1.4/	11/0.5 17/0.5 26/0.5	43
•	48.0 -		M&. pio1776 (2) elay	- - - - -	PB-11 2.46	end of the o.k., thin annular space around sample we = 6.03/6. wf = 21-61/6	
2010	<i>5</i> 3.0 -		48.5 qualing shiplify more plasfic fines	13-13	SPT-11 =	15/0.5 20/0.5- 25/0.5-	45-
ign draw	50.0 -	SM	5/10-54.5 SILTY SAND: as	- ,,	PB-12	end of tube o.k. We = 5.9616.	
16 des	57.0 -		disonierd for 40-245; w/some thin layer amorning man platte, mon fines as for 225 451.01; very dense; slightly moist.	<i>S-11</i>	2.5	wf=zz.4416. 32/0.5 42/0.5 45/0.5	
	540		ΑιλυνιυΜ/ςοιλυνιυΜ	B-14	1.4) 1.5	10/0 end shift @ 5:15 10/9 start shift @ 7:45	87
	56.0 -	SM	20-30% v. low-non-plostifines; 65-75% mod postly gwied sand, most fourward sawe u.f. and mil	S-12 - -	1.20	end of the ak., slowh on top of somple we = 6.11 16. when the word weighed	
	58.0 <u>-</u>	5//e	o-5% f. grovel abordant 94P. as distinct small entrals (some size as sand); perusing Caco out u. strong Rase to Hel-w/some carbonist brack on which as pour	B-15	PD :	35/0.5 60/0.3 Chd. qup. nodule in shoe) RD to char commend	857
	60-0 -	Serve > M	betting to 0.05-1.5 seen in set samples defined by sondition and of fines compostress,	5-13	2.42	end of tube slightly lenked small void a bottom and loose we = 6.08/6. - wf = 20.32/6	
	<u>.</u>		making moist to moist. "570-63.0 cont quides more - quietly from 10-15% to as much as 30-35% /pcolly quiets	B-16	5PT-14 0.6%.8	45/0.5 50/0.3 (1/4" saillit. @ 10p, Ast garally)	957
	62.0	SP SM	63.0-73.0 SILTPSAND:			end of the o.t.	
	64.0	SM	Roddish orn. to red (2.5 YR 4/5); 20-25% v. low-non-plostic fines; 70-75% poorly gooded, pred.	5-14 - B-17	2.36, 2.5 5PT-15	uf z (forgot to weigh) 50/0.5 (no gravel)	50+
	66-0 -	 - sm	perussive Cacor; gyp. acf.cyshir; small module: fairly common;		RD P3-16	Ro in traft sand, slow and of the o.k.	
	<i>69.</i> 0		Non-plastic fines, litta or no man-c. sand or quouel	5-15	2.03	we = 6.08 16. wf = 17.72/6. SHEET 3 OF 4	<u> </u>

PROJECT DIB-SCS Dams, Otah DATE DRILLED 10/8-9/81 HOLE NO. 57K-1 DEPTH CLASS. FIELD DESCRIPTION SAMPLE MODE REMARKS SM 630-73.0 SILTY SAND: (con't) (55/ 4.) 59/0.5 50/0.2 0.7/0.7 68.0 109+ 8-18 RO to char qualsbugh 70-0 PB-17 Fend of Whe O.K. we = 6.09/6. 5-16 uf= 19.44 16 7.33 72.0 -PB-18 Fend of tube o.k. 730-750 SILTY SANDY GRAVELI as for 545-630, but w/5000% gavet 5-17 kg, hd, subarg., pred. sandst., stirty to 11, most 12-12" We = 6.12/6. 74.0-2.47: Wf = 20.5/16. 5PT-17 446-5 536.3 0.76-8 (a/qavals to 1/2") 75.0-75.8 <u>SILTYSAND</u>: as for 54.5-65.0 bt slighty more more SM B-19 94+ Terminated Auto 875.81 76.0 B. H. @75.81 W/SUFFicient data acquired Backfilled w/mud & cuttings

SHEET 4 OF 4

DRILLING AND SAMPLING LOG

PROJECT 0/18,505 Dame Utah DATE DRILLED 10/9-10/8/ HOLE NO. 57K-Z
LOCATION Stock' Dam, wifa. 17+85 on & GROUND SURFACE ELEV. ~2814 (topa)
DRILLING CONTRACTOR Pitcher Onling Co. LOGGED BY DMY DEPTH TO GROUND WATER MOTHER AT 1978 HAMMER WEIGHT AND FALL 140/6. 30 SURFACE CONDITIONS SENT OF PORTH EMBRETER 1978 WHY. CABBLY WEATHER Char WARM

AD Auger w/6" flight auger SANDY GENET. I HAISE Arm. PART JET ARM. AUGERNAL JEWELT SAND SANDY SANDY GENETAL ARM. AUGERNAL JEWELT SAND SANDY JET ARM. AUGERNAL JEWELT SAND SANDY JEWELT SAND JEWELT SANDY JEWELT SAND SANDY JEWELT SAND JEWELT SANDY JEWELT SAND SANDY JEWELT SANDY JEWELT SAND JEWELT SANDY JEWELT SANDY JEWELT SANDY JEWELT SAND JEWELT SANDY JEWELT SAND JEWELT SANDY JEWELT SANDY JEWELT SANDY J	DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS
SAMOY GERVEL: wedshifted. SYRAM) 2075 40 Annylath the SYRAM) 2075 40 Annylath the SYRAM; 2075 40 Annylath the SYRAM; 2075 40 Annylath the Syram the Colored in Most sandit and y colored in Most 12 2 the colored in Most 12	0.0				AD	Augerw/6" flight auger
4.0 - LAC SICTY SAND: RISISTA Brown (SYR 4 (a); 20-25% rooms A lastic first count delatively, Ano - u low fustions; 75-30% B.O lower formal part of the lasticely, Ano - u low fustions; 75-30% B.O lower formal part of the lasticely, I poor for the string, part of the lasticely, I poor formal part of the lasticely part of the lasticely of	2.0		SANOY GRAVEL: REJAICH Drn.			e pmuel, coldas to
Destric Fines Laste Chlarvell Mo-v how Asylvines traggell B-0 Deorly reself sample offer Deorly reself offer Deorly	7-0		4.0-14.0 SILTY SAND: RSdish +	: - :	RD	Augenay smooths @ N4' - set up to be casing to 4' NO w/47/8" totons bit
stringer of supply children of the supply of	S.0 -		$1 + m n - 4 \times i \times i$ $i \times i $	- B-1	SPT-1:	Onve standard splitspoon 6/0.5 15/0.5 23/0.5") Escattend F. quarel to 12")
10.0 10.0 10.0	9.0	- -	and as small individual orestals; V.	B-2	0.3/1.0	3"Pitchir Barrel sample
14.0 - 26.5 CLAYEY SILTYSAND: SC- dk. RSdish bm. (s-YR 3/4); x-45x SM low ploster finer; mod. show dilot. Smooth of the sand, more from the sands. show dilot. Smooth of the sand, more from the sands. show dilot. Smooth of the sand. more from the sands. show dilot. Smooth of the sands. more from the sands. show dilot. Smooth of the sands. more from the sands. show dilot. show dilot. Smooth of the sands. more from the sands. show dilot. show d			delisa, v. Fnara, signify	-	1 1	lower part of sample lows—slipped from the and bagged we= 5.72/6
14.0 26.5 CLAYEY SILTY SAND: 14.0 26.5 CLAYEY SILTY SAND: SC- dk. Redich bin. (5-YR 3/4); 35-453; SM low plother fines; mod. slow dilot. SM low plother fines; mod. slow dilot. 16.0 mc.; 2-7% pred. f. qwall of Sandst, shale; w/qyp. as fr prev. wint, but less abondant; prev. plostic in their layers; med. dense; moist. N170-220 grader w/5-10% qual w/ cloits to 3/4" B-2 10% sandst. 90 w.1	10.0					17/0.5 22/0.5- 32/0.5- (trace 44-3/0"9muel)
14.0 26.5 CLAYEY SILTY SAND: SC- dk. Redish bm. (5-YR3/4); 3-45% SM low plostic fines; mod. slow dilat 1 10.5 9/0.5 11/0.5 11/0.5 9/0.5 11	z-0 -	- - - -		- - ک-2		2"clast amoved from bottom and, sample loose Une = 5.8916.
16.0 SM 1000 plastic trines; 1003. 5/000 scaled will 1000 toughiers; 50-60% mod will graded soud, more f. th. f. than graded soud, more f. th. f. than mc.; 2-7% pred. f. growl of sandst., shale; w/qyp. as for pou unit, but less abundant; mod. por. to Hol. grades more or rest plastic in their layers; med dense; moist. 18.0 18.0 18.0 18.0 18.5 18.5 18.5 18.5 18.5 18.6 18.7 18.7 18.7 18.7 18.8 18.9	14.0		dk. Rodish bm. (5-YR3/4):35-458		SPT-3	
18.0 mod. 201. To crclights mode or less plostic in thin layers; med 2.5 18.0 losse; moist. 17.0-220 grades w/5-10% 19.0 loss 23/0.5 1.4, 15-10% sandst. 90.21	16.0	- SM -	gooded soud, move f. 4. f than	- 8-5	<u> </u>	
17.0-220 grades w/5-10% = 1.4, 15-10% sandst. 9001			sander, shale; 0/940. as to +	5-3	1/2.5	
	σ.υ –	-	~17.0-220' quader w/5-10% qouel w/ clotts to 3/4"	B-6	SPT-4	20/0.5-15/0.5 23/0.5 (5-10% sandst.90W)

	PROJECT DILB-SCS Dams	, Utah	DATE DRILLED_	10/9-10	/81	HOLE NO. STK-
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DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS
50-0	50-	1 - 1/1	‡	PB-5" (con't)	end of tube body dented
	I SM	(33.1.7)	Ī 5-4.	'-	$w_0 = 5.71/6.$
ZZ:0 -	<u>‡</u> │	_	F(con'7)	2.48	,
	‡		‡	SPT-5	11/05 13/0.5 16/0.5
•	[₹ B-7	1.3/1.5	(trace f. qual)
<i>34</i> 0 -	<u>‡</u> .	_		PB-6-	end of the smally bent
	∄ │		<u> </u>]	sample looce a bottom
	‡		5-5-	1.36	$w_0 = 5.93/6$
Z6-0 -	<u>‡</u> │		‡	12.5	uf=13.8316.
		Z6.5-34.0 CLAYEPSICTY SAND +	Į ,	507-6	13/6.5 13/6.5 23/6.5
-	SC- SM	Fines 40-45% : YOW (01)-400.	* 8-8	141.5	- • •
<i>28-</i> 0 -	/ ₅	plastre, mod. quick to slow dilat,	<u> </u>	PB-7-	end of the dented, sample quoued full
	‡ 5	graded v.f. sand, some f. sand;	(‡ _ ,		- Ano Ha
•	F 1,1,	story was to Hollaco-dinent	<u> </u>	7.6,	$w_c = 5.9416$. $w_f = not weighed$
30.0 -	90000 S	strong veac. to Hely Good disent in small, 1/6-1/4" nodu bs; little visible gyp.; dense-very stiff; Slightly inolst.	+ -	(0.7.7	6/0.5 9/0.5 13/0.5
_		Slightly inolst.	Ī 8-9	1.4	- 10.5 1/0.5 13/0.5
	‡ '		‡		
32-0 -	<u>-</u>	-	-	PB-8-	end of the o.k., three amund
-	‡.		‡ <i>5</i> -7	2.57	annular space around bottom of sample
<i>-</i>	∄ │		<u> </u>		wf ≥ 22.15/6
<i>5</i> 4.0 -	Sc-	CLAYEY SAND: dt. AND bown	+	SPT-8	11/0.5 19/0.5 23/0.5
-	SM to	1 (~un > /n) . ~ ~ ~ /n - + 140-265	‡ B-10	1.4/1.5	70.0
フィカ	50	but with slightly move fines, locally grading move plastic; slightly less mi-c. rand; denie-	‡——		end of the at, amulus
38.0 -	Ŧ l	to underse very stiff to hard;	Ī	1 ' 1	as for MS-a somala
	<u>†</u>	to undense, very stiff to hard; slightly moist to moist.	‡ 5-8	2.54	60012 @ 60160m We = 5.9116.
	<u> </u>	38.0 grades to only 50-553	[2.5	wf= 21.86/6
3 0,0	‡	sand, no gravel; sandf-vif. grained; fines mod. plastic	‡	SPT-9	11/0.5 16/05 45/0.5
•	‡		Ŧ <i>\\</i> -11	1-3/5	Coothum a 3/w/sondst.
40-0-	‡	ALLUVIUM /COLLUVIUM 40.0-43.5 SILTY SAND: dk	<u> </u>		end of to be best over
,	SM	12 dolsh brown (5-487/4); 15-20 %	∄ .	1 1	Last 0.5 gawly
-	ŧ ˈ	poorly anded sand, put, innoxide	I 5-9	1.96	wf=not weight
R.O -	<u></u>	trace to few to south, shole	‡	2.5	19/0.5 31/0.5 42/0.5
	<u>‡</u>	cutifalling moduli to be and	B-12	SPT-10	(no gravel) 10/9 end shift 5:40
הונית	F	poorly graded sand, and, inon-oxide could, subang-submd gle. grams; trace to few to sorbit, shale gravel to be gravel, cupitalities and subject to be and and and controlling appropriate cows directly and course slightly moist.		11.5	SHEET OF
44.0	<u> </u>	Cree descrition on next sneet)	‡ <u>5-10</u>	P5-11	3HEE1_4_0F_3

F	· · · · · · · · · · · · · · · · · · ·	T DATE DRICE	1	1		- T
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS	
44.0	M	I willish bon (SYRAIL) to OK. OM.		PB-11 :	end of tube bady curked	
]	F	(7.5 YR 4/4): 10-20% non-plastic fines, quick dilotary, no toughts s; 80-70% mod poorly	5-10	- 1	- we - 3.80%	
1//		fines, good dilotary, no	(con'f.)	2.13/	wf≥ 17:93/6	
46.0		foughts st , 80-70% mod poorly	-	SPT-11	27/0.5 45/05 50/05 Ctap f. aprel)	
:		some v.f., less in -c. quains;	B-13	1.1/2	(trace f. grove)	95
] -		some v.f., less on-c. quains; 3-7% fm. provel, locally more,	<u> </u>		RO to charginal	
40.0		sandit, shale jewa bedding evident in proportion of sand			<u> </u>	
] 40.0		I SIZOC AND GLAND CONTRACT BOOK .	F		SNd slightly dinged, annulus as For PB-9	
		o.1-14 Prick ; Al. slow hace to HCL; exceptoully high qup.	5-11	2.09	$\omega_{\rm e} = 5.7216.$	
1		content to 10-15-% or more	ŧ	7.5	wf ≥ 17-11/b	
520		Missible as sugary small -		_ I		
:		I inducted from hit still common :	8-14	0.8/0.9	25/0.5 50/0.4	75+
-	-	throughout; compactions concertain-	!	00	(1/2" graval in shoe)	'` '
=		Lings of constants oritisment			Ro to char gravel	
57.0		throughout compactness uncertain- due to profil on spis oritis very friote, uncomercial slightly moist to moist.	-	PB-13_	and of the ak; very small void @ and	
		AKO 11/10-15% 91042/@	ŧ l		small void @ ond.	
-	, 3	46.0' w/10-15% 91048 @	5-12	7-16/	$w_e = 5.8016$	
54.0	Jasi]	Į	7.5	wf=17.49/6	
34.0-	J. Coar	540' w/60% 71002 pend -		SPT-13	22/0.5- 50/0.5-	
:	၂ ၂၂ န		B-15	1.9/1.0	(quelinbottom ofspoon)	72
1 =	send, m	54.0-550 conter sand, more		RD	20 tocker gravel	
56.0		qual =		PB-4	End of tibe at small	
	<u> </u>	56.5-58.5 SANDY STUTT TO STUTY		':	end of table o.t., small void @ end	
	E		5-17	7.37	we = 5.9016	
	E MIL	60% low plastic fines; 40-60%		2.5	wf= 18.3816	
58.0-	' -	5AND :25 YR 414) Redish Om.; 40- 60% low plastic fines; 40-60%; ARE four soird, no quivel-much sugary 940., inc. downward; very dence to hard; slightly whoist.		-		
	E	dense to hard; slightly whoist.	, L	SPT-14	22/0.5 37/0.5 50/0.3	874
-	E GM	ISRS-SILLSILTY GRAVELLY	B-16	171.3	Cabundant aypsom)	` `
, -	[SM	CAND to SANDY SEAUSL: 45		RD	RO to clear hole inquial	
දිං. 0 -		for 43.5-56.5 but w/ 9/2000 -		DR-10-	ection	
		varying 20-70%, 12-3+",	5-14	1.57	end of to be church up,	
	(kd nck?)	of PE-15 and in SPT-15 could	- '	1.5	$u_{2} = 5.7/16$	
62.0	1,40,000	-	1	<i>^</i>	could a bottom end us = 5.71/6 uf = nitweighod	50+
		too shallow leviewing scs toundation borings.	8-17	SPT-15		301
		Tounda Hori	<u> </u>]	50/0.2 0.1/0.2 FROSON ON SIFT. COBY2	
	<u> </u>	B.H. @ 61.7	†	-	Terminoled 106 @ 61.7!	
-	<u> </u>	J. 1. C 9/1/	<u> </u>	-	-maybe who bedock	
] =	<u> </u>	· =	<u> </u>		[
] =	-]	<u> </u>		Backfilled hob w/	
	<u> </u>		[embortings and	
-	<u> </u>	-	<u> </u>		L CAN ADM WILLIAM I MOMONAL	
]	<u> </u>		‡			
-	<u>-</u>	<u> </u>	[-	_	
	<u> </u>		†		SHEET 3 OF 3	

DRILLING AND SAMPLING LOG

PROJECT <u>01/8-SCS Dams</u>, Utah DATE DRILLED 10/10/81 HOLE NO. 57K-3
LOCATION Stucki Dam, wsta. 12+79; 97 d/s GROUND SURFACE ELEV. 2782 (John)
DRILLING CONTRACTOR Picher Dilling Co. LOGGED BY <u>DMY</u> DEPTH TO GROUND WATER MOT AND TYPE OF RIG Failing 1500 HOLE DIAMETER 478 HAMMER WEIGHT AND FALL 40 16. 30 "
SURFACE CONDITIONS 10050. hushy spoil downstram to WEATHER char, warm

URFACE	CONDITIO	NS 10052, MUSHY 58017 COWNSHOOM	<u>n 786</u>	WEATHER_	crar, warm	=
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS	8/0 F
<u>8-0</u>	≤M	0.0-6.5 SKTY SAND: reldish	8-1	1.0/	Drove standard splits poon 10/0.5 15/0.5 17/0.5 Pushed 3"shelby tube	-
20	_	bm. (2.5 YR 4/4) -15-25-% non- plastic fines, goick dilataucy, - v-low-no toughness; 75-80% poorly grobed, fv.f. End; 3-6% qual, 44-34", 1nd sandr, shalo;	(5-1)	(ST-1 2.31/2.5)	adjacent to 57K-3 From . 0.0-2.51 We=8.01/6 Wf=17.45/6)	
40 -	-	qual, 24-3/4" Ind sander, shalo; pervasive caco; and small (18") mod-soft nodules of caco; trace qup(?) as f. crystals: loose of sorface to mod dense (7) dry to		KO #	Set up tub, casing to qu	
-1.0	_	v. slightly carp.		P8-/	Sampled w/3" Pitcher Barrel end of to be slightly dings b	`0
s.o -	_	7.0-6.5' grider slightly more Fries downward	<i>S</i> -2	2.27	We = 5.94'16. Wf = 19.28 16.	
s.o_	SM	6.5-170 SILTY SAND: dk. Redish brn. (5483/4); 15-25% non-plastic fires; 70-80% poulf gaded, prd. f. sand, some mice, juins 5-locally 25% 92226, 14-1" prd. sandst, rhale; much	<i>\</i> 8−2	SPT-2	22/0.5 24/0.5 25/0.5 (fine % hd. to deeply worth. fine sondst- 9 muil)	4
,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,	- - -	Juins; 5-locally 25% 92201 - 14-1"/pred. sandst. chale; much pervasive Cacos culvistini pac. to Hol, small robbes of Cacos			end of hibe slightlydinged we = 6.00/b	
0.0		world fairly comment gyp. as individe control and sugary to powdery mothing case, irregion layerty station whomston	5-3	2.33	wf= 19.28/6	
		to dense; slightly moist to woist.	B-3	SPT-3	10/0.5-11/0.5-14/0.5- (5-10% = sandf.gnuel)	2
20 -	-	10.5-12.0 qual to 5-100%		PB-3	end dinged, sample began slipping from tube	
40	<u>-</u>		5-4	2.34	us= 6.02 lb	
-	-	14.5-16.0 W/parvasive mottling of sugary gyp. throughout	6-4	SPT-4-	33/0.5 44/0.5 50/0.5- (gravel in shoe)	9
16.0 T		17.0-25.0 (SKAVELLY) STLTY SAND		+	end of the dinged in two places	
78.0 -	SM	dr. eddish orn. to reddish om (2.5 YR 3/2/4); zo-25% non-platic Fines, locally groding v. low plastic; 55-60% mod well graded to mod!	5-5	2.23	We = 6.01 16 WF = 18.4416	
20.0		poorly quided sand, pred. fm., quained: 15-20% quivel, 14-12; and most 14-1/4, pred. sondit, some size.	8-5	517-5	22/0.5 37/0.5 50/0.5 //5-20% gavel) SHEET_1_OF_3	٤

PROJECT_	11/18-	SCS Dams, Utah DATE DRILL	ED 18/10/0	18/ HOLE NO. 57K-3
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE REMARKS
20.0 22.0 -	SM	17.0-25.0 (GRAVELLY) STLTYSAND: (con't). + truce black carbonized on debn's: tempactness uncertain, very frable, uncomented; stightly most to moist.	5-6	PB-5 = endof x be o.k. We = 6.0316 2.50 = Wf = 18.74/6
<i>24</i> .0 -		118.5' grades to silty sandy quiel, qually sond 21.0-25.0' grades less quielly	0-6	SPT-6 = 896.5 7/6.5 0.9/1.0 (Few f. gnwls) RD RD to clear gravels PR-6 = Pad of the o.k.
26.0 -	SM	25.0-34.9 INTEDD. SILTYSAND: reddish brn. (5 YR 4/3). And vary 15-30%, non-plastic; sand varies 50-80%, pad f. gained; some mc.; quarel varies 5% for locally 50-50%, ascually 3-7%;	s-7	2.43 Wf = 5.93/6 2.5 Wf = 19.67/6.
2 <i>8-</i> 0 -		locally so so %, according 3-7%, pred sandst., some sitest., in 1". persosive cacoz: pyp. Rl. common as simple crystals and small nodules in some inknows;	<i>B</i> -7	SPT-7 = 14/0.5 28/0.5 45/0.5 73 12/1.5 PB-7 = end of take body denked
30.0		26.5' very little fines to 5- 8%; non-plastic 26.5'-28.6 for untold units	5-8	2.34 $W_{e} = 5.9416$ $W_{f} = 18.7716$
32.0		Mored w/varing tofine and varying sand sizes, all sms- 24.5 grades to silty sondy gravel-gravelly sond w/ = clasts to 1" most /g-3/0" quading c. sand	B-8	SPT-8 = 29/0.5 &5/0.5 71/0.5 116 1.3/ (9 mods @ top to 1/2") 116 PB-8 = end of to be o.k.
340			5-9	2.33 = We = 6.02/6 2.5 = We = 6.02/6
<i>3</i> 6.0 -	5M- 5P	34.9-49.5 SILTYSAND-SAND; orange-red (2.5 YR 4/6). 10-15 90 101-plastic fires, gukk Libitually no-toughness; 85-90% v. poorly graded fv.f. sand, pred. suborg-sub-md. gtz, 1001-oxide stained; o-50/o f. gravel; little Cacoz wl v. slight rac. to HCI.	B-9 -	RD W/AT/O +nicora Nock 5it searching For bedook
38.0	-	v.slight has to HCI;	_	
40.0	-		-	Rattling hd @40! = 5 moother @ 40.5! PD-9 = and of the o.k.
42.0	grades sitter		-5-10	2.04 = 0.03/6 2.04 = 17.92/6
44.0		~43.6 grades slightly sittler and w/ca(oz mottling common	B-10	2.5= 37/0.5 67/0.5 SPT-10- (A0 9:04), A0+ Q:1RA \$3) 0.7/1.0= SHEET 2 OF 3

DEPTH CLASS. FIELD DESCRIPTION SAMPLE MODE REMARKS

DEPTH		ss.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS
44.0		SM- SP, SM	34.9-49.5 <u>SICTY SAND-SAND</u> :		RD	RD w/47/8° tutoro searching for bedrock
46-0 -		377	-			
43.0 -					-	
SO-0-		6M- 5M	49.5-51.5 SICTUSANDY GRAVECTOR SRAVELLY SAND: astor 17.0- 25.0 but w/20-50% fmull, DED. SANDST., SONG SAND, 14-16 25.0 SANDST., SONG SANDSTAX.		-	hard valling 49.5- 51.5' in quarel
52.0-	‡ - - - - - - -		8.H.@ 51.5'		-	Termittak hole@51.51 - W/sofficient da fa, did not encounter bedrock
-			<u>-</u>		-	Backfill w/ mud and conficient watered
-	± ± ± ±		-		-	‡ ‡ ‡
	‡ ‡		-	‡. ‡	-	‡ ‡ ‡
	‡ ‡		-		-	‡ ‡ ‡
	-		-	‡	-	1 .
			-	‡ ‡		‡ ‡ ‡
	1 1 1			‡. ‡.		
	#			<u> </u>		SHEET 3 OF 3

DRILLING AND SAMPLING LOG

PROJECT D118-565 Dams, CHAR _HOLE NO.<u> & YP-/</u> LOCATION Sypsum Wash Dam, Nota. 33+65 an & GROUND SURFACE ELEV.~ 2740 (topa) DRILLING CONTRACTOR Pitcher Drilling Co. LOGGED BY DMY ____DEPTH TO GROUND WATER not encountered TYPE OF RIG Failing 1500 HOLE DIAMETER 47/8" HAMMER WEIGHT AND FALL 140 16. , 30 SURFACE CONDITIONS CHIT of conth emborkment WEATHER beezy partly cloudy, mod. DEPTH CLASS. FIELD DESCRIPTION SAMPLE MODE REMARKS blows FF 0.0 SPT-1 tone standard splitspoon EMBANKMENT FILL SM 55 B-1 10/0.5 24/0.5 31/0.5 0.0-5.0 512TY SANO: Vellowsh Red (5 YRA16): 20-25% non-plastic fines, quick dilatonay, us toucher? Tools u. pourly graded f.-v.f. sand; 5-10% gravel, 14-34", prid. Sand; 14-14", prid. Sand, no mouse 16-19" Cocos nod; 400 not appoint; chase dry to v. slight, moist to chase dry to v. 1.5-30 w/15-3% gravel, few colose 5-6"

5.0-17.0 CIAVEV SILTY SAND: Auger w/s"flight auger 2-0 they section w/c.666s 10/10 end shift @ 5:50 133/0.5 50/0.5 517-2 83 0.7/1.0 *(refusal on cobble) B-S 4.0 RO FRO to 4.5' to observable PB-1 - Sample w/ 3" Pitcher Barrel 5.0-17.0 CLAYEY SILTY SAND: tend of tube curled over reddish tock. Addish bown (5 YR 6.0 8/2/4): 40-45-96 low plastice fries: mod. wick dilataxy, low townings; 50-60% mod. well anded, pred. fim. sand, gtz, sandt, volc. (?), shale; We = 5.95/6 SM 7.46, wf=19.71 16. SPT-3 20/0.5 23/0.5 30/0.5 1.4, (trace sondst growt to 3/4") gtz, sandrt. volc.(?), shale; o-5% gravel, /a-5/; mostly opay-green strt./shale, sand mod strong pac. to Hol; nowle to 1/4" and individ crystals of 2.0 **B-3** PB-2 Fend of take stightly flaved; sample v. loose in to be; some grawity shough on top dense to very stiff-hard, slightly 10.0 We = 5.97/6 WF = 19.90/6. moist; confains occasional V. fine 5-2 Z.26; rootlets coppears to be cambic horizon material from borrow). SDT-4 10/0.5 14/0.5 22/0.5 36 1.2/ = (Rw gravels to 1/2-3/4") 12-0 - B-4 PB-3 Fend danked fairly badly we = 5.97/6 14.0 · 5-3 Ufz 18.90 16 16.0 to 5-10% 9 ravel SPT-5= 1/0.5 11/0.5 22/0.5 16.0 16.5 grading redder color more B-5 coans sand, more plastic SC Tone ding on end of tube 17.0-19.5 CLAYEYSANDY GRAVELE GP-We = 6.02/6 as for 19.5-25.9 but w/ 80-70% of mull, sandit, stiff trivale; 200% or more clayey, 10w-mod plastic. Wf = 18.33 16. 18.0. 2.03/ 12.5 SPT-6 + (13/0.5 20/0.5 21/0.5 SPT-6 + (400.5 Way gauly) 1-3/ SHEFT 41 19.5-23.0 CLAYEY SAND: SC 20.0 B-6 SHEET_

CSES DESCRIPTION ON MEXT STREET)

I		CS Dams, 1stah DATE DRILLI		′ı ı	
DEP _. TH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS
20.0	50		(8-64.)	SNT-6 (conif.)	
#	<u>:</u>	yellowish red (5) R46) w/buff fints	<u> </u>	PG-5	one large dent in and of tube
22.0	:	30-35% low-mod plastic tines. mod slow dilatoney, low-mod.			$\omega_0 = 6.01/6$
26.0	-	toughness; 65-70% mod- ush graded sand, pred. f-m. grained;	5-5	7.53/	= Wf = 21.41 16 :
‡		0-5% 7/2021; PRIDOTICE CALUS OF	‡	12.5	
1	5M to 5C- 5M	andvis Entstals, one vein a 40° x	•		29/05 46/0.5 50/0.3
24.0	= +3//	l	F & 7	1.2/1.5	(15-20% f. gnavel)
<u> </u>	SM	10.0 33.6 CERYEVELITOSEEIN:	Ē -	RD	no thin gravelly slough
1		LOW Mastic fines: slow-wed quick		P8-6	end of tope rolled arean
26.0-	<u>.</u>	datoncy, low toughness, 55-65%	<u>-</u>	-	= we = 6.0316
‡		datoncy, low toughness; 55-65%; mod well graded sand, pred. f-m. quined, atc., sandst., sitest, state; 5-15% gravel, 1/4-3/4", lith. as-	5-6	2.27	Wf=19.2416.
‡	-	5-15% gravely 14-3/4", lith as		2.5	<u>-</u> · · · · · · · · · · · · · · · · · · ·
z8.0 -	<u>:</u>	for sond; penasive cacos w/ : strong lac. to Hel; typ. commin		SPT-8_	12/0.5 17/0.5 30/0.5
‡			B-8	1.3/1.5	(Fur fire mudst. grave !
1	-	(1/6") white notiles; dt. 6 nown =			end of take has one
30. o _		this oritionse; slightly moist to moist.		PB-7	small ding
300 <u>T</u>	- -	75 000157.	S-7	2.45	- We=5.8016
#	<u>:</u>	23.0-22.9 Assiplastic, w/ f. gmul to 15-20%	[2.5	wf= 20.61/6.
		<u> </u>		- 1 +	10/05 14/05 27/0.5
35-0-	-	0.000.00.00.00.00.00.00.00.00	8-9	_ _ '¬	
#	<u> </u>	Alluvium/COLLUVIUM	·	13/5	end of the church op on
Ŧ	60-	33.2-37.0' <u>CLAYEY SANOY BKAUEL</u> as for 23.0-33.2' but times, are	B-10	18-8 ±	end of the churd up on control of on, not weighed
34.0	- GM	less plastic locally; quovel to, - so-60%, sandst, quy-renstly,			RD to 35 1/2 through
₫		brown shale, to 2-3", most /2-11-		1 RD \$	goane gravely section.
ŧ		30-90% Fc. sand; waceneared;		I	nard, nttling
36.0 	<u>-</u>	frable.			end of the cloud up; sample, tagged, not
‡		BEDROCK	B-11	103/15	weighed wigges, not
. ‡	(CL-	37.0-43.7 SILTSTONE SHALE:	- -		
z&. o- <u>‡</u>	- MY	when least weath, granishyny	13-12	0.7	(hammening in weath. rock)
Ŧ		dark polish brows proje shale where least weath, granish project (566/1) sitst, dr. reddish orn. (54R3/4) shale where most weath;	<u>-</u>	11.5	MIX W/sack of beatonit to
Ŧ	- P&	predominantly MILLE INSE	- 1	1.58,	Assen gravel caving end of tibe dented we = 5.9210
40.0-		some inclusions. Hain stringers of	5-8	1.5	We = 5.92/b.
45.0-1	2000 s	harder sitst; contains abundant- turn gyp. mothes, veins along partings; v. slight reac to Cacos-	8-13	SPT-11	76/05:0.5/0.5
#	- of &	partings; V. slight Rac. TOCO'CO3-			Ro to all to clean hole
7. ‡	l a	Phys. Cond crushes (due towerth); shale soff-frieble, stist mas. hd-his;	<u> </u>	PB-11 =	end of the slightly belled up = 6.9816.
<i>₹</i> 2.0 	1/2 1	shale soft-friable, strit may not his; shale plartic to frighte, slift mad, strong; deeply weath shale, cel tash slist	5-9		_ wf= 20.71/6.
‡	SHALE	SITST. O RMO/ds to CL-ML	- '	2.25	Terminated hole @ 43.7' Backfilled hole w/ mud and cuttings
44.0	13 8	B.H. @ 43.7'	B-14)	1/2.57	and cuttings

47.0-24.0 5/0 = 200 = 7/00 = 40 low to 100 = 5

50/0.2 0.2/0.2 SHEET Z OF Z

50/0.2 0.2/0.2

(hamnering in nock)

DRILLING AND SAMPLING LOG

PROJECT DI/8-SCS Dams, Otah Date Drilled 10/11/8/ HOLE NO. 64P-1A

LOCATION GYASIM Wash Dam, wsta. 33+55 on to GROUND SURFACE ELEV. 2740 (topa)

DRILLING CONTRACTOR Ptcher Drilling Co. LOGGED BY DMY DEPTH TO GROUND WATER not encorted type of RIG Failing 1500 HOLE DIAMETER 478" HAMMER WEIGHT AND FALL 140 16. 370"

SURFACE CONDITIONS CAST of AD THE Ambank went WEATHER 6APRY town by party chady, party chady, cool

					, cos / //.
DEPTH CL	ASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS
00	SM	MEANK MENT FILL Note: See log for boring SYP-1 for more complete description of units			Auger w/6" Aight ouger
2.0		description of units -	‡ ‡		set up tub, casing to 1' -RD w/478 "Unione bit
40	2	-			-
6.0	?- SC- SM	-			_
8.0		-		1 '6-1 I	sample w/3"Pitcher Barkl end of tobe dinged we = 5.91/6
10.0		9.6 clayey sitty sand; fines low plastic to 30-40%; sand to 60-70%, truce gravel w/ 1 clast to 1/4" color is poldish to dk. pd. brn. (5 YR 3/2/4)	5-/	1 40 1	. wf = 16.16 16 RD until drilling smooths somewhat @ 10.5'
12.0				/1.5	- '
14.0		12.0 as above but less reddish color, trace c. rard, f.grave]		├	RD until drilling smooths. Somewhat @ 13.5 end of tube o.t.; thin Tannular space around
		15.5'as for 12.0' B.H.@15.5'	5-3	2.17	Sample @ end of hole . we = 5.98 16. . wf = 18.90 16. -Terminally hole @15.5
6.0				1	w/two additional useall samples of embank mant
+		-	+		Backfilled help w/mud and cuttings
<u> </u>		<u> </u>	‡		SHEET/_OF/

DRILLING AND SAMPLING LOG

PROJECT D1/8-SCS Dams, Otah DATE DRILLED 10/11/81 HOLE NO. 57P-Z

LOCATION KYPSTY: Wash Nam. NSta. 20to1, 89' U/S GROUND SURFACE ELEV. 2725 (tors.)

DRILLING CONTRACTOR Photos Drilling Co. LOGGED BY MY DEPTH TO GROUND WATER And PROSENTE 20

TYPE OF RIG Failing 1500 HOLE DIAMETER 47/2 HAMMER WEIGHT AND FALL 160/16. 30"

SURFACE CONDITIONS 1000 depins born sports: bixh WEATHER 1001; portly cloudy; wirdy

		1	, V	r ·	ESSI; PSTINI ENOUGE ST	- "ፖ ፐ ላ
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS	610 F
0.0	511 M2	0.0-5.5 SANDYSTLT-SILTY SAND:	B-1	1.2/	2/0.5 10/0.5 13/0.5	
2.0		redish om. (5 YR 1/2/3); 4000% V. 1000-1000 No 100 F-V. 5 JODO 1/3 V. posty grides f-V. 5 JODO 5000 Some Solo grabel; perusine Cacor; some TYP-, mes. conse-very stiff; cry. ALLUVIUM/COLLUVIUM		An	Auger w/6" flight auger	
4.0	- 55- 6M	3.5-6.5 CLAYET to SILTY SPAVEL:		RO	setup mud-tub, set, surface casing to 3. RD w/47/8"thicora bit thru gravels to rock 8 6.5"	
6.0		BEDROCK	-		Refum fluid clarges golor for brown & 0.5 Sample w/3" Pitcher Farrel	,
	(CL- ML)	dark reddish brn. (5483/3) least		PB-1	end of tube benton one side	I
8.0		weath to dk. rel. om (5 1/2 a/a) - w/912416 tiet when fishert; agnin folly fished stock, locally queles to blocky silfetore; w/	5-1	2.11	we = 5.80/6 wf = 18.73/6	
10.0	<u> </u>	Type criciality modules and grot quer (quoriensus) veins, storyant on partings (ocally uslight to no seos. to Hel; quales harden	- <i>B-</i> 2	11.5	38/0.5 45/0.5 58/0.5 (hammering in nuck)	10
12.0	TSTONE/SHALE	cushed to intermed the condi-	: -	2-13/	end of the o.k. we = 5.8016 wf = 18.9316	
14.0 -	127272	(partly due to weath, drain) spr); godes from soff to low hardress w/ depth; grades from plasifis to weak w/septh; grades deeply to mod weathered w/septh.	- P-3	1.2/1.5	(hammoring in nick)	7
16.0	-		- 5-3	P5-3	end of fibe could mon one side we = 6-0316 wf = 18.8216	
/9. s	-		B-4	1 // 5-4	Champary in mock)	E
<i>2</i> 0. g			5-4	PB-4 2.32/ 2.5	end of tykes p.k. wf = 19.75/b. SHEET_ OF Z	

PROJECT D118-SCS Dams, Wah DATE DRILLED 10/11/81 HOLE NO. 64P-Z CLASS. FIELD DESCRIPTION SAMPLE MODE DEPTH REMARKS PB-4 (con't) 5-4 6.5-33.5 SILTSTONE/SHALE: 20.0 (con't-) (con4.) SPT-5 = 4/0.5 9/0.5 16/0.5 Z1.0-ZZ.Z' SQUEDI 1/8-1/18" thick, approx. honzontal layou of quay-green 13-5 1.2/ 27.0. m Fend has one bad ding, this hard zone near (AMPSITIONE) dayey silt; may have resulted in reli 1000 blow count in spr=5 SKTSTONE/SHAL end of NM we = 6.0316 24.0 5.5 2.47 wf= 21.1016 12.5 597-6 40/0.5 50/0.2 0.5/0.7 (Hrusul in 109k) 90+ 8-6 RO TRO to char hole 26.0. PB-6 I and of the dinged We = 5.79'16 Wf= 18.79 16 5-6 Z8.0. 2.0 39/05 50/0-1, 0-6/06 (HUSOLIN MCK) B-7 894 30.0 RD - RD to slightly smoother toster willing a so. s PO-7 Fend of to be agg-shaped we * 6.01/6 5-7 wf = 21.63/6 32.0 2.43/ SPT-8 # 876.5 B-8 87*+* (Hotusal in nock) B.H. @ 33.5 34.OT Terminated hole @ 33.51 27 into bidrock, cul sufficient data acquired Backfilled hole w/ bushn deposits

SHEET 2 OF 2

DRILLING AND SAMPLING LOG

PROJECT <u>N118-SCS Dams</u>, Otah Date Drilled <u>10/12/8/</u> HOLE NO. <u>6YP-3</u>
LOCATION <u>Gypsum Wash Dam</u> <u>NSta</u>. <u>26+83 on & Ground Surface ELEV. NZ740 (Gond)</u>
DRILLING CONTRACTOR <u>Pitcher Drilling</u> Co. <u>LOGGED BY DMY</u> <u>DEPTH TO GROUND WATER NOT ENCOUNTED</u>
TYPE OF RIG <u>Failing 1500</u> HOLE DIAMETER <u>478"</u> HAMMER WEIGHT AND FALL <u>HO 16.30"</u>
SURFACE CONDITIONS <u>CRSf</u> of worth embank went WEATHER <u>Drewy</u>, partly cloudy, cool

DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS 6
0.0	SM	0.0-20 SILTY SAND: yel red.	B-1	SPT-1 1-2/1.5	Drove standard plitspoon 10/0.529/0.530/0.5
2-0 -	SM	(5784/5)25-20% n.p. tines; 75-90% v.ff. sand; 0-5% giovel; perusive co cos; dense; dry to slighty moist. 2.0-5.0 SILTY SAND: Redish bm.		A0-	Auger w/6"Aight auger 13/05 14/05 26/0.5
4-0 -		(5 YR 4/4); 25-35% non-plastic fines, 65-75% poorly graded, f- v.f. sano; 0-5% gravel; perasise cacos w/ strong space, to Heli gyp.	8-2	1.2/	(scattered f.gmwl) RD to coor scalled gmwls
	SC-	dense; slightly moist, moist, 20' some 3-5" cobbes, c. quoue!		1,2,1	Sample cu/3"Pitcher parroll
6.0	SM	5.0-17.9 <u>CLAYEY SILTY SAND:</u> Addish to Sk. Addish brown (5YR : 3/2/4); 35-45% low plastic fire;	S-1	2.48, 12.5	conside; small void @ Lottom and of tible we = 5.7816 uf = 18.4516
8.0	= = = = = = = = = = = = = = = = = = =	pred. f v.f. grained, less mc.; pred. f v.f. grained, less mc.; o-5% gravel, ky-b", sandst.graf	B-3	14/1.5	20/2 5 37/0.5 45/0.5 (Scottered f. gravel)
10-0	-	cacor: gyp. tairly common as small typ-14") nodiles some mottos; also as v. small clear crystals; contains some black carbonized organic debets some pl. fesh to partly carbonized nootets soms (only trace ant. as small pieces);	- <i>S-</i> Z	PB-Z 2.34 2.5	end of the curted in we = 6.011b wf = 20-0216.
12.0	-	med dense to dense; slightly moist to moist.	B-4	SPT-4.	7/0.5 13/0.5 15/0.5 [(10-15% c. sand, f. gaval)
· · · · · · · · · · · · · · · · · · ·		iens wie road in ST-3, others may be present locally anywise in onit		PB-3	Sample slipped From tube while pulling rods; end of tube not damaged
14.0				2.5	,
/6.0 -			5-3	2.19	one ding in end of tube we = 5.80 lb wf = 19.27 lb.
18.0	- sc (sc)	17.9-23.0 CLAYEY SAND: St. RUDISH OM. (5 YR 1/4); 25-35%	B-5	5.5- SPT-5-	19/0.5 21/0.5 33/0.5 (bottom o.9/15 000 H) weath. silfst. cobble)
Zo. 0 ¹	SC	nod will quality sand, willinger clasts of sonsit, quid-green state and dk. brn siols; 5-10% quiel	5-4	PB-3 = 2.5/2.5	end of the coved in, small void & end SHEET OF Z

PROJECT 0/18-SCS Dams, Otah DATE DRILLED 10/12/81 HOLE NO. 64P-3

22.0 27.9-230' EATER SAND (Conf.) 17.9-230' EATER SAND (Conf.) Some of a felt some off a star of the star of th

DRILLING AND SAMPLING LOG

PROJECT 0118-505 Dans Utah DATE DRILLED 10/12-13/81 HOLE NO. FH-1 LOCATION For Hollow Drm, NSta. 11+26 ON & GROUND SURFACE ELEV. 119 4 (tops.) DRILLING CONTRACTOR Picker Drilling Co. LOGGED BY DMY DEPTH TO GROUND WATER not excounted type of RIGENING 1500 HOLE DIAMETER 478 HAMMER WEIGHT AND FALL 140 16. 30 in.

SURFACE CONDITIONS Ap of embankment WEATHER clear, cool, 6 Rezy SAMPLE CLASS. FIELD DESCRIPTION MODE REMARKS DEPTH SPT-1 Drove standard split spoon EMBANKMENT FILL 0.0 -ايح (scattered f. gravel) 1/1.5 6/0.5 19/0.5 23/0.5- 42 ML 0.0-5.0 CCAY-SILT: RIGHT B-1 bra (5 YR 4/4); 95+ % 100 AD Flight auger to 21/2 plastic fines: slow-mod. guick plastic fines: slow-mod. guick dilat-, low toughness: 3-5-% f-v.f. sand; tace c. sand, f. quauel: strong rac. to HCI throughout: truce visible 94P. as whit specs, small (2 133) soft nodules: hard; dry to slightly moist w/depth. 2.0 . SPT-2 = 27/0.5 24/0.5 24/0.5 48 PB-1 sample 3/3 "Pitcher Barrel B-Z 4.0 One side cased in onend 2.0-5:0 godes sondier to 5-1 5-15% f-uf sand more visité app; moist trace carbonised organic debnis We= 5.96/6 -2-2.43 W4= 19.5016 ML 6.0 SPT-3 11/0.5 19/0.5 28/0.5 47 5.0-32.5 SANOY CLAY-SILT. 1.9/5 (Scatter & gave) dk. Addish brn (5 YR3/4);65= CL 75% low plastic fines, mod guick to slow dilatancy, low-8.0 PB-Z fend to be gaved in an mod. toughness: 25-35% poorly one side m.-c. sand variable; 0-15-20% 5-2
quavel locally, usually 23-5%
scattered madomly: unit is
interlayed as indicated on We = 6.07/b 2.34 WF=18.1216. 10.0 interlayered as indicated on log; truce to 1-2% locally of carbonized arganic makinal as mottles, Flacks small chips to 18", some pl. firsh twin, wot? frugments noted; modification to Hall throughout; agap as which, sugary irregulations from 132-12 pl. common some horder adulation SPT-4=1/0.5 13/0.5 15/0.5 28 0.75 (used spoon w/split CL-B-4 ML 12.0-PB-3 PB-3 cut smooth, even; sample slipped from tube while pulling mods 14.0. common, some harder nudules also (qyp. or qypsifenus stist?); very stiff to hard; slightly moist to moist. PB-4 DAR ding in end office.

part of sample pulled

2.51 off bottom, having n 6.5-7.1 quades to sitty f. Sand w/some fine bosatt quavel clasts 16.0 -1113 void 06 16. 4 Wf = 20.67/6. cH_{i} 10.5-11.2' contains a 0.05' thick CL-CH clay lens SPT-5 13/0.5 18/0.5 20/0.5 38 13-5 18.0-17.0-18.3 confains several CL-PB-5 and of the ak. Y8-1/2" CL-CH clay stams ML Wr = 20.25 16. 3 SHEET_ OF 3 5-4 20.01

* arbifrany datum

PROJECT DILA-SCS Doms, Utah DATE DRILLED 10/12-13/81 HOLE NO. FH-1								
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS			
20.0	2L ML	5.0-32.5' SANDY CLAY-SILT:	5-4 (0014)	PB-5- (con't)				
22.0-		22.0 quades move sandy) -	B-6	1.3/	8/0.5 11/0.5 19/0.5 (3/4" gravel in shoe)	30		
Z4.0 -		fine quavel locally cul common black clayey organic defins as mother, irregular, small pode to	5-5	2.37	end of tube o.k. We = 6.071b. W= 19.75 16.			
<i>z</i> 6.0 –	5M- ML	NZ4.0-Z5.5' grades to silty V.f. sand to U.f. sandy silt- W/ MON-plastic fines; some 94p.(2) visible	<i>B-</i> 7	597-7 131	17/0.5 15/0.5 19/0.5	34		
20.0	ML	mod hd. noder 18-12" rel-	5-6	18-7	end of tube o.k. We = 6.041b.			
28.0	,-	28 5-20 6 avoidor of a muelly	-	7.03	cuts quively @ 28/2			
30.0	waravel CL	28.5-29.6 quides to gravelly, sandy clay-sift w/cbslis of basalt, slist- to 1/2" in non-low plastic fines matrix	E-8		(165 14/05 20/05 (top 06' w/qna 1 ho /2")	34		
32.0 -	ML	29.0-30.4 contains thin stringers of sitty f. sand and cl-ch elay 32.5 w/some c.sand, f.gravel	<i>S-</i> 7	264	end of tube o.k. We = 6.0316 Wf = 21.6516 (driller purrent by o.14; depth not adjusted on log)			
34.0	- CL	37.5-48.5 SILTY CLAY: POSTUTE brown (5484/3): 95+46 10W- mod. plastic fires, 510W	13-9	SPT-9 1.51 1.5	- 8/0.5 13/0.5 18/0.5-	3/		
34.0	-	o-s-% v.f. sand; black organic clay to carbonicad organic debris Rl- common	5-8	PB-9	end of the o.k. we = 5.9816 uf = 19.8516			
39.0	- - - - - -	Patches, specks to 14" mox. dimen: mud. strong rac. to Hel throughout due to Cacas; occasional relifush fine root	B-10		8/0.5 13/0.5 17/0.5	30		
40.0		or twig remaint to 14" lorg- (only saw two); ayp. as fv. fine off-white specks al. common; very stiff to hard; slightly moist to moist.	5-9	PB-10 2.47	end of tube o.k. uk = 5.9716 uf = 21.0016			
42.0	- CL	41.4-41.8 quades to	B-11	1 //.5	8/05 16/0.5 ZZ/0.5	<i>38</i>		
44.0	CL	sandy silfy clay w/ - 10-20% fv.f. sand	5-10	PB-11 2.57	end dinged on one side cut quavely @ ~43' SHEET Z OF 3			

ROJECT_	0118-5	CS Dans, Wah DATE DRILLE	D 10/12-	-13/8/ HOLE NO. FH-/
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE REMARKS
44.0	<u> </u>	32.5-48.5 SILTY CLAY: (con'f)	-	(con't) We = 5 99 10.
24.0		of coars sond size quains, trace v. fine 944.	13-12	587-12 7/0.5 13/0.5 20/0.5 3
46.0 -	-			PB-12 end of two o.k. w= 5.9816.
48.0 -		485 grades slightly less plastic	5-11	2.25 Wf = 20.35 /6.
	CL- ML	485 grades slightly less plastic and slightly more Ane sand and scattered coarse grains ALLUVIUM KOLLUVICH 485-58.0 SANDY CLAY-SILTS	B-13	5PT-13 10/0.5 12/0.5 14/0.5 Z 1.2/ 11.5 10/12/01 end shift 05:05
50.0-		redish brown (54R4/3); 80-90%. 10w_mod. plastic fines: slow-mod. quick dilatancy, 10w-mod. toughruss; 10-20% skip-goded-sand, pred. V.f. trained cu/		RM?) 10/13/81 start shift @ 10:10
57.0 <u>-</u>	vois s	3-5% as ni -c. quinta sopreia		hok; chan hok to 50.5; w/no Hourn fluid; put PB into hok for run #
540	o de la companya della companya della companya de la companya della companya dell	bosait; trace f. amuel pred. basait; small off-while specks; nodules to 16" of ayp (3) rel. common-some irridisent black MHO(?) coatings on		to no resistance lowering sampler until ~55 Ino Heturn during sampling
3 -1-0 -	cL-	grains, blocky small soil peds: some black carbinited organic chans; occasional fine, relifiesh rootlets; pervasive cacos as		0.2(?) From; ROURR & 0.2-may 4.5 br sidewall or from bottom 4.5 of non; non PB-14, W/10 PB-14 Rourn fluid; total floid 1055 > 250 gals.
56.0	[ML 	indicated by strong reactor HC 1; conde, angular blocky structure. — (soil bed development?); very	5-12	1.43 - PB-14 - end of tybe o.k. We = 6.0016. We = (not weight)
58.0-	typamb Sunders	stiff; slightly moist to moist. 8.4. @ 58.0'	B-14	5PT-14 10/0.5 25/0.5 50/0.5 1.0/1.2= (gravel clast in shoe)
-		50.5-54.8 wid or very soft material -see amarks column for description		
-	-	of behavior during drilling/sampling	<u> </u>	
		56.8-58.0 grades increasing = 9 ravel down ward to ~10 - = 20% @ 57.4; clasts 14=		
-	- - - - -	slightly calcareous gtc. sandstone, It. gray-green (qypsiferous) siltstone;		
-	- - -	nent & top of sample; 0.15 basalt clast at bottom of sample, quides	-	<u> </u>
1	-	deeply weathered to bel. Fresh downward; top of bedrock(?)	·	
1				
"				\$HEET_3_OF_3_

DRILLING AND SAMPLING LOG

project 0/18, SCS Dams Utah Date Drilled 10/13,14/8/ Hole No. FH-Z
LOCATION From Hollow Dam. NSta. 10+52, 106'U/S GROUND SURFACE ELEV. N 89'K (topo)
DRILLING CONTRACTOR Pitcher Drilling Co. LOGGED BY DNY DEPTH TO GROUND WATER Not PACOUNT FOR OF PROJECT OF RIG Failing 1500 Hole DIAMETER 41/8" HAMMER WEIGHT AND FALL 146 16, 30"
SURFACE CONDITIONS SOFF, losse. Sediment in Asensir WEATHER partly Soudy, windy, cool

DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS
0.0	C1- ML	0.0-15.3 CLAY-SILT: dk.	8-1	SPT-1:	10/0.5 14/0.5 19/0.5
2.0		reddish brown (5 YR 3/4); 95+9 low-mod. plastic fines; mod. quick dilatancy, low-mod. toughness; <5-% U.f. sand;		RD-	set tub; surface casing
4.0		no gravel; trace of gyp. as c. sand-size soft to mod. Ad. nodules; occasional Hel. fresh.	<i>B</i> -z	SPT-2:	7/0.5 8/0.5 10/0.5 1 Samph W/3"Pitcher Barrel
4.0		toother's twigs, trace black carboniced dipris; some		PB-1	$end of fuble o.k.$ $w_{\ell} = 6.02/b.$
6.0	_	cacos disseminated throughout indicated by strong reac to Heli- stiff-v. stiff, med. dense; by	5-/	2.5	Wf=17.72 16.
8.0		to moist w/ depth.	B-3	SPT-3 1.2 1.5	1/0.5 8/0.5 14/0.5 2
	-		8-4	PB-Z	Losing drill fluid to formation, but maintain circulation @ ~6-8' All but 0.3' of sample
10.0	5M	10.5-11.2 lens of silty fine quarte sand		0.3	tube while compy out of hole (may be sue to
12.0 -	2L- ML	120-145 contains a few -	B-6	SPT-4:	3/0.5 7/0.5 8/0.5
11111		v. thin (1-2mm) silty f. sund lenses	 	-	Fortion of PR-3 slipped from two at surface- Rmander loose in two-
14.0	- - w/	14.5-15.3 contains common tain beds, lenses of silty f. san	8-8	2.5	Kept in plasfic bays
16.0	- ςΜ - ςΗ	15.3-16.0 CLAY: V. dk. gray to black (N3-NZ); modhigh plastic fines; no sand argravel; spongy-	B-10	1"3/2	2/0.5 4/0.5 5/0.5 S
- - -	SM CL-	feel, burnt-organic odor(may) be or-oH?); contains few feld fash vootleds, twigs, lense of partfally decayed org. debis; firm; moist. Alluvium Kolluviut	≤-2	17/1.0	wf=(slowh on top, not stop ps-4 wighted) not on
18.0	ML,	Firm; moist. ALLUVIUM KOLLUVIUM 16.0-17.0 SICTY SAND: mod. 6rn. (SYRA); 35-90% Tow-non- plastic fines; so-65% fm.	8-11	1.5	3/0.5 16/0.5 16/0.5 3
Z0.0		quained sand, some c. quains; no quaud; compostness uncertain; in-situ moistur uncertain.	5-3	PB-5	PRO to 19 to chan grave foot of hole still losing floid Thave lost ~ 500 gal SHEET OF

PROJECT DILA-SCS Dams	. Hah DATE	DRILLED 10/13, 14/	181 HOLE NO. F-4-2
	, -	_ , , ,	

ricolor	<u> </u>	DATE DRILL	I	1 1		- T
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS	
20.0	= GC	17.0-20.5 GRAVELLY, SANOY	5-3	PB-5	end bady cauld one side	
		LIAY-CUT to CLAYEY-SILTY	(con't)	(cont.)	$\omega_0 = 5.72/6$	
	E SM	GRAVEL: multi-colored clasts:	E		WG= 15.50 /b.	
ZZ-0 -	F	in dr. Reddish byn. (548514)	‡	SPT-7	6/05 16/0.5 Z1/0.5	37
22-0 -		matrix; matrix of low plastic -	B-12	1.1/	(a)/10-15% f-gravel	37
	F ML	1 1/2/200 /2-10 /2 20100 (119/1 4 2000)	‡	1.71.5	(w/10-15-4) f- gravel	
-	ŦΙ	quall vanes 10-60%, anch yill sandst., sitst., basalt /4-1/6, all moddeeply woothered; one w/Ca(Oz; some qyp. (2); some carbonized org. debris; Allow	<u> </u>	DO-4	end of the o.k. sample	
70 D	‡	sandst, sitst, basalt 14-118)	‡	'^ "	is notably soft wend, -some slough on top of	
29.0 -	Εl	W/Ca(Oz - 50 me 94D. (2); some	5-4	1.17	- some slough on top of sample	
	ĖΙ	carbonized org. debris; Rl. low	‡	2.5	$\omega_{e} = 5.5416$	
	Ŧ	COMPACTURIZZ / MISTON OF INCIDIO	₽ I	2.5	Wf=12.65/6	
Z6.0 -	‡	uncerfaid.		SPT-8		
20.0 -	Ŧ l	20.5-21.5 SILTYSAND: Similar	B-13	109	-6/05 7/0.5 K/0.5	23
	‡	to 16.0-17.0' but w/squent % c. sand, f. quavel; fines non-	‡	1.5	(w/soma f. gravel)	
	ŦΙ	plostic.		DF-7	end of the badk caued	
ZB.0 _	‡	21.5-32.0 SANDY CLAY-SILT:	‡	1,22,1	end of the bady caved in small wid at end	
20.0 =	F	(same as 48.5-58.0'in FH-1); Reddish 6.0 wn (548 4/3); 80-90%; low-mod-plastic fines; 10-70%; pRd. fm. grained sand, locally		2.28	Tot tube	
	ŧΙ	Addish 640 am (54R 4/3); 80-90%.	5-5	25	$w_{e} = 5.5416$	-
]	Ŧ	prd. fm. grained sand, locally	F	7.3	Wf=18.6716	
30.0	<u> </u>	less; locally trace f- growel; some black carrowings organics, -		SPT-9		
	-	I TENED AND THE MANUAL TOURS .	B-14		6/0.5 6/0.5 7/0.5	10
-	Ξ Ι	some Mno (3) contras on grain; - trace gyp. as white nodules to -	<u> </u>	1.5	. 70.0 70.3 70.3	13
:	†	Your gyp. as white nodures to	F 1	P5-8	end of tube o.k.	
37.0 <u> </u>		I TATA BLOCK DID. CLAY STUMEN I	<u> </u>	1,00		
	‡ •	I U-CNUR and Blacky STNUTURY	5-6	12.62	$w_{c} = 5.6216$	
-	- _C	grades hard to stiff w/ depth; grades slightly moist to mout, - locally wet.	1 1	1 /257	•	
:	± 1°2	locally wet.	<u> </u>		still losing fluid; have used another soo gals	
340-	<u> </u>	21.5-25.0 locally to 10-15-92.	Ł l	577-10	- Osto another 300 gais	
	‡	sandst., sitst. glavel	B-15	1.5/	8/0.5 10/0.5 12/0.5	22
_	<u> </u>	29.5' wet, soft end of	<u> </u>			-
:	‡	P8-7	‡	PB-9	10/14/81 start shift 08:30	
36.0	I .	black org. clay stans	<u> </u>	'-	Driff Fluid Standing (2)	
- 0.0	‡	1	5-7	~/95/	-zo, encounter slough @	
_	?-	37.0-57.0 SILTY CLAY: dk.	I ∣	/ <u>.</u> I	mul 7-216 ON H-00/	
	60	low-mod-plastic fines, slow	13-15	SPT-11	doug casin, to zq. RD	ļ,, , ,
38.0 -		dilatancy, mod. toughtessiss = < 5%, V.f. sand; very stiff;	<u> </u>	0.4/04	-up walls.	100+
	-	<5%, V.f. Sand; very stiff;	<u> </u>	1 +	PB-9-endofflife bant	
-		MOIST.	<u> </u>	RD]	- over chame of ab	
-	1 66	37.0-39.4 CLAYEY GRAVEL:	∄ l	1 ‡	we = 5.6716	
40.0	‡.	1 14 d. 6111. 5140 clay, clayey 201	<u> </u>	1 -	wf=(not unighed)	
] ,		<u> </u>	‡	SPT-11 100/0.4 (pounding	
	<u>† 15</u>	sand; basati clasts mod little weathered, clasts 44-2/2.	<u> </u>	-	on gravel or tock)	
	245457	389,39.4 giades more clay]	‡	RO. 37.6-38.8 nuch	
47.0-	18	39.4-43.0 BASAGT-401. FAS. A.	<u> </u>	-	- drilling, gravelly 38.8-39.4 smoother	
	‡ `	39.4-43.0 BASAGT-10/fiesh, - hd. basaff coffings, no clay	}		50.8-54.4 SMOOTRING 20.8-42.0 recky	
] 		‡	, ,	Losing some fluid duning Plus terminate hole @ 431 w/	
44.0	‡	B.H. @ 43.0'	1	1 ‡	SHEET Z OF Z	
<u> </u>	T L		t,l	_	UT UT]

sufferent as a regard Backfill w/ softing, mud

DRILLING AND SAMPLING LOG

PROJECT DII8 - SCS Dams, Wtah DATE DRILLED 10/14/81 HOLE NO. FH-3
LOCATION Fine Hollow Dam NSta. 1/+19 ON & GROUND SURFACE ELEV 119 * CHOPS.)
DRILLING CONTRACTOR Pitcher Onilling Co. LOGGED BY OMY DEPTH TO GROUND WATER NOT PROUMBLY
TYPE OF RIG Failing 1500 HOLE DIAMETER 4 1/8 " HAMMER WEIGHT AND FALL 140 16. 30 "
SURFACE CONDITIONS top of earth embankment weather char cool

		•			
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS
0.0	CL- ML	EMBANKHENT FILL 0.0-5.0 CLAY-SILT: see some Merval in FH-1 For detailed		AD	Begin set-up @ 11:30 -Auger to 4'w/5"flight
2.0	1	description		-	-
40 -		5.0-32.0' SANDY CLAY-SILT:		RO	set mud tub, set surface casing to 1' stort RD w/47/8" tricone bit e 11:54
6.0	- ML	see interval 6.0-32.5 in FH-1, for defailed discription; note variations below. 6.5 W/some gravel, some			slight chather of/z'
8.0		8.5 grades sittler w/some fine sand and black		 	- - -
10.0		omanic clay 9.5-10.0 gndes back more plastic 10.5 less stiff			drills fackr Onloss
12.0	1		-	-	- -
14.0					<u>-</u>
16.0	-	16.0' w/gypsam; trace of _ coarse sand		-	
18.0		18.0-19.0 grades siltier, - a little more stiff		-	- drills slower 18-19
20.0					SHEET 1 OF 3

PROJECT DIB-SCS Dams, Wah DATE DRILLED 10/14/81 HOLE NO. FH-3

\ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \	<u> </u>	-363 Vami, () 74 / DATE DRILL	1		
DEPTH	CLASS	. FIELD DESCRIPTION	SAMPLE	MODE	REMARKS
20.0	∠U M			RD	
<i>24</i> 0 -		28 gray-black clay lens 24 grades gravel, coards sand in clayer sitt, low plasticity matrix. clasts of yell 6 mm with		-	- ratting and slower drilling @ 24
Z6.0 -		sandst. 24.5-27.0 gades less c. sand figurel and more fire.f. sand		1	- - - -
zg.0 -		scattered f. gavel of garangay state, bown sandst.			:
<i>3</i> 0.0 -		30.5 quades more plastic withour fi-c. sand grains	‡ ‡		<u>-</u>
32.0-	C/2	37.0-46.0 SILTY CLAY: SEET INANAL 37.0-48.5' IN FH-1 for Sefailed discription; see blow for local variations.			- - drill rate slows @ 331/2"
<i>34</i> .0 -		32.0 gales mod. plasticul- aliw-black arganic clay mottling or lenses 33.5 grades stiffer			- bit balled-up out clay-dislosque 34
36.0 -		36.0 cuttings of gray wl- black charcoal tragments			-drill rate increases
38.0		37.0' some fibrous black organic debris incoffings 38.0' as e 36.0'			slightly © 37' -
40.0	 - - -	_			_
42.0	 - -	-			_
44.0		43.5 grading slightly sander without mc. sand	‡		SHEET Z OF 3

PROJECT 0118-SCS Dams, Ofah DATE DRILLED 10/14/81 HOLE NO. FH-3 N blious CLASS. FIELD DESCRIPTION SAMPLE MODE REMARKS DEPTH 32.0-40.0 SILTY CLAY: (CON'4.) 44.0 complete RO 6.45' @ CL Sambibi Sopither Barrel, end tube slightly dinged, slight chatter a 46 ALLUVIUM/COLLUVIUM PB-1 142 46.0-56.5 SANDY CLAY-SILT:
Addish brn. (5 YR4/3): 80-90 %

tow-mod. plastic fines slow to,

mod. quick dilatancy, low-mod.

toughtess; 10-20% sand, prd. t.

V.f. quanted variable of of m-c.

sand; tace f. queel as basatt,

stast, some sandst, some clear

to smoky quarte(?) quins, small

closts; locally culspect, some clear

Mnoc?) coatings wins, trace of

black carbonized opanic debris,

locally tack of tel. Fash mothets;

pervasive cocos w/ strong tochom

to ticl; crude and blocky structure

sosgesting soil ped develop ment

very stiff; slightly moist-moist
466-471 sevent cores to 46.0-56.5 SANOY CLAY-SILT: + 5-1 1.4 | We = 14.90 lb, stop rur

SPT-1 | Drive standard stit spoon

1.4 | 9/0.5 23/0.5 34/0.5

1.5 (w/coarre sand, f. gravel)

smoother & 48 winn

PE-2 | Chasing spT-1 winn

PE-2 | Chas 46.0 U-ML 53 B-1 48.0 one bad dent end of PB-Z 2 CLMI We = 5.63/6 256/ *5*-2 Wf = 20.70 16. fachtd interal 2.4-50.0 SPT-Z + 9/0.5 13/0.5 13/0.5 Z6 B-Z 17/5-CL-B-3 52.0 ML PB-3 and badly caved in on an side (lane gravel) 466-431 sown cores to 0.1' of >50 % c. sond fignish, w/white agate notable Z.zl/: We = 5.5916. 5-3 7.5 I WF= 18.49 16. 54.0 ~49.0-51.1 color to H. Rd: brown (25 YR 5/2/4) W/
Hacked appearance; V.
Strong Raction to HC/ SPT-3 = 8/0.5 12/0.5 13/0.5 25 51.1 color grailes to Action

Sim. (5 YR 3/2), thus more
Brown than previously

51.1-51.8 w/20-250/0 v.f. to

f. sand B-4 0.7/5 [(few gaves, one 34-1") 560-PB-4 = end mush nomed-publishly pushed cobble and in top front of function in 2.0 came to abript halfe 57.9 on bisalt C2 gravel 0.01 54.8' gndes more plastic, uf g 3/4-1" basalt clast some Agreen sitst fragments 58.0 56.5-57.9 STLTY CLAY: similar to above but less sand more plastic; w/scatters bazilt graveland-RO RD 57.9-60.0 in hard rock; no floid loss B.H. @ 60.0 Terminate hole e 60' 2-1'11to rock w/ sufficient data acquired 60. O BEDROCK 57.9-60.0 BASALT: gray to black; al. Fresh to unweathered Backfill hole w/mud 4 cuttings; hd; drills oniformly suggesting all unfraçtored (?); no day binder noted. - cultings

SHEET 3 OF 3

DRILLING AND SAMPLING LOG

PROJECT D118-SCS Dams Utah DATE DRILLED 10/14/81 HOLE NO. FH-4
LOCATION Frostbollow Dam, USta. 11+33 On & GROUND SURFACE ELEV. 119 * (topa)
DRILLING CONTRACTOR Picker Drilling Co. LOGGED BY DMY DEPTH TO GROUND WATER MOTERAL
TYPE OF RIG Failing 1500 HOLE DIAMETER 47/8 HAMMER WEIGHT AND FALL 140 16. 30"
SURFACE CONDITIONS top of farth Embank Ment WEATHER Skar, cool

			•	_	
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS
0.0	<2- ML	EMBANKMENT FILL 0.0-5.0 GLAY-SILT: SER SOME interval in FH-1 for detailed description.		AD	Auger drill w/6" Flight auger to 4' where c. quavel or couble is encountered
4.0		4.0 grayel or single cobble of -			=set-up tub; surface
۷٥ -	CL- ML	5.0-32.0 SANDY CLAY-SILT:- see interval 5-32 in FH-1 for chtailed description; not local uninations below.		RD .	= set-up tub; surface casing to 1; no mud mixed RD w/47/8 "tricane bit
8.0		6.0' quality more plastic, Ad- om. color locally 7.8' lowe quel clast _ 8.5' note Alack again clay			hd. natta @ ~7.8
10.0		straks in cottings 9.5' group clast, stringer of silty f. sand w/med= cours sand common			buef with, for Ar
12.0		12.0 quades locally to mod. Yel. brn. silty clay		-	
14.0		14.0' some med.c. sandin - cuttings	-		<u>-</u>
16.0		16.8 bocally gravelly	-		rattling @ 15.9'
18.0		19.0 culbbox, probably carbon- 120 organics charamount),			-
20.0	J	trace mc. sond giains	‡		+ arbitrary datum

PROJECT DIR-SCS Dans, Vtah DATE DRILLED 10/14/81 HOLE NO. FH-4

DEPTH	CLASS		FIELD DESCRIPTION	SAMPLE	MODE	REMARKS
Z0.0	- W	_	5.0-32.0'SANOY GLAY-SILT:		RO	<u> </u>
	M	4	(con4.)			
22.0	-		22.0 W/Aw coarre sand size white agates	-		
24.0	-		23.5-24.0 grades w/fv.f. sand, sither (Aciplastic), little mc. sand	-		-
26.0-	-		25:0 gades w/more mc.sand 26:0' w/some gay-great s/fst. frags	-		 - - - - - -
28.0	- - - -		37737.77.893	-		<u>-</u> - - - -
	<u>:</u>			-		<u>-</u> -
300	- - - -		29.5 wilmore cm. sand size - quains of sitst., basalt, trace f-growl size clasts	_		-
32.0			31.0' quart clast 32.0-47.5' SILTY GLAY: 59e			nth 8 31
34.0-		_	32-0-47.5 SILTY GLAY: See interval 32.0-20.5' in FH-1 for detailed description: see local variations described delow. 32.0' guides mod. plastic, ol- om-quy color w/6/k. om. clay mottling	- - -		
36.0				-		- - - - - - -
38.0	_			- -		
40.0			39.0 W/common black, Partially carbonized on the six (some identifiable as small twigs, shoots)	-		- - - - -
<i>\$2.0</i>	_			-		
44.0						SHEET 2 OF 3

EP,TH CL	ASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS	6
4.0	CL	32.0-47.5' SILTY CLAY (con't.)	2	· + <	ample al 3"Pitcher Barn	7
8.0			= 5-/	P23-1	ride, wo as of sample outlied off bottom of sample at surface	
‡ ‡		ACCOVION/COLLUVIUM	Ī	2.57	we = 5.60 lb ouf = 20.09/6 nue standard split spoon	
8.0	CL- ML	FH-1 and FH-3 IN Similar in Anals willocal variations as	‡ ,	15PT-17	19/05 32/25 30/0.5 some baself gravel to 34-1	
δ. 0 -		47.5-49.0 several basalt gravel clasts to max of		PB-2 5	end of the body cared in mail void a bottom of sample - encountry. In ill = 5 40 into run	é
· ‡		314-1" 49.0 W/appaciable f-v.f. Sand (Stringer or Ans.)		2.5	wf = 19.18 16	
7.0		51.5 cobr grades to dt. Ad. 6m (54R 5/4); w/ Buff caloued gyp. mottles A 10" w/mois to u conten	Ŧ <i>B</i> -Z	1////5	5/0.5 8/0.5 9/0.5 2 314" basatt clasts)	,
4.0		Successive moist to come	I	1 89, 1	nd of type ok.; some slough (?) on top of was = 5.6216	
<u>+</u> + ده		clasts; some pale-brn; inclusions to JI#" uf ligh ca (05 content light to Accept inkno from 49.0-511 in FH-3)		SPT-3	wf=17.47 16. 5/0.5 13/0.5 27/0.5 2/4-1"basalt clast in	
‡ ‡	CL	545' some v.t., R. + W.S. n rootlets 100ted	B-3	//.3 +	told drilling @ 57 a/ forcome nock bit and ess hard to 57.71, h	, es
₽° 	(6C)	white 94P. W/SPACKS A 14	<u></u>	PB-4 = 5	atting 87.7-53.1/HA	"
0.0 	grachs moi	mod to addish om (syr (13); To 90% mod plastic fires; 5-10% f-c. sand, put basalt; 5-25% pred basalt gavel; w/ both gyp locally common and your fain	5-4	RD =	for ping in the of the for pis-4 run; basilf at top of tube we = 5.99 lb. wif = not weight	2
7.0	BASALT	situ moistul confint uncertain 60.051.0 states to weatherd basatt Walay binder BEDROCK	<u>;</u>	1 +	wf = not weights 20 is slow in thing, from 61.0-63-0; no fluid los	1 5
‡ ‡	BAS.	61.0-630' BASALT: little weath for Fush, hard, no fluid loss.		+ + -	Terminak hob 865, Zinfo hurd bedrock	, :
1.0		B. H. @ 63.0'	‡	+	Backfilled hole w/ mud + cuttings	
#			‡	‡		

DRILLING AND SAMPLING LOG

PROJECT 0/18-5C5 Doms, Utah DATE DRILLED 10/15/81 HOLE NO. IV-1

LOCATION Diversion Dam #5 ~ Sta. 12+48 on & GROUND SURFACE ELEV. ~3189 (topo.)

DRILLING CONTRACTOR Pitcher Drilling Co. LOGGED BY OMY DEPTH TO GROUND WATER MOTE encountered by type of RIG Failing 1500 HOLE DIAMETER 478 HAMMER WEIGHT AND FALL 140 16. 30"

SURFACE CONDITIONS top of dirt embantment WEATHER partly cloudy, cool

SURFACE CON	ΙΟΙΤΊΟ	ins top of dirt embankment	\	WEATHER	partly cloudy, cool
DEPTH CL	. AS S.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS
0.0	SM	EMBANKMENT FILL:		SP7-1	Drove standard splitspion
‡	500	0.0-19.2' SICTY SAND: Rd	8-1	1.2/5	10/0.5 30/0.5 37/0.5 67 (trace gravel)
7-0		(10x4/6); 30-40% non-plastic;		AD	E CANGE ANNEW
		coase silt finas; 60-70% poorty graded, pred. fu.f. sand, w/ Rw olo mc, sand grains;			
1		+mace f. quaual; Cacoz dideniraled	8-2	377-2	40/0.540/0.5 7/0.5 77
4.0		specks dere to vidense dry to	3-2	17/1.5	(truce stest gravel clasts) Set up mud tub; casing to! Sample up 3" Pitcher Bornel Cond o. K.
		U. slightly moist (may be due to drill floid); some introls lightly		PB-1	
		cemented w/cacoz	5-1	2.54,	we = 5.62/6 wf = 20.24/6.
6.0			5-7	2.5	- 7 20.2470.
		-		SPT-3	:13/0.519/0.5 22/0.5 41
			B-3	1.2/	
8.0 +		7.8-8.0 yery abundant 94P		PB-2	end of the o.k.
‡		as white morning.		1 18-2	we = 6.5916
			2-5	2.34	Wf=19.4316.
10.0		10.3' 3 f. slfst. clasts to -		2.5	-
‡		11-0-11-6'much gyp. euident			10/0.5 17/0.5 35/0.5 50
12.0		17-0-17-0 Moon 9 4 p-10 Km	13-4	1.5	
/2.0				PB-3	sample v. loose in take
🗼		13.0 grades slightly more		=	we = 5.5116.
14.0 ‡		sandy, more grousland mc. grained sould Gusast	5-3	7.17	wf = 16.63/6
		,		2.5	140.5 30/0.5 21/0.5 51
]	B-5	1.1/	Custon-c. sand tace
16-0			<u>-</u>	1/25	= f.gavel)
1		<u> </u>		PB-4	tobe while pulling nos; end of tube not
		·	<u> </u>	0.0/	end of tube not damaged
18.0		18.5-19.2' grades to med. desse		2.5	Losing, a little water to
‡		GOLLOVIUM		SPT-6	hole starting NIG-17/
Zo-0	ML	19.2-21.0 SANDY SILT: as above but 50-50 To non-v. low plastic fines; med. denis.	B-6	1.2	3/0.5 6/0.5 12/0.5 18 SHEET OF 2
· · ·		,	<u> </u>	<u> </u>	

I	1	DATE DRILLE	<u> </u>		HOLE NO. 20-7
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS
20.0	ML gavel	19.2-21-0 SANDY SILT = (con'f.) 20.0 9 00 des 10-20% c. sand, f. gnovel 21.0-27.0 SILTSTONE: dk. Redish brown (2.5 YR 3/4); slfst braks down to low-mad. plastic	5-4	PB-5: 1.3aj 1.5	end of tube dinged; sample loose in tube 8.5916 W = 12.70/6 RO to ZZ.0 to Guargavel
	SILTSTANE	siffy clay-clayer sift; 74P- very prevalent as evertals and privasive strangers, mothles; slight	<i>₨</i> -8	1.5	10/0.5 25/0.5 27/0.5 52 (gravelor stist. Deottom of spoon) tube mod badly dinged
240 - -	38	angular blocky structure. Phy. Cond crushed; soft to Friable; plastic to friable; deeply weathered (hard in soils terms) 26.0' ang. blocky structure. 27.0' Gyp. crystals 14-1/2"	5-5-	1.83	on end we = 5-6916 wf = 16.3916.
26.0 -		B.H. @ 27.0	B-9	SAT-8:	13/0.5 71/0.5 60.5 5/000/2 en top: mol. Aard stfst. w/2412 @ End of span
28.0 -					Terminate Note 027 W/SOFFICIENT data, no Sinto belock Backfill hole w/sothing
					embankment fill "
-					
1				1	
-					<u>-</u>
-	<u>-</u>				,
			_ _ 		
-				-	SHEET 2 OF Z

DRILLING AND SAMPLING LOG

PROJECT <u>DI/8-565 Bms</u>, (Hah Date Drilled 10/15-16/81 Hole NO. <u>IV-2</u>
LOCATION <u>Diversion Dam #5</u>, <u>NSta. 20+00 on & Ground Surface Elev. N3189 (tops.)

DRILLING CONTRACTOR <u>Pitcher Drilling Co.</u> LOGGED BY <u>DMY</u> DEPTH TO GROUND WATER <u>MOTERIAL</u>

TYPE OF RIG <u>Failing 1500</u> HOLE DIAMETER <u>418</u> HAMMER WEIGHT AND FALL <u>140 16.</u> 30

SURFACE CONDITIONS top of dirt embant ment WEATHER saftered clouds, Gool</u>

URPACE	CONDITIO	NS 700 OT CIFT RMEAN MENT	v	VEAIREK.	57047471 810001, COOT
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS S
0-0	ML- SM	0.0-10.5 SANDY SILT 651LTY	B-1	5PT-1:	Prove standard solit spoor
2.0	-	SAND: Rd (10 R4/6); fines 45-35%, non-plastic, guick dilateral	-	A0 -	
		no toughness; sand, as-55%, very poorly quided, fivef. quained, two-oxide stained, subary to sub monded ato.; trace of c. sand to f. quart size sitest.	<i>G</i> -2	SPT-Z	26/0.5 28/0.5 406.5 6 (+mice sttst. gavel) 10/15-end shift @ 5:30
4.0		Rac. to Hel in disseminated GOO. I dense to v. dense (may be due to 1)		PB-/	sofup tubicasing to 1 sofup tubicasing to 1 somply 3" A form Barrell Part of sample slipped
6.0	<u>.</u>	slight convention by catos), dry to slightly moist (drill fluid involves sample relatively easily).	8-3	1.91	-Rmainder bagged
		6.5-10.5 940-, < 1-2% Visit & , several sm (18-14) Modoles	B-4	507-3	22/0.5 31/0.5 32/0.5 6 (scattent sltst. fings)
8.0	- C.				end of the o.k. $uk = 5.6716$. $uk = 18.7216$
10.0	gra des	ALLUVIUM/COLLUVIUM (?)	<i>S-</i> /	2.18	
12.0	SM	above, but with 40-45% tines	B-5	SPT-4-	18/0.5 21/0.5 27/0.5 4 (souther f. s. 14 st. gavel)
2.0	- 3/* - -			PS-3	end of the o.k. Probably washed port of sample during 18 run
14.0	<u>-</u>	140-180 nottled gyp. conc.	<i>S</i> -2	145	Ue = 5.72 lb - Wf = 14.13 lb
		14.8 thin lens w/215% - coard sand as sithet. chips	<i>B</i> -6	spr-5-	6/0.5 9/0.5 4/0.5 Z
К-O -		Sui b 2		P8-4	end of to be o.k. We = 5.62/b. Wf = 18.12/b.
18.0-	-	18.5-21.5' stratification	<i>√</i> -3	2.26	(o. z' slough on top amved- w/c. sand size sitst. chips common)
20.0		apparent with thin lenser- thom 14 to sev. inchens alternating more or less sand	<i>B</i> -7		2/0.5 4/0.5 10/0.5 1. SHEET 1 OF 2

PROJECT DILB -SCS Dams, Wah DATE DRILLED 10/15-16/81 HOLE NO. IV-Z DEPTH CLASS. FIELD DESCRIPTION SAMPLE MODE REMARKS 8/0.5 12/0.5 17/0.5" (C. sand s. Hst. frags to 1/8" = common) 10.5-22.0 SILTY SAND: GON'T) 20.0 SM 8-8 11.54 PBS end of toke o.k. wo.3

of sample stipped from

tube sould in 8-9;

lay gravel or gravelly shiph

2.5 an top of sample?

we = 5.5716

we = 16.2616 ZZ.O-Z4.1 SILTY SANDY GRAVEL ZZ-0 SPred (10 R 4K) with brownish fint; 10-15-96 NON-plastic Fines; 15-20% mod well graded sand; 65-75% ppd. Fine gravel from 1/4-1/2", mostly Flat shale and submunded to wunded sandst; : (8-9) 24-0 SPT-8 = 9/0.5 9/0.5 15/0.5 1.81 (c. sand as for SPT-7) SM SM-SP B-10 very pervious. MESM Terminate hole @25.5' 24.1-25.5 SICTY SAND: as -w/sufficient info. 26.0 for 10.5-22.0; medium dense; stratification apparent in Backfill w/cuttings variations shown in CLASS. see logs of IV-4 and column on log-IU-5 for more info. Z8-0 B.H. @ 25.5' this boring SHEET Z OF Z

DRILLING AND SAMPLING LOG

		DRILLING AND SAM		
PROJECT_	D118-5	CS Dams, Wah DATE DRILLE	D 10/16/	/8/ HOLE NO. ZV-3
LOCATION	Diversio	n Nam #5 ~ Sta. 36+11	GROU	UND SURFACE ELEV. ~3/89'(100)
DRILLING	CONTRA	CTOR Pitcher Drilling Co. LOGGED B	Y DMY	DEPTH TO GROUND WATER not encounted a
TYPE OF RE	G Failing	<u>500</u> hole diameter <u>47/8 "</u> ha	MMER WEI	IGHT AND FALL 140 16., 30 "
SURFACE	CONDITÍC	ins top of dirt embantment		WEATHER partly doudy, cool, basey
,				
DE D.T.I.	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE REMARKS
DEPTH	CLASS.	FIELD DESCRIPTION	JAMPLE	Dows standard split span Ft.
0.0	1	EMBANK MENT :		SPT-1 8/0.5 16/0.5 30/0.5 5Z
2.0	ML		B-1	1.21
	SM	0.0-10.8 SANDYSILT 651174	[1.2/ (c. sand to F good stist
2.0	;	SANO: 12 (10R4/6); 45-55%		AD = Shot
2.0		non-plastic fines: 45-55% V. poorly graded fv.f. grained; Iron-oxide stained gtz sand;	F	1 ²⁰ ‡
]		Inn-oxide stained etc. sand;		SPT-2 = 30/0.5 30/0.5 27/0.5 67
	;	1-3% c. soud to f. gravel size - SHST. Ishale Frogments; qyp. common	B-2	I <u>+</u>
4.0		I ar son he direction to the to		1.4/1.5 Sample w/3" Pitcher Born!
7.5		iraq mottled veins strong reac.	<u> </u>	PB-1 and of hibe dinged on
-	-	iraq. mottled veins: strong reac. to Hol w/abundant disseminated cacoz; lightly comented due to	_	
		I CX CO2: Now Of Wish A A A A A A A	5-/	2.50 = We = 5.70/6
6.0 -	-	stightly moist.	-	2.5 Wf = 20.13 16
	:	3.0-10,8 gyp-as soft to		<u> </u>
] -	-	low had nodules from hove	5 3	SPT-3 18/0.5 37/0.5 53/0.5 90
		1/2" fairly common, have white-pray color, sugary texture	5-3	1-9/ (Fau 1/4" sthot / shale gravels)
8.0	-	texture -		
				PB-2 out gravily
	-]		1
• • • • •	-		5-z	2.10 the not weighed
10-0-	-	COLLUVIUM (?)		2.5 (Ap a. 5' of sample is
]	- Soft	10.8-17.0 SILTY SAND and	- L	SPT-4 33/0.5 53/0.5 0.1/1.0 (deeply weath, stist, clost to 80
	SHST.	SILTSTONG: Rd(10R46) ::	B-4	1.0 Carepy whath strif clash to
12.0	- SM	matrix w/wak ad(10x4/) :	<u> </u>	PB-3 end of the baby and
'('		matrix is silty sand w/20-30%		in; top looks or
-		non-plastic fines, otherwise as above; ~50-70% of unit is	5-3	2.15 tue = 5.65/6
	;	deeply weathered sits. closts		12.5] Wf = 18.44 16
14.0 -	-	w/accordant Vat. 940 Crystok.		<u> </u>
]]		clarts are soft, weak, start	B-5	SPT-5 27/0.5 30/0.5 52/0.5 82
	-	borders, range 14-1/2 contains gyp. nodules and unothling, matrix eacts to Helivery denies to hard; v. sliphtly moist	[N-3	1-3/5 (as for SPT-4)
]		Racts to Helivery dence to hard in		PB-4 cuts V-slowly
16-0-	-	BEDROCK		PE-4-cuts V- slowly
<u> </u>	:	17.0-19.5 SILTSTONE: 4d -		1-75 = We = 5.57 16
	L _D	(Inpale) w/ number black fint.	5-4	2.5 Wf = 15.96 /6.
18.0-	- 3	RI. clean sitist was sand organial.		(Hose has a little slough
		small crystals, mod Rac to HC/;		SPT-6 (Roy Ad any State block 1819) 53
-		RI. clean strst. Wo sand or growth. clasts; abund gyp verning, mottling, small crystall mod lear to Heli crude and blocky structure; some mod he cares. Phycond constited weak to finable, plathe to triave; ude applyweath.	B-6	
	- "	10 4 more; udaspy weath.		1.2/ 5 Tormingte hole @ 19.50 Wheet OF
[20.0]	r I	B.H. @ 19.5	<u> </u>	
				Backfill hole of fulfings and embankment fill
				with A. American

DRILLING AND SAMPLING LOG

PROJECT DIB-SCE Dams, Ofah DATE DRILLED 10/16/81 HOLE NO. IV-4
LOCATION Diversion Dam #5 ~ sta. 19+62, 43' U/S GROUND SURFACE ELEV. ~3179 (tops.)

DRILLING CONTRACTOR Pitcher Drilling Co. LOGGED BY MY DEPTH TO GROUND WATER not excounter of type of RIGFailing 1500 HOLE DIAMETER 4/8 HAMMER WEIGHT AND FALL 140 16., 30"

SURFACE CONDITIONS 10058 Others basin deposits WEATHER SCOttlevel shule, cool

		NS 1002 & CHUND DOZIN CROSSITS		JEATHER SCOTTERS GRUNT, C	
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE REMARKS Drove stondard splits	be poon t
0.6	5M- ML	ON-ZO CHTUSAND tO SANDY	B-/	SPT-1 1/6.5 5/0.5 7/0.	5
2.0	SP. SM	SILT: Rd (10 R4/6); 45-55-60 MORT plastic fines; 45-55-40 V-poorly quaded fu.f. sand; frace c. sand to f. gravel; 44p. mothling; much disseminated Caco; med	-	AD Auger w/6" flight o	
4.0	219	2.0-30.0 SAND to SILTY SAND:	8-2	SPT-2 3/0.5 6/0.5 8/0.	
4.0		Fines: 85-95% v. poorly graded fines: 85-95% v. poorly graded fines to 1-30% widely scattered 5-500d to f.		PB-1 and 1/2 st bentonite rosing to so water to so	And rel,
6.0		quavel sized sitet sunder. clarks; subtle stratification locally doe to variations in the pister and south size; generally little usible ayp.;		o.o sample slipped from 12.5 while pulling rod	
8.0 -		strong reac. to Hal throughout; Uanes loose to mostly uned dense; V. slightly moist ceasify perstated by still fluid). by still fluid).	B-3	14.5	
-		oy drill thoras abound mothed ayp. 7.0-8.0 quides settler		PB-2 Anted on one sid We = 6.07 16.	
10.0		<u>-</u>	S-/ -	2.5	;-
12.0			B-4	0.7/.5	
			<i>S</i> -2	PB-3 Thicked a sand, f.g	nul f
14.0 -		4.5' some soft gyp. noted e bottom of PB-3		1-55 Slove on the of se 1-55 Slove on the of se 1-55 Slove on the of se We = 5.50 /6 Wf = 14.22 /6 SPT-5 5/0.5 8/0.5 11/0.5	
16.0			P5	0.7	
- - - -			<i>B</i> -6	PB-4 and of the dinge most of sample sting from tube or was away, sampling; Ra	ned l
18.0-		 		2.5 - sample tagg & c' we = not weight	३ ७
Zo.0	 		8-7	SFT-6 = 6/0.5 7/0.5 12/0.	

PROJECT.	· ·	DATE DRILLE	<u> </u>	, , ,	HOLE NO	ı
DEPTH	CLASS.	FIELD DESCRIPTION	SAMPLE	MODE	REMARKS	
20-0	SP- SM	2.0-30.0'SAND to SILTY SAND: (con't.)	5-3		end of tobs o.k. ; a little slough on top of sample; either wasted formation or part slipped from tols pulling vods Luc = 5.51 16. Luf = 12.07 16	
24-0-			<i>E-8</i>	1.1/	wif = 12.07 16 5/0.5 9/0.5 11/0.5 NI' of IN-situsond was RHO away while roing No clean hole	20
Sq.0 -		- -	-	PB-6	Sampl from PB-s either washed away and or slipped from tube white pulling 10 ds	
z8.0 -			B-9	SPT-8 0.8/	wf = nothing toweigh 9/0.5 14/0.5 23/0.5 (sample recovered as liquefied soury-sees-9)	37
30-0-	C.L.	35.0-37.5 SANDY SILTY CLAY F Rd(10 R4/6) w/brownish fint; 80% 10w-mod. plastic fines, slow- dilatancy, mod. toughness, zolov v.f. sand; sattered sitst. from the va": gyp. mottling common; little to no rac. tottel; have; slightly moist.	5-4	~1.55 2.0	end of tube at; may be some slough on tup we = 5.9916. Wf = 15.77/6	
37-0-		V. f. sond; scattered sitst. trous to ya" gyp. Innottling common; little to no reac to Hel; hand; slightly moist. B.H. @ 32.5	B-10	0.8/	8/0.5 11/0.5 25/0.5 Cw/scathed stat. Frags) Terminated hole @ 32.5	38
34-0 -				-	Terminated hole @ 32.51 w/ hole through sand unit -sufficient data acquired Backfilled hole w/mud, cultings	
- - -				<u>-</u>		
-			-	-		
- -			-	-	SHEET 2 OF Z	

DRILLING AND SAMPLING LOG

PROJECT DIA SC DAMS STAD DATE BRILLED MORE/BOST HOLE NO. IN-S LOCATION DIVERSON DAM ST B 562 20 to 2:42 S/S BROUND SURFACE ELEV. 23 175 (App.) DRILLING CONTRACTOR PICKER DIFFIDINGS. LOGGED BY DMY DEPTH TO GROUND WATER ACCORDING WATER AC	DRILLING AND SAMPLING LOG									
DEPTH CLASS. FIELD DESCRIPTION SAMPLE MODE REMARKS OF CONTROLLY OF CHILD STATE AND SAMPLE REMARKS OF CHILD STATE AND SAMPLE REMARKS OF CONTROLLY OF CHILD STATE AND SAMPLE REMARKS OF CHILD STATE AND SAMPLE SAMPLE STATE AND SAMPLE	PROJECT_	0/18-50	S Nams Utah DATE DRILLE	D 10/K/S	P/ HOLE NO. <u>IV-5</u>					
DEPTH O GROUND WATER AT A MAMMER WEIGHT AND FALL (A) (B. 31) SURFACE CONDITIONS dirt access and; bush, DEPTH CLASS. FIELD DESCRIPTION SAMPLE MODE DIAMETER AT A MAMMER WEIGHT AND FALL (A) (B. 31) DEPTH CLASS. FIELD DESCRIPTION SAMPLE MODE DESCRIPTION SAMPLE MODE DESCRIPTION DEPTH CLASS. FIELD DESCRIPTION SAMPLE MODE DESCRIPTION SAMPLE MODE DESCRIPTION AND A SOLICIAN DESPOTTS SULTY FROM SAMPLY SULTY FROM SAMPLY SULTY FROM SAMPLY AD makinal cavis dashy A makinal cavis dashy SP and SULTY FROM SAMPLY SULTY FROM SAMPLY SULTY FROM SAMPLY AD makinal cavis dashy A makinal cavi	LOCATION	Diversio	n Dam #5 @ sta. 20+02;42'0	S GROUN	ND SURFACE ELEV. ~3(75 (Apo.)					
TYPE OF RIGINITIONS JUST ACCESS IN A STATE OF RIGHT AND FALL LED 16. 31 MINERAL COOLS. 31 MINERAL COOL	DRILLING CONTRACTOR Pitcher Drilling Co. LOGGED BY DMY DEPTH TO GROUND WATER 104 PROUMER L									
DEPTH CLASS. FIELD DESCRIPTION SAMPLE MODE REMARKS DEPTH CLASS. FIELD DESCRIPTION SAMPLE MODE REMARKS DOWNSHAND DEPOSITS O.O. SM. ABOUTH DEPOSITE DEPOSITED DEPOSITE DE	TYPE OF R	16Failinal	500 HOLF DIAMETER 47/8" HA	MMER WEIG	HT AND FALL 140 16 30"					
DEPTH CLASS. FIELD DESCRIPTION SAMPLE MODE REMARKS ONE SHARKS ON	SURFACE	CONDITIO	ons dirt access mad; brushy	w	EATHER Scattered slowls, cool, buif					
DEPTH CLASS. FIELD DESCRIPTION SAMPLE MODE REMARKS BINGS OF THE CLASS. FIELD DESCRIPTION SAMPLE MODE REMARKS BINGS OF THE CLASS. THE CLASS OF THE CL					to overcast 'H. shower					
0.0 SH ABSLIAN OSPOSITE ML 2.0-5.0 SIRTYSAND SANDY SIT: ROGOR 4(5), 45-55-6 ROTTED STATES AND SANDY SIT: ROGOR 4(5), 45-55-6 ROTTED STATES AND SANDY SIT: ROGOR 4(5), 45-55-6 ROTTED STATES AND SANDY AD marked course body AD marked course body AD marked course body ROGOLOVIAL OSPOSITS STATES AND ROSE STATES AND SANDY AD MARKED STATES AND SANDY STATES AND ROSE COMMON AND STATES A	050711	01.400	FIELD DECORIDATION	0.0457.5						
2.0 SH ABSLAND OSPASITS ML 2.0-5.0 SETTSAND SANDY TO STITE RESCORA 4(5) 45-55-50 FOUND HOSPASITS ROUTH FIRST, 45-55-50 FOUND HOSPASITS FOUND HOSPASITS AND MORRAL CAVE WAIT AND MORRAL CAVE WAIT AND MORRAL CAVE WAIT SOFT 2 1/0.5 4/0.5 5/0.5 AND MORRAL CAVE WAIT AND MORRAL CAVE WAIT AND MORRAL CAVE WAIT FOUND HOSPASITS SOFT 2 1/0.5 4/0.5 5/0.5 AND MORRAL CAVE WAIT AND MORRAL CAV	DEPIH	CLASS.	FIELD DESCRIPTION	SAMPLE	1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1					
ML 0.0-5.0 SELTYSHID SANDY SUT: ROGING 4(5); 45-5546 PROPORTY FIRST 1855506 AD MORAL CONSTRUCT FOR THE SELECT FOR THE SELECT FIRST 185506 SO THE ESTATE OF THE SELECT FOR THE SELECT FIRST 185506 SO	40	• 4	Ocaucit Acascian	- 1						
2.0 First Solice 416, 185536 First 185636 First 18563 Fi			7		1 +					
2.0 PROPRIOTICE FINELY 45 536 by pounts, grades of the property of the pounts, grades of the pounts, grades of the pounts of the	-	- 1000		[] /	11/15					
4.0 4.0 4.0 4.0 4.0 4.0 A MEL PRINCIPLE STATE SOLUTION AND SOLUT]		1516/ : ROCIOR 419); 45-55 .0.		Auger of 6" Flight auger;					
4.0 4.0 4.0 4.0 4.0 4.0 4.0 4.0	Z-0 -	-	poorly graded, fi-v.f. sand;	<u>-</u>	AD + makinal cause badly					
COLLUMAL DEASITS COLLUMAL DEASITS SO -20.0' XNTERPRESIDEN SAMO S. D20.0' XNTERPRESIDEN SAMO S. S]		0-5% scattged + gravel to	<u> </u>						
COLLUMAL DEASITS COLLUMAL DEASITS SO -20.0' XNTERPRESIDEN SAMO S. D20.0' XNTERPRESIDEN SAMO S. S	-	-	soft spots, nodules; strong reac-	<u> </u>	l ————————————————————————————————————					
COLLUNAL DEPOSITS COLLUNAL DEPOSITS SO -20.0 INTERPOSICE SAND AND SITTY FRIND PORTOGORADO SP. AND SITTY FRIND PORTOGORADO MASH SO -50.75% AND SO]]	-	to Hal throughout boss to med.	<i>B</i> -2	1 /4 mT					
6.0 SP THE WAY 10-15% in SP sord, to 25-50% in Rest common SM, non-plastic; sand wants 85-30d in SP to 70-75% in SM, pad sast glz., inan-axide spained sast glz., inan-axide spained sast glz., inan-axide spained cand, f. gravel as sandst and B-3 stiff frags gyp. variable from some inequality saves and from most claim from powers to several as a sandst and the polling and some inequality cands from the damaged; some inequality save graves from well damage to very dance with inc. dopth; v. slightly soft from the damaged from some inequality sources from the polling and some	4-0 -	-								
6.0 SP THE WAY 10-15% in SP sord, to 25-50% in Rest common SM, non-plastic; sand wants 85-30d in SP to 70-75% in SM, pad sast glz., inan-axide spained sast glz., inan-axide spained sast glz., inan-axide spained cand, f. gravel as sandst and B-3 stiff frags gyp. variable from some inequality saves and from most claim from powers to several as a sandst and the polling and some inequality cands from the damaged; some inequality save graves from well damage to very dance with inc. dopth; v. slightly soft from the damaged from some inequality sources from the polling and some		[ZOCCOVIAL VEFOSITS	<u> </u>	RD + RO 70 3.0. 70 CRAN NOW.					
6.0 SP THE WAY 10-15% in SP sord, to 25-50% in Rest common SM, non-plastic; sand wanes 85-90% In SP to 70-75% in SM, pad sast glz., inan-axide stained, cosand, f. gravel as sandit and B-3 Stist frags; gyp. varioth from Anne visible to speepel of as some inequality sand; some inequality sand; some inequality sand 12.0 We = 5.48% Ing. 10-516/5 in SM, pad SPT-3 10/1.5 11/0.5 12/0.5 23 SPT-3 10/1.5 11/0.5 12/0.5 23 SPT-3 10/1.5 11/0.5 12/0.5 23 SPT-3 10/1.5 11/0.5 12/0.5 25 SPT-3 10/1.5 11/0.5 12/0.5 27 SPT-3 10/1.5 11/0.5 12/0.5 28 SPT-3 10/1.5 11/0.5 12/0.5 28 SPT-3 10/1.5 11/0.5 12/0.5 28 SPT-3 10/1.5	-	-	5.0-20.0 INTERESCOED SAND	<u>-</u>	End of tibe egg-shaped					
10-0 SP 18-0 SP 18-0 SP 18-0 SP 18-0 SP SP SP SP SP SP SP SP SP S		5P,	and SILTY SAND: 12d(10124/6):	E	PB-1 = otherwise o.k.					
mon-plasfic; sand wars 85-70% mash to 70-75% in SM, pad sas of 872. inn-oxide stained of sas for 972. inn-oxide stained of c sand, figurel as sandst and B-3 start frogs; qyp. varioth from start from oxide for squal of some interpretation of some interpretation of some interpretation of start from the first from sity quality sand 18.0 specially startified in special or successfully complete are probably startified for successfully complete are successfully startified for successfully complete are succes	6-0	Elsm	10 25 20% in 1950 common 5he		+ We = 5-48/6					
Sa-sr gfz., ironoxide stained of most is f. quart as sansta and B-3 SM most is f. quart as sansta and B-3 Start rags; qyp. variety from some visible to several of as sansta and B-3 SM mothing, small nodules, mod. SM mothing small nodules, mod. Some ineq. zones frognests at some stample presumably sipped as from med. Gonz comented; quades from med. Gonz comented; quades from med. Gonz comented; quades from med. Gonz to be dead and B-4 12.0 with inc. depth; v. singthly SP most (drill flub presence from the polling rads for most (drill flub presence from the polling rads for most (drill flub presence from the polling rads for most (drill flub presence from the polling rads for most (drill flub presence from the polling rads for most form the polling rads for mo		[],	1 -0 30 10 11 1/31 COMMON 31/1,	[5-/	1.18/ = UF= 12.02/6					
SAM Sass of first first first for sond of the pulling nots of the			in sp to 70-75% in SM, PRA	;	1 +					
c. sand, f. gravel as sandit and B-3 stiff trops; gyp. variable from none visible to several elo as none visible to several elo som nothing, small nodoles, med. strong leac to hel through out; some rieg, zones, formented; grades from need denne to ery denne inc. depth; v. slightly spr. 12.0 wife no hecovery spr. 12.0 wife no hecovery spr. 13.0 listed in several dented to held the several elo spr. 13.0 listed in several elo spr. 14.0 listed in several elo spr. 16.1 listed in several elo spr. 16.1 listed in several elo spr. 16.2 listed several elo spr. 16.3 listed fine sand spr. 16.4 sithy fine sand spr. 18.5 listed fine to elo spr. 18.6 listed in several elo spr. spr. 18.6 listed in several elo spr. s			sa-sr qtz., iron-oxide stained,		10/10/10/05/12/0.5 23					
10.0 SP, None visible to several 9/0 as SP,	8-0 =	- ISM	most 15 f v.f.; locally 5% m-	- - -	I I					
10.0 SP, None visible to squemol 9/0 as mothing, small nodules, mod. Mothing, small nodules, mod. Strong back to the trinshiph out. Some ineq. tones, forments are triggered as the model sample of sample presumably strong host. Sp. None ineq. tones, forments are triggered as the model sample of sample presumably strong as the following in distribution and the pulling in distribution and the sample presumably strongered as the model sample of the sample probably sample are the pulling in distribution and the pulling in distribution a		:	c. sand, f. gravel as samult and	[\(\lambda^{-5}\)						
12.0 Lighty Cacos comented; grades from med. Hence to very dence from med. Hence to very dence wiff, inc. depth; v. slightly wiff, inc. depth; v. slightly easily). 12.0 Light moist (drill fluid penetroles to uny easily). 13.5-10.8 ckon SPs and 11.5-11.8 ckon SPs and 11.5-1] [HOME wisible to several % as:		end of the not damaged;					
12.0 Lighty Cacos comented; grades from med. Hence to very dence from med. Hence to very dence wiff, inc. depth; v. slightly wiff, inc. depth; v. slightly easily). 12.0 Light moist (drill fluid penetroles to uny easily). 13.5-10.8 ckon SPs and 11.5-11.8 ckon SPs and 11.5-1	,, , ;	· '	motting, small modules, mod		PB-2 Idvilled, smooth, quiet -					
Tom med stanse to very dense from med stanse to very dense sp. with no. depth; v. slightly moist (and fluid penetices fairly easily). SP. SP. SM. 11.5-11.8 ckon SP sand 11.8-12.3 quites to slightly sitty quality sand 15.5-16.1 clean SP sand SP. SM. 16.1-16.4 silty fine sand SP. SM. Note: bedding, tratification identified in SPT samples; not successfully samples and inknown of the samples and inknown of th	/0-0		Some inco Tones fromments are	<u> </u>	- Sample presumably suppod					
12.0 with inc. dopth; v. slightly worst (smill fluid printings fairly easily). SP, SP, 11.5-11.8' ckon SP sand 11.8-12.3' qwides to slightly sitty quality sand 11.5-16.1' clean SP sand SP, SM, Note: bedding stratification identified in SPT samples; not succertibilly complete are probably similarly stratified 18.0 FB samples and inkerols not succertibilly complete are probably similarly stratified FB-4 end of two and sand SP sand sitty sand SP sand sitty fine sand RO same graphs should SPT S T/o.5 16/o.5 21/o.5 S			I (rollita) Ca ca > demented: 9/00003 .		20 1 / NM Note politing hads					
14.0 SP 16.0 SP 16.1-8.4 silfy fine sand SP SP SP SP SP SP SP SP SP S			I think med dence to her danie -	-	15.3 = mt = NO HECONON					
14.0 SP 16.0 SP 16.1-8.4'silfy fine sand SP SP SP SP SP SP SP SP SP S	1,20		moist can'll flind penetrales tainly		SPTA = 9/05 16/05 21/05 37					
14.0 16.0 18.5-18.1 gustes to slightly sitty quartly sand 15.5-16.1 clean SP sand 16.1-16.4 sitty fine sand SP, SM, Note: bedding, tratification intervals and inter	/2-0	- Aprilos	4 C arc'(a)	R-A	0.81 = (w/common stst-/ss.					
14.0 16.0 18.5-18.1 gustes to slightly sitty quartly sand 15.5-16.1 clean SP sand 16.1-16.4 sitty fine sand SP, SM, Note: bedding, tratification intervals and inter]		7.5-9.0 2Wdely bedded SM:		1.5 + grows p.3-1.0 into					
16.0 16.0 16.0 16.1-8.4' silty fine sand 18.0 18.		SM	11.5-11.8 GRONSPSOND		PB-3 Feed Italy at a scan					
16.0 SP 16.1-16.4' silfy fine sand SP, SM Note: bedding, tratified profiles and interpole are proceed are proceed are proceed are proceed are proceed as succentially completed are proceeding similarly stratified. 18.0 SP, SM Note: bedding, tratified are proceed are proceed are proceed by similarly stratified. 18.0 SP, SM Note: bedding, tratified are proceed are proceed by similarly stratified. SP-6 13/0.5 25/0.5 30/0.5 55 55 55 55 55 55 55 55 55 55 55 55 5	14.0	<u> </u>	Lydeilant 29 males to slightly		1 10/10 0/ /010 01/1/02/07/10					
16.0 SP 15.5-16.1' clean SP sand 15.5-16.1' clean SP sand 16.1-16.4' silty fine san	; ;	;	sitty quautity sand	:	tube pulling rods					
18.0 SP 16.1-K.4'silty fine sand B-5 16.1-K.4'silty fine sand B-6 16.1-K.4'silty fine sand B-7 16.1-K.4'silty fine s		[<u> </u>	<u> </u>	$\omega_{e} = 5.461b$					
18.0 SP. SM 16.1-K.4' silty fine sand B-5 1/0.5 16/0.5 21/0.5 37 SP. SM Note: bedding, tratification RD some gauges slowed some gauges slowed some gauges slowed hole some gauges slowed as succertfully complete are probably similarly stratified R-6 13/0.5 25/0.5 30/0.5 55 PB-4 end of tube o.k.			155-161'clean SPSand		1 2 10 Requiry					
18.0 SP, SM Note: bedding, stratification RD some gravels slowled of the samples and intervals not successfully completed are probably similarly stratified RD some gravels slowled and intervals not successfully completed are probably similarly stratified RD some gravels slowled and intervals not some gravels slowled and indicated are probably similarly stratified RD some gravel) SPT-6 13/0.5 25/0.5 30/0.5 55 Successfully completed are probably similarly stratified RD some gravels slowled SPT-6 13/0.5 25/0.5 30/0.5 55 Successfully completed are probably similarly stratified RD some gravels slowled SPT-6 13/0.5 25/0.5 30/0.5 55 SPT-6 13/0.5 25/0.5 50/0.5 50/0.5 50/0.5 50/0.5 50/0.5 50/0.5 50/0.5 50/0.5 50/0.5 50/0.5 50/0.5 50/0.5 50/0.5 50/0.5 50/0.5 50/	10.0		-	_	SPT-5= 7/05 16/05 21/05 37					
18.0 SP, SM Note: bedding, stratification RD some graces slowled of the samples and intervals not successfully completed are probably similarly stratified RD some graces slowled stratified SPT-6 13/0.5 25/0.5 30/0.5 55 Successfully completed are probably similarly stratified PE-4 end of tube o.k.		- ISM	16.1-16.4 silfy fine sand	13-5	0.9 = (no a sand or gracel)					
succertailly compled are probably similarly stratified PE-4 end of tibe o.k.	=			·	1 // /= 1					
succertailly compled are probably similarly stratified PE-4 end of tibe o.k.		SM	Note: bedding, tratitication		RO = some gravels stonked					
probably similarly stratified PE-4 and of tibe o.k.	18.0-	<u>-</u>	PR ample and inknow not	<u> </u>	SP-6-13/6-25/65 30/65 55					
PE-4 = end of tube o.k.				8-6	as they remained modes					
	-	-	probably similarly stratified		115] of sand to sith sand)					
+ 2-0 SHEET OF - OF -	20 1	;]		1.44] CHEET / - 7					
	[20.0	<u> </u>]	٠ ٢ ٢	7.0] SHEET OF					

PROJECT DII8-SCS Dams Otah DATE DRILLED 10/16/81 HOLE NO. IV-5 DEPTH CLASS. FIELD DESCRIPTION SAMPLE MODE REMARKS We= 5.5016 20.0 PB-4: 200-28.5 <u>SAND</u>: orangey Hd (10R4/6) to (10R4/8); <5% SP (con't) I wf= 13.53/6. (co14) non-plastic fines; >95% v.poor4 quaded, non-oxide stained gtz; trace of scattered sub-nounted SPT-7= 12/0.5 30/6.5 38/0.5 = (as for SPT-6) Z2-0 rrace of scaffeed sub-nounted sandst. clasts from 44-1": mod.
hac. to Hel thomphout; acasimon
trace of gyp; dense to very dense
(seems lightly centerted in some
comes w/cacos); slightly moist
(penation to be easily by drill fluid). RD to 23'to envis a ban 507-8 14/0.5 25/0.5 36/0.5 Llas for SFT-7) 24-0 RO to 25' to ensure clean RD 14/0.5 21/0.5 26/0.5 comented nodule a top of sample)
RO to Z7 to ensurcle an hole for SPT B-9 Z6.0 RD 597-10= 156.5 33/0.5 43/0.5 B-10 (as for SAT-8) 28.0 RH. @ 28.5' Terminate 116 @ 28.5' Finto high 610w-count sand: 7=30 p.m. -too dark to continue dilling safly 30.0 Backfilled hole w/ cottings, surficial

SHEET Z OF Z

APPENDIX C

TEST PIT EXCAVATIONS

Appendix C TEST PIT EXCAVATIONS

A total of 32 test pits were excavated on the dam embankments and in their foundations to investigate the following:

- 1. To provide information on the type, zoning and classification of embankment and foundation materials.
- 2. To perform in situ density tests.

The test pit excavations were performed at the same time as the drilling operations. Excavation of test pits was subcontracted to Ziegler's Backhoe Service of Cedar City, Utah. Logging of the test pit excavations and in situ density tests were performed by Richard Morris of Earth Sciences Associates.

Three to five test pits were excavated at each dam, using a Case 580C backhoe with a 24-inch bucket. The locations of test pits are shown in Figures B-1 through B-8 and are summarized in Table C-1. Logs of the test pits are given in Figures C-1 through C-18 along with the locations of the in situ sand cone density tests. In many cases, attempts were also made to obtain undisturbed Shelby tube samples by pushing with the backhoe bucket. All such samples were sealed and shipped to the ESA laboratories in Palo Alto, California.

Pit locations were chosen so as to provide information on embankment conditions at both the crest and the toe of each dam, usually on the upstream side. This was the case at all three Green's Lake dams and at Stucki, Gypsum Wash, and Frog Hollow dams. At Warner Draw and Ivins dams, the crest pit was dug on the downstream face of the embankment; at Green's Lake Dam No. 2, Green's Lake Dam No. 3, and Stucki Dam, crest pits were dug on both the upstream and downstream faces. Toe pits were also dug on the downstream face at Gypsum Wash and Frog Hollow dams. All pits were located to be at or near points of maximum section or of embankment curvature. Those pits excavated at the embankment toe were stepped so that one level of the pit was in the embankment and a second level was in the foundation.

Large bag samples of material excavated from the pits were collected from all crest pits, and two bag samples—one from each level—were collected from the toe pits. These samples were used for laboratory compaction testing and material classification purposes.

Sand cone density tests were performed in about three-fourths of the test pits. Although not all pits were tested, the density tests were disturbed among the pits to provide a sampling of both embankment and foundation densities at each dam. Testing was done in general accordance with ASTM method D1556-64 and in specific accordance with U.S. Bureau of Reclamation method E-24. Deviations from the standard methods included the use of preweighed sand volumes to avoid complications associated with the field weighing. Most tests were performed using an ASTM standard six-inch cone apparatus; a twelve-inch cone was used in two pits where very large rock particles were encountered in the fill. To improve the accuracy of the density determinations replicate tests were made in all cases. Typically, three density tests were performed when using the six-inch cone, and two tests were performed when using the twelve-inch cone. Moisture contents were obtained either by ESA personnel in the field or by the Civil/Earth Consulting Group of Fort Collins, Colorado. All samples obtained during the density testing were retained and shipped to the ESA laboratories in Palo Alto, California.

Table C-2 presents information on the sampling and testing for the test pits at each dam site.

Table C-1
Locations of Test Pit Excavations

<u>Dam</u>	Pit No.	Crest Station	Offset from Dam Centerline	Approx. Bottom Elevation	Material in Trench
Green's Lake	GL2-TP1	12+75	15' upstream	6064'	Shell
No. 2	GL2-TP2	11+30	40' "	60531	Shell
	11	11	48' "	6050'	Foundation
	GL2-TP3	9+05	14' downstream	60631	Shell
	GL2-TP4	4+00	15' upstream	60631	Shell
	GL2-TP5	2+80	21' "	6059'	Shell
			29' "	6056'	Foundation
Green's Lake	GL3-TP1	17+65	15' upstream	6058'	Shell
No. 3	GL3-TP2	17+15	31' upstream	6053'	Shell/Found.
	II	11	38' "	60491	Foundation
	GL3-TP3	14+40	33' "	6052'	Shell/Found.
	ţŗ	11	40' "	6048'	Foundation
	GL3-TP4	5+70	18' "	6057'	Shell
	11	11	25' "	60531	Foundation
	GL5-TP5	6+40	13' downstream	6060'	Shell
Green's Lake	GL5-TP1	2+07	15' upstream	5933'	Shell
No. 5	GL5-TP2	1+98	62' "	5923'	Shell/Found.
	Ħ	11	67' "	5920'	Foundation
	GL5-TP3	1+19	61' "	59221	Shell/Found.
	11	11	681 11	5919'	Foundation
Warner Draw	WD-TP1	13+67	15' downstream	2979'	Shell
	WD-TP2	13+25	128' upstream	2945'	Shell
	11	11	136' "	2942'	Foundation
	WD-TP3	18+92	143' "	2940'	Shell
	11	11	150' "	2937'	Foundation
Stucki	STK-TP1	16+43	15' upstream	2807†	Shell
	STK-TP2	18+90	14' downstream	2805'	Shell
	STK-TP3	16+03	88¹ upstream	27831	Shell
	ff	11	94' "	2780'	Foundation
	STK-TP4	18+96	39' "	2799'	Shell
	tt	Ħ	44' ''	2796'	Foundation
Gypsum Wash	GW-TP1	21+56	36' downstream	2720'	Shell
	11	ŧī	42' "	2715'	Foundation
	GW-TP2	15+48	37' upstream	2725'	Shell
	ŧŧ	Ħ	431 "	27221	Foundation
	GW-TP3	25+86	14' "	2733'	Shell
	GW-TP4	35+99	43' "	2724'	Shell
	11	11	50' "	2720'	Foundation

Table C-1 (Continued)

Locations of Test Pit Excavations

<u>Dam</u>	Pit No.	Crest Station	Offset from Dam Centerline	Approx. Bottom Elevation	Material in Trench
Frog Hollow	FH-TP1 FH-TP2 " FH-TP3 " FH-TP4	12+97 11+48 " 12+49 " 16+10	17' upstream 28' downstream 35' " 88' upstream 95' " 87' "	111' 106' 103' 93' 89' 93' 90'	Shell Shell Found./Old fill Shell Foundation Shell Foundation
Ivins Diversion No. 5	IV-TP1 IV-TP2 IV-TP3 IV-TP4	42+35 " 21+53 " 8+46 " 15+31	22' upstream 29' " 19' " 25' " 41' " 47' " 15' downstream	3179' 3176' 3180' 3176' 3172' 3169' 3181'	Shell Foundation Shell Foundation Shell Foundation Shell

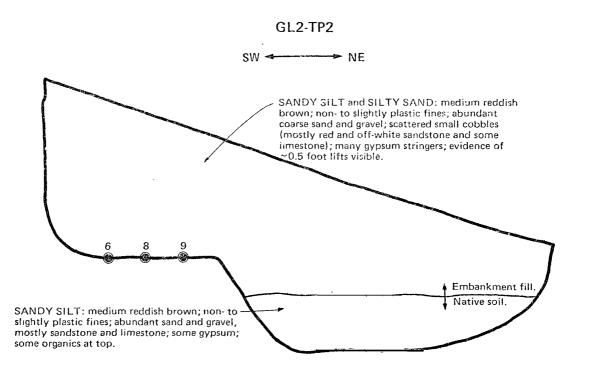
Notes: Stations and offsets refer to approximate center of pit. Elevation refers to pit floor. Two values for a pit indicate data for both benches in pits. In general, tests were run one or more days after excavation of test pits.

Table C-2

Sampling and Testing of Test Pit Excavation

Dam	Number of Test Pits	Number of Pits Tested	Number of Bulk Samples	Number of Sand Cone Tests	Number of Shelby Tube Samples
Green's Lake No. 2	5	3	7	8	5
Green's Lake No. 3	5	5	8	15	4
Green's Lake No. 5	3	3	5	9	2
Warner Draw	3	2	5	6	2
Stucki	4	3	6	9	2
Gypsum Wash	4	3	7	7	3
Frog Hollow	4	3	7	7	2
Ivins Diver- sion No. 5	<u>4</u>	<u>3</u>	7	9	<u>3</u>
TOTALS	32	25	52	70	25

GL2-TP1 SW NE SILTY SAND: medium reddish brown; non- to slightly plastic fines; abundant coarse-grained sand, gravel, and sulfate; fragments of sandstone and limestone. SANDY SILT: medium reddish brown; nonplastic fines; abundant coarse sand and gravel. SILT: medium reddish-grayish brown; slightly plastic fines; scattered coarse-grained sand; some



All materials in test pit excavation GL2-TP1 are embankment fill.

2. (a) indicates location of sand cone density test.

Notes

0 1 2 feet

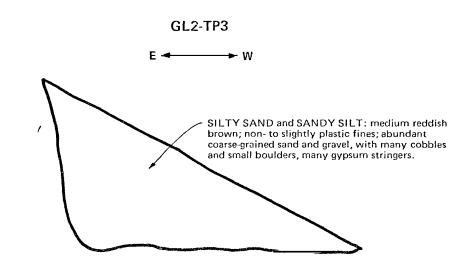
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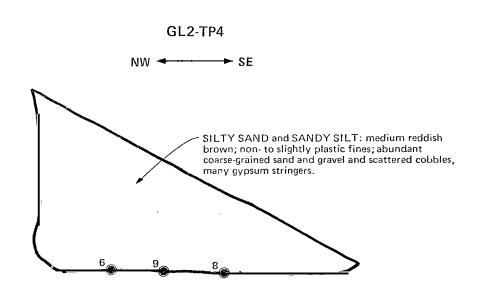
Earth Sciences Associates

Palo Alto, California

SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS LOGS OF TEST PIT EXCAVATIONS GREEN'S LAKE NO. 2

Checked by MLT Date 3/27/82 Project No. Figure No. Approved by EA Way Date 27 May 82 D118 C-1





All materials in test pit excavations are embankment fill. indicates location of sand cone density test.

Notes

0 . 1 2 feet

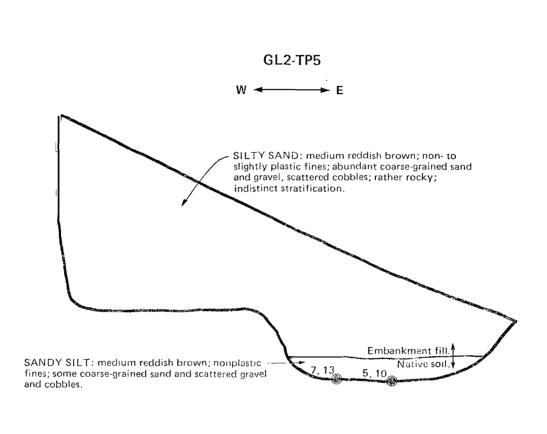
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Palo Alto, California

SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS LOGS OF TEST PIT EXCAVATIONS GREEN'S LAKE NO. 2

Checked by MLT Date 5/27/82 Project No. Figure No. Approved by Author Date 7/May & D118 C-2



Note indicates location of sand cone density test. 1 2 feet

0 1 2 fee

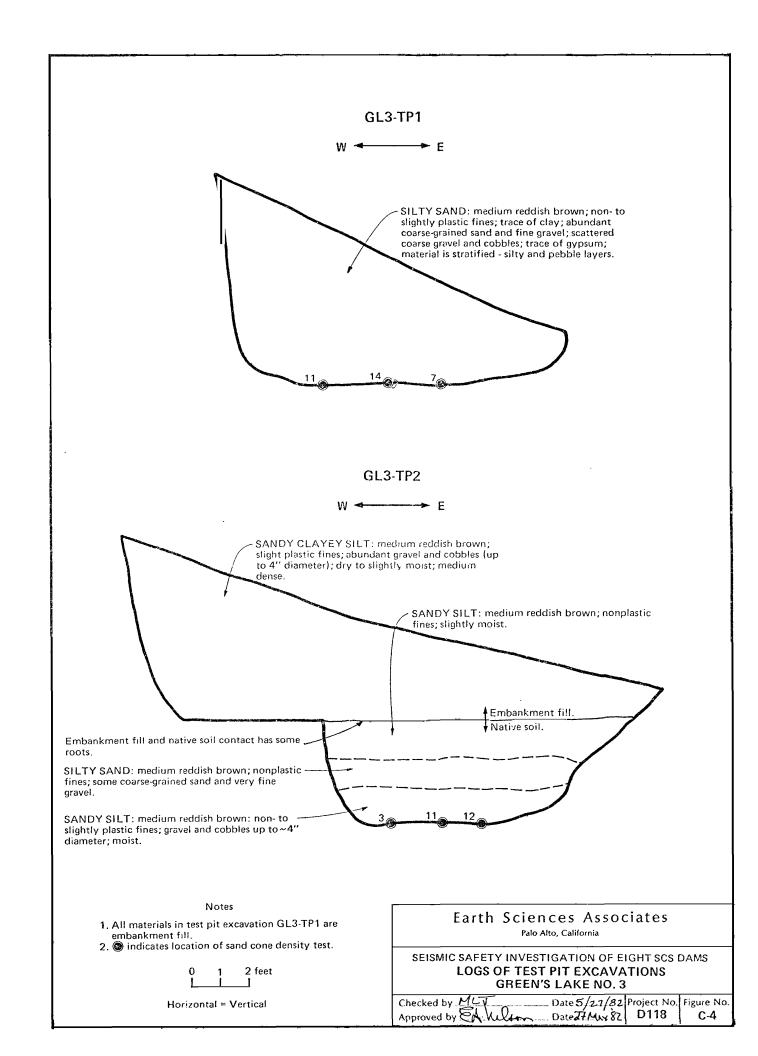
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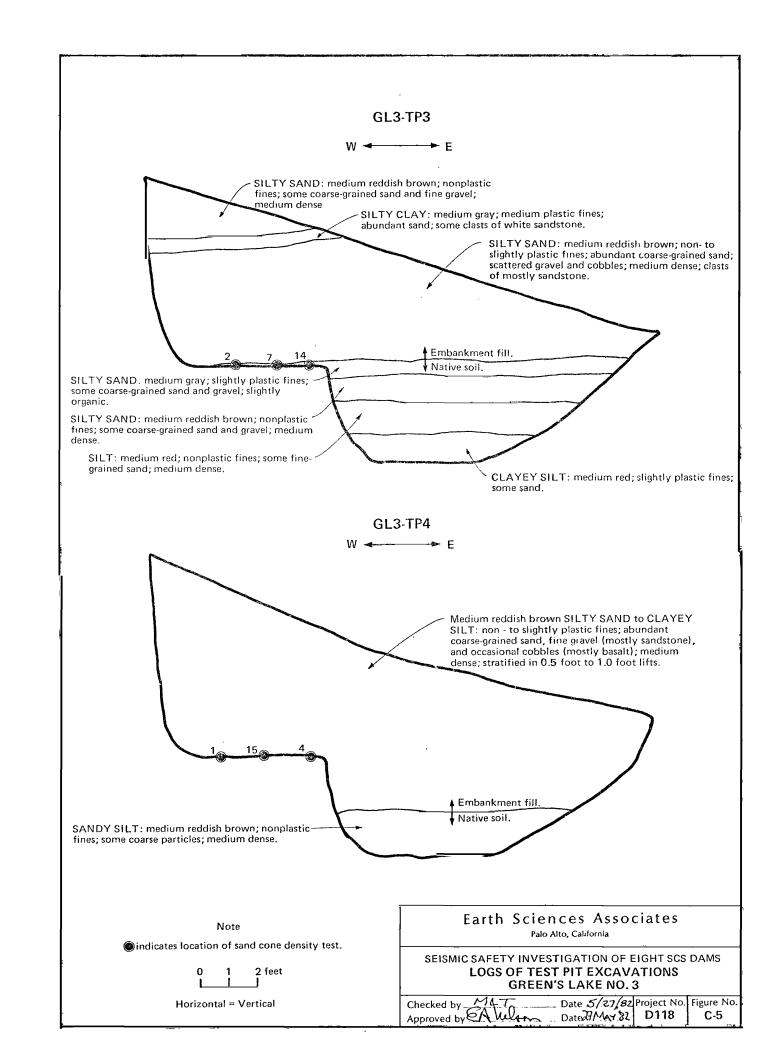
Earth Sciences Associates

Palo Alto, California

SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS LOG OF TEST PIT EXCAVATION GREEN'S LAKE NO. 2

Checked by MUT Date 5/27/82 Project No. Figure No. Approved by PAWA Date MAY 82 D118 C-3





GL3-TP5 SILTY SAND: light reddish brown; non- to slightly plastic fines; abundant coarse-grained sand and fine gravel; scattered cobbles; abundant gypsum. SILTY SAND: medium reddish brown; slightly plastic fines; abundant coarse-grained sand and fine gravel; abundant gypsum. 13 CLAYEY SILT: medium grayish brown; slight plastic fines; some coarse-grained sand and fine gravel; abundant gypsum, calcareous. SANDY SILT: medium reddish brown; nonplastic fines; some sand, gravel, and cobbles.

Notes

- 1. All materials in test pit excavation are embankment fill.

 2. indicates location of sand cone density test.

Horizontal = Vertical

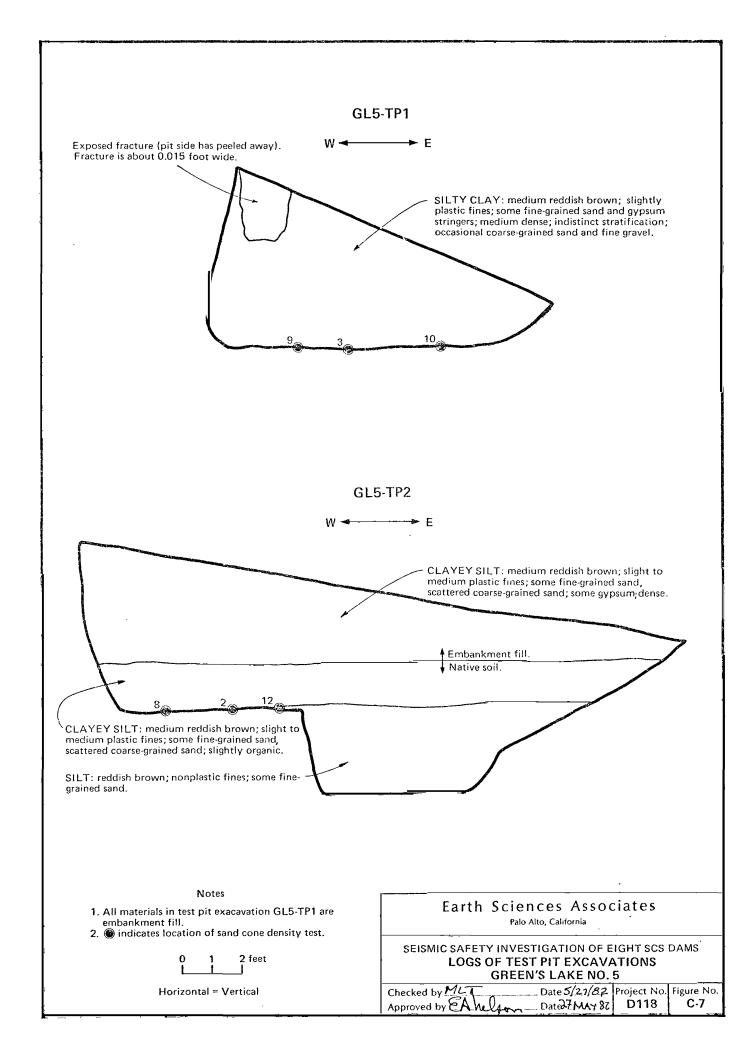
Earth Sciences Associates Palo Alto, California

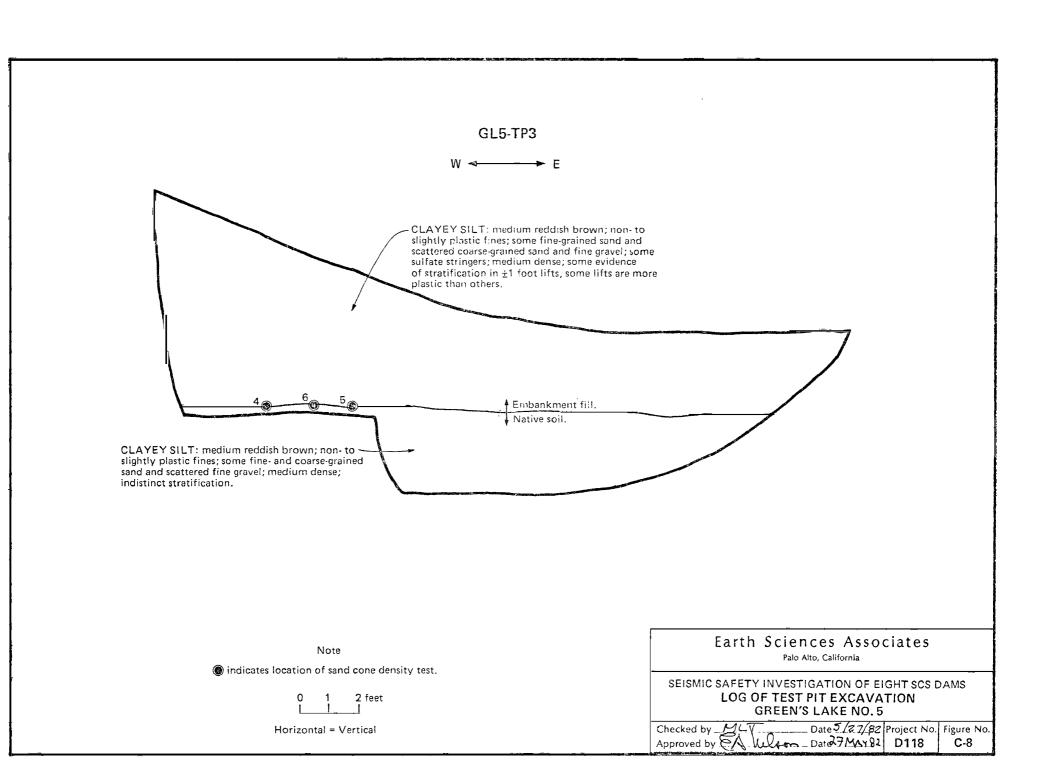
SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS LOG OF TEST PIT EXCAVATION GREEN'S LAKE NO. 3 .

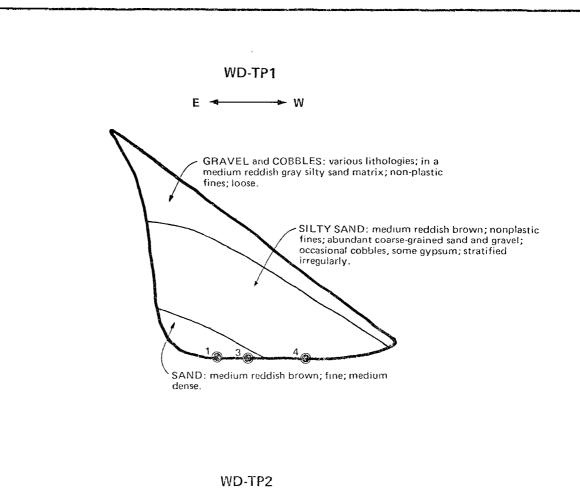
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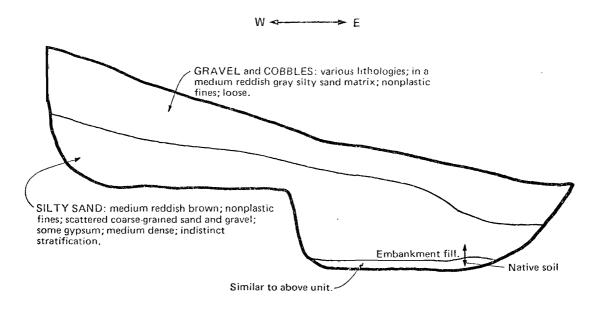
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C-6

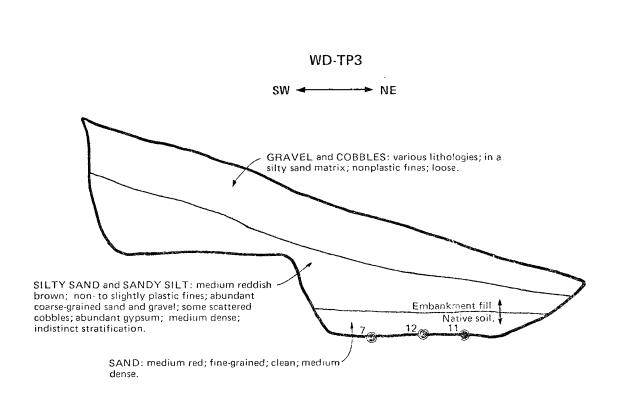








Notes Earth Sciences Associates 1. All materials in test pit excavation WD-TP1 are Palo Alto, California embankment fill. 2. Oindicates location of sand cone density test. SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS LOGS OF TEST PIT EXCAVATIONS 2 feet **WARNER DRAW DAM** Approved by ____ Date 5/27/82 Project No. Figure No. Horizontal = Vertical Date 27My 82 D118 C-9



Notes

- All materials in test pit excavation are embankment fill.
- 2. indicates location of sand cone density test.

0 1 2 feet

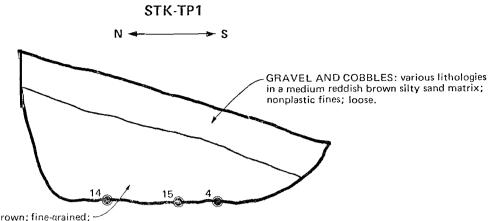
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Earth Sciences Associates

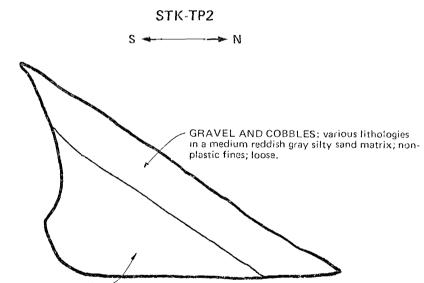
Palo Alto, California

SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
LOG OF TEST PIT EXCAVATION
WARNER DRAW DAM

Checked by Mary Date 5/27/8,2 Project No. Figure No. Approved by Aulan Date 1 May 82 D118 C-10



SILTY SAND: medium reddish brown; fine-grained; scattered coarse sand and gravel, non- to slightly plastic fines; medium dense; some gypsum stringers.



SILTY SAND: medium reddish brown; nonplastic fines; with abundant coarse sand and gravel; scattered cobbles; medium dense; indistinct stratification; some gypsum stringers.

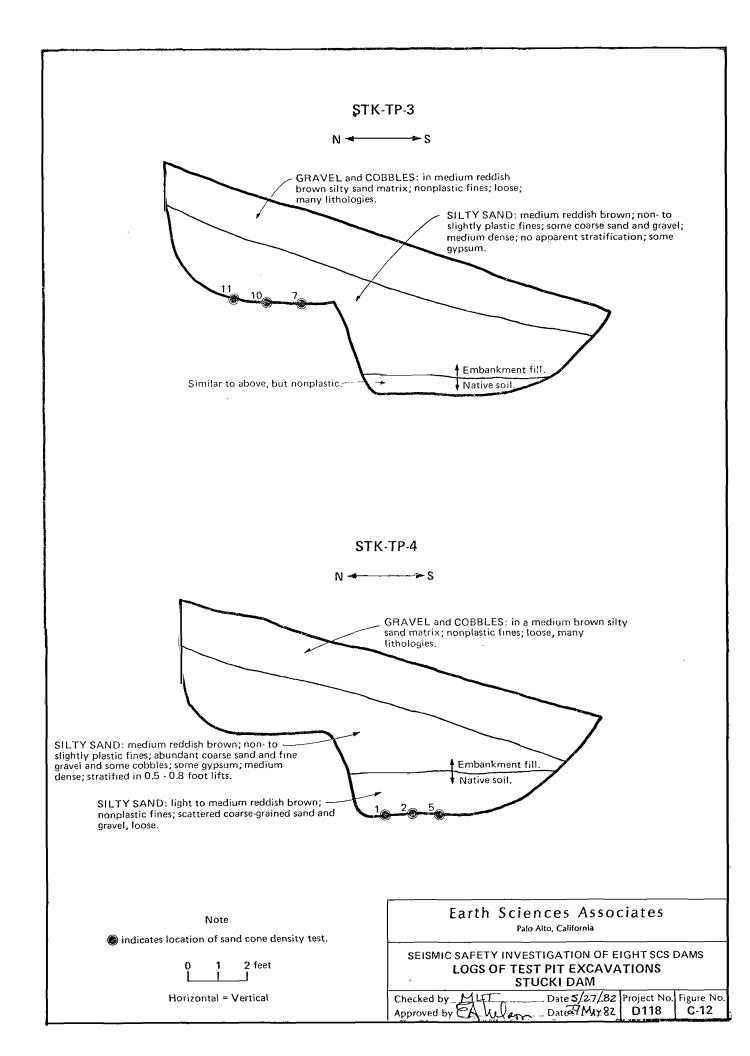
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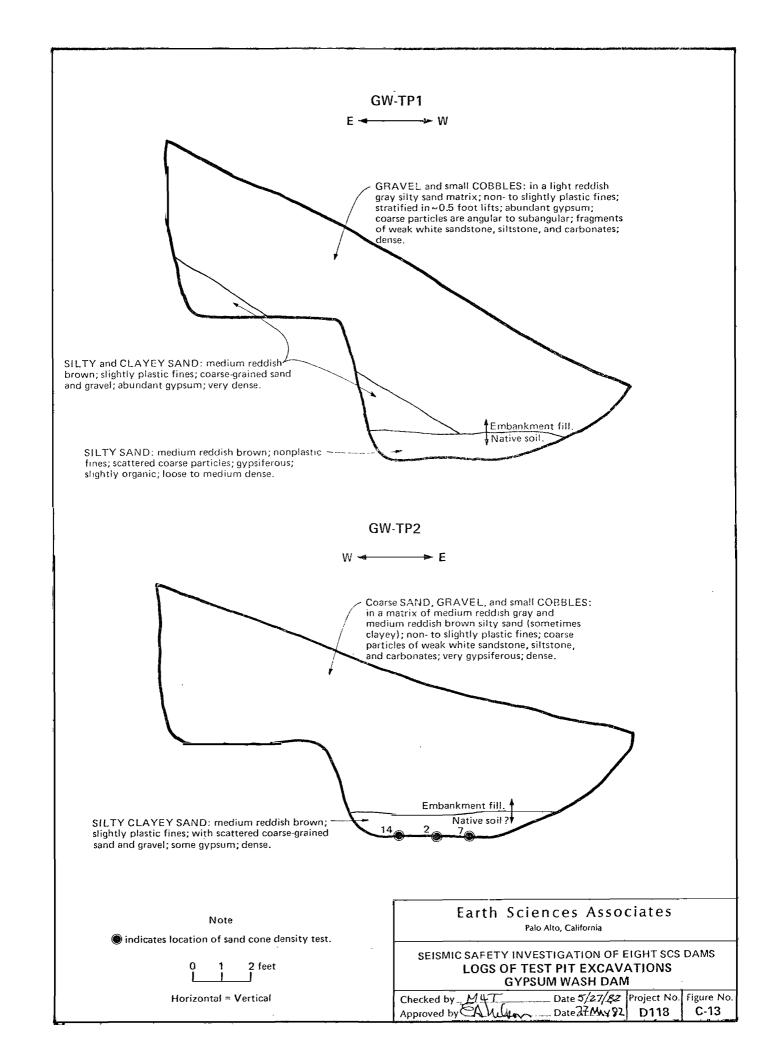
- All materials in test pit excavations are embankment fill.
- 2. indicates location of sand cone density test.

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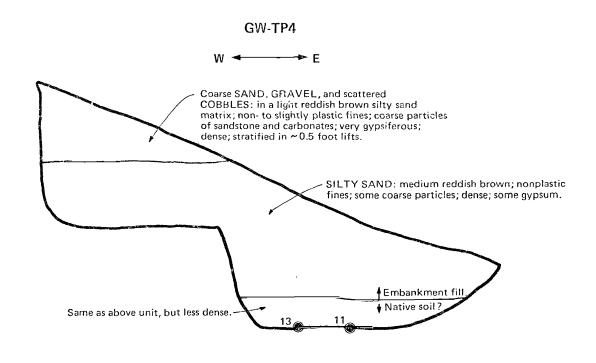
SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS LOGS OF TEST PIT EXCAVATIONS STUCKI DAM

Checked by MAT Date 5/27/82 Project No. Figure No. Approved by Date Date 7/482 D118 C-11





GW-TP3 W SILTY and CLAYEY SAND: Medium reddish brown slightly plastic fines; some coarse-grained sand and gravel; abundant gypsum; very stiff. SILTY CLAY: medium grayish brown; fragmented, weathered claystone. Same as above unit, but with highly gypsiferous materials (gray with white streaks).



All materials in test pit excavation GW-TP3 are embankment fill. indicates location of sand cone density test.

Notes

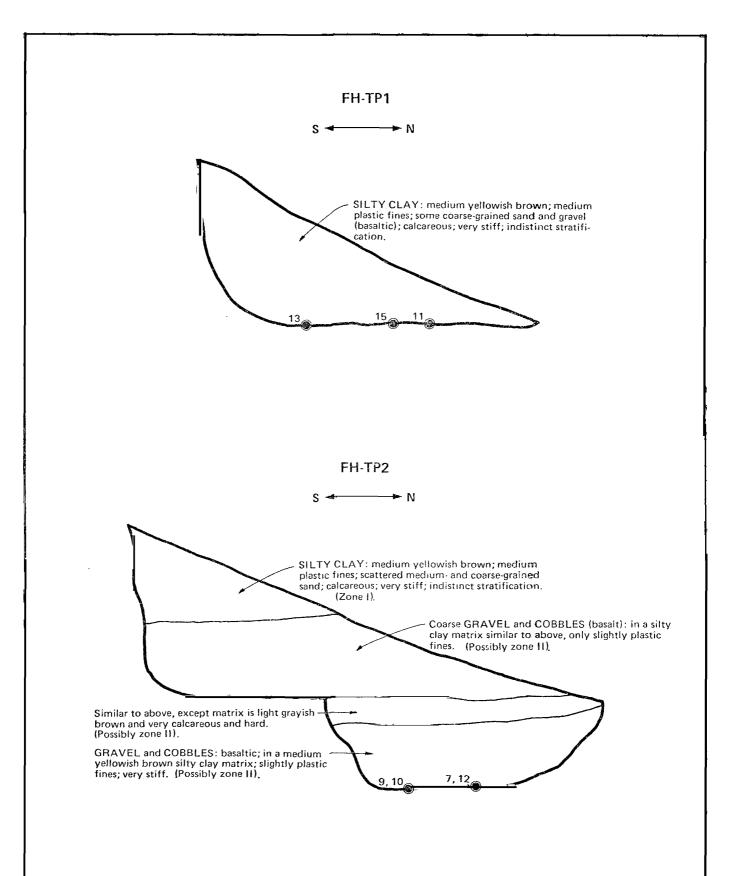
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Earth Sciences Associates Palo Alto, California

SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
LOGS OF TEST PIT EXCAVATIONS
GYPSUM WASH DAM

Checked by MLV Date 5/21/82 Project No. Figure No. Approved by A Lusan Date AMAY 82 D118 C-14



Notes

- 1. All materials in test pit excavations are embankment fill.
- 2. indicates location of sand cone density test.

0 1 2 feet

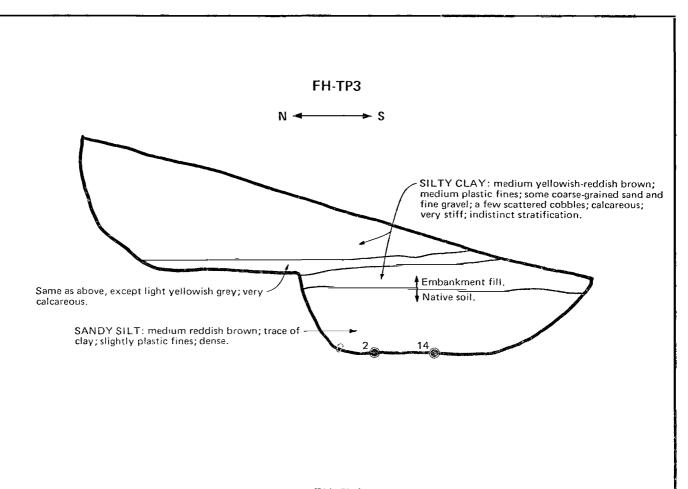
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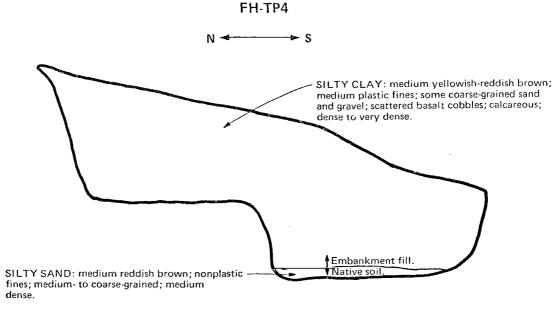
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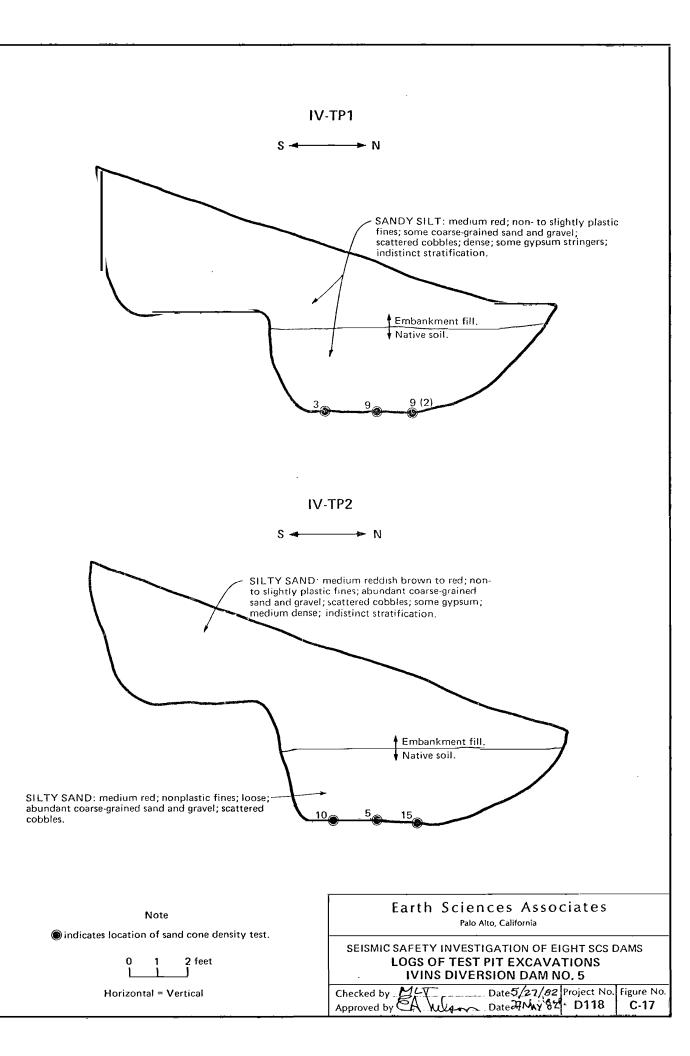
SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS LOGS OF TEST PIT EXCAVATIONS FROG HOLLOW DAM

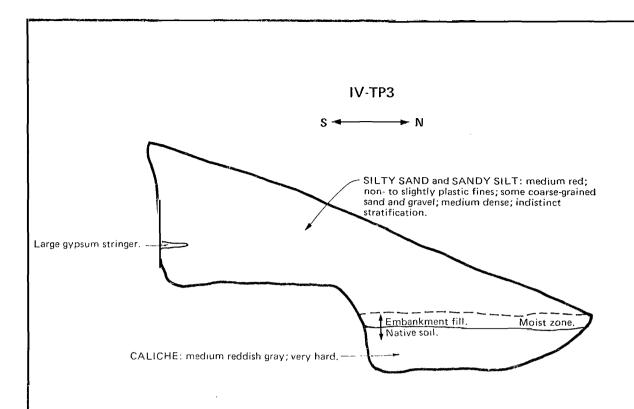
Checked by MLT Date 5/27/82 Project No. Figure No. Approved by EA War Date 27 MAY 82 D118 C-15

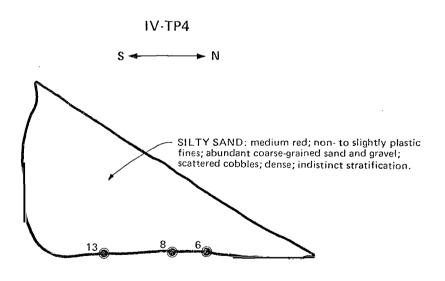




Note Earth Sciences Associates Palo Alto, California SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS LOGS OF TEST PIT EXCAVATIONS FROG HOLLOW DAM Checked by Man Date 3/27/62 Project No. Figure No. Approved by A Who Date 2 May 81 D118 C-16







Notes

- 1. All materials in test pit excavation IV-TP4 are embankment fill.
- 2. indicates location of sand cone density test.

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SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
LOGS OF TEST PIT EXCAVATIONS
IVINS DIVERSION DAM NO. 5

Checked by MLT Date 5/27/82 Project No. Figure No. Approved by Date 27May 22 D118 C-18

APPENDIX D

LABORATORY TESTING

Appendix D LABORATORY TESTING

Laboratory tests consisting of grain-size determination and compaction tests were performed to establish the relative compaction of the embankment and foundation materials. In addition, Atterberg limit tests were performed to help facilitate the classification of the various materials encountered. Results of Atterberg limit tests are shown in Figures D-1 and D-2. Sieve analyses and compaction tests were run on all of the bag samples obtained from the test pits. Sieve analyses were also performed on some of the sand cone test samples to establish their similarity with the corresponding bulk sample. Gradation curves for all of the samples tested are shown in Figures D-3 through D-10 for each of the dam embankments. Compaction curves for the materials tested are shown in Figures D-11 through D-18 for the various materials tested along with the corresponding maximum dry densities and optimum moisture contents. compaction tests were performed using a 4-inch diameter mold in accordance with ASTM procedure D698-70 method C (standard Proctor and 3/4 inch maximum particle size).

The degree of relative compaction at each sand cone test location was computed using the appropriate value of maximum dry density established in the laboratory for the corresponding bulk sample. These results are tabulated in Tables D-1 through D-8 for each test performed at each dam location.

Table D-1

Relative Compaction Evaluation

Green's Lake Dam #2

			Unified Soil Class	sification	Field T	Cest	Labo	ratory	Relative	
Dam/Test Pit	Can #	Zone Tested	Field	Lab	Yd (pcf)	w/c (%)	Ydmax (pcf)	Opt w/c (%)	Compaction (%)	Remarks
GL2-TP2	#6	Shell	SM-ML w/gravel,	SM-SC	111.3	10.1			90.4	
GL2-TP2	#8	11	11	11	106.5	7.9	123.1	10.7	86.5	
GL2-TP2	#9	11	11	11	110.4	8.5			89.7	
GL2-TP4	#6	Shell	SM-ML w/gravel	SM-SC	101.7	5.7		•	85.8	
GL2-TP4	#9	11	tt	ŧr	101.0	6.0	118.5	13.2	85.2	
GL2-TP4	#8	11	. 11	tt	101.6	8.4			85.7	
GL2-TP5	#7,13	Founda- tion	SM w/gravel, cobbles	SM-SC	116.4	6.9	121.8	11.0	95.6	12" sandcone
GL2-TP5	#5,10	11	11	11	106.7	7.0	721.0	21.0	87.6	ττ

Table D-2

Relative Compaction Evaluation

Green's Lake Dam #3

		7000	Unified Soil Classif	ication	Field T		Labor Ydmax	ratory	Relative Compaction	
Dam/Test Pit	Can #	Zone Tested	Field	Lab	(pcf)	w/c (%)	(pcf)	Opt w/c (%)	(%)	Remarks
GL3-TP1	#11	Shell	SM w/gravel, cobbles	ML	102.1	9.6			82.5	Sandcone sample
GL3-TP1	#14	11	11	ft.	117.6	7.6	123.7	9.9	95.1	has finer sand
GL3-TP1	#7	ττ	tt	11	116.2	6.5			93.9	than bulk sample
GL3-TP2	#3	Founda- tion	ML w/sand, gravel,	SM	106.3	11.6			86.6	
GL3-TP2	#11	11	TŤ	11	107.3	10.8	122.8	9.7	87.4	
GL3-TP2	#12	11	11	11	108.7	11.5			88.5	
GL3-TP3	#2	Founda- tion	SM w/gravel	SM-ML	111.3	9.5			88.5	
GL3-TP3	#7	17	11	11	117.3	8.4	125.7	10.6	93.3	
GL3-TP3	#14	11	ft	11	115.9	10.6			92.2	
GL3-TP4	#1	Shell	SM-ML w/gravel,	SM-ML	112.1	7.6	,		89.0	
GL3-TP4	#15		cobbles	11	105.2	7.9	126.0	11.6	83.5	
GL3-TP4	#4			11	109.9	7.3			87.2	
GL3-TP5	#5	Shell	SM w/gravel, cobbles	SM-ML	98.8	8.6			80.2	
GL3-TP5	#13	11	11	11	103.9	9.6	123.2	10.9	84.3	
GL3-TP5	#10	11	17	11	97.2	11.7			79.3	

Table D-3

Relative Compaction Evaluation

Green's Lake Dam #5

			Unified Soil Class	ification	Field T	`est	Labo	ratory	Relative	
Dam/Test Pit	Can #	Zone Tested	Field	Lab	γ <mark>d</mark> (pef)	w/c (%)	Ydmax (pcf)	Opt w/c (%)	Compaction (%)	Remarks
GL5-TP1 GL5-TP1 GL5-TP1	#9 #3 #10	Shell "	CL w/sand gravel	CL "	79.3 86.8 92.5	10.8 10.6 11.6	113.0	13.5	70.2 76.8 81.9	Field test water contents are not in situ values due to time between opening opits and testing (rain).
GL5-TP2	#3	Founda- tion	ML w/sand	SM-SC	98.7	12.3			82.9	
GL5-TP2	#2	11	11.	TŤ	99.4	11.3	119.1	11.0	83.5	
GL5-TP2	#12	11	11	††	108.6	10.3			91.2	
GL5-TP3	#4	Founda- tion	ML w/sand, gravel	CL-ML	98.5	11.4			88.3	
GL5-TP3	#6	11	11	11	97.0	11.6	111.6	15.5	86.9	
GL5-TP3	#5	11	ΤΤ	11	96.3	10.2			86.3	

Table D-4

Relative Compaction Evaluation

Warner Draw

			Unified Soil Class	ification	Field T	est		ratory	Relative	
Dam/Test Pit	Can #	Zone Tested	Field	Lab	γ _d (pcf)	w/c (%)	Ydmax (pcf)	Opt w/c (%)	Compaction (%)	Remarks
WD-TP1	#1	Shell	SP-SM w/gravel, cobbles	SM	120.3	7.1			103.0	
WD-TP1	#3	11	ŤĒ	11	120.2	5.6	117.8	7.0	104.0	
WD-TP1	#4	11	īī	Ħ	108.8	3.6			92.4	
WD-TP3	#7	Founda- tion	SP	SM	127.5	9.6			106.0	
WD-TP3	#12	11	tt -	11	124.1	6.5	119.8	9.0	104.0	
WD-TP3	#11	tt	tt	11	114.7	5.3			95.7	

Table D-5

Relative Compaction Evaluation

Stucki

			Unified Soil Cla	assification	Field T	est	Labo	ratory	Relative	
Dam/Test Pit	Can #	Zone Tested	Field	Lab	Yd (pcf)	w/c (%)	γdmax (pcf)	Opt w/c (%)	Compaction (%)	Remarks
STK-TP1	#14	Shell	SM w/gravel	SM	126.8	9.0	-		102.6	
STK-TP1	#15	tt .	11	11	119.9	9.2	123.6	9.4	97.0	
STK-TP1	#4	11	11	11	114.4	8.9			92.6	
STK-TP3	#11	Shell	SM w/gravel	SM	125.5	6.4			98.3	
STK-TP3	#10	17	11	††	120.8	7.1	127.7	9.7	94.6	
STK-TP3	#7	11	11		135.2	7.4			105.9	
STK-TP4	#1	Founda- tion	SM w/gravel	SM	113.6	7.9			92.0	
STK-TP4	#2	11	tt	11	113.2	7.2	123.5	10.8	91.7	
STK-TP4	#5	tt	††	tt	120.3	6.4			97.4	

Table D-6

Relative Compaction Evaluation

Gypsum Wash

			Unified Soil Class	sification	Field T	est	Labor	ratory	Relative	
Dam/Test Pit	Can #	Zone Tested	Field	Lab	γ _d (pcf)	w/c (%)	γdmax (pcf)	Opt w/c (%)	Compaction (%)	Remarks
GW-TP2	#14	Founda- tion	SC-SM w/gravel	SM	+95.9	8.9		,	76.4	Water contents
GW-TP2	#2	11	††	11	114.4	7.5	125.6	10.1	91.1	
GW-TP2	#7	TT	11		97.4	6.3			77.5	11
GW-TP3	#15	Shell	CL w/silt, sand, gravel	SM	103.1	16.7	122.0	11.0	84.5	11
GW-TP3	#9	11	11	11	107.6	13.3	1220	11.0	88.2	11
GW-TP4	#13	Founda- tion	SM w/gravel	SM	125.3	5.2	120.4	7.3	104.1	11
GW-TP4	#11	tt	tt .	11	127.9	4.6	12011		106.2	tt

Note: + Field tests required rock correction.

Table D-7

Relative Compaction Evaluation

Frog Hollow

		_	Unified Soil Class	sification	√ Field T		Labor		Relative	
Dam/Test Pit	Can #	Zone Tested	Field	Lab	γ _d (pcf)	w/c (%)	γ _{dmax} (pef)	Opt w/c (%)	Compaction (%)	Remarks
FH-TP1	#13	Shell (Zone I)	CL w/silt, sand, gravel	CL-ML	111.6	13.3			95.1	
FH-TP1	#15	11	11	11	116.8	12.8	117.4	13.5	99.5	
FH-TP1	#11	TT .	11	tt	115.5	12.0			98.4	
FH-TP2	#9,10	Shell (Zone II)	GC-GM w/cobbles	GC-GM	72.2+	14.4			79.3	12" sandcone.
	`	· · · · · ·					116.2	15.4		
nii mpo	" 7 1 0			••	1000	400				
FH-TP2	#7,12	,	II.	11	106.8+	10.0			91.9	12" sandcone.
FH-TP3	#2	Founda- tion	ML w/clay, sand	ML-CL	90.2	12.5	112.1	16.2	80.5	
FH-TP3	#14	ii V	11	11	99.9	11.8	112.1	10.2	89.1	

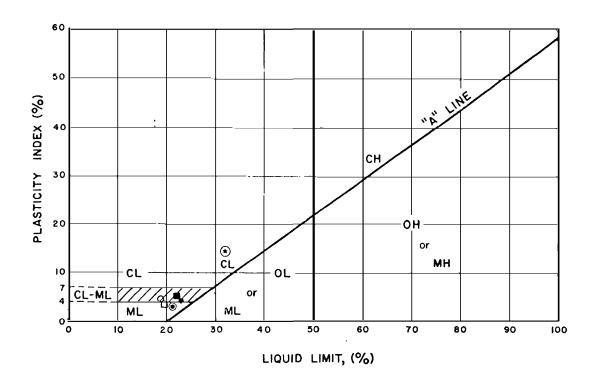
Note: + Field test required rock correction.

Table D-8

Relative Compaction Evaluation

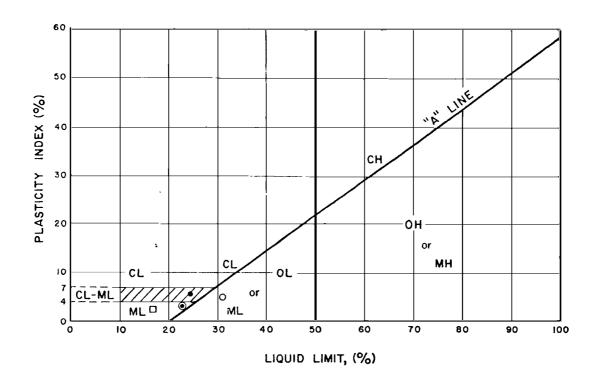
Ivins Diversion No. 5

			Unified Soil Cla	ssification	Field T	est	Labor	atory	Relative	
Dam/Test Pit	Can #	Zone Tested	Field	Lab	γ _d (pcf)	w/c (%)	γ _{dmax} (pcf)	Opt w/c (%)	Compaction (%)	Remarks
IV-TP1	#3	Founda- tion	ML w/gravel,	SM	113.1	10.2			95.4	-
IV-TP1	#9	11	11	11	102.8	10.9	118.6	13.1	86.7	
IV-TP1	#9(2)	11	π	11	106.2	10.5			89.5	Sandcone sample is coarser than corresponding, bulk sample.
IV-TP2	#10	Founda- tion	SM w/gravel, cobbles	SM	92.1	8.9			83.3	
IV-TP2	#5	11	11	rt	91.0	9.3	110.5	11.5	82.4	
IV-TP2	#15	11	11	tt	93.8	8.6			84.9	
IV-TP4	#13	Shell	SM w/gravel	SM	116.3	6.1			95.7	
IV-TP4	#8	tt	11	tf	110.7	5.6	121.5	10.2	91.1	
IV-TP4	#6	11	tt .	I†	128.1	4.5			105.4	



SYMBOL	BORING NO.	MATERIAL	LIQUID LIMIT, %	PLASTICITY INDEX, %	USC SYMBOL
0	GL3 - TP1	SHELL	19.3	3.7	ML
•	GL3 - TP4	SHELL	21.0	3.3	SM - ML
0	GL2 - TP2	SHELL	18.8	4.2	SM - SC
•	GL2 - TP5	FOUND	23.4	4.6	SM - SC
•	GL5 - TP1	SHELL	32.1	14.4	CL
	GL5 - TP2	FOUND	21.9	4.7	SM - SC
	1			<u> </u>	

		_							
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SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS SUMMARY OF ATTERBERG LIMITS									
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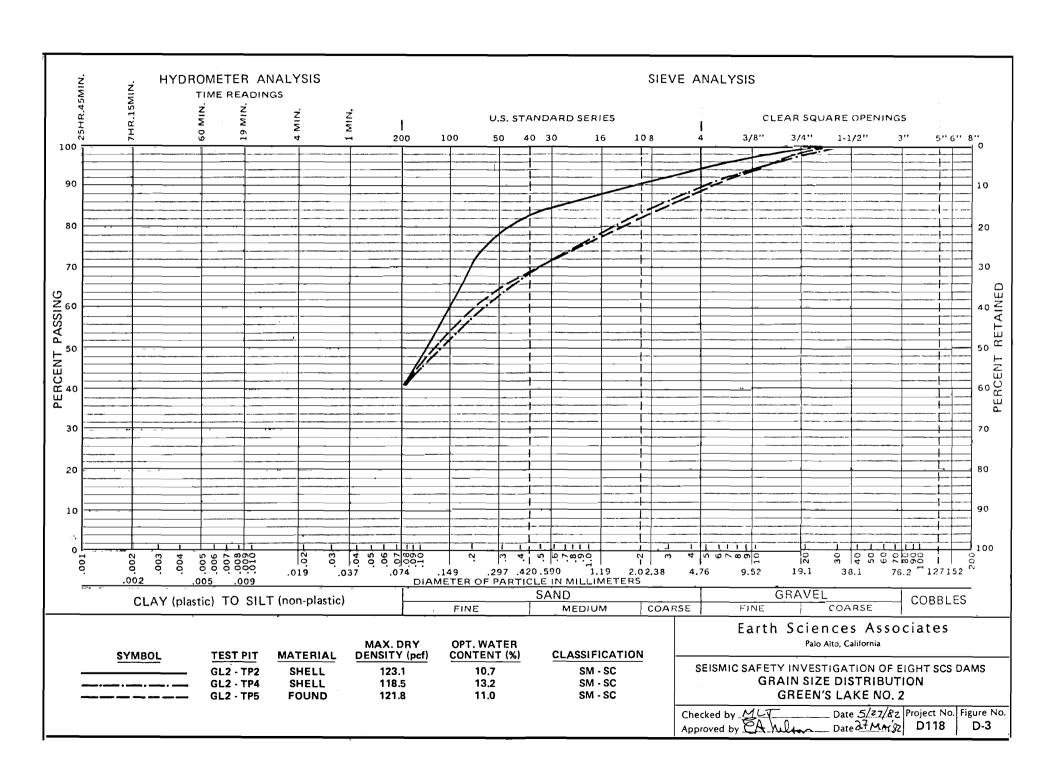


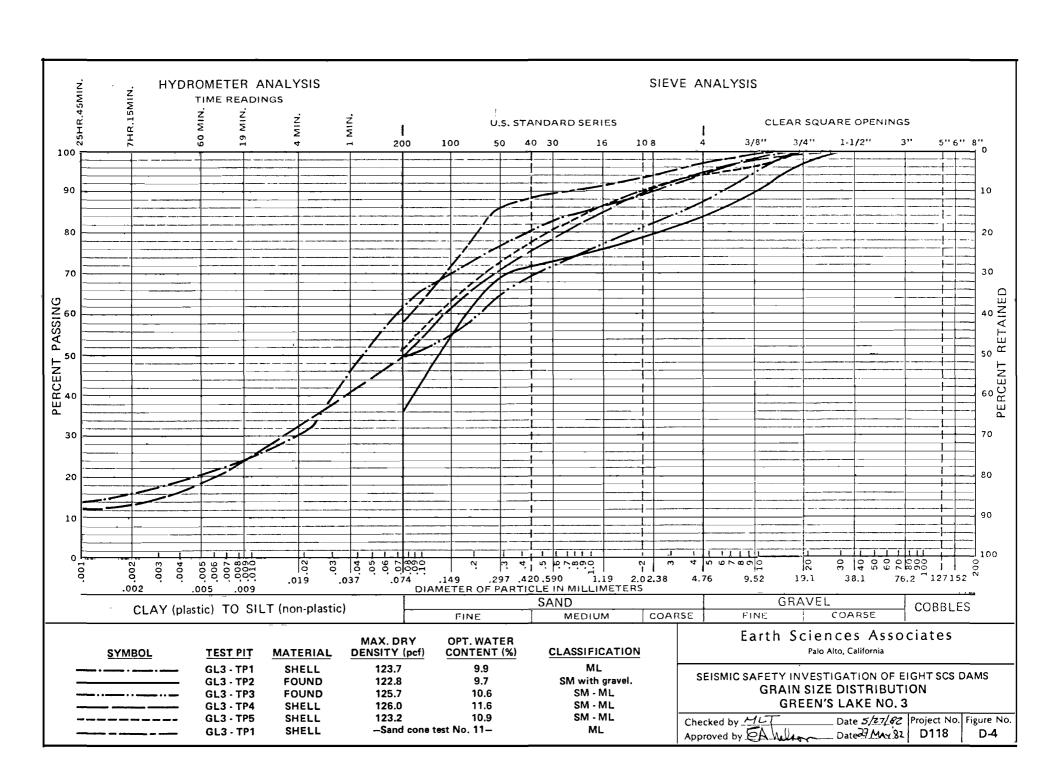
SYMBOL	BORING NO.	MATERIAL	LIQUID LIMIT, %	PLASTICITY INDEX, %	USC SYMBOL
•	FH - TP1	SHELL	24.4	5.5	CL - ML
0	FH - TP2	FOUND	32.1	5.3	GC - GM
0	STK - TP3	SHELL	16.4	2.5	SM
•	GW - TP3	SHELL	22.3	3.1	SM
			1		
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			ı		

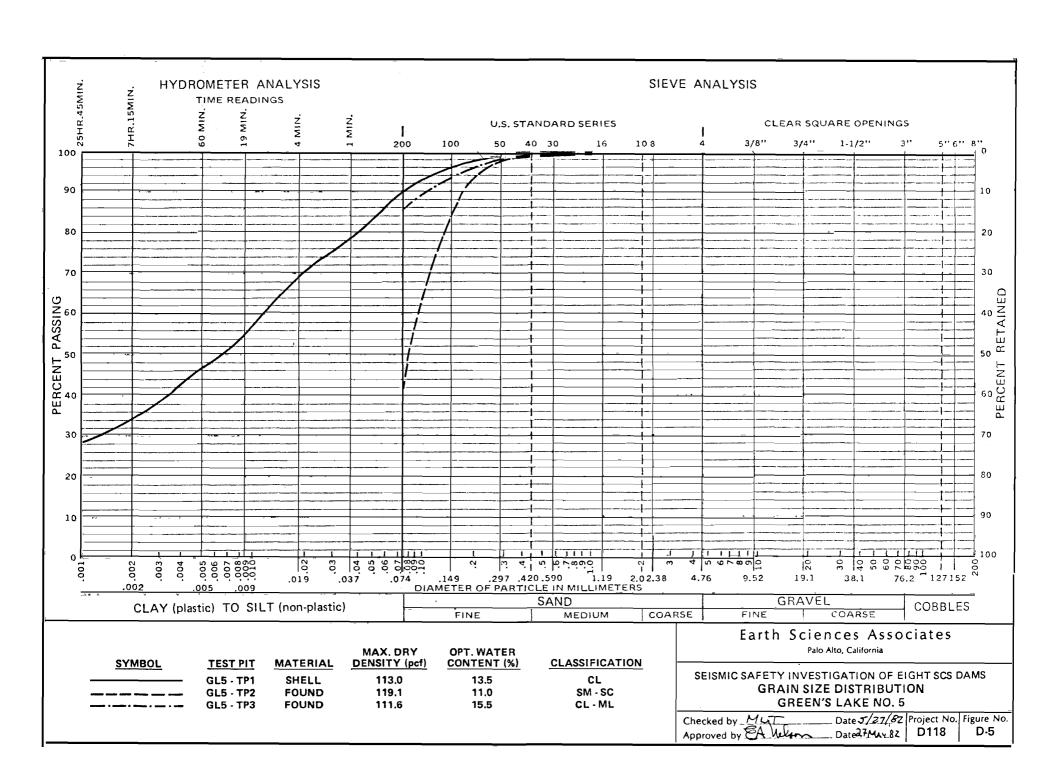
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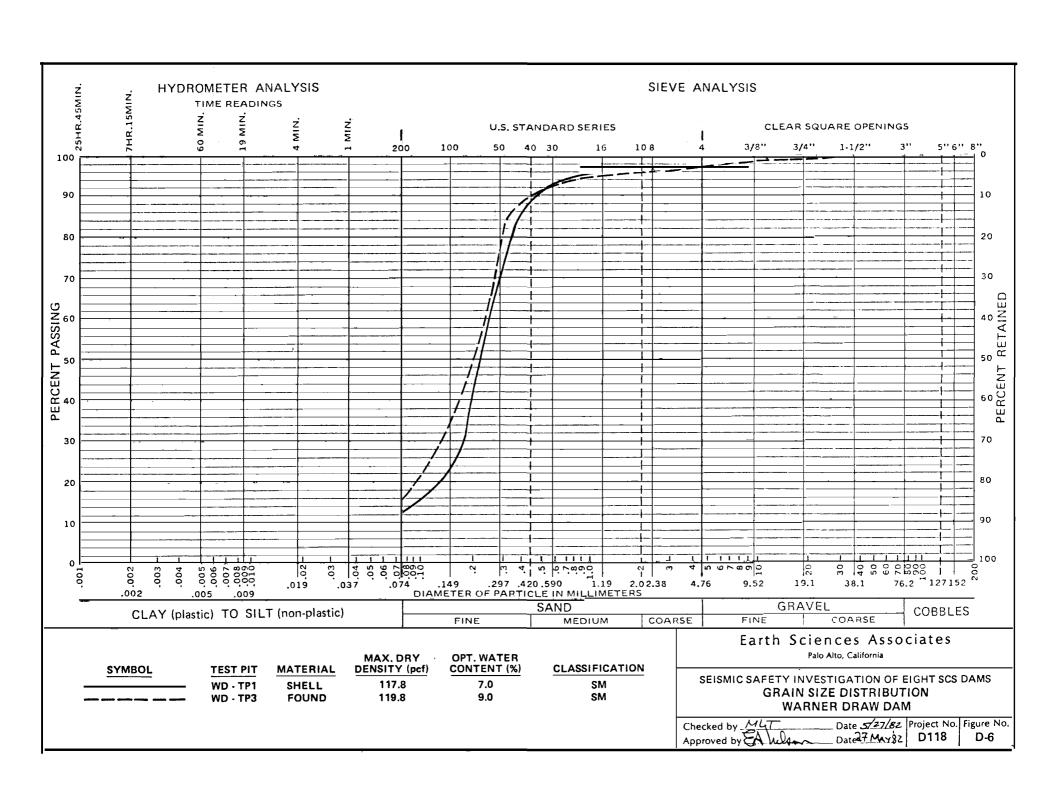
SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS SUMMARY OF ATTERBERG LIMITS

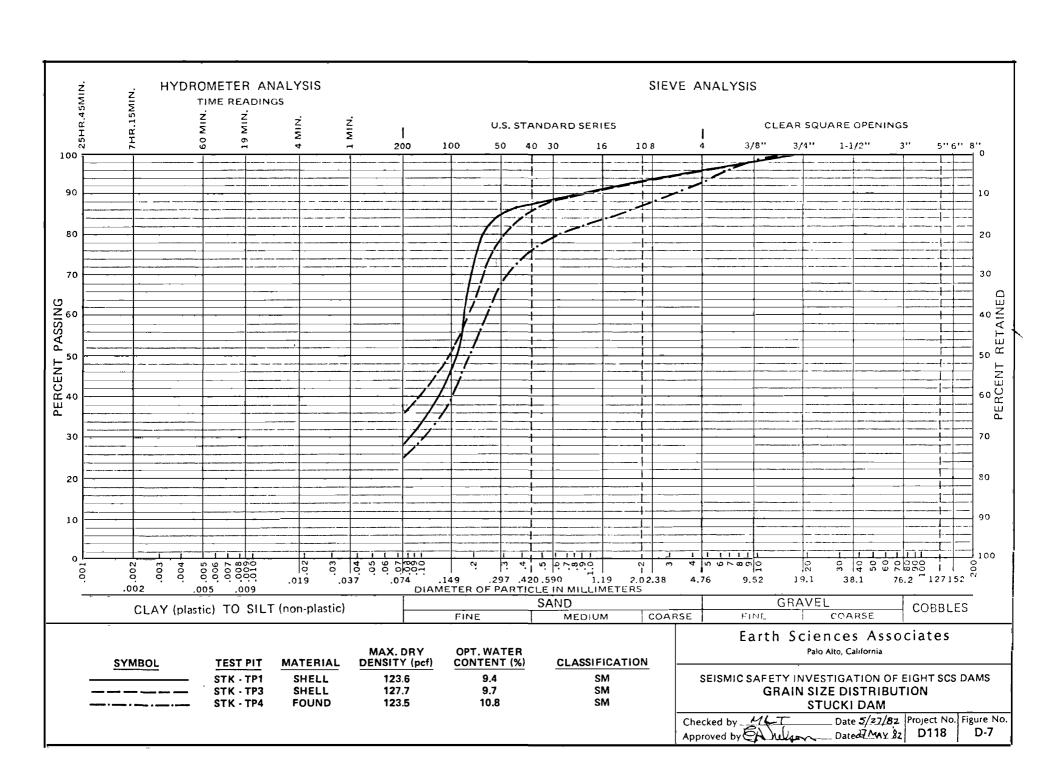
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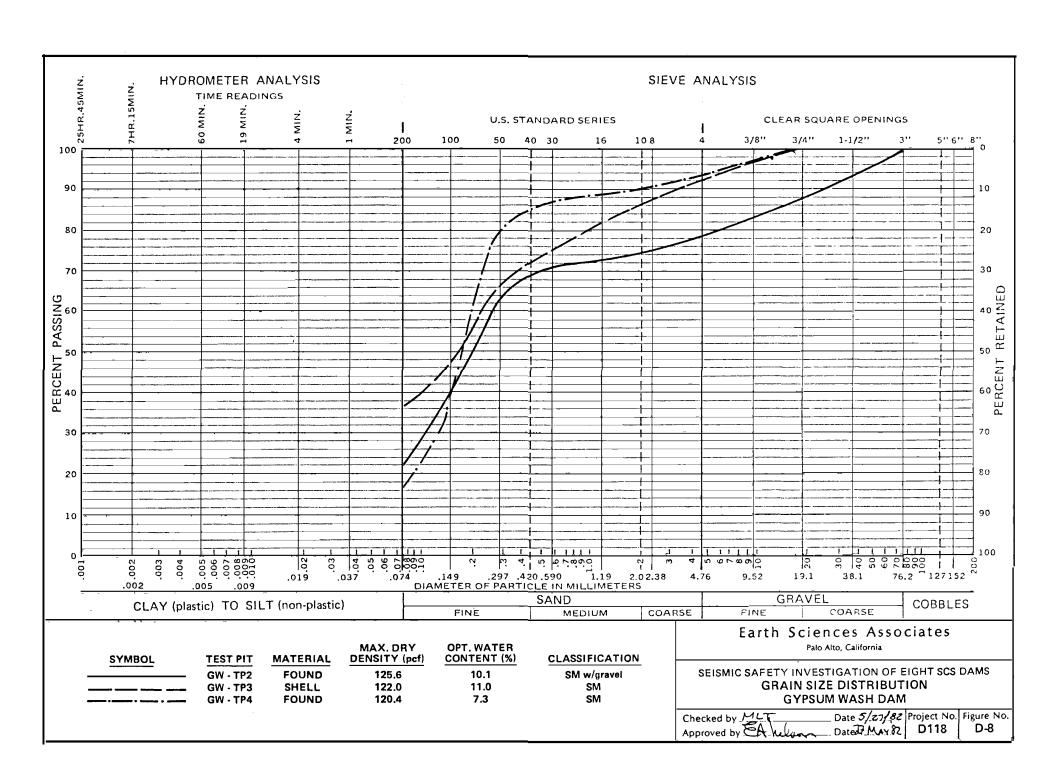


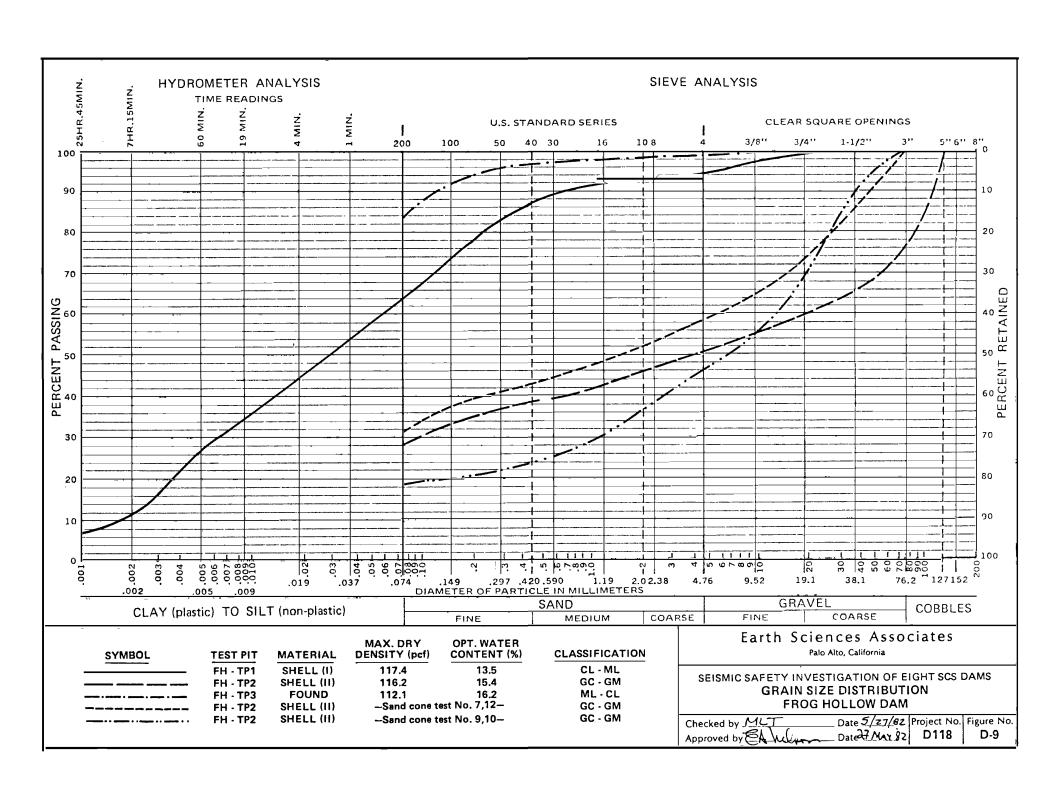


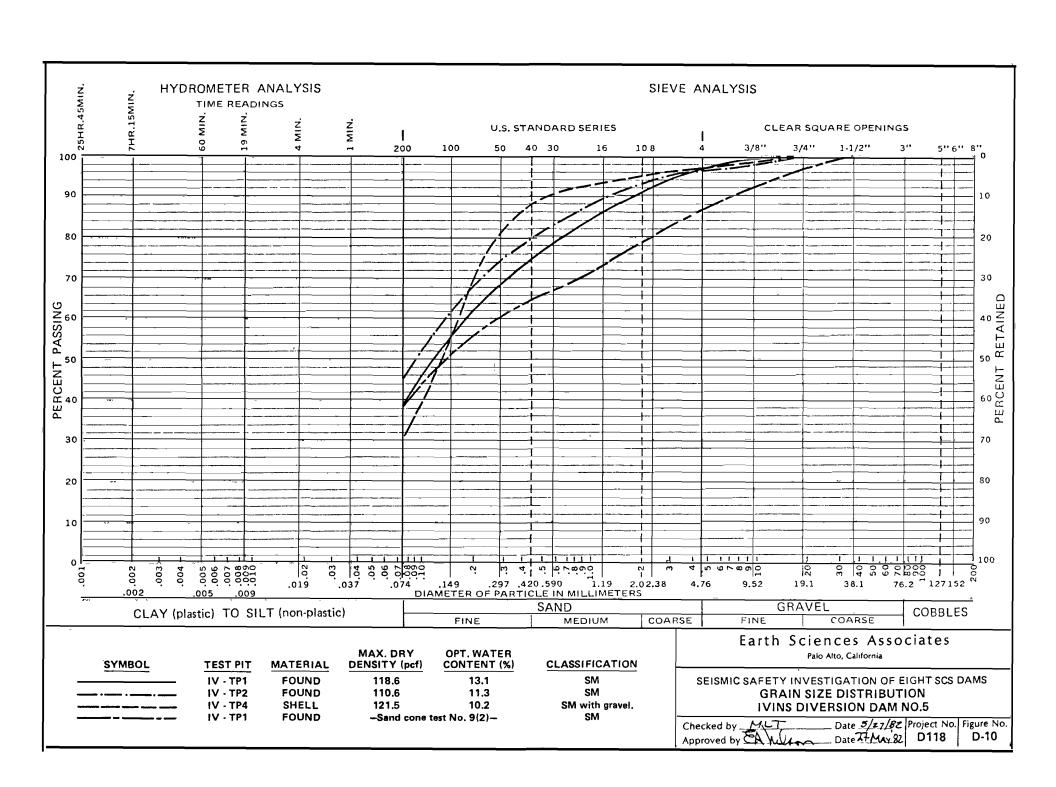


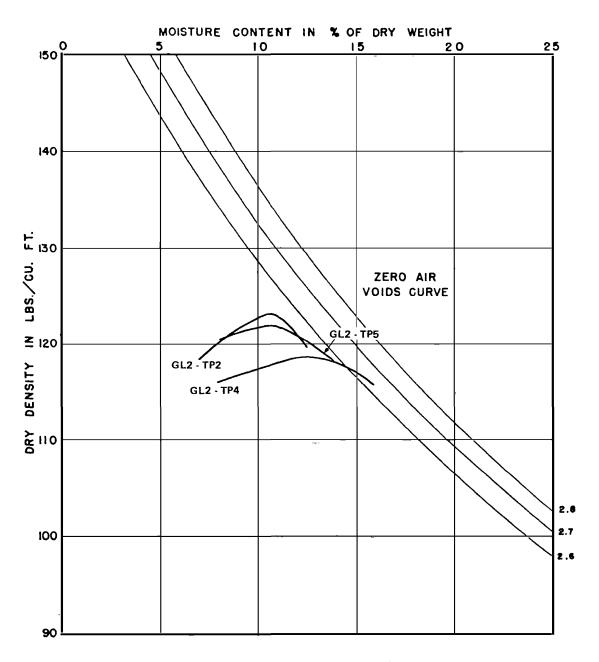












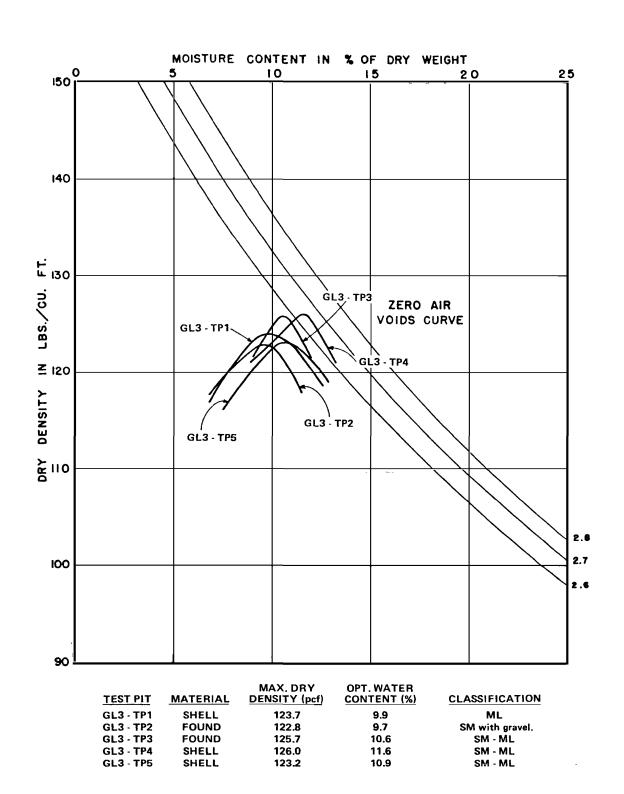
TEST PIT	MATERIAL	MAX DRY DENSITY (pcf)	OPT. WATER CONTENT (%)	CLASSIFICATION
GL2 - TP2	SHELL	123.1	10.7	SM - SC
GL2 - TP4	SHELL	118.5	13.2	SM - SC
GL2 - TP5	FOUND	121.8	11.0	SM - SC

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SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
MAXIMUM DENSITY CURVES
GREEN'S LAKE NO. 2

Checked by MAT	Date 5/27/82	Project No.	Figure No.
Checked by MAT Approved by A Alfan	Date 27 May 82	D118	D-11

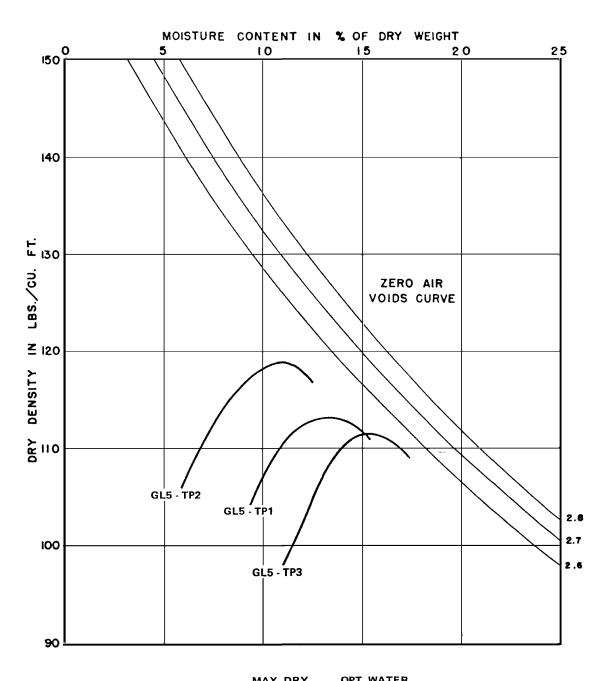


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SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS **MAXIMUM DENSITY CURVES GREEN'S LAKE NO. 3**

Checked by MUT Date 5/27/82 Project No. Approved by A W. 1400 Date 7 May 82 D118	Figure No.	Project No.	Date 5/27/82	Checked by MUT
Approved by Car March Dates 11-183 122	D-12	D118	Date MAY 82	Approved by EA William



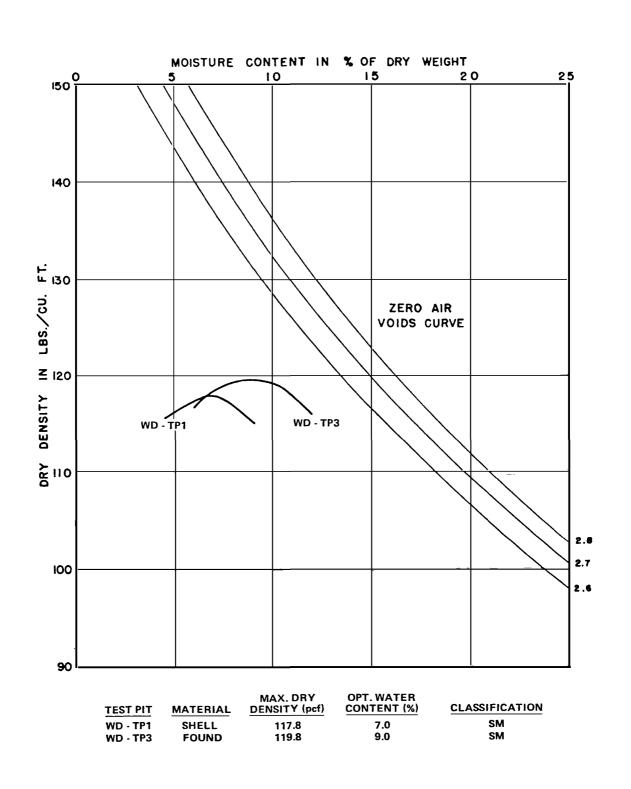
TEST PIT	MATERIAL	DENSITY (pcf)	CONTENT (%)	CLASSIFICATION
GL5 - TP1	SHELL	113.0	13.5	CL
GL5 - TP2	FOUND	119.1	11.0	SM - SC
GL5 - TP3	FOUND	111.6	15.5	CL - ML

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SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
MAXIMUM DENSITY CURVES
GREEN'S LAKE NO. 5

Checked by MLT Approved by A We 4000	Date 3/27/82	Project No. D118	Figure No. D-13
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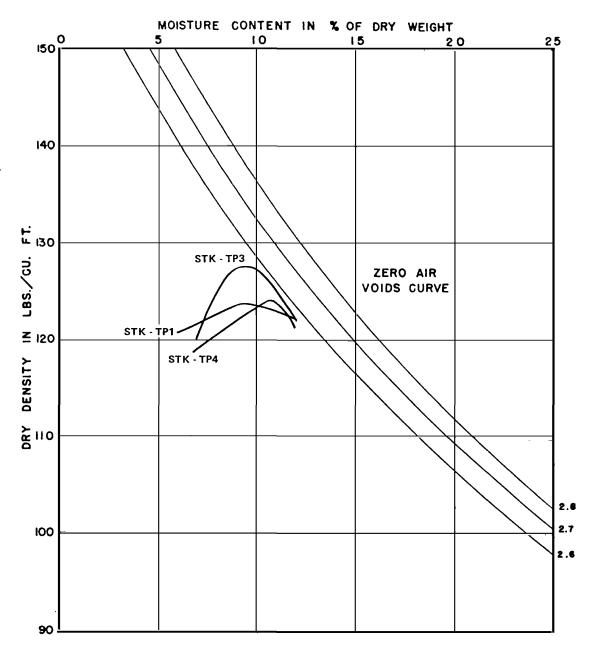


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SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
MAXIMUM DENSITY CURVES
WARNER DRAW DAM

ı	Checked by MUT	Date 5/27/8	2 Proje	ect No.	Figure No.
1	Checked by MUTAN	Date 5/27/87 Date 27 MAY 8	82 D	118	D-14
	Approved of Carting Otto		7		1



TEST PIT	MATERIAL	MAX. DRY DENSITY (pcf)	OPT. WATER CONTENT (%)	CLASSIFICATION
STK - TP1	SHELL	123.6	9.4	SM
STK - TP3	SHELL	127.7	9.7	SM
STK - TP4	FOUND	123.5	10.8	SM

Earth Sciences Associates

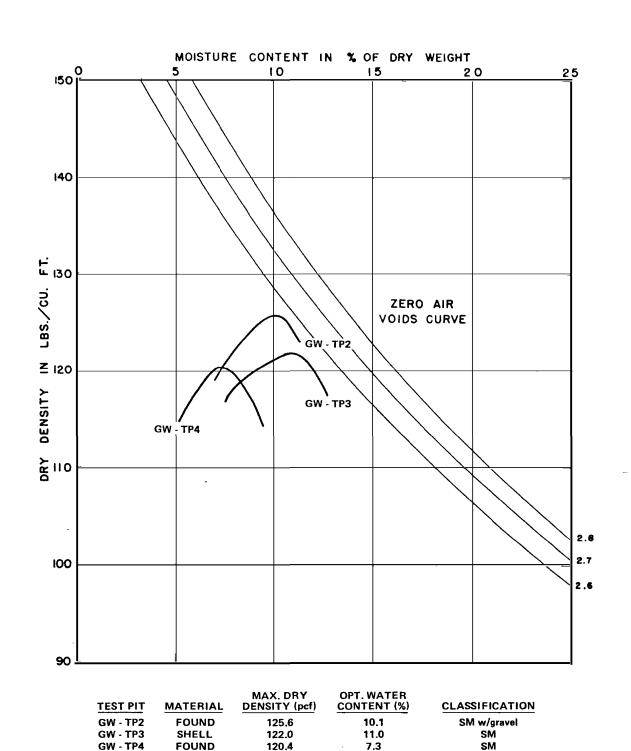
Palo Alto, California

SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS

MAXIMUM DENSITY CURVES

STUCKI DAM

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7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7			1



Palo Alto, California

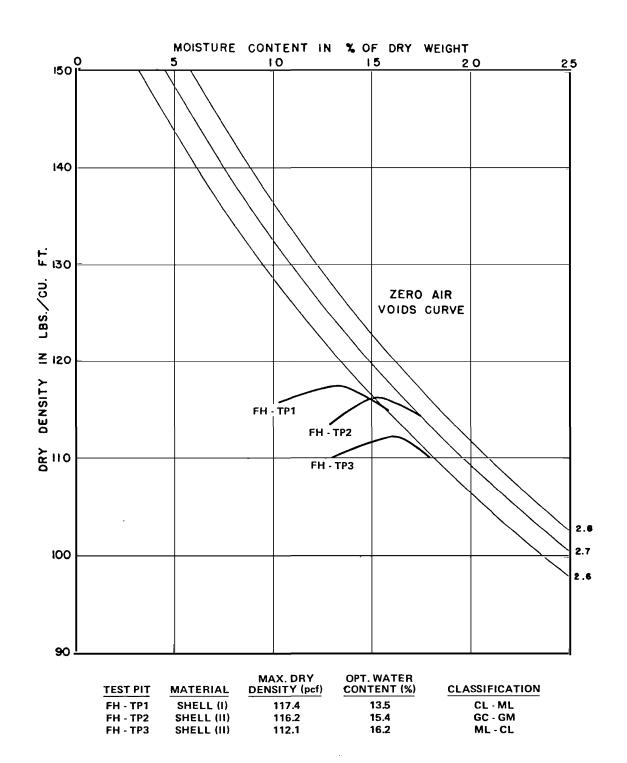
Test procedure: ASTM D698-70

METHOD C.

SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS **MAXIMUM DENSITY CURVES GYPSUM WASH DAM**

Earth Sciences Associates

Checked by MLT Date 5/21/82 |
Approved by A VIII Date 77/MY 82 Date 5/21/82 Project No. Figure No. D-16 D118



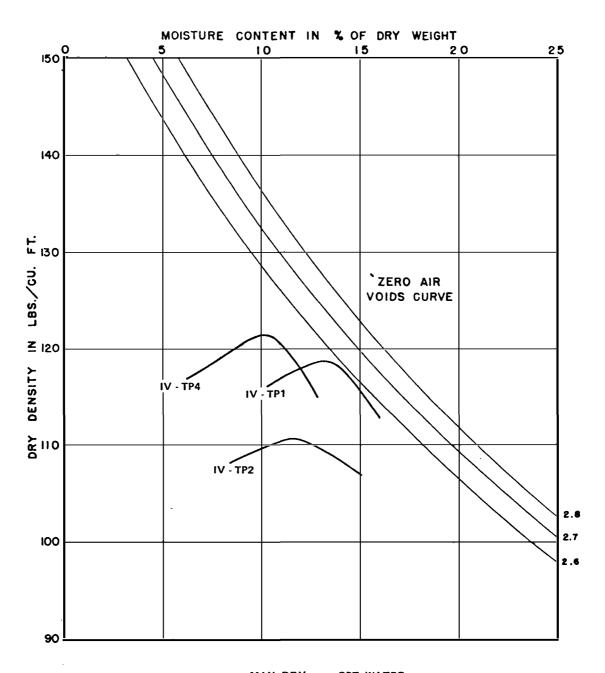
Test procedure: ASTM D698-70 METHOD C.

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SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
MAXIMUM DENSITY CURVES
FROG HOLLOW DAM

Checked by Date 5/27/82 Project No. Figure No. Approved by Date Date 27/May 87 D118 D-17



TEST PIT	MATERIAL	MAX. DRY DENSITY (pcf)	OPT. WATER CONTENT (%)	CLASSIFICATION
IV - TP1	FOUND	118.6	13.1	SM
IV - TP2	FOUND	110.6	11.3	SM
IV - TP4	SHELL	121.5	10.2	SM with gravel.

Test procedure: ASTM D698-70 METHOD C.

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SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
MAXIMUM DENSITY CURVES
IVINS DIVERSION DAM NO.5

Checked by MLT	Date <u>5/27/82</u> Project No	Figure No.
Checked by MLT Approved by Aulan	Date HMAY 82 D118	D-18

APPENDIX E

REGIONAL GEOLOGY AND TECTONICS

APPENDIX E

REGIONAL GEOLOGY AND TECTONICS

- E.1 Tectonic Framework
 - E.1.1 General Setting
 - E.1.2 Tectonic Processes and Crust-Mantle Framework
 - E.1.3 Tectonic Provinces
 - E.1.3.1 Northern Basin and Range (Great Basin) Province (NBR)
 - E.1.3.2 Southern Basin and Range (Sonora Desert) Province (SBR)
 - E.1.3.3 Colorado Plateau Province (CP)
 - E.1.3.3.1 High Plateaus Section of the Colorado Plateau Province (CP-HPS)
 - E.1.3.3.2 Grand Canyon Section of the Colorado Plateau Province (CP-GCS)
 - E.1.3.4 Basin and Range-Colorado Plateau Transition (BR-CP Trans.)
 - E.1.3.5 Great Basin Sonoran Desert Transition (GB-SD Trans.)
- E.2 Geologic History of Southwest Utah
 - E.2.1 Pre-Cenozoic Geologic History
 - E. 2.1.1 Precambrian-Paleozoic Geology
 - E.2.1.2 Mesozoic History
 - E. 2. 2 Cenozoic Geologic History
- E.3 Structural Geology
 - E.3.1 General Structure of Study Area
 - E.3.2 Structure in the Cedar City Area
 - E.3.3 Structure in the Hurricane-Frog Hollow Area
 - E.3.4 Structure in the St. George Area
 - E. 3.5 Structure in the Ivins Area

FIGURES

No.	<u>Title</u>
E-1	Tectonic Sketch Map of Western North America
E-2	Crust-Upper Mantle Structural Profile Across Basin and Range-Colorado Plateau Transition
E-3	Tectonic Provinces and Subprovinces
E-4	Physiographic Map of Southwestern Utah
E-5	Sketch Map of Major Pre-Cenozoic Tectonic Features Affecting Southwestern Utah
E-6	Representative Stratigraphic Column for Southwestern Utah
E-7	Synthesis of Cenozoic Tectonic Pattern in Western United States
E-8	Physiographic Features and Major Structural Features of Southwestern Utah
E-9	Geologic Sketch Map Showing Structural Overview of Southwestern Utah
E-10	East-West Structural Sections Across Southwestern Utah
E-11	Map of Cenozoic Structural Elements East of Cedar City
E-12	Geological Map of the St. George Area
E-13	Geological Map of the Ivins Area

E. REGIONAL GEOLOGY AND TECTONICS

E.1. TECTONIC FRAMEWORK

E.1.1 General Setting

The Cedar City-St. George area, located in the southwestern corner of Utah, lies along the boundary (more properly, within a transition zone) between the Basin and Range and Colorado Plateau physiographic-geologic provinces (fig E-1). The area is in a tectonically active intraplate setting about 400 km northeast of the San Andreas transform fault, which forms the boundary between the Pacific and North American plates (fig.E-1.

The eastern margin of the Great Basin, or northern portion of the Basin and Range province, is well known to be tectonically active. It coincides spatially with a lengthy segment of the Intermountain seismic belt (Smith and Sbar, 1974; Smith, 1978; Arabasz and Smith, 1981)—schematically depicted infigure E-1-a major zone of intraplate seismicity within western North America that follows the boundary between thin weak crust and lithosphere of the eastern Basin and Range province and thicker more stable crust and lithopshere of the Middle Rocky Mountains and the Colorado Plateau. The Intermountain seismic belt is characterized by extensive late Quaternary normal faulting, diffuse shallow seismicity, and episodic scarp-forming earthquakes (6½<M<7½) (Arabasz and Smith, 1981).

In considering the tectonic setting of the Cedar City-St. George area along the eastern margin of the Basin and Range province, it should be noted that the area is situated close to an important boundary at roughly lat 37°N that separates the northern, or Great Basin, section of the Basin and Range province from the southern, or Sonoran Desert, section (see fig.E-1). The boundary, whose deep-seated origin is not fully understood, is marked by a change in Bouguer gravity of nearly 100 mgal (Eaton and others, 1978).

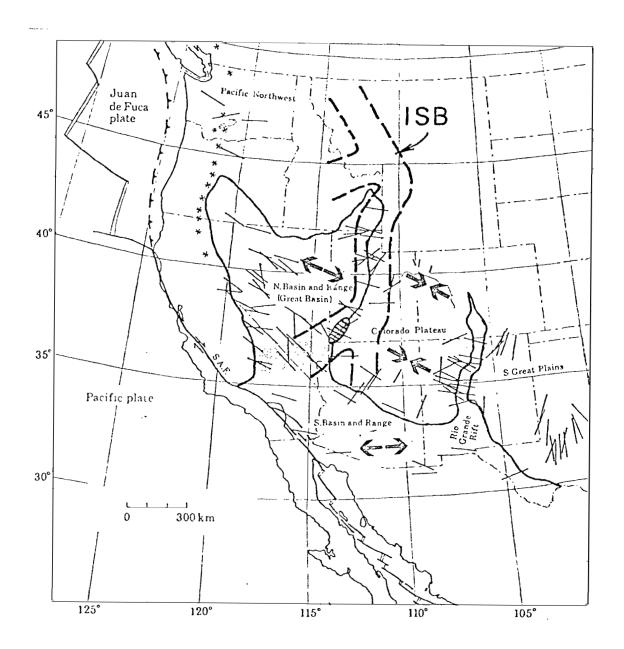


FIGURE E-1-Tectonic sketch map (adapted from Zoback and others, 1981) showing setting of the Cedar City-St. George area (dashed ellipse) in southwestern Utah. Heavy dashed lines outline Intermountain seismic belt, ISB (after Arabasz and Smith, 1981); heavy arrows, generalized stress directions: either least (outward directed) or greatest (inward directed) principal horizontal compression (from Zoback and Zoback, 1980). Shaded region in southern Nevada and Utah marks transitional

boundary between northern and southern Basin and Range provinces.

Importantly, the two sections of the Basin and Range province, which are physiographically different, have had significantly different late Cenozoic histories (Eaton and others, 1978; Eaton, 1979; Zoback and others, 1981), a point we return to later.

The Cenozoic history of the Cedar City-St. George area is closely associated with tectonism along the Great Basin-Colorado Plateau boundary. However, proximity of the area to the Great Basin-Sonoran Desert transition must be emphasized. Generalizations applicable to the contemporary seismotectonics of southwestern Utah will not generally be applicable to north-western Arizona--to the south of the area of interest.

E. 1.2 Tectonic Processes and Crust-Mantle Framework

A host of hypotheses have been proposed to explain the late Cenozoic tectonics of the Basin and Range province, the Colorado Plateau, and the Intermountain West in general. Abundant information has recently been summarized in Memoir 152 of the Geological Society of America (Smith and Eaton, eds., 1978) and in subsequent papers by Thompson and Zoback (1979), Eaton (1979, 1980), Zoback and Zoback (1980), and Zoback and others (1981). Stewart (1978) and Davis (1980) review the chief categories of models that may pertain:

(1) oblique intraplate fragmentation caused by differential motion along the North American-Pacific plate boundary; (2) backarc extension and mantle upwelling, typically above a subducting plate; (3) major changes in dip or configuration of a subjacent subducted plate (or plates); (4) subduction of the East Pacific rise; and (5) mantle plumes.

For the present site-specific study, postulated tectonic models become chiefly academic (in the face of continuing debate and on-going evolutionary

studies), so we emphasize some fundamental observations. (Appendix J includes a selected bibliography relevant to the "big picture" tectonics of western North America). Later, however, we describe aspects of some tectonic models for which there seems to be good academic agreement, and which provide a useful conceptual framework for discussion. Ultimately, the record of late Quaternary faulting and observed seismicity—discussed in Appendix F—provide the basis for evaluating contemporary tectonics within a time frame of engineering relevance.

The eastern margin of the Great Basin has been interpreted to define a subplate boundary between the Great Basin and the Colorado Plateau-Middle Rocky Mountains (e.g., Smith, 1978); in any case, it clearly is a locus of see fig. E-2a tectonic instability reflected, in part, by thin crust (\sim 25 km), anomalous upper mantle ($7.4 \le P_n \le 7.9$ km/sec), high heat flow (>120 mW/m²; 1 HFU = 41.9 mW/m²), high regional elevation (>1,500 m), and a crustal low-velocity layer in the 5-15 km depth interval (see fig. E-2b) (Smith, 1978; Eaton and others, 1978; Blackwell, 1978).

Figure E-2c illustrates the anomalous nature of crust-upper mantle structure along the eastern Great Basin; it also depicts transitional change from the Basin and Range province to the Colorado Plateau. The boundary between the Basin and Range province and the Colorado Plateau province in central and southwest Utah is well known to be a transitional one--not only physiographically (e.g., Stokes, 1977), but also in terms of surficial geology (Burchfiel and Hickcox, 1972; Anderson and Mehnert, 1979), crustal velocity structure (Smith and others, 1975), heat flow (Bodell and Chapman, 1982), lithospheric thickness (Thompson and Zoback, 1979), and other geophysical parameters (see reviews by Thompson and Zoback, 1979; Smith, 1978; Keller and others, 1979). Typically, the geophysical parameters tend to

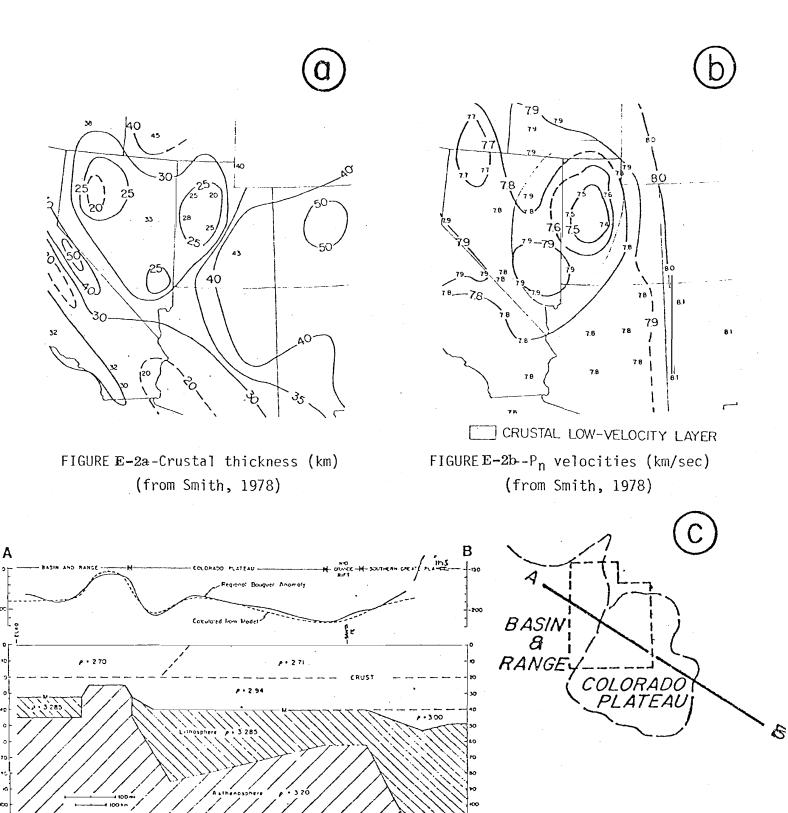


FIGURE E-2c-Crust-upper mantle structure along NW-SE profile that crosses Basin and Range-Colorado Plateau transition (see location map on right) (from Thompson and Zoback, 1979).

have transitional values that extend several tens of kilometers east of the classical physiographic boundary between the two provinces.

In the Cedar City-St. George area, crustal thickness is ~ 30 km and subcrustal P_n velocity is ~ 7.9 km/sec (fig. E-2a, b). Smith and others (1975) and Smith (1978) summarize evidence for anomalous teleseismic P-wave residuals in southwestern Utah (see also Anderson, 1978, p. 3), supportive of relatively thin crust and anomalous upper mantle. It should be emphasized that the observed P-wave anomalies were not unique to southwest Utah--because of limited spatial sampling--but rather they would presumably typify a more extensive belt along the eastern Great Basin margin.

Contemporary stress is predominantly extensional in a WNW-ESE direction along the eastern Great Basin (fig.E-1), but there are spatially rapid changes in stress orientation in local areas, and there is a notable change to a compressional stress regime (with roughly a WNW-ESE maximum principal horizontal stress direction) within the interior of the Colorado Plateau (Zoback and Zoback, 1980). The source of the observed stress field in the Intermountain West is not completely understood, but it likely involves contributions both from differential motion along the San Andreas margin and sublithospheric tractions (Zoback and Zoback, 1980; Thompson and Zoback, 1979; Smith, 1978).

Residual effects of early Cenozoic subduction along the western margin of the North American plate are thought to be of great importance to understanding the stress and thermal fields (and hence the contemporary tectonics) of the Intermountain West (Thompson and Zoback, 1979; Zoback and Zoback, 1980; Bodell and Chapman, 1982). Bodell and Chapman (1982), for example, discuss how transitional changes in observed heat flow across the Basin and Range-

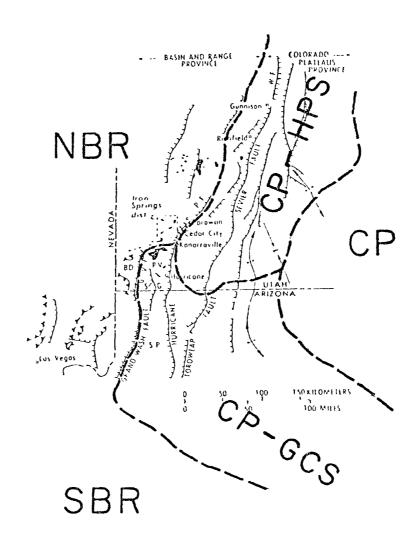
Colorado Plateau boundary (ranging from >100 mW/m² in the eastern Great Basin to >50 mW/m² within the interior of the Colorado Plateau) may be a significant clue to understanding the late Cenozoic tectonics along the province boundary. Their thermal modeling suggests that lateral warming and weakening of the Colorado Plateau lithosphere began at the Basin and Range boundary some 20 m.y. ago. Lithospheric thinning under the Plateau would have quickly resulted in surface uplift, but corresponding surface heat flux would predictably lag in time by as much as 15-20 m.y.--in agreement with current observations (Bodell and Chapman, 1982). Long-continued thermal weakening apparently had earlier resulted in anomalously thin crust and lithosphere beneath the Basin and Range province as seen in figure E-2q(see Eaton, 1979).

To adequately describe the active tectonic processes in southwest Utah, we need to consider the complicated Cenozoic history of the region, as well as its contemporary seismotectonics, seismic flux, and strain rate--all of which are elaborated in later sections. From what we have seen so far, we can anticipate that (1) because of its intraplate setting, southwest Utah will expectedly be associated with lower rates of secular deformation than along the San Andreas plate margin, (2) crustal extension and thermal processes will have a fundamental influence on earthquake generation, and (3) normal faulting associated with deformation of the Great Basin-Colorado Plateau boundary will be of primary importance.

E.1.3 Tectonic Provinces

In figure E-3,tectonic provinces are outlined in the region surrounding the Cedar City-St. George area. This figure and the following brief descriptions are intended to place the study area in regional perspective on a scale of a few hundreds of kilometers. Various workers differ in their delineation of some of the province boundaries, or perhaps in their outline of subprovinces--depending upon their emphasis of physiography, geologic structure, or geophysical attributes (see Eaton, 1979). We have attempted to follow well-established usage, allowing for updating by more recent geological work. Also, we portray and depict two transitional regions that do not have precisely determined boundaries, but whose general location has important tectonophysical implications. The tectonic provinces outlined in figure E-3 are chiefly based on geological structure and physiography. They are not uniform seismogenic zones; we discuss such zones separately in Appendix F.

E.1.3.1 Northern Basin and Range (Great Basin) Province (NBR). The northern Basin and Range province encompasses a broad region of Nevada and western Utah characterized by basin-range physiography—i.e., elongate ranges with a length-to-width radio of ~4-8 and crest-to-crest spacing of ~25-35 km, separated by basins filled with Tertiary-Quaternary fluvial-lacustrine sediments (Zoback and others, 1981). Mention has already been made of separation of the Basin and Range province into northern and southern sections. The northern or Great Basin section stands mostly above about 1,500 m (5,000 ft), is marked by interior drainage, and has predominant NNE-trending physiography that has chiefly developed within the past 10 m.y. (Eaton 1979; Zoback and others, 1981). Basin-range structure, which has led to the development of



TECTONIC PROVINCES AND SUBPROVINCES

NBR	Northern Basin and Range (Great Basin) Province
SBR	Southern Basin and Range (Sonoran Desert) Province
CP	Colorado Plateau Province.
CP-HPS	High Plateaus Section of the Colorado Plateau Province
CP-GCS	Grand Canyon Section of the Colorado Plateau Province
	Basin and Range-Colorado Plateau Transition*
	Great Basin-Sonoran Desert Transition*

^{*}not shown on this figure

normal-fault-controlled alluviated basins and mountain blocks by crustal extension, "is commonly inferred to represent either (1) blocks tilted along downward-flattening (listric) faults in which the upslope part of an individual rotated block forms a mountain and the downslope part a valley or (2) alternating downdropped blocks (grabens) that form valleys and relatively upthrown blocks (horsts) that form mountains" (Stewart, 1978, p. 1).

- E.1.3.2. Southern Basin and Range (Sonoran Desert) Province (SBR). In contrast to the Great Basin, the Sonoran Desert section of the Basin and Range province lies mostly below about 900 m (3,000 ft) and has a general NW-trending physiography that developed by block faulting, chiefly in the period 13-10 m.y. ago (Zoback and others, 1981; Eaton, 1979). Earlier timing (and a different stress field) for the development of basin-range structure in the southern Basin and Range province, compared to that in the Great Basin section, accounts for its present low elevation, extensively eroded ranges with broad range-bounding pediments, paucity of active faulting, and relatively low seismicity (Eaton, 1979; Zoback and others, 1981).
- E.1.3.3 Colorado Plateau Province (CP). The Colorado Plateau, sometimes referred to as the Colorado Plateaus province, defines a crudely circular region covering adjoining regions of Utah, Arizona, New Mexico, and Colorado. The province coincides with a relatively coherent block surrounded on three sides by the extensional regimes of the Basin and Range province and the Rio Grande rift (seefig. E-1) that has experienced nearly two kilometers of vertical uplift during the last 20 m.y. while remaining relatively undeformed (Thompson and Zoback, 1979; Hunt, 1956). Cenozoic uplift of the Plateau is

generally ascribed to upper-mantle thermal processes (see McGetchin, 1979; Bodell and Chapman, 1982). The interior of the Colorado Plateau is characterized by a 40 km-thick crust, a P_n velocity of about 7.85 km/sec, a compressive stress field, and an average elevation of about 1,500-2,000 m (Thompson and Zoback, 1979); its average heat flow is about 50 mW/m² (milliwatts/square meter) (Bodell and Chapman, 1982). The pre-Tertiary geology of the Colorado Plateau is dominated by stable shelf deposits; on the other hand, early Tertiary fluvial-lacustrine sediments, late Tertiary intrusive and extrusive rocks (particularly around the margin of the Plateau), and a scarcity of late Cenozoic sedimentary rocks all relate to instability during the Cenozoic (Hunt, 1956). Peripheral regions of the Colorado Plateau province generally are deformed as a result of monoclinal warping and block faulting. We next describe two of these peripheral subprovinces, delineated by Hunt (1956), which form the western and southwestern part of the Colorado Plateau province.

E.1.3.3.1 High Plateaus Section of the Colorado Plateau Province (CP-HPS):

The High Plateaus of central and southwest Utah form a series of high-standing (2,700-3,300 m), NNE-trending mountain blocks that abruptly rise above the eastern Great Basin, separating it from the interior of the Colorado Plateau (fig. E-4). The mountain blocks are generally flat-topped (capped by lower Tertiary formations or volcanic rocks) with typically steep sides and intervening valleys defined by <u>en echelon</u> normal faults of late Tertiary and Quaternary age (Anderson and Rowley, 1975; Hunt, 1956). Extensive volcanism occurred during much of the Cenozoic in the southwestern High Plateaus and adjoining Great Basin (Rowley and others, 1979; Anderson and Rowley, 1975). The distribution of regional ash-flow tuffs leads Rowley and others (1978) to conclude that differential uplift of the High Plateaus

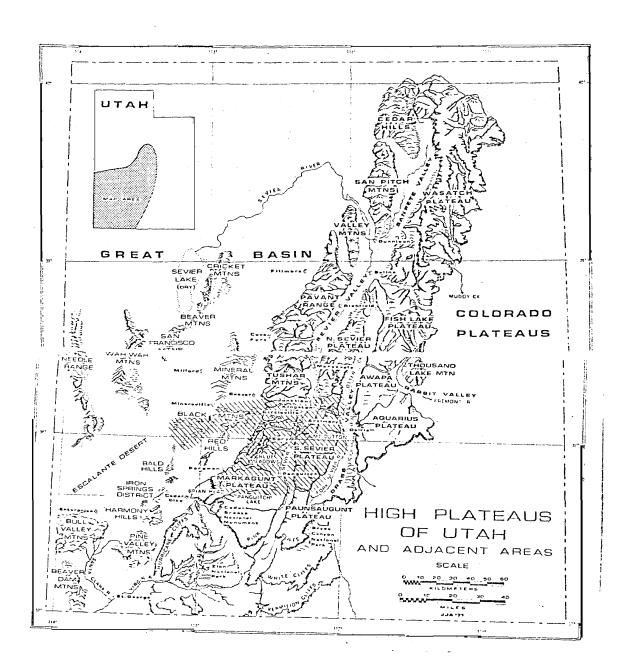


FIGURE E-4-Physiographic map of southwestern Utah showing relation of Cedar City-St. George area to High Plateaus section of the Colorado Plateau province (from Anderson and Rowley, 1975).

to form the Great Basin-Colorado Plateau boundary occurred some time after 29 m.y. ago; structural differentiation was apparently underway by 26 m.y. ago and had produced significant topographic contrasts by 24 m.y. ago. Complexity of the Great Basin-Colorado Plateau boundary in the Cedar City-St. George area is discussed in a later section on structural geology.

E.1.3.3.2 Grand Canyon Section of the Colorado Plateau Province (CP-GCS):

The Grand Canyon section of the Colorado Plateau in Arizona, as delineated by Hunt (1956), encompasses relatively high-standing blocks of nearly horizontal strata that form the southwestern rim of the Plateau. Maximum elevations are less than 2,000 m, significantly lower than in the High Plateaus section. The northern part of the Grand Canyon section is dominated by NNE-trending, down -to-the-west normal faults (Grand Wash, Hurricane, and Toroweap faults) that continue northward into Utah (Hunt, 1956; Hamblin, 1970). Evidence for Quaternary displacements on NNE-trending faults in this area is discussed by Huntoon (1977, 1979; see also Anderson, 1979). Southward, the structural grain and province boundary of the Grand Canyon section become parallel to NW-trending structure of the Sonoran Desert section of the Basin and Range province.

- E.1.3.4 <u>Basin and Range-Colorado Plateau Transition (BR-CP Trans.)</u>. Because the boundary between the Basin and Range and Colorado Plateau provinces in central and southwest Utah is transitional (see Section E.1.2), demarcation of the two provinces is not straightforward. Features of both provinces overlap: Cenozoic normal faulting extends, with diminishing displacements, approximately 50 km into the High Plateaus section of the Colorado Plateau—while many of the easternmost basin-range mountain blocks have a general plateau structure (e.g., Stokes, 1977; Anderson and Rowley, 1975; Spieker, 1949). The Basin and Range-Colorado Plateau transition zone depicted in figure 3 for central and southwest Utah is based on Stokes (1977). Extrapolation of this transition zone outside of Utah is uncertain. West of the Utah-Nevada border the transition zone becomes mixed with the Great Basin-Sonoran Desert transition; in northwestern Arizona, the boundary between the Basin and Range and Colorado Plateau provinces is more clearly demarcated by the Grand Wash fault (Lucchitta, 1979).
- E.1.3.5 <u>Great Basin-Sonoran Desert Transition (GB-SD Trans.)</u>. Division of the Basin and Range province into northern and southern sections was discussed in Section E.1.1. Atransition zone between the two sections of the province, following Zoback and others (1981) and Eaton (1979), is sketched in figure E-3. Comparison with figure 1 shows that the Great Basin-Sonoran Desert transition roughly coincides with a NE-SW trending branch of the Intermountain seismic belt that extends across southern Nevada--following a major structural corridor of late Tertiary deformation that transects a northerly-trending structural grain (Anderson, 1978). The transition zone defined a relatively amagmatic corridor during late Cenozoic time (Anderson, 1981). Also, the zone displays complex structural ties between the northern and southern sections of the

Basin and Range province relating to an episode of thin-skinned extensional tectonics about 15 m.y. ago (Anderson, 1981), although modern basin-range structure in the Great Basin is distinctly younger (<10 m.y.) than in the Sonoran Desert section (13-10 m.y.) (Zoback and others, 1981).

E.2 Geologic History of Southwest Utah

E.2.1 Pre-Cenozoic Geologic History

The post-Archean, pre-Cenozoic geologic history of southwest Utah (i.e., its history between 2,400 m.y. and 65 m.y. ago) is dominated by (1) marine sedimentation and a westward-facing, passive continental plate margin during Late Precambrian-Paleozoic time, and (2) intraplate, eastward-directed, compressive orogenic activity during the Mesozoic resulting from convergent plate tectonics farther west. Figure E-5, adapted from Stokes and Heylmun (1963) provides a useful framework for discussion.

Of fundamental importance is the concept of the Wasatch-Las Vegas Line (see Stokes and Heylmun, 1963; Stokes, 1976, 1979), also referred to as the Cordilleran Hingeline (e.g., Hill, 1976, p. 6)—the term we adopt here for simplicity. This feature marks a persistent zone of weakness in the North American plate, apparently inherited from late Precambrian continental rifting. It defines an axis of differential crustal movement that controlled late Precambrian to Mesozoic-age stratigraphy in the Cordilleran geosyncline, and also defined where eastward-moving thrust plates broke to the surface during Cretaceous-early Tertiary time (Stokes, 1976). Because of its important control on Mesozoic and Cenozoic tectonism, the Cordilleran Hingeline separates regions of contrasting geologic history.

E.2.1.1 <u>Precambrian-Paleozoic Geology</u>. The Transcontinental Arch, shown in figure E-5;s a major northeast-trending element of the North American Precambrian craton that formed an important positive feature in the early Paleozoic (see Hintze, 1973, p. 99). The oldest stratified rocks in the western United States--including the Vishnu Schist, which crops out in the Grand Canyon region and in the Beaver Dam Mountains west of St. George (Cook, 1960; Hintze, 1973, p. 10)--form part of the crystalline terrain of this belt.

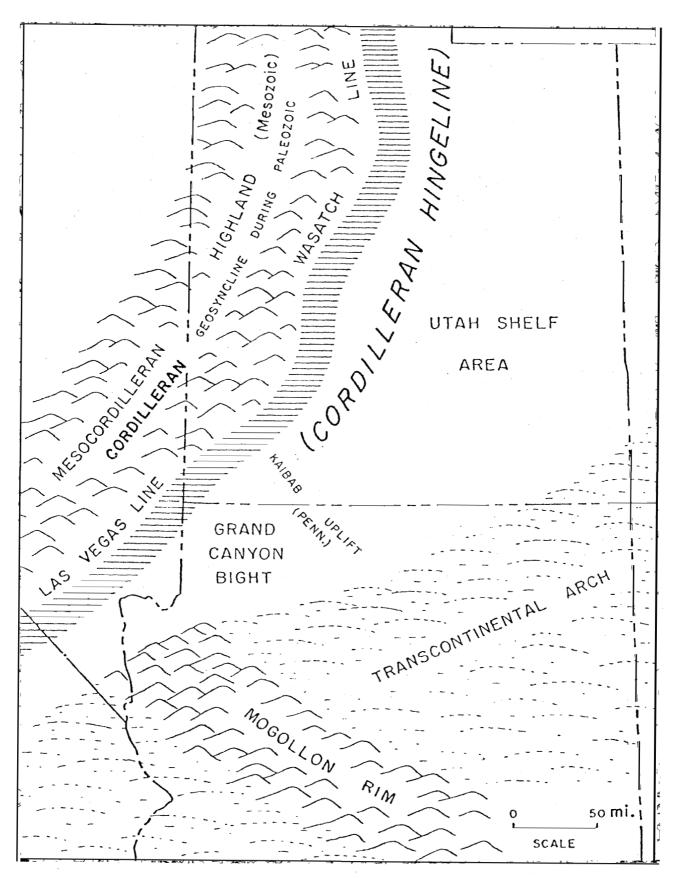


FIGURE E-5. Sketch map of some major pre-Cenozoic tectonic features affecting southwestern Utah (adapted from Stokes and Heylmun, 1963).

According to extensive radiometric dating and field studies by Silver and others (1977), these rocks represent marine strata, deposited in a major eugeosyncline 1,730-1,770 m.y. ago, that were subsequently metamorphosed, intruded by orogenic batholithic rocks 1,700-1,740 m.y. ago, and later intruded by anorogenic granites 1,410-1,465 m.y. ago.

In late Precambrian time, miogeoclinal deposition (i.e., sedimentation along a passively subsiding plate margin) began along the western margin of the North American craton. Deposition of a thick sequence (up to 7,600 m [25,000 ft]) of upper Precambrian and Lower Cambrian diamictites (conglomeratic mudstones), tholeitic basalts, and marine sedimentary strata has been interpreted by Stewart (1972) as resulting from rifting of the North American protocontinent (<850 m.y. ago). Creation of a new westward-facing continental margin marked the beginning of the Cordilleran geosyncline. Isopachs of upper Precambrian-Lower Cambrian strata indicate abrupt downwarping and probable continental separation roughly along the Cordilleran Hingeline (see Stewart, 1972; Stokes, 1979). Aeromagnetic data are consistent with such framentation of the Precambrian craton; an order-of-magnitude difference in the amplitude of magnetic anomalies in eastern Utah, as opposed to "quiet" magnetic data over the Basin and Range province, appears to be related to the regional distribution of buried Precambrian rocks (Zietz, 1980).

Figure E-6 (from Hintze, 1973), which summarizes the general stratigraphy of the Cedar City-St. George area, indicates the predominance of shallow-water marine sedimentation throughout the Paleozoic, reflecting deposition in the vicinity of the Cordilleran Hingeline--with facies associated with either the continental shelf or adjoining, mildly negative, platform (see Stokes, 1979, and Hintze, 1973, for details of paleogeography and paleo-plate tec-

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FIGURE E-6. Representative stratigraphic column for the Cedar City-St. George area (from Hintze, 1973).

tonics). For example, Lower Paleozoic rocks (Prospect Mountain quartzite, Pioche shale, and overlying carbonates) exposed in the Beaver Dam Mountains west of St. George, which are important exposures because of their isolation from correlative rocks by considerable distances, represent a transgressive assemblage along the miogeocline-craton hingeline (Hintze, 1963).

During Late Devonian and Mississippian time, the Antler orogeny produced compression, uplift, and eastward thrusting in the western part of the Cordilleran geosyncline, chiefly affecting the area of western and central Nevada; only mild folding occurred in western Utah and parts of the Colorado Plateau (Shawe and others, 1978; Stokes, 1979). During the Pennsylvanian period and continuing into the Permian, areas of the eastern Great Basin were affected by the development of northwest-trending downwarps and uplifts that accompanied the Ancestral Rocky Mountains orogeny, the effects of which were most intense to the east and southeast of the Utah region (Stokes, 1979; Hintze, 1973). The Pennsylvanian Kaibab uplift (fig. E-5)formed in northern Arizona and southern Utah at this time (Stokes and Heylmun, 1963).

E.2.1.2 Mesozoic History. The unconformity (or more correctly, the disconformity) shown in fig. E-6 at the base of the Triassic Moenkopi Formation reflects a major change from marine deposition, which persisted in Utah into Early Triassic time, to chiefly continental deposition by Late Triassic time. Profound change accompanied reversal of relief of the Cordilleran geosyncline into the Sevier orogenic belt during the Mesozoic era, with most pronounced orogenic effects in Utah during the latest Jurassic and Cretaceous periods (Armstrong, 1968; Hintze, 1973). Thus areas to the west of the Cordilleran Hingeline became positive while those to the east became negative, receiving sediments. In fig. E-5 the uplifted area is identified as

the Mesocordilleran Highland (also referred to as the Sevier Geanticline, e.g., Hintze, 1973). A companion uplifted area, the Mogollon Rim, appears to have originated in the middle Triassic at about the same time as the Mesocordilleran Highland; the Mogollan Rim became the southwestern margin of Triassic, Jurassic, and Cretaceous sedimentary basins—while a relatively low area between the Mogollan Rim and the Mesocordilleran Highland (the Grand Canyon Bight in fig. E-5)received thick accumulations of Upper Triassic, Jurassic, and Cretaceous sediments (Stokes and Heylmun, 1963).

Continental deposition in southwest Utah in Late Triassic time involved both eastward-spreading detritus from orogenic highlands in western Nevada and California, as well as detritus from uplifts to the east and southwest (Shawe and others, 1978; Hintze, 1973, p. 57). Aridity in the lee of the Mesocordilleran Highland, combined with a surrounding of the Colorado Plateau region by uplifts, resulted in deposition of Late Triassic fluvial-lacustrine red beds (Stokes, 1979).

The dominance of continued continental deposition in southwest Utah through most of the early Jurassic is reflected by the Navajo sandstone (fig. E-6) with its large-scale cross-bedding of sand-dune origin. Middle and Upper Jurassic marine strata of the Carmel Formation are interpreted to be related to marine invasion by a southerly extending tongue of a shallow epicontinental sea that was connected to a Pacific Ocean through Canada or Alaska; the extensive Upper Jurassic Morrison Formation, of continental origin, did not extend into southwest Utah (see Hintze, 1973, p. 59-67).

During the Cretaceous, the last epicontinental sea in Utah, which had spread northwestward from the Texas coastal plain, was bounded on the west by the Mesocordilleran Highland. Cretaceous strata in southwest Utah reflect

typical marginal-marine, coastal-plain conditions immediately east of the Sevier orogenic belt. The thick Cretaceous stratigraphic section (fig. E-6), for example, includes marine sandstones and shales, coal-bearing strata, and occasional intercalations of locally coarse deposits derived from the Sevier orogenic highlands to the west.

The Mesocordilleran Highland depicted in fig. E-5 resulted from strong folding and large-scale eastward thrusting during the Sevier orogeny, probably involving foreshortening of several tens of kilometers (Armstrong, 1968). According to Armstrong (1968) the Sevier orogeny occurred throughout the Cretaceous period. Sevier thrusting in southwestern Utah had probably terminated by latest Cretaceous time (Shawe and others, 1978; Armstrong, 1968; Stokes and Heylmun, 1963), although elsewhere in the so-called Idaho-Wyoming-Utah foreland thrust belt there is overlap with compressional tectonics generally ascribed to the Laramide orogeny during latest Cretaceous-Paleocene-Eocene time (Burchfiel, 1980). Effects of the Laramide orogeny extended well inland (eastward) of the fold and thrust belt.

E.2.2 Cenozoic Geologic History

Chief elements of the Cenozoic history of southwestern Utah, i.e., its history from 65 m.y. ago to the present, include (e.g., Shawe and others, 1978, p.3): "extensive volcanism, plutonism, and continental sedimentation in local basins, and in later stages regional uplift, block faulting, and attendant continental sedimentation." Another important element is the structural differentiation between the Colorado Plateau and Basin and Range provinces. Fortuitously, some excellent detailed summaries of these various aspects of the Cenozoic geology of southwestern Utah have been recently published. These notably include: (1) a comprehensive summary of the stratigraphic and structural framework of southwestern Utah by Rowley and others (1978); (2) a guidebook edited by Shawe and Rowley (1978) that focuses on economic mineral deposits in southwestern Utah, but which includes an informative geologic overview; and (3) various publications by R.E. Anderson and colleagues that focus on the Quaternary tectonics of southwestern Utah, particularly as relevant to problems of active faulting and earthquake hazards (e.g., Anderson, 1981,1979, 1978; Anderson and Mehnert, 1979; Anderson and Bucknam, 1979; Zoback and others, 1981).

Figure E-7 from Zoback and others (1981), provides an up-to-date synthesis of Cenozoic plate tectonic elements and magmatic patterns in the western U.S. that is useful for placing the history of southwestern Utah in a regional perspective. One introductory comment is required. During the Laramide orogeny (ca. 80-40 m.y. ago), there was a prominent absence of calc-alkaline volcanism throughout most of the western U.S. landward of a west-coast subduction zone; this absence of magmatism, together with the great inland extent of compressional Laramide tectonism, have been attributed to long-continued, very low-angle subduction along the western plate boundary (see Zoback and others, 1981).

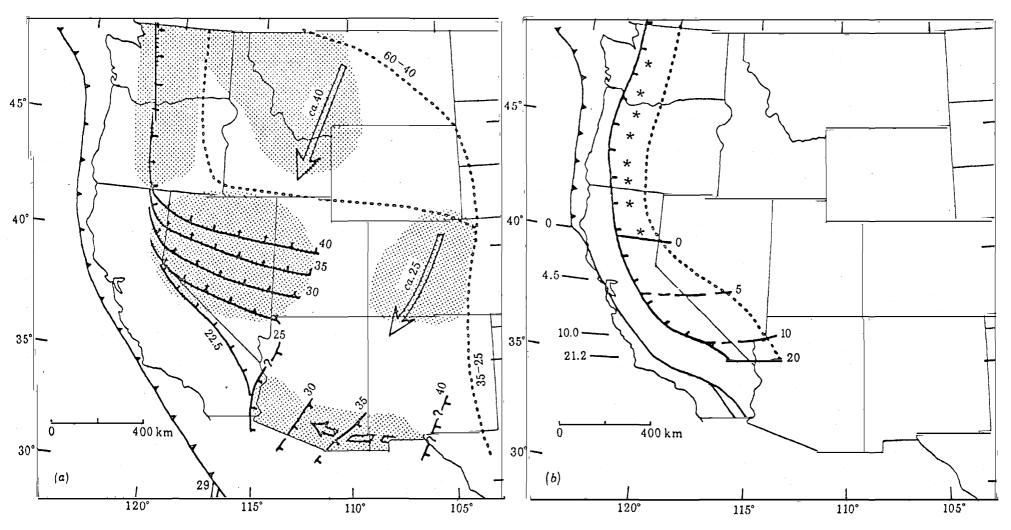


FIGURE E-7. Synthesis of Cenozoic tectonic patterns in western U.S. (from Zoback and others (1981) involving (a) migration of calc-alkaline volcanism across the Laramide magmatic gap, and (b) northward migration of a triple junction along the western North American plate. Hachured lines indicate westernmost extent of calc-alkalic volcanism; small circles, the easternmost extent. Numbers identify time stages (in megayears).

Between about 40 and 20 m.y. ago after cessation of the Laramide orogeny-coinciding with changes in world-wide plate motions (Coney, 1978)--calc-alkaline volcanism systematically swept west to southwestward across the Laramide magmatic gap (fig. E-7a) Zoback and others (1981) present evidence for early extensional tectonism concurrent with the cala-alkaline volcanism in an "intra-arc" or back-arc" setting.

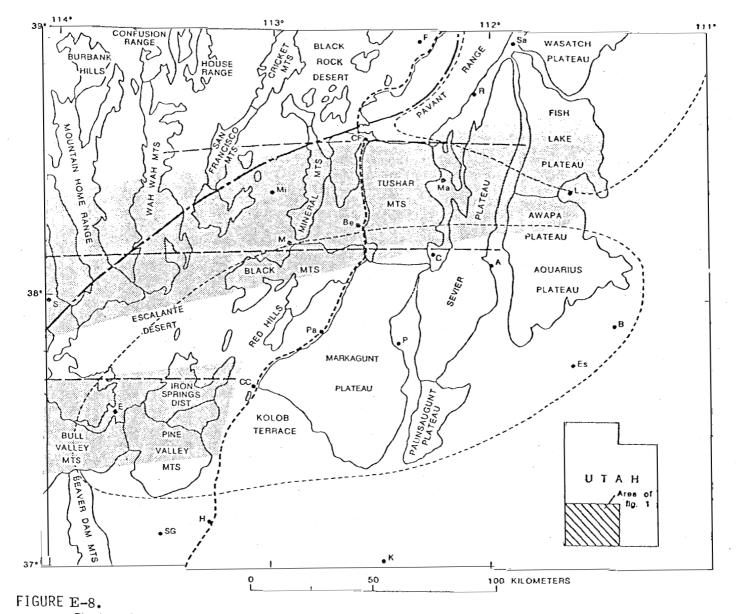
Between about 20 and 30 m.y. ago, a ridge-trench collision along the western North American plate caused major tectonic changes in the western U.S., including formation of the San Andreas transform boundary (Atwater, 1970).

Northward migration of a triple junction along the San Andreas boundary after about 20 m.y. ago (fig.E-7b)produced the following important changes in the Basin and Range region (Zoback and others, 1981; Eaton, 1979): (1) replacement of calc-alkaline volcanism by basaltic volcanism in the southern Basin and Range region by about 20 m.y. ago; (2) initiation of significant extension in the southern Basin and Range region by about 20 m.y. ago, with intensive NW-trending basin-range block faulting in that region 13-10 m.y. ago; (3) widespread initiation of basaltic volcanism throughout the northern Basin and Range region at about 17 m.y. ago; and (4) development of modern NNE-trending basin-range block faulting in the northern Basin and Range region after 10 m.y. ago.

With the above-outlined scenario in mind, much of the complex Cenozoic geology of southwestern Utah can be related to three general stages: (1) early Tertiary (ca. 65-35 m.y. ago) sedimentation in broad basins formed by local warping east of the Sevier orogenic highlands during Laramide time, (2) middle Tertiary (ca. 35-20 m.y. ago) calc-alkaline magmatism along ENE-trending igneous belts, and (3) late Cenozoic (<20 m.y. ago) bimodal basalt-rhyolite volcanism and extensional tectonism.

Figure E-8, from Rowley and others (1979), provides a simplified overview of some elements of the Cenozoic geology of southwestern Utah. More than 300 m of fluvial-lacustrine strata accumulated in a broad easterly-trending depositional basin in southwest Utah during latest Cretaceous(?) to perhaps middle Oligocene time (Rowley and others, 1979). The Paleocene-Eocene Cedar Breaks Formation (fig. E-6) exposed near Cedar City was deposited in this basin, which did not extend as far south as St. George.

In middle Tertiary time, the southward migration of calc-alkaline magmatism that swept through Nevada and western Utah (fig.E-7a)produced extensive andesitic and rhyolitic volcanism (accompanied by granitoid intrusions) in roughly east-trending belts, each successively younger to the south (Stewart and others, 1977). Calc-alkaline igneous rocks in southwest Utah, which nearly all range in age from about 35 m.y. to 19 m.y. (Rowley and others, 1979; see also fig.E-6) define one of these belts--a broad zone of Oligocene-Miocene magmatism between roughly 37°N and 39°N latitude (see Shawe and others, 1977). Within this broad zone in southwest Utah, two distinct ENE-trending belts of mineralization are recognized (see fig.E-8): the Pioche-Marysvale belt and the Delamar-Iron Springs belt (Shawe and others, 1978).



Physiographic features and major structural features of southwestern Utah. Basin and Range — Colorado Plateaus boundary shown by heavy short-dashed line; Pavant Range, Tushar Mountains, and named plateaus belong to the High Plateaus subprovince. Central part of Pioche-Marysvale igneous belt on north, and of Delamar-Iron Springs igneous belt on south, patterned. Dot-dash line, east edge of major Sevier thrust faults, modified from Crosby (1973). Heavy long-dashed lines, axes of lineaments, consisting from north to south of the Black Rock, Blue Ribbon, and Timpahute lineaments. Fine dashed line on the north, approximate minimum boundaries of depositional basins of the North Horn to Crazy Hollow Formations; to the south, the Claron Formation and related rocks. Tushar highland is between them. A, Antimony; B, Boulder; Be, Beaver; C, Circleville; CC, Cedar City; CF, Cove Fort; E, Enterprise; Es, Escalante; F, Fillmore; H, Hurricane; K, Kanab; L, Loa; M, Minersville; Ma, Marysvale; Mi, Milford; P, Panguitch; Pa, Parowan; R, Richfield; S, Stateline mining district; Sa,

Salina; SG, St. George. (From Rowley and others, 1979)

Ash-flow tuff sheets originating from volcanic centers in the Pioche-Marysvale and Delamar-Iron Springs igneous belts were widespread in southwest Utah during middle Tertiary time. Indeed, their distribution provides critical evidence for timing the structural differentiation between the Colorado Plateau and Basin and Range provinces. Regional dating of ash-flow tuffs suggests that the provinces became structurally separate some time after 29 m.y. ago, and that vertical movements on boundary faults along the west side of the Colorado Plateau began about 26-18 m.y. ago in late Oligocene-early Miocene time (Rowley and others, 1978, 1979; Best and Hamblin, 1978).

Movement on the Hurricane fault may not have been so early. Although Rowley and others (1979) associated the Hurricane fault with the Basin and Range-Colorado Plateau boundary (see fig. E-8) Andreson and Mehnert (1979) argue otherwise and suggest that major movement on the Hurricane fault was probably post-Miocene.

During late Cenozoic time (<20 m.y. ago), basaltic volcanism, regional uplift, and extensional tectonism dominate the geology of southwestern Utah. Some important aspects of basaltic volcanism in the general region surrounding the Cedar City-St. George area include (Best and Hamblin, 1978; Rowley and others, 1979): (1) a hiatus of several million years between the end of mid-Tertiary calc-alkaline volcanism and the inception of basaltic volcanism; (2) the occurrence of relatively isolated, low-volume basaltic fields, particularly in the boundary region between the Basin and Range and Colorado Plateau provinces, without large central volcanic complexes; (3) representative ages ranging from about 13 m.y. to 0.5 m.y. for basaltic volcanic fields in the eastern Basin and Range province to be tholeiitic, whereas basalts on the Colorado Plateau are more alkalic and undersaturated--possibly

reflecting differing regimes of upper-mantle magma generation. Basaltic fields in the immediate vicinity of St. George are discussed by Hambin and others (1981) and by Hamblin (1970b).

The history of extensional faulting and basin-fill sedimentation in southwestern Utah is clearly complex. There seems to be agreement that the present-day topography of the northern Basin and Range province did not develop before 10 m.y. ago (Zoback and others, 1981; Stewart, 1978). However, pre-basin-range extension (now recognized by faulted and tilted strata exposed in uplifted range blocks) was under way locally in the Basin and Range province by at least 30 m.y. ago (Zoback and others, 1981). A major episode of pre-basin-range, thin-skinned extensional normal faulting affected rocks now exposed in the Beaver Dam and Bull Valley Mountains west of St. George between about 15 and 11 m.y. ago (Anderson, 1981).

In the High Plateaus subprovince of the Colorado Plateau, major block faulting was under way by 20 m.y. ago, and from about 20 m.y. ago to 7 m.y. ago the High Plateaus were characterized by both faulting and broad warping (Rowley and others, 1979). Basin-fill from erosion of fault blocks has been dated to be at least as old as 14 m.y. in the High Plateaus, and as old as 10 m.y. in the northern Basin and Range province—although the oldest ages of basin-fill might be about 20 m.y. and 15 m.y., respectively, in the two regions (Rowley and others, 1979; Zoback and others, 1981). Pliocene basin-fill in the Cedar City-St. George area is identified in the stratigraphic column of fig. E-6 as the Muddy Creek Formation.

Relative vertical motions between the eastern Basin and Range province and the Colorado Plateau introduce additional complexity. Observed uplift and northeastward tilting of the western margin of the Colorado Plateau may relate to a broad upwarping that involved the entire Basin and Range province during

late Cenozoic time; if so, then relative down-dropping of the eastern Basin and Range province must have accompanied collapse of the upwarp by fault fragmentation (see Best and Hamblin, 1978). In any case, both provinces are now rising, but the western Colorado Plateau is rising at a significantly faster rate (Hamblin and others, 1981).

E.3 STRUCTURAL GEOLOGY

E.3.1 General Structure of Study Area

Figure E-9gives an overview of the general structure of the Cedar City-St. George area, which is dominated by N- to NNE-trending faults that deform a complex boundary region between the eastern Great Basin and the higher-standing western Colorado Plateau to the east. East-west structural sections in figure E-10 (from Best and Hamblin, 1978) illustrate differences across the province boundary in northwestern Arizona (section B-B') compared to southwestern Utah (section A-A'), as one passes from the Grand Canyon section to the High Plateaus section of the Colorado Plateau. Section B-B', 50-60 km south of the Utah-Arizona border, shows the relatively simple structure of the Grand Staircase-a series of northeastward-tilted fault blocks bounded by major down-to-the-west normal faults. Northward into Utah, section A-A' (about 30 km north of Cedar City) shows more "ragged" faulting across the province boundary involving horsts, grabens, and ramp structures with scissor-like displacement (Best and Hamblin, 1978).

In northwestern Arizona, the Grand Wash fault forms the boundary between the Basin and Range and Colorado Plateau provinces (section B-B', fig.E-10). To the north of the Utah border, the province boundary is less well defined.

Cedar Pocket Canyon—
Anderson and Mehnert (1979) interpret that it follows the Gunlock-Veyo fault (the northern continuation of the Grand Wash fault), then passes north of the Pine Valley Mountains to join the Hurricane-Parowan/monocline-fault system near Cedar City (see fig.E-9). Thus en echelon stepping of the province boundary Cedar Pocket Canyon—

Paragonah from the Gunlock-Veyo fault to the Hurricane-Parowan/ structure at roughly lat 37.5°N forms a structural block encompassing the Pine Valley Mountains and the St. George Basin. This structural block (fig.E-9) is mildly warped, cut

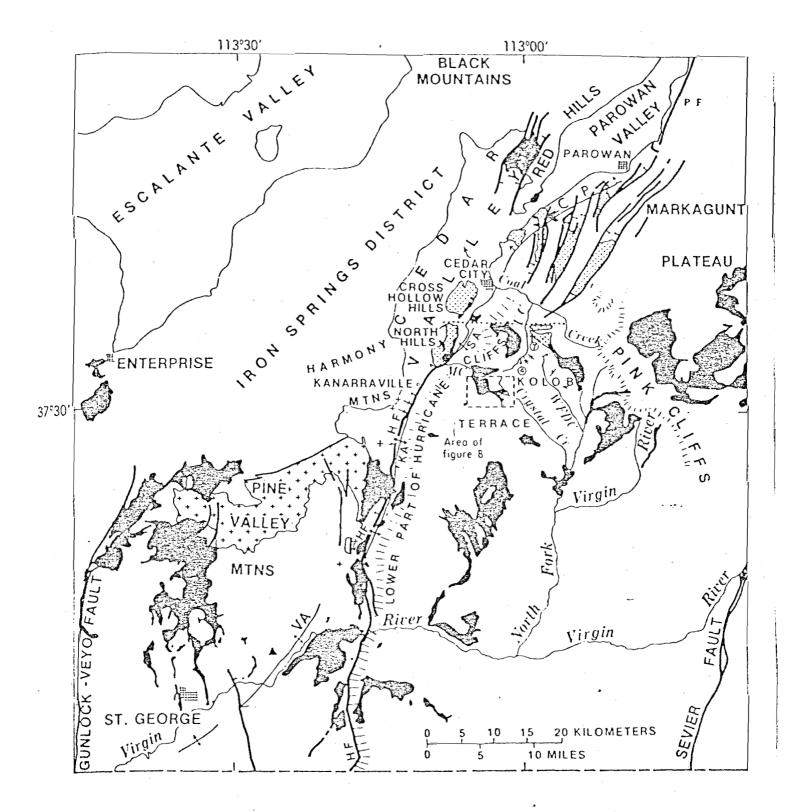


FIGURE E-9. Geologic sketch map showing structural overview of Cedar City-St. George study area (from Anderson and Mehnert, 1979). HF indicates Hurricane fault; PF, Paragonah fault; VA, Virgin anticline; KA, Kanarra anticline; SA, Shurtz Creek anticline; CP, Cedar City-Parowan monocline. Distribution of upper Cenozoic basaltic lava flows shown in black.

BASIN & RANGE

KILOMETERS

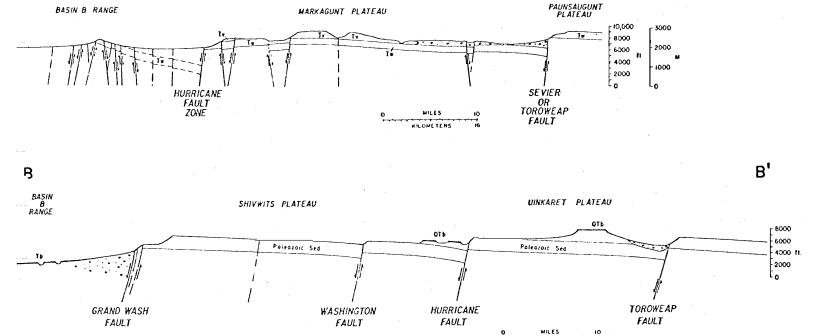


FIGURE E-10. East-west structural sections (from Best and Hamblin, 1978) illustrating changing structural style across the Basin and Range-Colorado Plateau province boundary along strike in southwestern Utah (A-A') and northwestern Arizona (B-B'). Section A-A' located 30 km north of Cedar City, Utah; B-B', 50-60 km south of Utah-Arizona border. QTb, Tb, basaltic flows; Tv undifferentiated Oligocene andesitic volcanic fields; Tw. Eocene Wasatch Formation.

by faults of relatively small dispacement, and is marked by the absence of any prominent basin-range graben structure--suggesting affinity with the Colorado Plateau province (Anderson and Mehnert, 1979).

The Hurricane fault is the most prominent structure in the study area. Anderson and Mehnert (1979) provide an extensive review of geological data relevant to the Utah segment of the fault. Their interpretation, however, that the total normal displacement on the fault in Utah is only 600-850 m has been modified (R.E. Anderson, personal communication to W.J. Arabasz, January 1982) after access to confidential oil-company data that require the total displacement to be at least 2,000 m. Given current displacement rates of the order of 300-500 m/m.y. along the northern Hurricane fault (Appendix F.3.4), such total displacement can still be accommodated chiefly in Quaternary and Pliocene time.

Eastward-directed thrust faults and related folds of Cretaceous age associated with the Sevier orogeny have affected pre-Cenozoic rocks in southwestern Utah. The eastern limit of major thrusting (see Rowley and others, 1979) passes near the northwestern extremity of fig. E-9. The leading edge of the Sevier disturbed belt, however, directly affected the study area where NE- or NNE-striking overturned and open folds, some locally thrust faulted, were formed in a belt parallel to the front of major thrusting (Rowley and others, 1979). These include the Virgin and Kanarra anticlines (fig.E-9) Anderson (1980) and Anderson and Mehnert (1979) review the influence of pre-Quaternary structure on the NE-trending fault fabric of the study area.

E.3.2 Structure in the Cedar City Area

SCS dams Greens Lake Numbers 2, 3, and 5, located south of Cedar City, lie within a complicated structural region bridging the northernmost, well defined segment of the Hurricane fault, 15 km south-southwest of Cedar City, with the Parawon-Paragonah fault, about 35 km northeast of Cedar City (see fig. E-9) (Threet, 1963; Anderson and Bucknam, 1979; Anderson and Mehnert, 1979; Averitt, 1964; Averitt and Threet, 1973). Threet (1963) describes the situation:

"....the Neogene Hurricane fault cannot be extended properly along the plateau margin between Cedar City and Paragonah; instead, a Neogene monocline controls the plateau margin geomorphology in that segment. The northwestward-throwing monocline intersects the east flank of the Laramide Kanarra fold, a few miles north of Cedar City, in a remarkably well documented case of oblique unfolding...." (Threet, 1963, p. 110).

Figure E-11, combining illustrations from Threet (1963) and Anderson and Bucknam (1979) shows the combination fault-monocline structure.

A detailed geologic map (scale 1:24,000) of the area of interest surrounding Cedar City has been published by Averitt and Threet (1973). As part of this study, detailed photointerpretation and field mapping and exploration were conducted in the immediate Cedar City area. The results of the photointerpretation, part of a larger study of the Hurricane fault in southwestern Utah conducted as part of this study, is discussed in detail in ChapterIV of the report and shown on figure IV-2 (scale 1:62,500). The results of the field mapping and exploration in the Cedar City area is discussed in detail in ChapterVIII-A and B of the report and shown on figure VIII-1 (scale 1:24,000).

In summary, the results of detailed photointerpretation and field mapping and exploration in the immediate Cedar City area indicate the structure described above is further complicated along the Hurricane Cliffs below Lone Tree Mountain by large scale, complex landsliding, and possibly some basalt flows, which underlie the escarpment in this area. As shown on figure VI-1, recent traces of the Hurricane fault apparently cut the landslide deposits and lie somewhat upslope from the present topographic break-in-slope at the base of the Hurricane Cliffs south of Cedar City. These traces apparently exit the escarpment in the vicinity of Greens Lake Dams Numbers 2 and 3, but cannot be traced on air photos to the north across the recent alluvial fan deposits that underlie the dams.

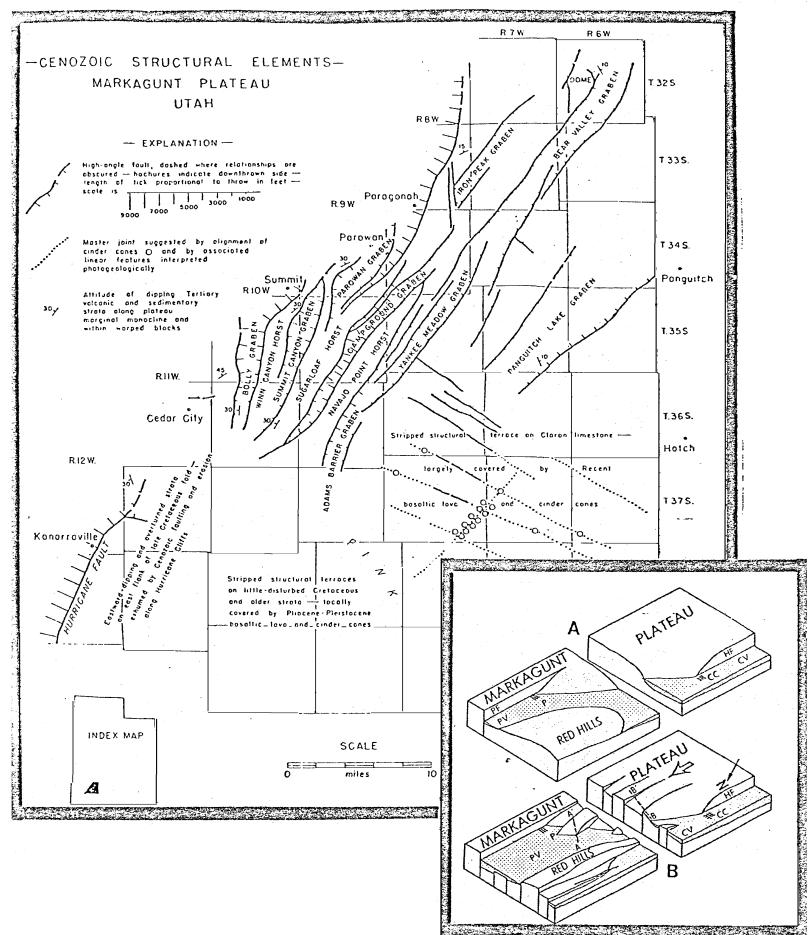


FIGURE E-11. Map of Cenozoic structural elements of the Markagunt Plateau near Cedar City (from Threet, 1963). Inset (from Anderson and Bucknam, 1979) illustrates NE-trending monoclinal flexure that serves as a structural bridge between the Paragonah fault (PF) and the Hurricane fault (HF). CC = Cedar City; P, Parowan.

Southwest of Cedar City, fault traces cut Pleistocene alluvium and locally overlying basalt, and may cut late Pleistocene colluvium along the east side of Cross Hollow Hills (see figure VIII-1). These fault traces and associated air photo lineaments can be mapped discontinuously to the southwest along the east side of North Hills to a projected intersection with the main Hurricane fault near Murie Creek (see figure IV-2). This relationship suggests that the alluvial valley(s) southwest of Cedar City between the Hurricane Cliffs and Cross Hollow - North Hills may be a late Pleistocene graben, herein called the Shurtz Creek Graben (see figure IV-2), similar to those located northeast of Cedar City along the southern part of the Paragonah fault (figure E-11).

E.3.3 Structure in the Hurricane - Frog Hollow Area

The SCS Frog Hollow Dam is located southeast of Hurricane in the relatively flat upland area of the Grand Canyon Section subprovince of the Colorado Plateau Provine, east of the Hurricane Cliffs. The structure of interest in this area is the Hurricane fault which passes approximately 3 km west of the dam. Localized Quaternary basaltic volcanism has occurred west and southwest of the dam.

As part of this study, detailed photointerpretation and field mapping were conducted in the immediate Frog Hollow area. The results of these investigations are discussed in detail in Chapter VIII-F of the report and shown in Figure VIII-3 (scale 1:24,000).

E.3.4 Structure in the St. George Area

SCS dams Gypsum Wash, Warner Draw, and Stucki are located southeast of St. George along the southeast side of the alluviated Washington Fields adjoining the western edge of Warner Ridge. The structure of primary interest in this area is the Washington fault, a NNW-trending, down-to-the-west normal fault that crosses the eastern part of the St. George basin (fig. E-12) passing within 0 to 1225 m of one or more of the dams. The Washington fault extends approximately 68 km from just north of the town of Washington, southward with increasing displacement into northwest Arizona (Dobbin, 1939; Cook, 1960; Cook and Hardman, 1967; Hamblin, 1970a, b) Best and Hamblin, 1978; Wilson and others, 1969). Vertical displacement on the Washington fault has been estimated to range from several hundred feet (a few hundred meters) near the Virgin River, where the fault breaches the northeast-trending pre-Quaternary Virgin anticline, to 2,500 feet (760 m) at the Arizona state line (Dobbin, 1939; Cook and Hardman, 1967).

As part of this study, detailed photointerpretation and field mapping and exploration were conducted in the immediate area of the SCS Dams. The results of these investigations are discussed in detail in Chapter VIII C, D, and E of the report and shown on Figure VIII-2 (scale 1:24,000). In summary, the results of the detailed photo-interpretation and field mapping and exploration in the Gypsum Wash - Warner Draw - Stucki area indicate that the Washington fault displays probable early Holocene, normal, east-side-up displacement in the vicinity of Gypsum Wash Dam.

E.3.5 Structure in the Ivins Area

The Ivins area, which includes the Ivins Bench Diversions, lies in the St. George Basin within a structural block bounded on the west by the northern extension of the Grand Wash fault (designated the Cedar Pocket Canyon - Gunlock fault in fig. E-13 and the Gunlock-Veyo fault in fig. E-9) (Cook, 1960; Dobbin, 1939; Anderson and Mehnert, 1979; Best and Hamblin, 1978) is part of a major fault zone which is mapped as extending to the vicinity of Grapevine Wash in northwestern Arizona (Wilson and others, 1969).

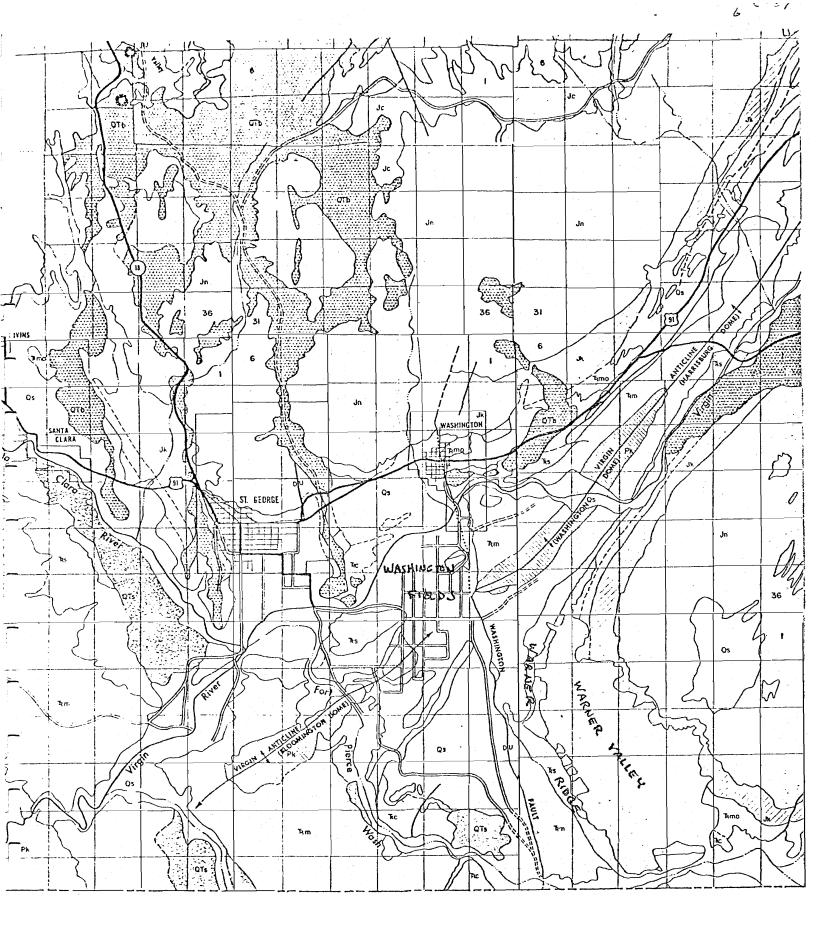


FIGURE E-12. Geological map of the area of the Washington fault east of St. George, Utah (from Cook, 1960). Section lines define scale.

The total mapped length of the Grand Wash - Cedar Pocket Canyon - Gunlock fault zone is approximately 159 km. The structural block between the Hurricane and Grand Wash faults in the St. George area has been uplifted at least 64 m/m.y. in late Cenozoic time (Hamblin and others, 1981). In addition to the Grand Wash and Hurricane faults, other faults of interest in the Ivins area include the Washington fault, located approximately 16 km east-southeast and a 20 km long zone of suspected Quaternary faulting (see Appendix F, figure F-7) traced by Anderson and Miller (1979) on the basis of unpublished mapping made available to them by W. K. Hamblin, located approximately 10 km east-northeast.

As part of this study, detailed photointerpretation and field mapping and exploration were conducted in the immediate Ivins area. The results of these investigations are discussed in detail in Chapter VIII G of the report and shown on figure VIII-4 (scale 1:24,000).

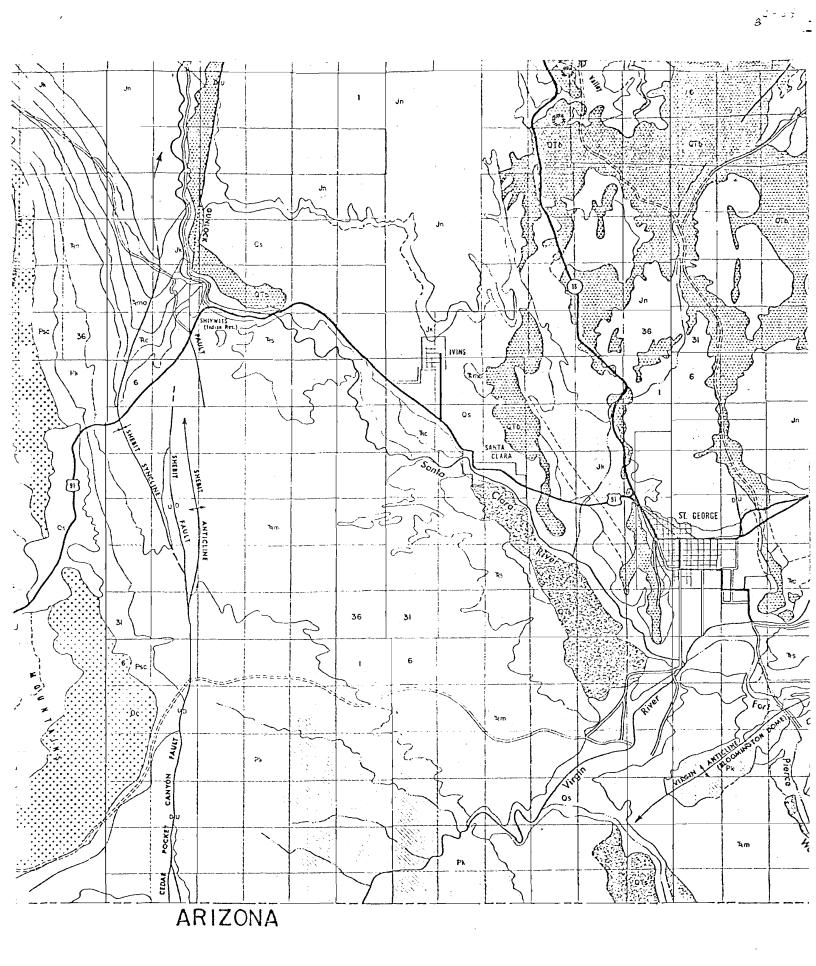


FIGURE E-13. Geological map of the Ivins area northwest of St. George, Utah (from Cook, 1960). Section lines define scale.

APPENDIX F

HISTORIC SEISMICITY AND LATE QUATERNARY FAULT ACTIVITY

APPENDIX F

HISTORIC SEISMICITY AND LATE QUATERNARY FAULT ACTIVITY

- F.1 Regional Seismicity
 - F.1.1 General Features
 - F.1.2 Historical Earthquake Record for Southwestern Utah
 - F.1.2.1 Largest Historical Earthquakes in Southwestern Utah and Vicinity
 - F.1.2.2 Swarm Seismicity in Southwestern Utah
 - F.1.3 Detailed Instrumental Seismicity in Southwestern Utah
- F.2 Seismotectonics and Earthquake Hazard Evaluation
 - F.2.1 Implications of Current Geological Studies
 - F. 2.1.1 Identification of Active Faults
 - F.2.1.2 Late Quaternary Versus Holocene Faulting
 - F.2.2 Implications of Current Seismological Studies
 - F. 2. 2.1 Listric Faulting and Problematic Correlation of Seismicity with Geologic Structure
 - F.2.2.2 Implications of Swarm Occurrence
 - F.2.2.3 Focal Mechanisms and Stress Orientation
 - F.2.2.4 Measurements of Ground Response in the Cedar City Area
- F.3 Earthquake Recurrence and Fault Activity Rates
 - F.3.1 General Statement
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FIGURES

No.	Title
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F-2	Schematic Outline of Seismicity Domains Related to Compilations of Regional Seismicity in Vicinity of Various Study Area
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F -5	Epicenter Map of the Utah Region, July 1962 to June 1978
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F-12	Earthquake Recurrence Data for the Utah Region and Subregions
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F. HISTORIC SEISMICITY AND LATE QUATERNARY FAULT ACTIVITY

F.1 REGIONAL SEISMICITY

F.1.1 General Features

The Cedar City-St. George area lies directly within the Intermountain seismic belt (ISB) (fig. F-1), an extensive linear zone of current intraplate deformation in the western United States that is characterized by late Quaternary normal faulting, diffuse shallow seismicity, and episodic scarpforming earthquakes ($6\frac{1}{2} \le M \le 7\frac{1}{2}$) (Arabasz and Smith, 1981; Smith, 1978; Smith and Sbar, 1974).

As summarized by Arabasz and Smith (1981), the notable features of the ISB are: (1) its length (>1300 km) and breadth (100-200 km); (2) its segmentation into several sectors with divergent trends; (3) the diffuseness of seismicity, with focal depths almost exclusively shallower than 15-20 km, and weak correlation with major active faults (based on dense-network monitoring with both portable and fixed microearthquake networks); (4) a general predominance of normal faulting reflecting an extensional stress regime, but with spatially rapid changes in stress orientation in some areas; (5) the common occurrence of swarm sequences at various localities throughout the ISB; (6) relatively low rates of crustal strain (∿10⁻⁸ mm/mm/yr or less) compared to those at active plate boundaries; (7) moderate background seismic flux, which, for comparison is lower by a factor of 4-6 than that in the California-Nevada seismic zone (Algermissen, 1969); (8) apparent stresses that are within the range of values (0.01-100 bar) computed similarly for intraplate earthquakes elsewhere (Doser, 1980); (9) relatively long (>1000 yr) recurrence intervals for surface faulting (excepting parts of the Wasatch fault) (Swan et al., 1980; Wallace, 1980);

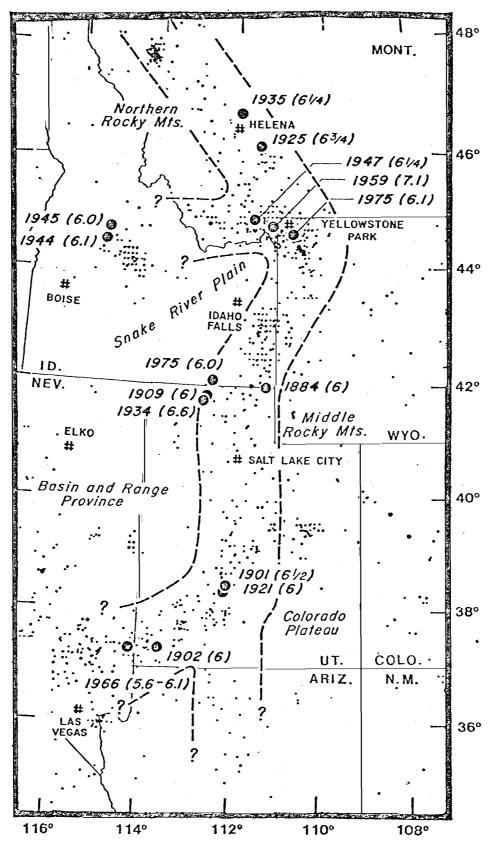


FIGURE F-1.--Map of Intermountain seismic belt (from Arabasz and Smith, 1981). Epicenters of historical mainshocks (M 6.0) shown as large circles, NOAA epicenters through 1974 as smaller circles. Schematic dashed outline of seismic belt based on various seismotectonic studies as well as depicted seismicity. Magnitudes (in parentheses) are estimated or measured values of $\rm M_{\tilde{L}}$, except values for 1925-1959 (attributed to Pasadena), which are either $\rm m_{\tilde{b}}$ or $\rm M_{\tilde{s}}$. Epicenter of 1966 shock (M_L 5.6, $\rm m_{\tilde{b}}$ 6.1) added for reference.

and (10) a historical paucity of large (M>7.0) surface-faulting earth-quakes--despite abundant late Quaternary and Holocene fault scarps throughout the ISB (e.g., Bucknam et al., 1980; Swan et al., 1980; Witkind, 1975).

In southern Utah at about lat 38°N, there is an apparent southward bifurcation of the ISB (fig.F-1). The most active belt of seismicity extends southwestward into southern Nevada, but there is a weak continuation of seismicity southward into northern Arizona (fig.F-1). Smith and Sbar (1974) originally considered that the main ISB continued southward into Arizona, with the ENE-trending zone that passes through southern Nevada and southwest Utah as distinct. However, Smith (1978) more recently has identified the weak seismicity continuing into Arizona as part of a zone surrounding the entire Colorado Plateau, and he traces the main ISB into southern Nevada.

The branch of the ISB diverging southwestward into southern Nevada-herein referred to as the SW Utah-S. Nevada seismic zone, is without question far more seismically active than the zone continuing southward into
Arizona. Earlier (Section C.1.3.5) we discussed how the SW Utah-S. Nevada
seismic zone correlates with a major structural corridor (see also Anderson,
1978, p. 3)--perhaps reflecting a profound upper-mantle discontinuity (Eaton
and others, 1978).

One important observation that should be emphasized with regard to figure F-1 is the absence of any historical earthquake in the southwest Utah region larger than about magnitude 6½. Here, as throughout much of the ISB, the historical paucity of scarp-forming earthquakes contrasts with abundant geologic evidence for late Quaternary faulting.

F.1.2. Historical Earthquake Record for SW Utah

To evaluate historical and instrumental seismicity in the general vicinity of the Cedar City-St. George study area, we have placed fundamental reliance on earthquake data from the University of Utah. The University of Utah Seismograph Stations has compiled the most authoritative information on the seismicity of the Utah region dating from 1850, and also carries out the most comprehensive instrumental monitoring of current seismicity in the Intermountain area (Arabasz and others, 1979, 1980; Richins and others, 1981).

The following systematic approach was taken to document seismicity in the area of interest (see figure F-2):

- 1. Compilation and plotting of <u>regional</u> seismicity, including all earthquakes of magnitude 3.0 or greater within 200 km of the study area. (For convenience, the 200-km radial distance was measured from a point half-way between Cedar City and St. George at lat 37°25'N, long 113°15'W.) These are tabulated in Appendix H.
- Listing of seismicity (M>3.0) within 150-km radial distance of four specific points, representative of the various dam sites (see Appendix I for partial listing):

Cedar City: 37°39.00'N, 113°04.30'W

Frog Hollow: 37°07.00'N, 113°15.30'W

St. George: 37°03.30'N, 113°29.00'W

Ivins: 37°10.30'N, 113°40.00'W

3. Compilation and plotting of the <u>entire</u> University of Utah earthquake catalog for the period: 1850-December 31, 1981 within a rectangular

area encompassing St. George and Cedar City (lat 36.75°N-38.00°N, long 112.00°W-114.25°W) for detailed evaluation of <u>local</u> seismicity (see Appendix J).

The catalog of documented earthquakes in the Utah region, as elsewhere in the western United States is a mixed one variously relying upon early reports and newspaper accounts, and later upon seismographic recordings during several stages of evolving coverage. The documentation of earthquakes pre-dating July 1962—the beginning of the University of Utah's instrumental catalog—combines both historical (i.e., based on felt reports) and early instrumental seismicity—determined for example by the U.S. Coast and Geodetic Survey or by the Seismological Laboratory in Pasadena. Arabasz and McKee (1979) outline with extensive annotations the 1850–1962 catalog for the Utah region, whose compilation involved the careful checking and correlation of numerous sources. The term "Utah region" as used by the University of Utah, signifies the rectangular area extending from lattitude 36°45'N to 42°30'N, and from longitude 108°45'W to 114°45'W—which defines the areal bounds of the University of Utah earthquake catalogs, unless specified otherwise.

The University of Utah master catalog for the Utah region comprises three parts (e.g., Arabasz and others, 1979, 1980): (1) the 1850-June 1962 historical catalog; (2) a catalog for the period July 1962-September 1974, consisting of systematically revised, instrumental earthquake locations and magnitudes; and (3) a catalog of instrumental seismicity since October 1974 based upon data from an extensive network of telemetered seismic stations. Data from the various time blocks have differing levels of uncertainty.

It should be noted that local magnitudes (M_L) for all earthquakes located in the Utah region since July 1, 1962, have been systematically estimated by either direct or indirect relation to Wood-Anderson-type torsion seismographs at three (currently two) widely spaced sites within Utah. The importance here is that a uniform earthquake catalog is essential for meaningful estimates of earthquake recurrence.

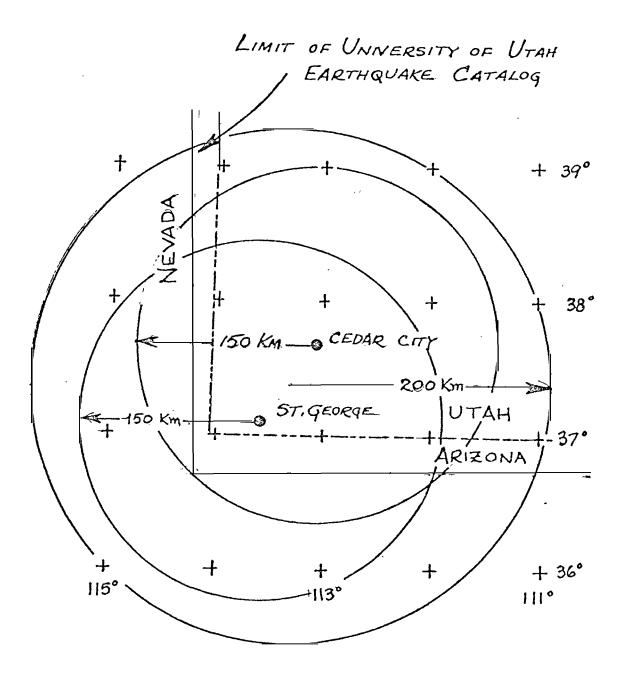


FIGURE F-2.--Schematic outline of seismicity domains related to compilations of regional seismicity in vicinity of various study areas.

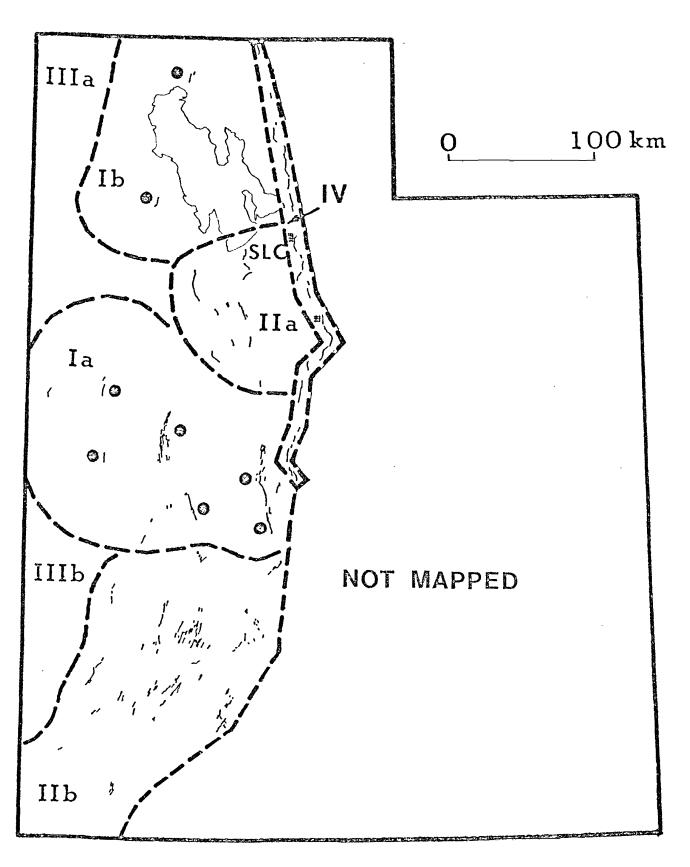


FIGURE F-8.--Map from Bucknam and others (1980) showing late Quaternary fault scarps in unconsolidated sediments (fine lines), Holocene fault scarps in unconsolidated sediments (fine lines with adjacent dot), and seismic source regions I, II, III, IV (heavy dashed lines).

TABLE F-1. LARGEST EARTHQUAKES WITHIN 200 KM OF STUDY AREA, 1850 THROUGH 1981

Date (GMT)	Lat. (°N)	Long. (°W)	<u>I</u> 0	Magnitude <u>(M</u> L)	Location	Distance (km)
1887 Dec 05 1891 Apr 20	37.1 37.1	112.5 113.6	7 6	(5½) (5)	Kanab, Ut. Washington Co., Ut. (St. George)	77 34
1901 Nov 14	38.8	112.1	8 - 9	(6½+)	Richfield, Ut. Pine Valley, Ut. Pine Valley, Ut.	182
1902 Nov 17	37.4	113.5	8	(6)		24
1902 Dec 05	37.4	113.5	6	(5)		24
1908 Apr 15	38.4	113.0	6	(5)	Milford, Ut. Elsinore, Ut. Elsinore, Ut. Williams, Ariz. Elsinore, Ut.	110
1910 Jan 10	38.7	112.1	6-7	(5-5½)		171
1910 Jan 12	38.7	112.1	6	(5)		171
1912 Aug 18	36.5	111.5	6-7	(5½)		186
1921 Sep 29	38.7	112.1	8	(6)		171
1921 Sep 30	38.7	112.1	7	(5½)	Elsinore, Ut. Elsinore, Ut. Parowan, Ut. SE Nevada Cedar City, Ut.	171
1921 Oct 01	38.7	112.1	8	(6)		171
1933 Jan 20	37.8	112.8	6	(5)		60
1934 Apr 15	38.0	115.0	-	5.0		167
1942 Aug 30	37.7	113.1	6	(5)		34
1942 Sep 26 1945 Nov 18 1952 May 24 1959 Feb 27 1959 Jul 21	37.7 38.8 36.1 38.0 37.0	113.1 112.0 114.7 112.5 112.5	6 6 6 6	(5) (5) 5.0 (5) 5½-5 3/4	Cedar City, Ut. Glenwood, Ut. ArizNev. border Panguitch, Ut. Kanab, Ut.	34 186 195 93 81
1966 Aug 16	37.5	114.2	6	5.6	NevUt. border	80
1967 Oct 04	38.5	112.2	7	5.2	Marysvale, Ut.	158

¹Summarized from compilation of regional seismicity (M>3.0) within 200-km radial distance of study area. Table includes earthquakes of estimated Richter magnitude 5 or greater. I = maximum Modified Mercalli intensity. Magnitudes in parentheses estimated from I (M=1+2/3 I o). Distance measured from a point half-way between Cedar City and St. George.

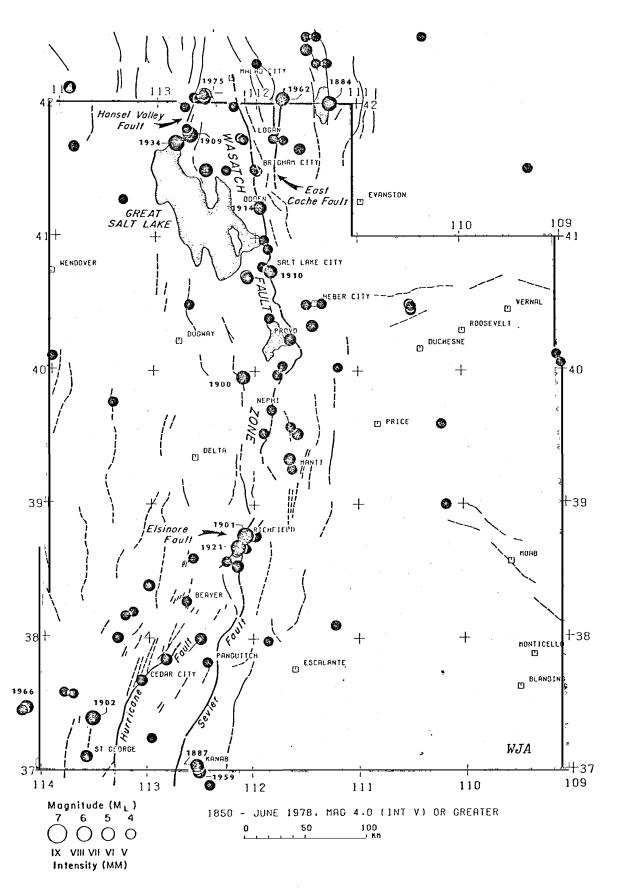


FIGURE F-3.--Epicenter map of the largest historical earthquakes in the Utah region, 1850-1978. For coincident epicenters, only the largest event is shown. Earthquakes of magnitude $5\frac{1}{2}$ or greater are dated by year (see table F-1) (from Arabasz and Smith, 1979).

- (3) An aftershock of the 1902 Pine Valley earthquake that occurred on December 5, 1902 ($I_0=6$, $M_L\sim5$).
- (4) An earthquake of January 20, 1933 ($I_o=6$, $M_L\sim5$) near Parowan, 30 km northeast of Cedar City. Minor damage occurred in Parowan and Minersville (Williams and Tapper, 1953).
- (5) Two earthquakes, apparently as part of a swarm sequence, near Cedar City on August 30, 1942, and September 26, 1942 (each with $I_o=6$, $M_L\sim5$). Minor damage occurred in Cedar City during both of these shocks (Williams and Tapper, 1953).
- F.1.2.2 <u>Swarm Seismicity in SW Utah</u>. Episodic earthquake swarms have been observed since the 1920's in southwestern Utah and appear to be characteristic of the area (e.g., Arabasz, 1979, p. 425; Richins and others, 1981). Earthquake swarms are a general type of earthquake sequence characterized by many events of nearly the same size without a distinct main shock, and they are common worldwide in areas of recent volcanism (e.g., Sykes, 1970).

Smith and Sbar (1974) discuss the occurrence of several earthquake swarms throughout the ISB, but they do not convey the relative abundance of swarm seismicity in southwestern Utah. As outlined in table F-2, at least seven swarm sequences with maximum magnitudes in the range of 3 to 5 have been documented in southwestern Utah since 1926.

The distribution of observed swarms in southwestern Utah is shown in figure F-4. There appears to be clustering in the vicinity of the Hurricane and Sevier fault zones; however, direct association with

TABLE F-2. SWARM SEISMICITY IN SOUTHWESTERN UTAH

Approx. Date	Location	Max. Magnitude (M _L)	References
1926-1927	Orderville	∿ 3	(Williams and Tapper, 1953; Arabasz and McKee, 1979)
Feb-Apr 1937	Panguitch	√4-4½	(Williams and Tapper, 1953; Arabasz and McKee, 1979)
Aug-Sep 1942	Cedar City	∿5	(Williams and Taper, 1953; Arabasz and McKee, 1979)
Dec '63-Feb '64	NevUt. border	3.2 (m _b 4.0)	(Smith and Sbar, 1974; Arabasz and others, 1979; "United States Earthquakes", 1963,1964)
Nov 1971	Cedar City	3.7 (m _b 4.5)	(Arabasz, 1979)
Dec 1978	Cove Fort	3.3	(Olson, 1976; Richins and others, 1981)
Dec '80-May '81	Kanarraville	4.5	(Richins and others, 1981)
Mar 1981	Panguitch	2.3	(Richins and others, 1981; Univ. of Utah unpub. data)
Jun-Aug 1981	Mineral Mts.	1.5-2.0 (⊳1,000 shocks)	(G. Zandt, Univ. of Utah, unpub. data)

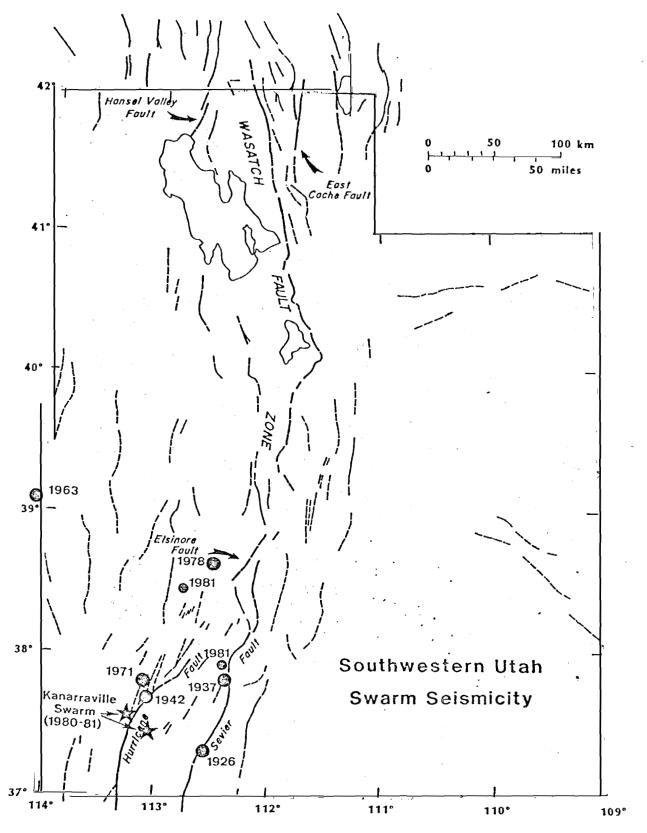


FIGURE F-4.--Location map of documented earthquake swarms in southwest Utah (see table F-2) (from Richins and others, 1981).

these faults has yet to be established (see Section F.2.2). A more general association of most of the swarm seismicity might be made with a broad zone of late Cenozoic volcanism in southwestern Utah, along the Basin and Range-Colorado Plateau transition zone (cf. Best and Hamblin, 1978, fig. 14-7).

F.1.3 Detailed Instrumental Seismicity in SW Utah

Figure F-5, from Arabasz and others (1979) gives an overview of the instrumental seismicity of the Utah region since 1962. A larger scale epicentral plot, complete through December 31, 1981, is presented in figure F-6 for a rectangular region encompassing the Cedar City-St. George study area.

A persistent feature of the seismicity of southwestern Utah (fig. F-5) is a broad band of diffuse, but locally intense, seismicity trending to the southwest below about lat 38.5°-39°N and extending into southern Nevada to form the SW Utah-S. Nevada seismic zone (Smith and Arabasz, 1979; Richins and others, 1981). Cedar City lies within the central, most seismically active part of this belt, whereas St. George lies along its southern fringe. Diffuse seismicity and problematic correlation of seismicity with the traces of major active faults is typical throughout the ISB (Arabasz and Smith, 1981). Perhaps the most striking example of this in the Utah region is the persistent low level of earthquake activity along most of the trace of the major Wasatch fault zone in central and north-central Utah (fig. F-5; Arabasz and others, 1980). In southwest Utah, the scattered occurrence of earthquake swarms contributes to the regionally diffuse epicentral pattern. Significantly, the general NE-SW trend defined by the diffuse seismicity transects the NNE-trending structural grain of the region.

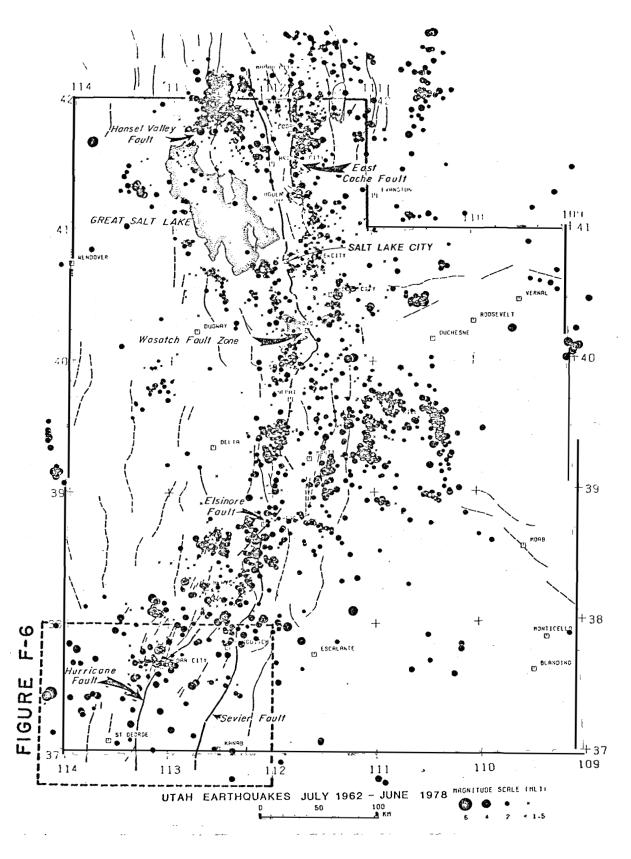
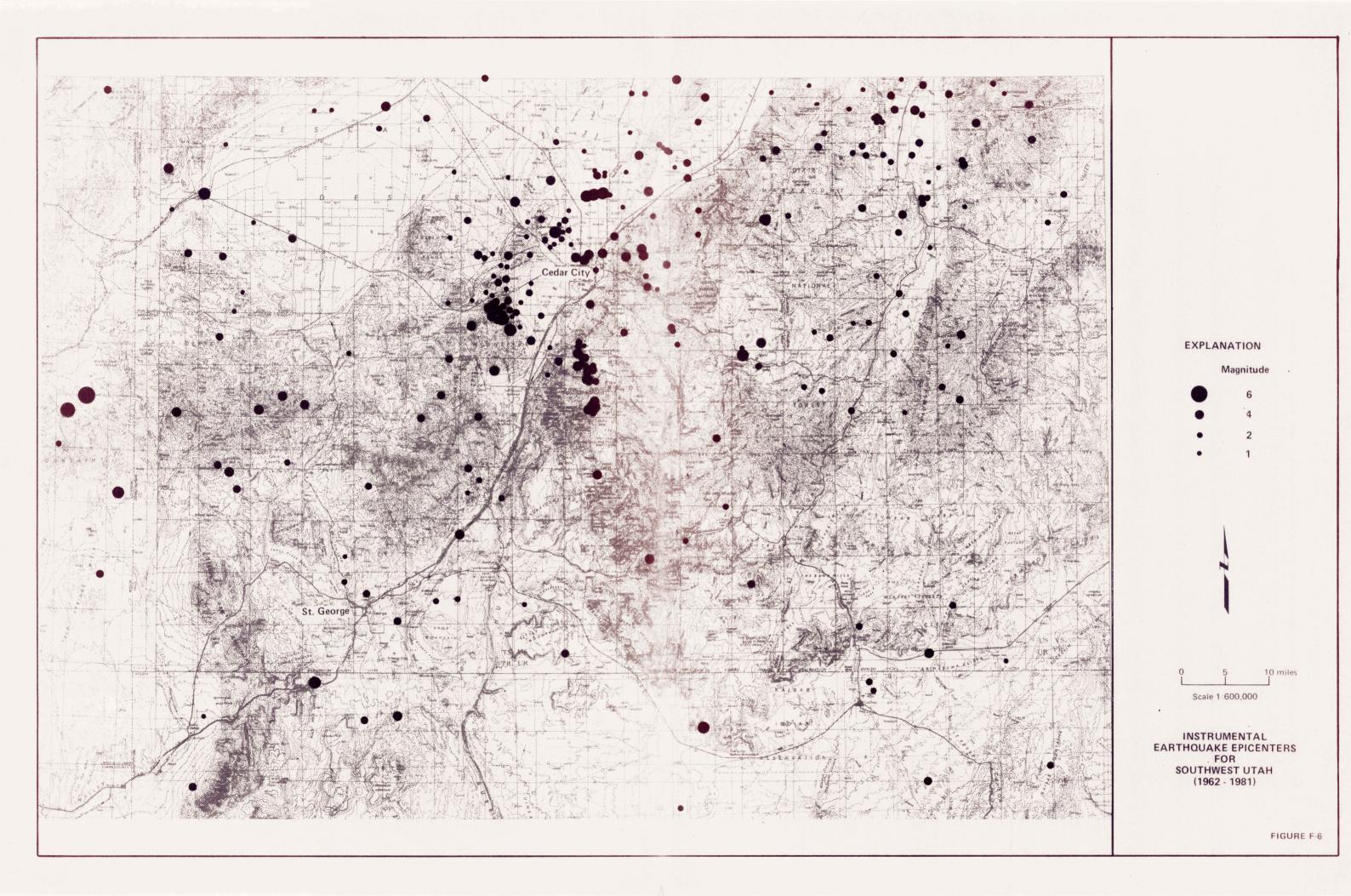


FIGURE F-5.--Epicenter map of the Utah region, July 1962-June 1978 (from Arabasz and others, 1979). Dashed box shows location of Figure F-6.

In the large-scale plot of the Cedar City-St. George area (fig. F-6), instrumental seismicity determined by the University of Utah is plotted for the period from July 1, 1962 to December 31, 1981. Before analyzing details of the epicentral patterns, some preliminary comments on the completeness and quality of earthquake locations in southwest Utah are appropriate (see Arabasz and others, 1979, p. 83, 119; Richins and others, 1981).

Prior to July 1962, earthquake epicenters in southwest Utah were based chiefly on felt reports and are assumed to have an accuracy of about ±25 to ±50 km. Assigned epicenters for non-instrumental locations correspond to the location of a town or city where felt effects were strongest. The availability of felt reports at widely distributed small towns tends to constrain the location of most of the pre-instrumental earthquake locations, particularly those of small magnitude (M_L<4). For the post-1962 period, instrumental locations greatly improve the accuracy of earthquake locations. However, the fairly wide spacing (mostly greater than 50 km) of seismograph stations in southwest Utah--even through 1981 (see Arabasz and others, 1979; Richins and others, 1981)--limits the precision with which earthquakes in the area can be located. Accuracy and precision are certainly lower than in other more densely instrumented areas of the Utah region such as the Wasatch Front area.

For the Wasatch Front area, epicentral precision is estimated as ± 5 and ± 2 km for the 1962-1974 and post-1974 periods, respectively (Arabasz and others, 1979). For southwestern Utah, precision expectedly would be less--



perhaps at best about ± 10 km and ± 5 km for the same two instrumental periods. Richins and others (1981) estimate an epicentral precision of ± 2 km for the earthquake locations near Kanarraville in 1981 as a result of special data processing.

The majority of earthquakes plotted in figure F-6 have been located with a restricted focal depth of 7.0 km. Reasonably accurate resolution of focal depth requires a recording station to be roughly as close as the depth of the earthquake--a condition not generally met in southwest Utah.

Post-1962 instrumental earthquake locations in southwest Utah are believed to be systematically complete above about magnitude (M_L) 2.5 (Arabasz and others, 1979; Richins and others, 1981).

Notable features of the intrumental seismicity depicted in figure F-6 include the diffuse scatter of epicenters, the greater abundance of earthquakes in the northern half of the sample area compared to the southern half, and a significant temporal clustering of earthquakes (evaluated from corresponding listings). An area within about 25 km radius of Cedar City displays abundant earthquake activity during the 1962-1981 period.

Earthquakes during the 1962-1974 period were widely scattered throughout the sample area. The largest shock was an event of magnitude (M_L) 5.6 (M_b =6.1) that occurred on August 16, 1966 close to the Nevada-Utah border at the western extremity of the sample area (see Smith and Sbar, 1974). The most significant temporal clustering during this time period was associated with the November 1971 Cedar City earthquake swarm

(Arabasz, 1979). Thirteen of the circles clustered immediately to the north of Cedar City correspond to located earthquakes (2.1 \le M_L \le 3.7) associated with that swarm.

Earthquakes during the 1974-1981 period scatter throughout the sample area, but the sample for this time period is dominated by intense clustering to the northeast and southeast of Kanarraville-- associated with an earthquake swarm ($M_L \le 4.5$) during the period December 1980-Mav 1981 (Richins and others, 1981). We discuss this earthquake sequence further in Section F.2.2.

Clustering to the northwest and southeast of Panguitch also relate to an earthquake swarm—a sequence in March 1981 ($M_L \le 2.3$) that displayed the same curious pattern of paired epicenter clusters observed for the Kanarraville swarm (Richins and others, 1981). Earthquakes in the St. George area are clearly less abundant than to the north. The largest shock in the immediate vicinity of St. George during the entire 1962–1981 period was an earthquake of magnitude (M_L) 3.6 on August 5, 1979, about 15 km southwest of St. George.

F.2 SEISMOTECTONICS AND EARTHQUAKE HAZARD EVALUATION

F.2.1 Implications of Current Geological Studies

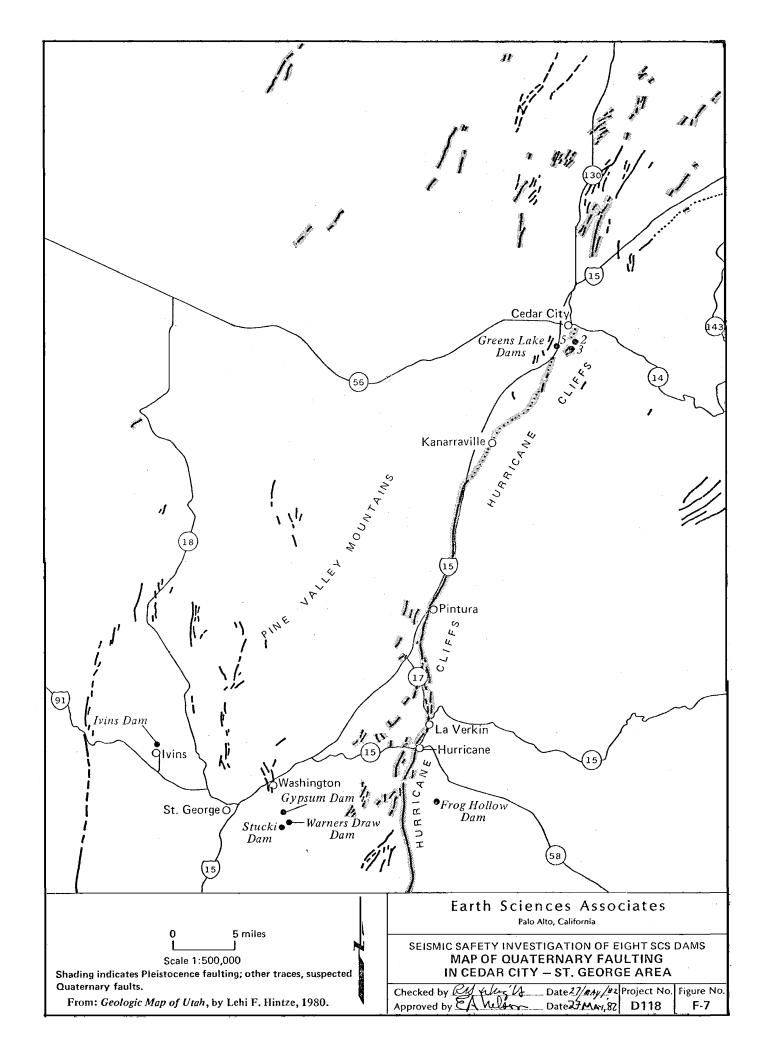
Since about 1975, as part of the National Earthquake Hazards Reduction Program, the U.S. Geological Survey (USGS) has carried out active geological studies in the Utah region aimed at earthquake hazard evaluation. Of particular relevance to this study is a USGS project entitled "Southwestern Utah seismotectonic studies," headed by R. E. Anderson; the goals of the project are: "(1) to define the distribution in space and time of Quaternary faulting and deformation in southwestern Utah and (2) to provide an improved understanding of the late Cenozoic tectonic history and framework of the area" (Anderson, 1981).

R. E. Anderson and R. C. Bucknam of the USGS have jointly carried out mapping of late Quaternary fault scarps in unconsolidated deposits throughout western and southwestern Utah, systematically covering 1° x 2° quadrangles (see Anderson, 1980). Unfortunately, no compilation for the Cedar City 1° x 2° sheet (1:250,000 scale), which encompasses the Cedar City-St. George area, is yet available. However, abundant information is directly relevant to the seismotectonics of the study area has been published by R. E. Anderson, R. C. Bucknam, and colleagues at the USGS (e.g., Anderson, 1978, 1979, 1980, 1981; Anderson and others, 1978; Anderson and Mehnert, 1979; Anderson and Bucknam, 1979; Bucknam and Anderson, 1979; Bucknam and others, 1980; and Zoback and others, 1981). Much of this section involves a distillation of on-going USGS work to identify key information relevant to the present SCS project.

F. 2.1.1 Identification of Active Faults. Although results of current USGS fault studies in the Cedar City - St. George area are not yet available in regional map form, the "Quaternary Fault Map of Utah" (1:500,000 scale) recently compiled by L. W. Anderson and D. G. Miller (1979) provides the following preliminary summary, which is modified as noted by the results of the detailed photogeology and geological field work carried out as part of this project and discussed in detail in Chapters IV and VIII of the report and Appendix E. 3. 2. 5.

Figure F-7 shows part of the Anderson and Miller (1979) fault map relevant to the present study indicating fault traces associated by Anderson and Miller with either late Pleistocene (10,000 to about 500,000 yr B.P.) movement or suspected Quaternary (0 to about 2 or 3 m.y. B.P.) movement. No Holocene (<10,000 yr B.P.) faulting was indicated in the study area before trench exposures developed during this study were examined (see Chapter VIII of this report).

The main active faults shown in figure F-7 are the northernmost part of the Grand Wash fault (comprising segments variously named the Cedar Pocket Canyon fault and the Gunlock-Veyo fault) and the Hurricane-Parowan fault system. Fault traces of probably late Pleistocene age that are indicated in the northern part of figure F-7 in the Escalante Desert region displace Quaternary basalts and alluvial deposits, and are based on personal communications to Anderson and Miller (1979) by W. K. Hamblin and R. E. Anderson. Only the northwesternmost part of the Washington fault near the town of Washington is indicated in figure F-7 as being active during the Quaternary, although subsurface exploration conducted as part of this study across the fault at Gypsum Wash Dam revealed probable early Holocene faulting (see Appendix E. 3.4 and Chapter VIII).



F.2.1.2 <u>Late Quaternary Versus Holocene Faulting.</u> A critical problem prior to the current study was whether there is convincing evidence anywhere in the study area for Holocene faulting (i.e., the occurrence of scarp-forming earthquakes during the last 10,000 years). The absence of any historic surface rupture in southwest Utah has led to emphasis on Quaternary studies relating to the geomorphic evolution of fault scarps (Anderson, 1978; Bucknam and Anderson, 1979), Lake Bonneville history and its correlation with glacial deposits and soil stratigraphy (Anderson and Bucknam, 1979), and the stratigraphy of faulted Quaternary extrusive rocks (Hamblin and others, 1981).

Published statements by R.E. Anderson (1978, p. 5; 1978a) indicating Holocene faulting northeast of Cedar City relate to a graben structure at Braffits Creek (see fig. F-7) which Anderson has subsequently interpreted as a tensional collapse structure along the rising and spreading mountain front of the eastern Markagunt Plateau (Anderson and Bucknam, 1979; Anderson, 1980, p. 528). Another area of probable Holocene deformation in southwestern Utah described by Anderson and Bucknam, 1979) is in Escalante Valley, more than 30 km northwest of Cedar City, where Lake Bonneville shorelines have been deformed by regional uplift. Anderson (1980, p. 525) also describes deformation within the Enoch graben, about 10 km north-northeast of Cedar City (see fig. F-7), where Holocene faulting may have occurred; however, the interpretation of faulting is equivocal.

At Shurtz Creek to the south of Cedar City (see fig. F-7) there is a prominent alluvial scarp along the trace of the Hurricane fault that may not be much older than Holocene in age. Anderson (1980), p. 536) summarizes the following key information:

"Averitt (1962) described and illustrated a conspicuous scarp at the mouth of Shurtz Creek about 8 km south-southwest of Cedar City. The scarp is at least 20 m high, is formed on very coarse bouldery alluvium, has a slope angle of 29°, and is deeply incised by active streams. On the basis of its profile, the scarp appears to be pre-Holocene, but its age is probably close to the Pleistocene-Holocene boundary. The surface that is displaced at this scarp was referred to by Averitt (1962) as the Schurtz Creek pediment. Recent studies have shown that the Quaternary gravels that mantle the pediment serve as parent material for a soil that is definitely pre-Holocene and is probably more than 50,000 years old."

A second, more subdued and probably older scarp in alluvial fan deposits located upslope to the east of this location was identified during this investigation (see Chapter IV and figure IV-2).

Bucknam and others (1980) have summarized the results of USGS studies of late Quaternary faulting in western and southwestern Utah, including a preliminary map of seismic source zones defined on the basis of geologic data. That source zone map is reproduced here as figure F-8. (Note that the fault scarps shown in the figure are only those formed in unconsolidated deposits.) Most of southwestern Utah lies within seismic zone IIb (fig. F-8) defined by "the occurrence of late Quaternary fault scarps and the absence of Holocene fault scarps" (Bucknam and others, 1980, p. 304). Adjacent region IIIb is defined by the absence of fault scarps in unconsolidated deposits, whereas region Ia is characterized by abundant Holocene fault scarps. Faults at Braffits Creek and near Enoch are not included in the compilation because of major uncertainties; the interpretation of faulting within the Enoch graben is equivocal, and faulting at Braffits Creek may be associated with an aseismic style of deformation (Anderson, 1980; Bucknam and others, 1980).

In the absence of definitive dates on samples of materials offset by fault traces exposed in trenches excavated as part of this study, the characterization of most of southwestern Utah as without Holocene faulting must be considered a working hypothesis; it seems to be consistent with the morphology of fault scarps encountered in the area (e.g., Bucknam and Anderson, 1979; Anderson, 1980, 1979). The position of USGS workers such as R. E. Anderson remains that unless it can be proven that dated Holocene sedimentary strata have been faulted, the assignment of a Holocene age to any

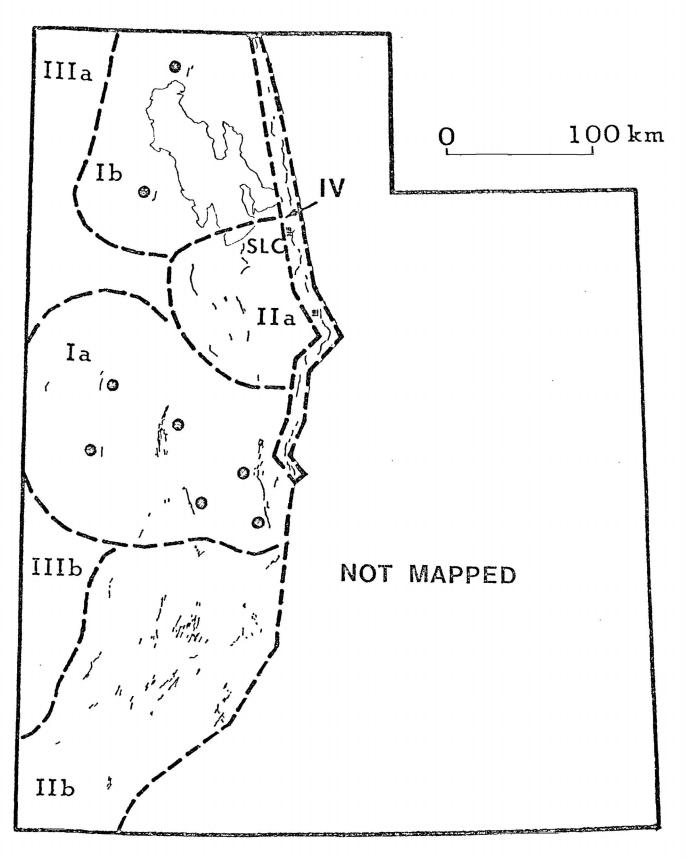


FIGURE F-8.--Map from Bucknam and others (1980) showing late Quaternary fault scarps in unconsolidated sediments (fine lines), Holocene fault scarps in unconsolidated sediments (fine lines with adjacent dot), and seismic source regions I, II, III, IV (heavy dashed lines).

faulting in southwestern Utah should be considered uncertain (e.g., see Anderson, 1979, and Huntoon, 1979).

It should be emphasized that the absence of Holocene faulting in southwestern Utah, if valid, does <u>not</u> imply no likelihood of a future scarp-forming earthquake in that region. Fault-scarp data in region IIb (fig. F-8) over the last 100,000 years or so lead to the expectation of about two earthquakes in the magnitude range 7-7½ somewhere in region IIb during Holocene time--and the apparent absence of any such event is not rigorously inconsistent with a Poisson process (Bucknam and others, 1980). Bucknam and others (1980, p. 306) conclude: "We believe that the observed [relatively high rate of] historic seismicity of the region IIb is consistent with strain release primarily by numerous relatively small earthquakes and comparatively infrequent large events (M = 7.0 to 7.6)."

F.2.2 Implications of Current Seismological Studies

F.2.2.1 <u>Listric Faulting and Problematic Correlation of Seismicity With</u>

<u>Geologic Structure</u>. Critical problems facing earthquake researchers in the

Utah region include uncertainties about the subsurface structure of major

normal fault zones in the area, the correlation of diffuse seismicity with

geologic structure, and evaluation of the relative importance of low-angle,

listric (downward-flattening) normal faulting (see Figure F-9) as part of the

seismogenic process (Arabasz and Smith, 1981).

During the past few years, seismic reflection surveys have been conducted in seismically active areas of the Intermountain region as part of oil industry efforts to explore and extend knowledge of the foreland "overthrust belt" into the Great Basin (e.g., MacDonald, 1976; Royse and others, 1975; Effimoff and Pinezich, 1981). New high-resolution seismic reflection data indicate that the seismogenic upper 10 km of the crust along the Intermountain seismic belt is more deformed than previously envisaged, and the style of normal faulting is more complicated than perceived from surface mapping. Converging evidence suggests that the structural style along the eastern Great Basin is that of "thin-skinned" tectonics involving low-angle detachment surfaces. Low-angle thrust faults, formed during the Sevier and Laramide orogenies fundamentally influenced the development of Cenozoic basin-range structure, and pre-Cenozoic low-angle faults may now be accommodating normal dip-slip displacement in an extensional stress field (Arabasz and Smith, 1981; Arabasz, 1981; Smith, 1981).

Older low-angle faulting in parts of the Great Basin, including the study area (Anderson, 1981) resulted from a pre-basin-range extensional episode in about Miocene time. The important point here is that active normal faults

mapped at the surface within the study area probably do not have a simple subsurface geometry, and the deformational behavior of these faults is poorly understood. For example, little information is available about the subsurface geometry of the Hurricane fault. However, conspicuous reverse-drag flexing of basin-fill strata observed along the Hurricane fault (Hamblin, 1965; Hamblin and others, 1981) is now known to typify downward-flattening listric normal faults (e.g., Anderson and Zoback, 1981). Confidential data acquired by Mobil Oil Corporation suggests that the Hurricane fault probably is listric at depth (R.E. Anderson, personal communication to W.J. Arabasz, January 1982).

Complicated upper-crustal structure undoubtedly has a fundamental bearing on the significance of diffuse background seismicity in the Intermountain seismic belt and on how infrequent large earthquakes are generated in the area. California-type models that assume high-angle faulting and the planar clustering of small-magnitude earthquakes on simple fault planes clearly are inapplicable to southwest Utah.

Recent detailed microearthquake studies in central and south-central Utah (Arabasz, 1981; see fig. F-9) suggest that diffuse seismicity in areas where listric normal faulting has been identified in the subsurface does not simply reflect seismic slip on low-angle normal faults; rather, seismic slip seems to predominate on relatively high-angle fracture planes—either the upper portion of west-dipping downward-flattening faults, related east-dipping antithetic faults, or moderately dipping secondary faults within the major fault blocks. Research on such problems is underway, but is still in an early stage.

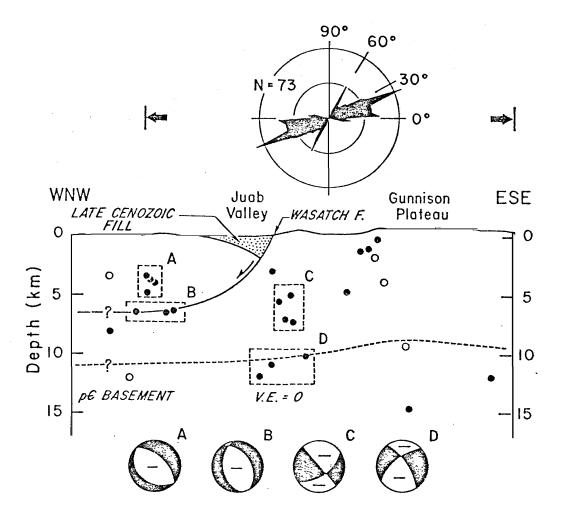


FIGURE F-9.-- Cross-section depicting results of detailed study of microseismicity, focal mechanisms, and correlation with geological structure near the southern end of the Wasatch fault (from Arabasz, 1981). Earthquake hypocenters, shown as small circles (open for lesser quality), are from field study by McKee and Arabasz (1981). Geology adapted and "sanitized" from confidential oil-company information. Rose diagram in upper part of figure summarizes fault dips from subsurface structure chiefly deleted from diagram. Faults within area of section delineated by heavy arrows were discretized into 1-km-long segments, and the average orientation of the discrete segments was then measured and summarized. Focal mechanisms A, B, C, and D were determined for clustered hypocenters in the correspondingly labeled boxes by iteratively inverting SV/P amplitude ratios (Kisslinger et al., 1981) -- with constraints from P-wave first motions. Despite the predominance of low-angle faulting inferred by oil-company scientists in the subsurface, the focal mechanisms do not indicate seismic slip on low-angle (listric) faults. Focal mechanisms are lower-hemisphere projections, equatorial plane.

F.2.2.2 <u>Implications of Swarm Occurrence</u>. A recent earthquake swarm sequence $(M_L \le 4.5)$ occurred during December 1980-May 1981 near Kanarraville, 20 km southwest of Cedar City, and was studied in detail by Richins and others (1981a). Epicenters located to an estimated precision of $\frac{1}{2}$ km by a joint-hypocenter-determination technique (e.g., Dewey, 1979) separate into two clusters, 5-10 km northwest and southeast, respectively, of the main trace of the Hurricane fault (fig. F-10). During the swarm there was a temporal migration of activity from the southwest cluster to the northwest cluster involving "the complete cessation of earthquakes ($M_L \le 1.5$) in the first cluster prior to activation of the second cluster 10 days later" (Richins and others, 1981a).

The Kanarraville earthquake sequence typifies swarm seismicity characteristic of southwest Utah (see Section F.1.2.2) as well as the absence of any clear correlation with mapped geologic structure. Richins and others (1981a) interpret that the sequence reflects deformation on subsidiary structures, related in some complicated way to the Hurricane fault.

The tectonic significance of swarm seismicity in southwest Utah is poorly understood. One might argue that the apparent predominance of swarm seismicity—with maximum events of only moderate size (M_L 4½ to 5)— perhaps diminishes the likelihood of large mainshock earthquakes in this area. Such a position would be difficult to maintain in view of geological evidence for large, albeit infrequent, earthquakes (see Section F.2.1). Dewey and others (1973) caution that swarm—producing tectonic conditions may be highly localized and do not necessarily exempt a site from large "non-swarm" earthquakes nearby. For example, they note that a magnitude 6 3/4 earthquake occurred in 1925 east of Helena, Montana, only 60 km from areas such as Flathead Lake, Montana, and

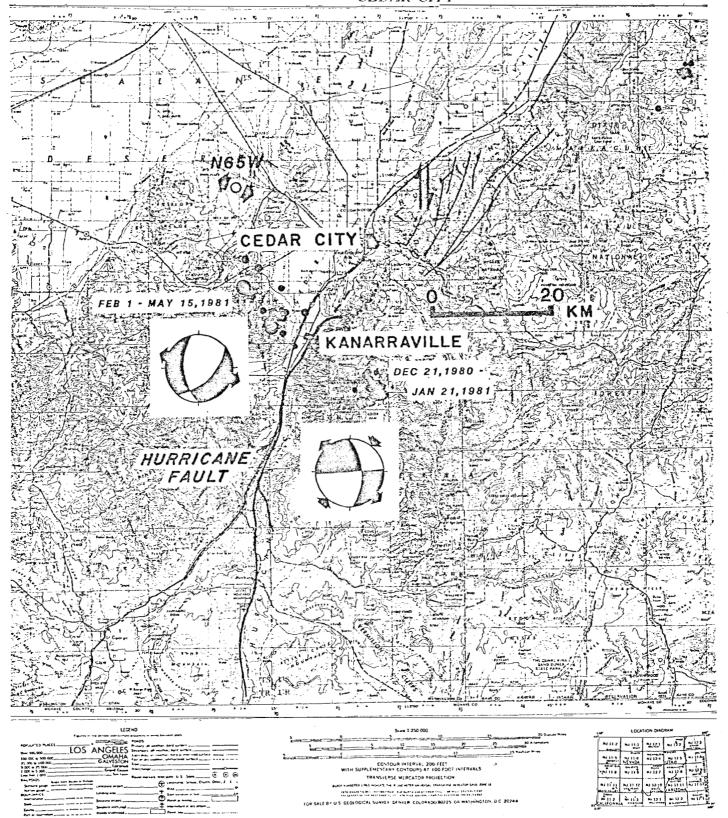


FIGURE F-10.--Summary map of swarm seismicity (M ≤ 4.5) located by the University of Utah near Kanarraville, Utah, during December 1980-May 1981 (after Richins and others, 1981a). Earthquake epicenters (solid circles) cluster northwest and southeast of the trace of the Hurricane fault. Beach balls indicate composite focal mechanisms (lower-hemisphere, equal-area projections; compressional quadrants shaded) for the two clusters. Outward-directed arrows indicate extensional direction; extensional direction shown with open circle is from Anderson (1980), inferred from late Quaternary geology.

Helena itself—areas characterized by historical swarm seismicity (e.g., Smith and Sbar, 1974).

In summary, the significance of earthquake swarm occurrence in southwestern Utah is:

- 1. There is no observable correlation of such activity (which includes earthquakes up to Magnitude $4\frac{1}{2}$ to 5) with known major structures, and they are not thought to be associated with volcanic activity, and
- 2. Therefore, it is possible that renewed swarm activity could occur anywhere in southwestern Utah on a relatively frequent basis, and such activity could be located near any of the dam sites (especially those in the Cedar City area).

F.2.2.3 Focal Mechanisms and Stress Orientation. Available focal-mechanism information for the main seismic belt in central and southwest Utah is summarized in figure F-11, which shows a predominance of normal faulting with extension predominant in a general east-west direction. There clearly are local complications, however, as reflected by reverse faulting (solution 6) or a large component of strike-slip displacement (solutions 9, 10, 13, 17, 19, 21).

The mean direction of the T-axes (the direction of least principal horizontal stress) for the earthquake focal mechanisms shown in figure F-11 is in the 093° - 273° direction ($^{\pm}$ 29°), i.e., approximately east-west. The modern regional stress field in the Basin and Range province is characterized on a regional scale by a horizontal least principal stress trending approximately WNW-ESE (Zoback and Zoback, 1980). Along the eastern Great Basin there is some variation. In north-central Utah the least-principal-stress orientation is approximately N75°E, a direction consistently indicated by average T-axes from focal mechanisms, moment-tensor summation, hydrofrac measurement, and measurements of the orientation of young faults and slickensides (Zoback, 1981).

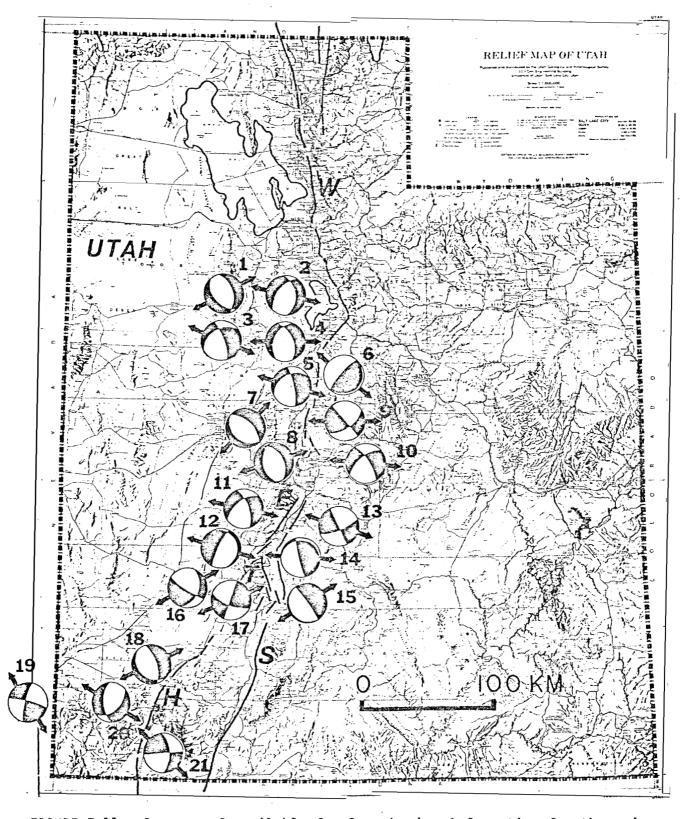


FIGURE F-11.--Summary of available focal mechanism information for the main seismic belt in central and southwest Utah (after Arabasz, 1981). Focal mechanisms are lower-hemisphere, equal-area projections, with compressional quadrants shaded. Outward-directed arrows indicate extensional direction. Sources: 5, 14, 18, 19 (Smith and Sbar, 1974); 6, 15 (Sbar and others, 1972); 11, 12 (Olson, 1976); 20, 21 (Richins and others, 1981a); others (unpub. data, University of Utah). H = Hurricane fault, S = Sevier fault, E = Elsinore fault, W = Wasatch fault, T = Tushar fault.

For the data in figure F-11 the mean least-principal-stress orientation is significantly different from N75^OE, as verified by applying Student's t test. The N75^OE trend is intriguing, however, because there appear to be local areas where NE-SW trending T-axes represent local deviations from more general WNW-ESE extension typical of the Great Basin on a regional scale. In southwestern Utah, for example, mechanisms 15, 16, 17 and 18 lie in an east-west belt coincident with ENE-trending geophysical anomalies (Stewart and others, 1977). Each mechanism comes from a different source of data, yet together they display a remarkable consistency. A measurement of <u>in situ</u> stress near Cedar City reported by Smith (1978, fig. 6-12) shows a direction of minimum compressive stress that also trends NE-SW. These observations raise the possibility that local stress orientation near Cedar City may be different than that to the south near St. George--perhaps somehow related to observed differences in background seismicity in the two areas.

F.2.2.4 Measurements of Ground Response in the Cedar City Area. W.W. Hays and K.W. King of the USGS have recorded nuclear explosions with portable broadband velocity seismographs in various urban areas in Utah to evaluate ground-shaking hazard (e.g., Hays and others, 1980). Characteristics of ground response determined by these studies are of value because of the existence of only one strong-motion record in Utah--that for the 1962 Cache Valley earthquake (M₁5.7) in northern Utah (Smith and Lehman, 1979).

King (1982) reports results of ground-response measurements recently made in the Cedar City area at ten sites, where values of mean transfer functions (velocity spectra), relative to rock sites, show differences of a factor of 5 in the period band 0.2-1.0 sec. According to K.W. King (personal communication to W.J. Arabasz, March 1982), one of the ten sites occupied during the Cedar City ground-response study was at one of the SCS dam-site embankments at the southeastern corner of Cedar City; at this site, ground response compared to rock sites was only higher by about a factor of two.

F.3.1 General Statement

Rates of earthquake recurrence in any seismically active region can be variously estimated using both seismological and geological data. If a catalog of historical and/or instrumental seismicity is available, then relationships of earthquake frequency versus magnitude can be used to estimate the average recurrence interval or inter-event time for shocks of a specified magnitude. Recurrence estimates can also be made from seismic moment rates determined from available geologic data on Quaternary faulting (e.g., J.G. Anderson, 1979; Molnar, 1979). Recurrence of surface faulting events and fault displacement rates can be estimated from trenching studies or other special geologic studies of specific fault segments.

Each of the above-mentioned approaches has been recently applied to different seismically active areas in Utah. In this section we summarize available information on earthquake recurrence in southwestern Utah and evaluate parameters of earthquake generation and surface fault displacement specific to the Cedar City-St. George area.

The region considered for this study is shown in Fig. F-12a as subregion III. This region of southwest Utah encompasses an area of approximately 56, 160 km 2 .

F.3.2 Earthquake Frequency Versus Magnitude

The relationship of earthquake frequency (of occurrence) versus magnitude is commonly expressed by the well-knowm equation:

$$\log N(M) = a - bM \tag{1}$$

where N(M) is the number of earthquakes of specified magnitude M per unit time. Here we will designate N(M) as either n, the annual frequency of occurrence of earthquakes of a given magnitude, or as $N_{\rm c}$ (i.e., "cumulative N"), the annual frequency of occurrence of earthquakes equal to or greater than a given magnitude.

Equation (1) defines a linear relationship between log N(M) and M in

which the slope coefficient <u>b</u> determines the relative proportion of small to large earthquakes. The constant <u>a</u> depends both on the temporal and spatial distribution of earthquakes within the region being considered. It is important to note that calculations of recurrence always imply a distribution of earthquake activity within the region that must be clearly specified to compare recurrence intervals in a meaningful way. In most seismic risk studies this distribution is usually assumed to be uniform throughout the entire region.

The following factors pose problems for establishing frequency-magnitude relationships for the study area: (1) the non-uniformity of the regional 200-km-radius seismicity sample surrounding the study area (fig. F-2), and (2) relatively low local seismic activity that precludes, say, the establishment of a reliable frequency-magnitude relationship for the St. George area separate from the Cedar City area--much less separate relationships for individual faults. Even for the rectangular region encompassing the Cedar City-St. George area (fig. F-6), it is doubtful that the post-1962 instrumental seismicity provides a good basis for estimating the frequency of occurrence of dàmaging earthquakes. For the 19.5-yr period July 1962-December 1981, the latter sample includes only 36 earthquakes of magnitude 3.0 or greater and is dominated by swarm earthquakes.

In our judgment the most meaningful approach is to follow the analyses of seismicity in the Utah region carried out by Arabasz and others (1980) and Smith and Arabasz (1979), restricting attention to the University of Utah earthquake catalog. Smith and Arabasz (1979) performed analyses specific to the southwest Utah region (region III, fig. F-12a). This region encompasses an area of 56, 160 km² and reflects the seismotectonics of the study area more appropriately than the 200-km-radius sample.

Only independent mainshocks (or the largest event of a swarm sequence) were included in the following analyses, in order not to invalidate the assumption of a Poisson distribution. Accordingly, predicted numbers of earthquakes based on the various recurrence relationships likewise refer to independent shocks.

Figure F-12 summarizes recurrence relations for the defined southwest Utah region. Figure F-12b shows the annual frequency of occurrence versus maximum Modified Mercalli epicentral intensity in southwest Utah, based upon the 129-yr record from 1850 through 1978. Stepp's (1972) technique was used to correct the historical catalog for sample incompleteness. Assuming the Gutenberg and Richter (1956) relationship: $M_L = 1 + 2/3 I_o$, as justified for the Utah region by the U.S. Geological Survey (1976), the results of figure F-12b can be transformed to express either n or N_c in terms of M_L :

$$\log n = 1.64 - 0.54 M_{\parallel}$$
 (2)

$$\log N_{c} = 2.38 - 0.65 M_{L}$$
 (3)

For the 16.0-yr sample of instrumental seismicity (1962-1978), the following relationships apply to the southwest Utah region (fig. F-12c; Smith and Arabasz, 1979):

$$\log n = 2.38 - 0.74 M_{\odot}$$
 (4).

$$\log N_{c} = 2.73 - 0.74 M_{L}$$
 (5)

Note that equations (2) through (5) apply to the frequency of occurrence of earthquakes for the entire southwest Utah region (region III, fig. F-12a which has an area of 56, 160 km²).

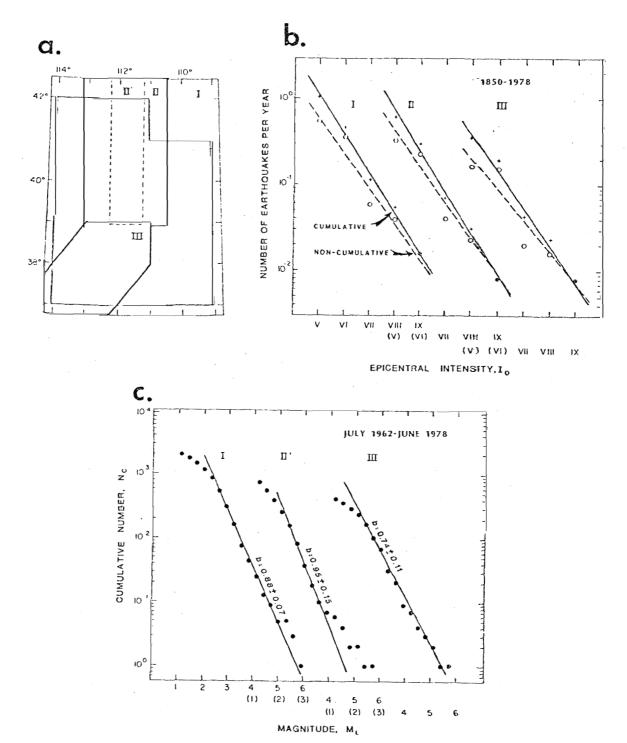


FIGURE F-12.- Earthquake recurrence data for the Utah region and subregions outlined in (a). (b) Annual frequency of occurrence versus Modified Mercalli epicentral intensity, corrected for sample incompleteness, based upon the 129-yr record: 1850 through 1978. Letting $N_{\rm C}$ = the annual number of earthquakes equal to or greater than a given magnitude, and n = the annual number of a given magnitude, the regression coefficients are: (I) log $N_{\rm C}$ = 2.34 - 0.46 I₀, log n = 1.81 - 0.41 I₀; (II) log $N_{\rm C}$ = 2.26 - 0.48 I₀, log n = 1.72 - 0.42 I₀; and (III) log $N_{\rm C}$ = 1.73 - 0.43 I₀, log n = 1.10 - 0.36 I₀. (c) Cumulative number of earthquakes versus magnitude for the 16.0-yr sample of instrumental seismicity compiled in section 8. Values of the slope coefficient b at the 95 percent confidence limits computed from non-cumulative distributions by the method of maximum likelihood.

The method of extreme values (e.g., U.S. Geological Survey, 1976; Arabasz and others, 1980) was also applied to the historical data set for 1850-1978 in southwest Utah to minimize errors for sample incompleteness; the results (fig. F-13) are close to those computed by the relationship: $\log N_{\rm C} = 1.73 - 0.43 I_{\rm O}$ outlined in figure F- b. The extreme-value distribution of maximum intensities in each year of the historical record is estimated by:

$$F(x) = \exp\{-\exp[-(x - 3.50)/ 1.14]\}, -\infty < x < +\infty$$
 (6)

where F(x) is the probability that the largest earthquake in a year will have intensity less than or equal to x. The return period R, the interval in which an earthquake of a given intensity or greater has a 63 percent probability of occurrence, is estimated by (see U.S. Geological Survey, 1976):

$$R = 1/[1 - F(x)]$$
 (7)

Here also, maximum epicentral intensity can be transformed to magnitude using: $M_L = 1 + 2/3 I_o$.

Estimates of the frequency of occurrence for a range of earthquake sizes—using the various relationships derived from seismicity—are outlined in table F-3. Values for estimated recurrence within 50-km radius of any site within the defined southwest Utah region simply assume uniform seismic activity throughout the region (a reasonable assumption for the larger earthquakes (see fig. F-3)), and they account for the proportionality of a 50-km-radius area to the total southwest Utah sample area.

A fundamental assumption regarding the estimation of earthquake recurrence is that the data used to calculate the constants \underline{a} and \underline{b} in equation (1) accurately represent the long-term seismicity of a region. Ideally the data should represent a long enough period of time that includes at least one repeat interval of the largest earthquake. This clearly is not the case in southwest Utah for earthquakes larger than about 6, so there is considerable

uncertainty in extrapolating to earthquakes of larger size. Cluff and others (1980), for example, discuss this problem at length for the Wasatch Front area of north-central Utah where extrapolation of the historical seismicity does not adequately represent the rate of occurrence of moderate to large earthquakes. Prior to the present investigation, the apparent absence of Holocene fault scarps in southwest Utah (Bucknam and others, 1980) seemed inconsistent with the estimated recurrence intervals of $\sim\!200\text{--}700~\text{yr}$ for $\text{M}_L\geq7.5$ indicated in table F-3. The identification of Holocene and probable Holocene faulting during this investigation (see Section VIII) removes some of this inconsistency and suggests that rather than being absent, Holocene faulting in southwest Utah is probably present in many areas, but not yet identified by site specific studies.

Rates of earthquake occurrence in southwest Utah can be estimated fairly well for magnitudes up to about $6-6\frac{1}{2}$ on the basis of past seismicity. Comparing the data of tables F-1 and F-3, it can be seen that for the 129-yr historical record of southwest Utah the observed number of independent earthquakes in that region is $1 \geq M_L$ 6.5, $3 \geq M_L$ 6.0, $8 \geq M_L$ 5.5, and $18 \geq M_L$ 5.0-whereas the respective numbers of predicted earthquakes is ~ 2 , ~ 4 , 7-8, and 14-17.

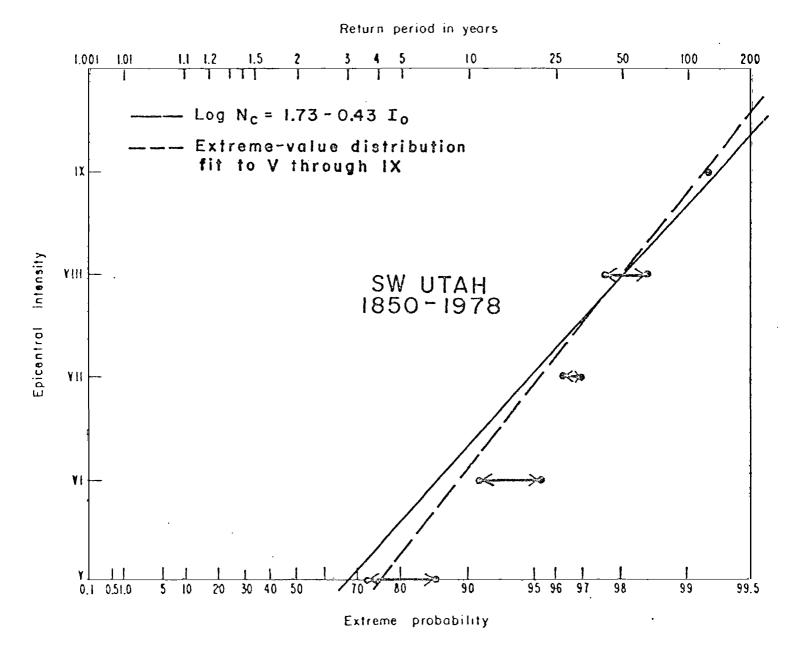


FIGURE F.-13.--Extreme-value distribution (dashed line) showing the probability that a given intensity is the largest intensity in a given year. Also shown is the return period or the interval in which a given intensity or greater has a 63 percent probability of occurrence. The solid line indicates the recurrence relationship derived by Smith and Arabasz (1979) for southwest Utah, transformed to extreme probability.

Table F-3--Estimated Frequency of Occurrence of Damaging Earthquakes (in years)

Earthquake	Recurrence	for Entir	re SW Utah	Recurrence	Within 50 k	m of any Site
Size	Case I	Case II	Case III	Case I	Case II	Case III
M _L ≥7.5	313	241	660	2,240	7,720	4,720
M _L ≥7.0	148	125	282	1,060	894	2,020
M _L ≥6.5	(1) [2] 70	65	120	500	465	858
M _L ≥6.0	(3) [4] 33	34	51	236	243	365
M _L ≥5.5 (9)[3-8] 16	18	22	114	129	157
M _L ≥5.0 ([18] [14-17] 7.4	9.5	9.3	53	68	66

Case I: Historical data, 1850-1978: log N_c = 2.38 - 0.65 M_L

Case II: Historical data, 1850-1978: R = 1/[-F(x)]

where

 $F(x) = \exp \{-[-\exp [-(x-3.50)/1.14]\}, -\infty < x < +\infty$

Case III: Instrumental data, 1962-1978: log N_c = 2.73 - 0.74 M_L

F.3.3 Seismic Moment Rates

The frequency of occurrence of moderate-to-large earthquakes can also be estimated from geologic data by relating geologically determined slip rates on individual faults to seismic moment rates (e.g., J.G. Anderson, 1979; Molnar, 1979). Such an approach usefully complements analysis based on the historic record of seismicity, which is generally too short to evaluate with the confidence /long-term seismic flux over hundreds or thousands of years.

Seismic moment M_0 is a fundamental parameter now used by seismologists to describe the "size" of an earthquake. In practice, M_0 is generally determined from spectral measurements of seismic-wave recordings. An important result based on elastic dislocation theory is that seismic moment is proportional to the average slip on a fault during an earthquake: $M_0 = \mu A \bar{u}$, where μ is the shear modulus, M_0 is the fault area, and \tilde{u} is the average slip on the fault during an earthquake. Seismic moment rate M_0 , the rate of occurrence of seismic moment, would then directly relate to the time derivative of \tilde{u} , which is the average slip rate along a fault. In simple terms, moment rate can be calculated from geological information on fault slip rate and fault length, assuming some estimate of the depth of faulting. For a region, a moment rate can be estimated by summing the moment rates of major faults throughout the region, or by using a relationship involving the average rate of deformation of a region (e.g., J.G. Anderson, 1979).

Assuming that all slip on faults occurs seismically, a theoretical expression can be derived for $N(M_0)$, the number of earthquakes per unit time of moment M_0 or greater (Molnar, 1979). The complicated expression is a function of geologically-determined moment rates, the moment of the largest possible earthquake in the

region, the slope coefficient \underline{b} from the frequency-magnitude relationship appropriate for the region, and another slope coefficient defining an empirical linear relationship between seismic moment and magnitude in a region (Molnar, 1979). The expression provides an upper limit for $N(M_0)$ if some aseismic slip occurs on faults.

Doser and Smith (1982; see also Doser, 1980) have applied the moment rate method to various regions in Utah, including the southwest Utah region (region III, fig. F-12a) defined by Smith and Arabasz (1979) for recurrence analysis of seismicity. They first determined a moment-magnitude scale for Utah based on the spectral analysis of 19 earthquakes in the Utah region in the magnitude (M_L) range 3.7 to 6.6. Moment rates were calculated from a compilation of geologic information on slip rates of Quaternary faulting in southwest Utah. Earthquake recurrence rates were then calculated using the \underline{b} -values of Smith and Arabasz (1979) and an appropriate coefficient defining the moment-magnitude relationship. Results for southwest Utah are shown in figure F-14 (from Doser and Smith, 1982).

For southwest Utah, the frequency of earthquake occurrence based on geologically determined moment rates is essentially in agreement with that calculated from historical earthquake data (fig. F-14). The estimated recurrence interval for an earthquake of $7.0 \le M_L \le 7.5$ somewhere in the southwest Utah region is between 200 and 600 years (Doser and Smith, 1982), where the range results from an uncertainty in the moment-magnitude scale for Utah. For the same magnitude range and area, extrapolated seismicity (equations 3, 5, 6, and 7) would predict a recurrence interval between 200 and 500 years. The discrepancy between such expectation of a scarp-forming earthquake somewhere in southwestern Utah every few to several hundred years and the apparent absence of Holocene

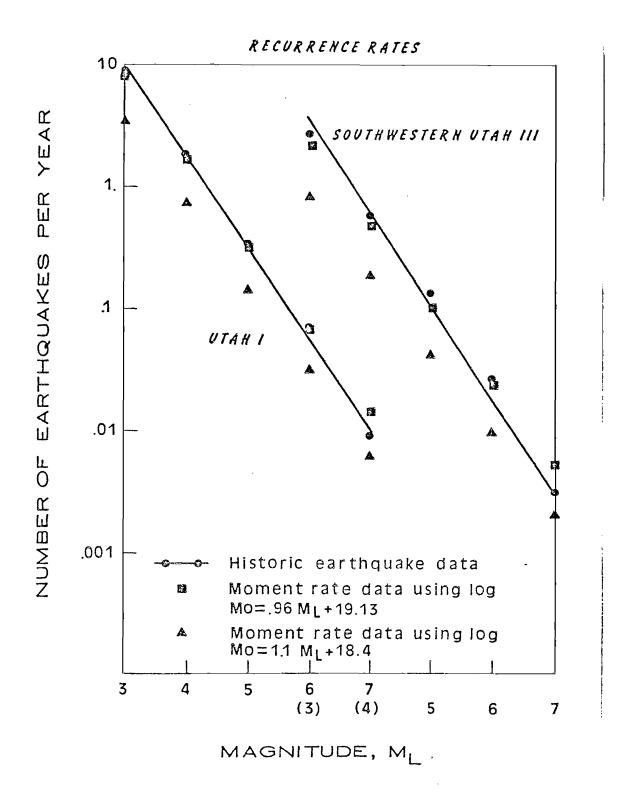


FIGURE F-14.--Comparison of earthquake recurrence data for the southwestern Utah region (as well as the entire Utah region) based on seismic moment rates and historical seismicity (from Doser and Smith, 1982).

fault scarps in southwest Utah is unresolved. Anderson (1980) and Bucknam and others (1980) consider various interpretations relevant to this problem and believe that there has been no radical change in the rate of occurrence of large earthquakes during Holocene time compared to the late Quaternary record in southwest Utah. Two implications are: (1) the recurrence interval of surface faulting on individual faults in southwest Utah may be very long, and (2) there may be different frequency-magnitude relations for earthquakes smaller than and larger than, respectively, about magnitude $6-6\frac{1}{2}$ in southwest Utah.

Greensfelder and others (1980) have used independent data to estimate moment rates for various parts of the Great Basin. The moment-rate analysis of Doser and Smith (1982) is judged more reliable for application to southwest Utah because of more careful attention to extensive local geological information and, particularly, because of the derivation of a moment-magnitude relationship specific for the Utah region--as opposed to assuming a relationship derived for California.

APPENDIX G

EMBANKMENT AND FOUNDATION CONDITIONS

Appendix G EMBANKMENT AND FOUNDATION CONDITIONS

G-1. Green's Lake Dam No. 2

Green's Lake Dam No. 2 was built in 1958. It is a zoned compacted earth fill embankment with a height of approximately 20 feet, a crest length of 1,315 feet and a total fill of about 53,700 cubic yards. The dam has an estimated total capacity of 45 ac-ft. A representative cross section of the embankment is shown in Figure G-1.

"As built" construction drawings of the embankment specify the core (Zone II) materials as compacted select fill. Construction records and corresponding laboratory tests suggest that the material used to construct the core is a calcareous sandy clayey silt or sandy silty clay with varying amounts of gravel. Similar materials were encountered in one exploratory borehole drilled through the core of the embankment during this investigation (see Appendix B). It appears that these materials were to be originally compacted to at least 95 percent relative compaction as determined by the Standard Proctor test procedure. It should be noted that the maximum particle size of the materials used in the available construction compaction tests could not be determined. In addition, the compactive effort specified for these test differs somewhat from the ASTM D698-70 testing standard. Results of some of the in situ density tests performed during construction indicate relative compactions much less than this and some of the materials which did not meet the specified compaction requirement do not appear to have been rechecked or recompacted. Relative compactions ranging from 81 to 111 percent were achieved during construction in the core materials yielding an average relative compaction of about 99 percent.

Standard Penetration Test (SPT) results obtained from the borehole drilled through the core of the embankment range from 25 to over 100 blows/foot. In most cases, however, the higher blow counts were the result of pushing large pieces of gravel or small cobbles. Nevertheless, the materials comprising the core of the embankment are probably dense and well compacted.

The Zone I material which makes up the outer shell of the embankment is specified on the "as built" construction drawing as a compacted soil, gravel and cobble fill. It appears that only one in situ density test was performed on these materials during construction. This test yielded a relative compaction of 105 percent. In situ density tests performed in the test pits excavated as part of this investigation gave relative compactions ranging from 85 to 90 percent as determined by test procedure ASTM D698-70 method C. The materials logged in the test pits were primarily silty sands and sandy silts with varying amounts of gravel and cobbles (see Appendices C and D). Based on the results of the in situ density tests, these materials are probably medium dense and are moderately compact.

The foundation at Green's Lake Dam No. 2 consists of alluvial and colluvial deposits. These deposits are stratified sands, silts, clays and gravels which contain varying amounts of gypsum. Cobbles and boulders are also commonly present. SPT blow counts taken in the boreholes drilled during this investigation range from 12 to more than 82 blows/foot with typical values around 20 to 40 blows/foot. In situ density tests performed in one test pit excavation gave an average relative compaction of approximately 92 percent. These results and the SPT blow count data indicate that the foundation soils are probably dense and are moderately compact.

The Green's Lake Dam No. 2 has performed satisfactorily since its construction. However, some minor caving of the embankment into the reservoir area has been reported to have occurred in 1980. At the time of the field investigation, some differential settlement along the axis of the embankment and some cracks which run parallel to the embankment centerline were observed. Cracks were also observed in the foundation in the vicinity of the embankment.

G-2. Green's Lake Dam No. 3

Green's Lake Dam No. 3 was built in 1958. It is a zoned compacted earth fill embankment and is similar in construction to Green's Lake Dam No. 2. The embankment has a height of approximately 17 feet, a crest length of 2,030 feet and a total fill of about 74,600 cubic yards. A representative cross section of the embankment is shown in Figure G-2.

"As built" construction drawings of the embankment show the core (Zone II) materials as compacted select fill. Records indicate that these materials are similar to those used to construct the core of Green's Lake Dam No. 2; that is, they consist of clayey sandy silts and sandy silty clays with some gravel. These types of materials were present in the exploratory borehole drilled through the core of the embankment during this investigation (see Appendix B). However, the materials appear to be generally more cohesive than the core materials of Green's Lake No. The core materials were to be originally compacted to at least 95 percent relative compaction as determined by the Standard Proctor test procedure. As in the case of Green's Lake Dam No. 2, limitations exist in the available compaction test data used in construction. In addition, some of the in situ density tests performed during construction gave relative compactions less than the specified compaction criterion. Out of a total of 9 in situ density tests performed in the core of the embankment during construction, three tests had relative compactions less than 95 percent. Relative compactions ranged from 93 to 102 percent. Standard Penetration Test (SPT) results obtained from the one borehole drilled through the core of the embankment range from 20 to 52 blows/foot. This data and the in situ density tests performed during construction suggest that the core of the embankment to be medium dense to dense, stiff to hard and moderately to well compacted.

The shell (Zone I) materials consist of a compacted soil, gravel and cobble fill. A total of five in situ density tests were performed in these materials during construction. Relative compactions from these tests range from 98 to 106 percent. Nine in situ density tests were performed in the shell materials in the test pits excavated as part of this investigation. Relative compactions obtained from

these tests range from 79 to 95 percent as determined by test procedure ASTM D698-70 method C. These tests yield an average relative compaction of 86 percent. This data suggest that the upstream and downstream shells of the embankment are probably medium dense and poorly to moderately compacted.

The foundation at Green's Lake Dam No. 3 consists of alluvial and colluvial deposits. The foundation materials are stratified sands, silts, clays and gravels which contain considerable amounts of lime and gypsum. Cobbles and boulders are also commonly present. Overall, the foundation conditions at the site are unfavorable due to the permeable soils which are high in gypsum.

SPT blow counts taken in the boreholes drilled in the foundation materials during this investigation range from 9 to over 100 blows/foot, however, the higher blow counts were recorded in the gravelly foundation soils present at depth. SPT blow counts recorded in the finer-grained foundation soils are typically around 15 to 30 blows/foot. In situ density tests performed in the foundation materials during this investigation yield relative compactions which average around 86 percent. These data indicate that the foundation soils are stiff to very stiff and medium dense.

The Green's Lake Dam No. 3 has experienced some operational problems since its construction. Subsidence was first noted in the basin area following a flood runoff which occurred during the summer of 1963. The area that subsided was circular in shape and was restricted to the east part of the reservoir adjacent to the upstream toe of the embankment. Total vertical displacement along the fractures that developed amounted to approximately 2 feet. The embankment itself did not experience any cracking or settlement and there appeared to be no seepage through the foundation.

Subsidence in the reservoir became progressively worse after each storm and in 1965 the area was repaired by placing a compacted blanket and by grading the area to provide drainage toward the outlet. In 1967, flood water was retained in the reservoir for about 3 months. This caused extensive settlement in the reservoir basin and in the dam. The reservoir basin near the east end of the dam subsided more than 5 feet and the subsided area extended through the dam. Cracks widened to depths of 15 to 20 feet by erosion and piping. Some block rotation

was also observed along the crest of the dam. The dam was considered ineffective as a flood-control structure at this time and extensive repairs had to be made to make the structure operational. Consolidation of water-sensitive soils and piping along cracks through the dam and the foundation are the suspected causes of these problems.

The cracks in the embankment and reservoir area were repaired during the spring of 1969 by grouting them with a soil-slurry mixture. In all, approximately 580 cubic yards of slurry were pumped into the cracks in the dam and reservoir bottom. The cracks in the dam were filled using about 465 cubic yards and the cracks in the reservoir bottom took 115 cubic yards. Most cracks in the dam were found to be interconnected while the cracks in the reservoir were found to be generally quite shallow and not usually connected.

At the time of the field investigation conducted as part of this study, cracks along the upstream face and transverse to the crest on the embankment were observed. Settlement along the dam crest was also noticeable.

G-3. Green's Lake Dam No. 5

Green's Lake Dam No. 5 was built in 1958. It is an uniform compacted earth fill dam with a height of 22 feet at maximum cross section, a crest length of about 235 feet and a total fill of 27,100 cubic yards. A representative cross section of the embankment is shown in Figure G-3.

The logs from one exploratory borehole and three test pits excavated as part of this investigation show that the embankment materials consist primarily of clayey silts and silty clays (see Appendices B and C). Available construction records indicate that these materials were to be compacted to at least 95 percent relative compaction as determined by the Standard Proctor test procedure. Of nine in situ density tests performed during construction of the embankment, four tests did not pass and results of one test are questionable and are probably in error. The relative compactions of the remaining four tests which passed the compaction criterion range from 95 to 107 percent. In situ density tests performed in the test pit excavated in the embankment during this investigation gave relative compactions ranging from 70 to 82 percent as determined by test procedure ASTM D698-70 method C (see Appendices C and D). Standard Penetration Test (SPT) blow count data obtained from the one borehole drilled through the centerline of the embankment range from 17 to 63 blows/foot (see Appendix B). Most of the SPT blow count data indicate, however, that the sampled embankment materials are very stiff/medium dense. Based on these data, it is prudent to assume that the embankment was not uniformly compacted during construction and zones of poorly compacted materials probably exist within the embankment. Despite this, the embankment has performed satisfactorily up to the present time. Some minor cracking was observed in the embankment, however, it shows no obvious signs of structural distress.

"As built" drawings of Green's Lake Dam No. 5 show the foundation of the embankment to consist of clayey sands, sandy clays and clayey silts overlying stratified clayey sands, sandy clays, gravel and boulders. Most of the foundation materials contain considerable amounts of lime. The available drawings show that the cutoff trench of the embankment was excavated to the lower coarse-grained soil units. The logs of the boreholes drilled during this investigation encountered

similar types of foundation materials. SPT blow counts taken in the foundation soils are typically very high (>100 blows/foot). These measurements are probably not a good basis upon which to estimate density or strength of the soil due to the high percentage of gravel present. In situ density tests were performed in test pits excavated into the finer-grained foundation soils. Relative compactions of these materials range from 84 to 91 percent indicating that these soils are medium dense and moderately compact.

G-4. Gypsum Wash Dam

Gypsum Wash dam was built in 1974 and 1975. It is a zoned compacted earth fill dam with a height of approximately 30 feet, a crest length of 3,128 feet and an estimated total fill of 238,600 cubic yards (Margheim, 1972). The dam has an estimated total capacity of 440 ac-ft. A representative cross section of the embankment is shown in Figure G-4.

According to the design report for Gypsum Wash Dam (Margheim, 1972), the materials that were to be used in the construction of the core (Zone I) consist of silty sands and sandy silts, with 20 to 65 percent passing the No. 200 sieve. Logs of exploratory boreholes drilled through the core of the embankment during this investigation show this to be generally true, however, some clayey sands and gravels were also encountered. The core materials were originally compacted to at least 95 percent relative compaction as determined by test procedure ASTM D698-70 method A (Standard Proctor). Relative compactions ranging from 95 to 105 percent were obtained from density tests performed at the time of construction. An average relative compaction of 99 percent was computed from 77 available in situ density tests. Standard Penetration Test (SPT) blow count data obtained from the boreholes drilled through the core of the embankment (see Appendix B) range from 23 to well over 100 blows/foot with values typically being about 30 to 40 blows/foot. These results, along with available construction records, show that the Zone I materials are dense to very dense and are well compacted.

According to the design report for Gypsum Wash Dam, the upstream and downstream shells (Zone III) of the embankment were to be constructed of poorly graded to silty gravels and silty sands with less than 25 percent fines. Coarser gravels were to be routed to the outside of the shell to provide protection against rilling and wind erosion. Test pits excavated in the upstream and downstream shells of the embankment during the course of this investigation show this to be generally true. The Zone III materials were compacted to at least 95 percent relative compaction as determined by test procedure ASTM D698-70 method C. Construction records indicate that relative compactions between 95 and 104 percent were obtained during construction and an average relative compaction of 99 percent was obtained from all the in situ density tests. Two in situ density

tests performed in the test pits excavated as part of this investigation have relative compactions of 85 and 88 percent. These values are not as high as those given in the construction records (see Appendix D), however, these materials were logged as being dense and compact.

The left and right abutments and foundation of the embankment consists of gypsiferous interbedded shale and siltstone which are moderately fractured and weathered. The bedrock was originally overlain with a relatively shallow veneer of gypsiferous alluvial soils consisting of low plasticity silts, clays and silty sands. This material was removed prior to the construction of the embankment. A cutoff trench was excavated into the firm shale deposits. SPT blow counts recorded in the foundation materials during this investigation range from 25 to well over a 100 blows/foot. Overall, the foundation conditions at Gypsum Wash Dam are good and the performance of the dam has been satisfactory. During the field investigation, some bulges were noted along the downstream slope of the embankment, however, there was no evidence of any settlement or excessive cracking along the crest of the embankment.

G-5. Warner Draw Dam

Warner Draw Dam was built in 1974. The embankment has a height of approximately 60 feet at its maximum cross section and a crest length of about 1,300 feet. The structure was constructed using about 164,000 cubic yards of fill material with a total capacity of approximately 1,048 ac-ft (Deming and Bridges, 1971).

"As built" drawings and construction records of Warner Draw dam indicate that it is a zoned earthfill embankment. A representative cross section of the dam is shown in Figure G-5. The Zone I core materials of the embankment consist primarily of clayey sands and clayey silty sands. The core materials were compacted to at least 95 percent relative compaction as determined by test procedure ASTM D698-70 method A (Standard Proctor). Relative compactions ranging from 95 to 107 percent were achieved during construction in the Zone I materials. An average relative compaction close to 100 percent was computed from the available construction records.

Standard Penetration Test (SPT) results obtained from the exploratory boreholes drilled through the core during this investigation (see Appendix B) were found to range from 15 to 74 blows/foot. SPT blow counts were typically within 40 to 60 blows/foot indicating that the Zone I materials are dense to very dense and are well compacted.

The upstream and downstream shells of the embankment (Zone III) are constructed of silty and clayey sands with some gravel and cobbles. The Zone III materials were also compacted to at least 95 percent relative compaction as determined by test procedure ASTM D698-70 method C. Relative compactions ranging from 95 to 101 percent were achieved during construction and an average relative compaction of 98 percent was computed from the available construction records.

The SPT blow counts obtained in the Zone III materials during this investigation averaged about 50 to 60 blows/foot indicating that these materials are dense to very dense. In situ density tests performed in a test pit excavated near the

crest of the embankment (see Appendices C and D) verified the dense nature of the Zone III embankment materials.

Foundation conditions at the Warner Draw Dam site are generally good. The right abutment of Warner Draw Dam consists primarily of sandstone with some interbedded siltstone and shale. The left abutment consists of alluvial materials. Logs of boreholes available in the geologic report for Warner Draw Dam (Deming and Bridges, 1971) indicate that the alluvial materials are very dense, interbedded sands, silty sands and some clayey sands. These soils are weakly to moderately cemented with lime and overlay shale and sandstone bedrock. The foundation of the dam embankment near its maximum cross section consists of weathered siltstone and shale.

G-6. Stucki Dam

Stucki Dam was built in 1974. The embankment has a height above the surrounding ground surface of approximately 30 feet. However, because of the poor foundation conditions that were anticipated from geologic investigations, deep foundation excavations were made and the actual filled embankment height is about 45 feet. The crest length of the embankment is about 1,400 feet. A preliminary estimate of the amount of fill material required to construct the embankment is approximately 63,000 cubic yards. The estimated capacity of the reservoir is about 126 ac-ft.

Stucki Dam is a zoned earthfill embankment. A representative cross section of the embankment is shown in Figure G-6. Exploratory boreholes drilled as part of this investigation in the Zone I core of the embankment indicate that it consists of silty sands, clayey silty sands and clayey sands. These materials were compacted to at least 95 percent relative compaction as determined by test procedure ASTM D698-70 method A (Standard Proctor). Relative compactions ranging from 95 to 103 percent were achieved during construction in these materials yielding an average relative compaction of about 99%. Standard Penetration Test (SPT) results obtained from the boreholes drilled within the core as part of this investigation range from 20 to 84 blows/foot with typical values on the order of 30 to 50 blows/foot (see Appendix B). These results, along with construction records, indicate that the Zone I core materials are very stiff to dense and are well compacted.

The upstream and downstream shells of the embankment (Zone III) consist primarily of silty sands which contain varying amounts of gravel and cobbles. The Zone III materials were compacted to at least 95 percent relative compaction as determined by test procedure ASTM D698-70 method C. Available construction records indicate relative compactions ranging from 96 to 105 percent were achieved and an average relative compaction of about 99% was calculated for these materials.

The SPT blow counts obtained in the Zone III materials during this investigation averaged about 40 to 50 blows/foot. Relative compactions obtained from in situ density tests performed in test pit excavated along the upstream shell of the

embankment were found to range from 93 to 106 percent (see Appendices C and D). These results are in good agreement with the available construction records indicating that the Zone III materials are dense to very dense and are well compacted.

The geologic report for Stucki Dam reported that geologic conditions at the site are poor (Deming and Bridges, 1971). Pervious, gypsiferous alluvial and colluvial deposits were found to underlie the foundation and right abutment. Soft, fractured silty sandstones were also discovered in the left abutment. In order to prevent settlement and piping in the foundation, nearly all of the highly gypsiferous and pervious deposits were removed prior to the construction of the embankment and rather deep cutoff excavations were constructed. These measures appear to have yielded a satisfactory design since the embankment does not show any significant signs of distress at the present time. Logs of boreholes available in the geologic report indicate that the foundation of the embankment primarily consists of sand, silty sands and clayey sands. Some of these deposits contain varying amounts of gypsum. SPT blow counts indicate that the foundation soils are dense to very dense. SPT blow counts data and in situ density tests performed during this investigation generally agree with the findings of previous investigations.

G-7. Frog Hollow Dam

Frog Hollow Dam was constructed in two stages. The original embankment was constructed in 1956 and had a height of approximately 33 feet at maximum cross section and a crest length of 600 feet. Major portions of the old embankment were removed and used as construction materials in the 1978 reconstruction of the dam. The reconstruction effectively raised the old dam about 16 feet. The present dam has a total height of 48 feet at the maximum cross section and a crest length of 1,900 feet.

"As built" drawings and construction records of Frog Hollow dam indicate that the portion of the embankment built in 1978 consisted of zoned earth fill. A representative cross section of the embankment is shown in Figure G-7.

No information was available at the time of writing of this report on the composition and geometry of the original embankment. Logs of three exploratory boreholes drilled along the crest of the embankment and four test pits excavated during this investigation (see Appendices B and C) show that the Zone I (see Figure G-7) materials consist of sandy and silty clays, sandy silts and silty sands with varying amounts of gravel and a few scattered cobbles. The Zone I materials in the portion of the embankment constructed in 1978 were compacted to at least 95 percent relative compaction as determined by test procedure ASTM D-698 method A (Standard Proctor). Relative compactions ranging from 94 to 128 percent were achieved during construction in the Zone I materials. A number of in situ density tests performed during construction yielded relative compactions less than the specified compaction criterion (i.e., 95 percent relative compaction). For these tests, it appears that additional compactive effort was applied to these lifts before additional fill was placed.

Standard Penetration Test (SPT) results obtained from one exploratory borehole drilled through the Zone I materials during this investigation range from 26 to 48 blows/foot with typical values being around 30 blows/foot. Logs from three exploratory boreholes drilled through the Zone I materials during this investigation do not indicate any changes in material type or consistency at the depths where the original embankment materials (i.e., the embankment constructed in 1956) should have been encountered. It was discovered after the completion of the field exploration program that the borings drilled along the crest were located

in the vicinity of the old corrugated metal pipe which was part of the principal spillway of the original embankment. The original embankment was removed in this area in order to remove the pipe and was probably backfilled with Zone I fill materials. Therefore, materials comprising the original embankment were never encountered in any of the boreholes. The log from one shallow borehole drilled along the crest of the original embankment during the geologic investigation for the new embankment (Deming, 1976) indicates that the old embankment materials consists of gravelly silty sand which contains some cobbles.

In situ density tests performed in one test pit excavated in the Zone I materials of the new embankment during this investigation yield relative compactions of 95 and 100 percent (see Appendices C and D) as determined by test procedure ASTM D-698 method C (Standard Proctor). These results and the SPT blow count data suggest that the Zone I materials are very stiff to hard and are well compacted.

Construction records and the "as built" construction drawings of the embankment do not specify the types of materials used to construct the downstream shell (Zone II). The log from one test pit excavated during this investigation shows that the shell material consists of coarse gravel and cobbles in a silty clay matrix. Two in situ density tests performed in the test pits gave relative compactions of 79 and 92 percent. Construction records suggest that in situ density tests were not performed in the Zone II shell materials during construction. The materials were probably compacted using the same techniques as were used to place the Zone I embankment materials. Based on the limited amount of available data, it is prudent to assume that the Zone II shell materials are probably stiff/medium dense and are moderately compact.

The geologic report for the Frog Hollow dam site indicates that the site is geologically a poor location for a flood control structure (Deming, 1976). Most of the site's foundation consists of basalt flows and gypsiferous unconsolidated alluvial and colluvial deposits. The left abutment of the dam is formed by two basalt flows which are separated by a three to five foot thick deposit of highly compacted sandy silt which can be classified locally as a siltstone. Approximately 10 to 20 feet of alluvium originally covered the basalt flows at the left abutment, however, most of this material was removed in the cut-off and drain trench excavations.

The basalt flows are dark gray to black in color, fractured, very fine-grained and highly vesicular along the top and bottom cooling zones. The permeability of the rock is high in the fractured zones and along most joints. Piping of fine-grained embankment materials and the materials separating the two basalt flows was considered a possibility because of the highly permeable basalt. Construction drawing of the portion of the embankment built in 1978 show that a protective blanket of Zone I embankment materials was placed along the left abutment to prevent piping.

The younger basalt flow is absent in the right abutment of the dam, however, the upper cooling zone of the older basalt flow was found to be highly permeable. The inner portion of the flow (below the upper cooling zone) is dense and lightly fractured (Deming, 1976) with small nonconnecting vesicles. The joints in this portion of the basalt are generally tight with some joints being filled with silt and clay. The right abutment basalt was overlain by approximately 5 to 10 feet of calcareous, gypsiferous, sandy and clayey alluvium, however, this material was removed prior to the construction of the embankment. Piping between the vesicular, upper cooling zone found in the right abutment and the fine-grained embankment materials was considered a possibility at the time of the geologic investigation.

The valley foundation of the embankment is a narrow channel cut in the older (deeper) basalt flow. Both abutments slope steeply towards the channel foundation. Erosion has removed the upper cooling zone of the flow and the basalt remaining in the valley is only "lightly fractured and intensely (Deming, 1976) weathered". Joints in the rock were reported to be generally tight and permeability rates were low.

Frog Hollow dam has experienced some cracking problems since the construction of the raised portion of the embankment in 1978. In early 1981, the cracks were investigated by SCS by digging a series of trenches along the centerline of the embankment. The cracks that were noted were mostly transverse to the centerline of the embankment and ranged from 3 to 9 feet in depth. Many cracks were found to extend through the entire embankment fill. The width of the cracks ranged from hairline fractures to $1\frac{1}{2}$ inches and averaged about $\frac{1}{2}$ inch. All cracks were

open at the surface and gradually narrowed at depth. Many of the cracks were found to be filled with grass and roots.

In addition to the transverse cracking of the embankment, one longitudinal crack was also found. It extended from about station 11+00 to 12+00 and was 35 feet upstream of the centerline. Additional longitudinal cracking was noted along the upstream face of the embankment during the field investigation conducted as part of this study. No explanation has been offered in the available studies that have been conducted to date as to the cause of this cracking. However, "as built" construction drawings of the portion of the embankment built in 1978 show that the cut-off trench in the vicinity of the cracks was probably placed in old debris basin deposits and/or alluvial and colluvial deposits. Consolidation of these materials could, therefore, be a possible cause of these cracks.

During the drilling operations conducted as part of this investigation, a significant zone of apparently low density materials was encountered in the interval between $50\frac{1}{2}$ feet to 55 feet of borehole FH-1. This hole was drilled to a depth of $50\frac{1}{2}$ feet on the first day of the drilling operations at the dam site. The drill rods were pulled out of the hole at the end of the day and the water level in the hole was very near the surface. At the beginning of the next day, there was no water in the hole to the depth of $50\frac{1}{2}$ feet. Drill rods were let down to this depth and subsequently "dropped" to a depth of 55 feet. No circulation was established below a depth of $50\frac{1}{2}$ feet and no soil samples were obtained between $50\frac{1}{2}$ to 55 feet. A Pitcher barrel soil sample was obtained from 55 to 56.8 feet and was found to be a sandy silt/sandy clay.

Because of the anomalous water loss and the apparently very weak soils encountered, two additional boreholes were drilled on either side of boring FH-1. The soils encountered in the two additional boreholes at depths comparable to the anomalous water loss zone of boring FH-1 consisted of sandy clay/sandy silt. SPT blow counts taken in these soils were somewhat less than those recorded in the surrounding soils.

The anomalous water loss experienced in borehole FH-1 could be the result of several factors. First, the 50 feet of hydraulic head acting on the soils overnight could have softened them to the extent that they could not be sampled. Second,

borehole FH-1 was drilled at the approximate location of the old 24 inches corrugated metal pipe which was part of the principal spillway structure of the original embankment, as was previously noted. This pipe was removed and the trench was backfilled prior to the construction of the new embankment raise. It is possible that the backfill of the trench was poorly placed resulting in the weak zone of soils encountered in borehole FH-1.

The causes of the embankment cracking of Frog Hollow dam and remedial measures to repair the embankment are presently being investigated by the SCS.

G-8. Ivins Diversion Dam No. 5

The Ivins Diversion Dam No. 5 was built in 1977. The embankment is an uniform compacted earth fill dam with a crest length of about 5,300 feet and a height of 20 feet at its maximum cross section. A representative cross section of the embankment is shown in Figure G-8.

Information obtained from the exploratory boreholes and test pits excavated as part of this investigation suggest that some attempt might have been made during construction to place the finer-grained fill materials near the centerline of the embankment. The logs of three exploratory boreholes drilled along the crest of the embankment indicate that the center of embankment consists of silty sands and sandy silts containing some medium to coarse sand and fine gravel. Standard Penetration Tests (SPT) performed in these materials range from 33 to 90 blows/foot (see Appendix B) indicating that these materials are dense to very dense.

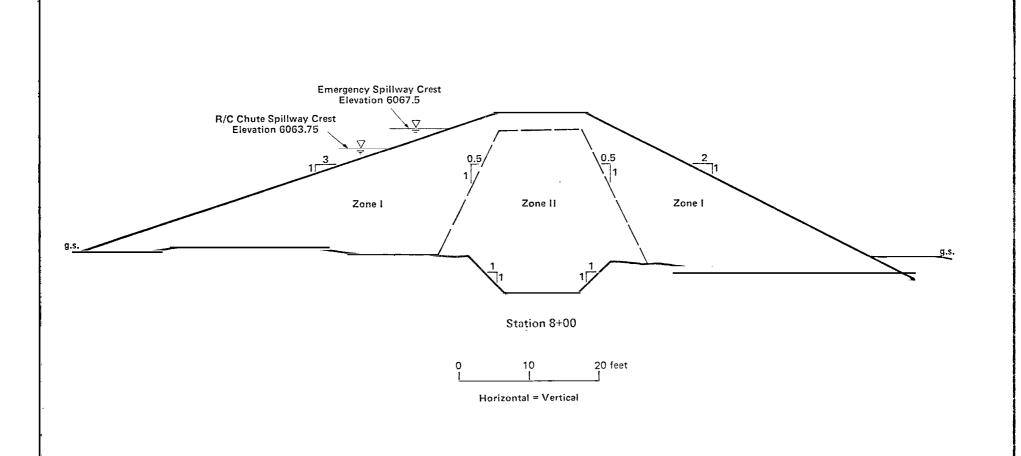
Logs of the test pit excavations show that the materials near the slopes of the embankment are sandy silts and silty sands with varying amounts of coarse sand, gravel and scattered cobbles. Three in situ sand core density tests performed in these materials during this investigation show that they are generally dense and have a relative compaction averaging about 97 percent (see Appendices C and D) as determined by test procedure ASTM D-698 method C (Standard Proctor). In situ density tests performed at the time of construction of Ivins Diversion Dam No. 5 are limited. However, they do show that the embankment was compacted to at least 95 percent relative compaction as determined by test procedure ASTM D698-70 method A.

The geologic report for the Ivins site (Bridges, 1972) states that the site is geologically a poor location for this type of flood structure. The foundation and abutments are underlain by gypsiferous, pervious, water-sensitive soils and weathered, gypsiferous bedrock. Vugs and small solution channels were reported in the upper part of the bedrock which consists of interbedded siltstones and sandstones. The foundation and abutments are reported to be subject to settlement

and/or piping because of the gypsiferous water-sensitive soils and bedrock. The geologic report recommended that the loose materials in the foundation should be removed or consolidated to prevent excessive settlement and that a cutoff should be dug into the bedrock to prevent possible piping.

Available advanced copies of the "as built" drawings of the dam suggest that the location of the centerline and the configuration of the embankment were modified several times from those reported in the geologic report. The preconstruction drawings show a cutoff trench only 4 feet deep. The available drawings do not show whether or not the cutoff trench was excavated into the bedrock.

Logs of the exploratory boreholes and test pits excavated during this investigation indicate that the foundation soils consist of medium dense to very dense sandy silts and silty sands which contain varying amounts of gypsum. SPT blow counts range from 14 to over 80 blows/foot with typical values around 40 to 50 blows/foot. In situ density tests performed in two test pit excavations yield relative compactions ranging from 82 to 95 percent. Siltstone bedrock was encountered in two of the exploratory boreholes and was found to be deeply weathered and to contain abundant gypsum. Despite the poor geologic conditions, only some minor cracks and erosion in the embankment was observed at the time of the field investigation.

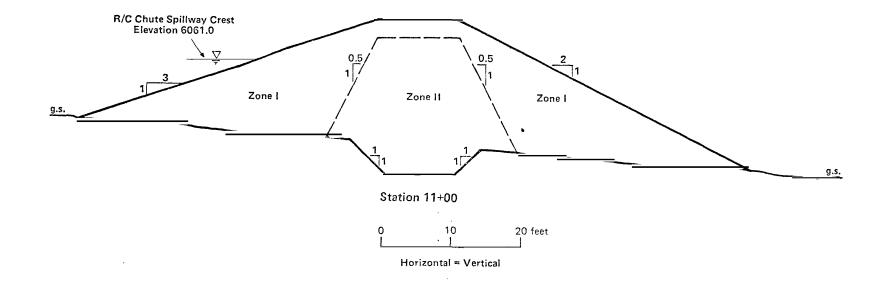


Zone	Material
i	Compacted soil, gravel and cobble fill.
11	Clayey silt, sandy silty clay with some gravel.
Note:	g.s. indicates approximate ground surface.

Palo Alto, California

SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS REPRESENTATIVE EMBANKMENT CROSS SECTION GREEN'S LAKE DAM NO. 2

Checked by MAT Date	/21/8	2 Project No.	Figure No.
Approved by Date 2	MAY 8	Z D118	G-1

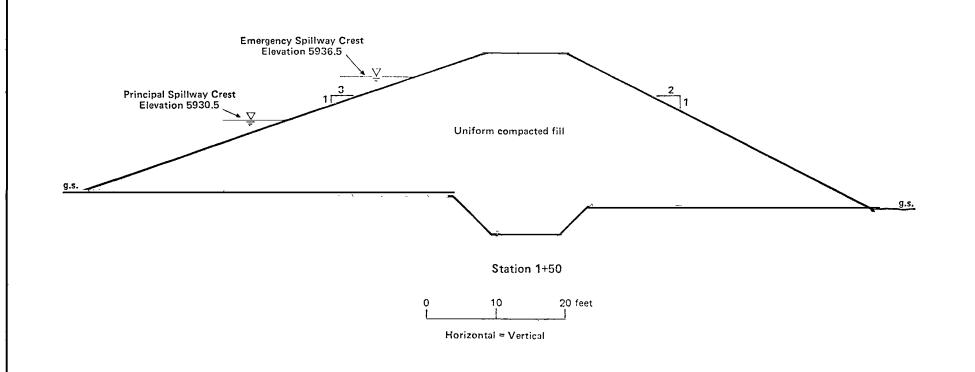


Zone	<u>Material</u>
ı	Compacted soil, gravel and cobble fill.
11	Clayey sandy silt, sandy silty clay with some gravel.
Note:	g.s. indicates approximate ground surface

Palo Alto, California

SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS REPRESENTATIVE EMBANKMENT CROSS SECTION GREEN'S LAKE DAM NO. 3

	Checked by MLT Date 5/27/82	Project No.	Figure No
_	Checked by MUT Date 5/27/82 Approved by Awley Date 7/MAY 82	D118	G-2



Material

Uniform compacted fill consists of clayey silt and silty clay with trace gravel.

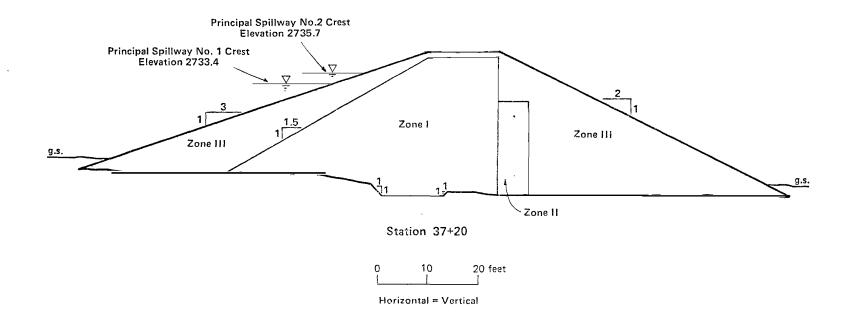
Note: g.s. indicates approximate ground surface.

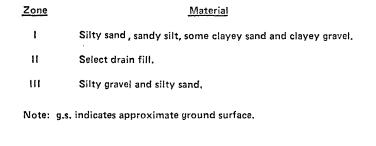
Earth Sciences Associates

Palo Alto, California

SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
REPRESENTATIVE EMBANKMENT CROSS SECTION
GREEN'S LAKE DAM NO. 5

Checked by MUT Date Approved by Date	5/27/87	Project No.	Figure No.
Approved by A William Date	JMAY 82	D118	G-3

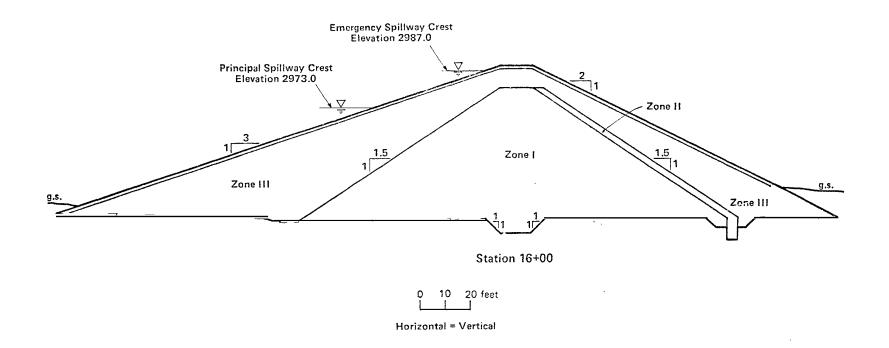




Palo Alto, California

SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
REPRESENTATIVE EMBANKMENT CROSS SECTION
GYPSUM WASH DAM

Checked by MyT Date 5/27/82	Project No.	Figure No.
Checked by MCT Date 5/27/82 Approved by Date 3/27/82	D118	G-4



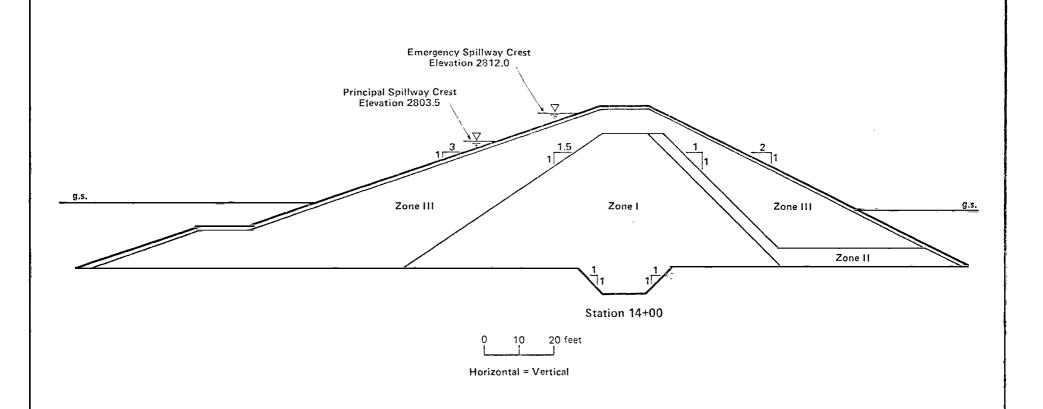
Zone	<u>Material</u>
1	Clayey sand and clayey silty sand.
н	Select drain fill,
111	Silty sand and clayey sand with gravel and cobbles.
Note:	g.s indicate approximate ground surface.

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SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
REPRESENTATIVE EMBANKMENT CROSS SECTION
WARNER DRAW DAM

Checked by Mary Date	5/27/82	Project No.	Figure No.
Checked by Approved by Approved by	ZFNYY8Z	D118	G-5

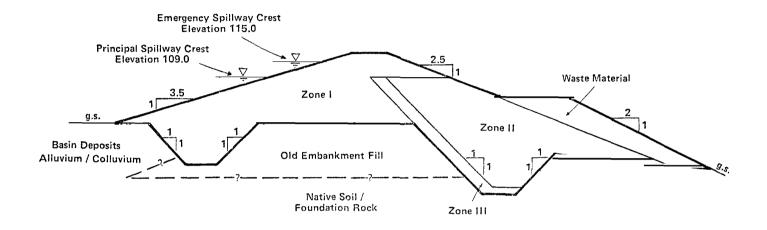


Zone	<u>Material</u>
1	Silty sand, clayey silty sand and clayey sand.
11	Select drain fill.
Ш	Silty sand with gravel and cobble.
Note:	g.s indicates approximate ground surface.

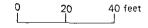
Palo Alto, California

SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
REPRESENTATIVE EMBANKMENT CROSS SECTION
STUCKI DAM

С	necked by MUTA	_ Date 5/z	7/82	Project No.	Figure No.
A	pproved by A Welson	_ Date271	SERA	D118	G-6







Horizontal = Vertical

Zone	<u>Material</u>
1	Sandy clay, silty clay and silty sand with varying amounts of gravel and scattered cobbles.
11	Coarse gravel and cobbles in a silty clay matrix.
ш	Select drain fill.
Old Embankment Fill.	Graveley silty sand with some cobbles.
Note: g.s indi	cates approximate ground surface.

Palo Alto, California

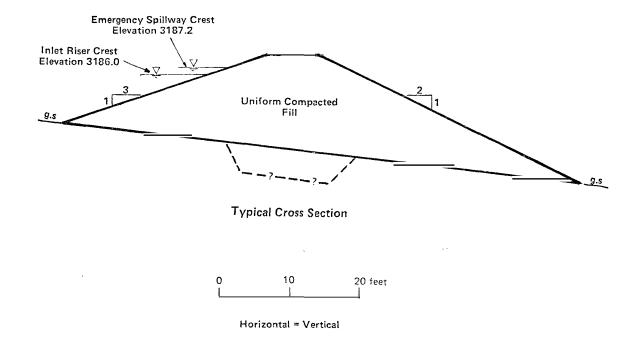
SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS REPRESENTATIVE EMBANKMENT CROSS SECTION FROG HOLLOW DAM

Checked by ALTICAL Approved by

Date 7/82 Project No. Figure No. Date 7 MAY 82

D118

G-7



Material

Uniform compacted fill consists of silty sand and sandy silt with varying amounts of gravel and scattered cobbles.

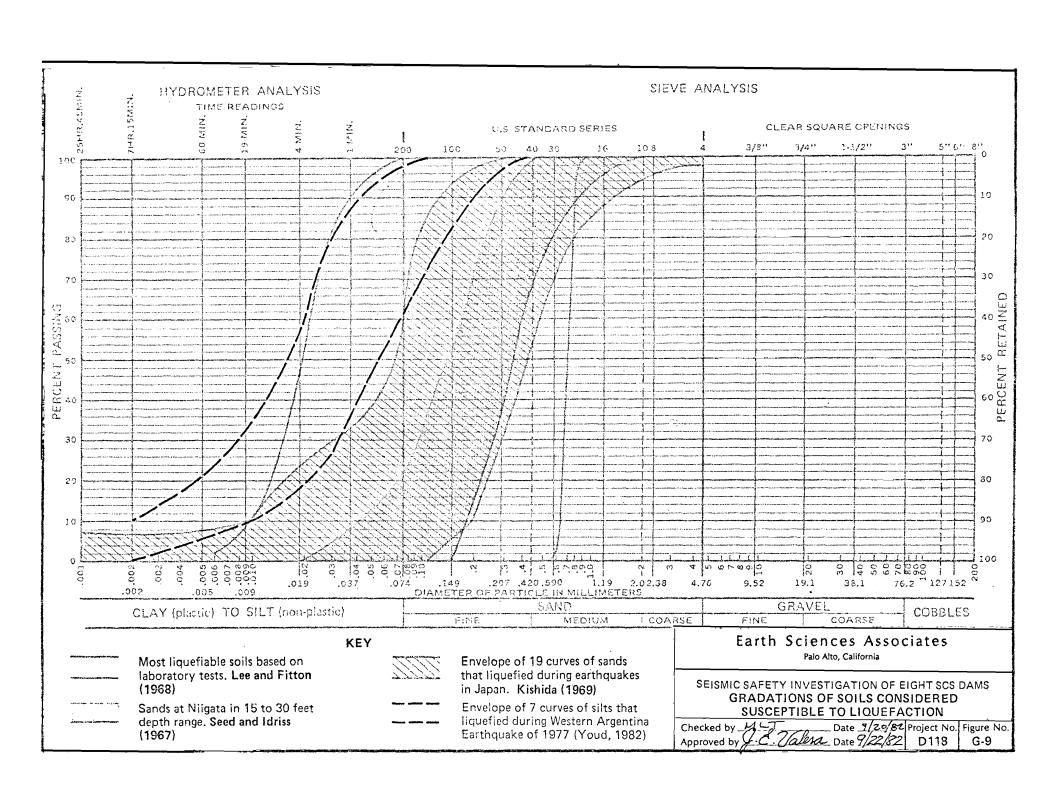
Note: g.s indicates approximate ground surface,

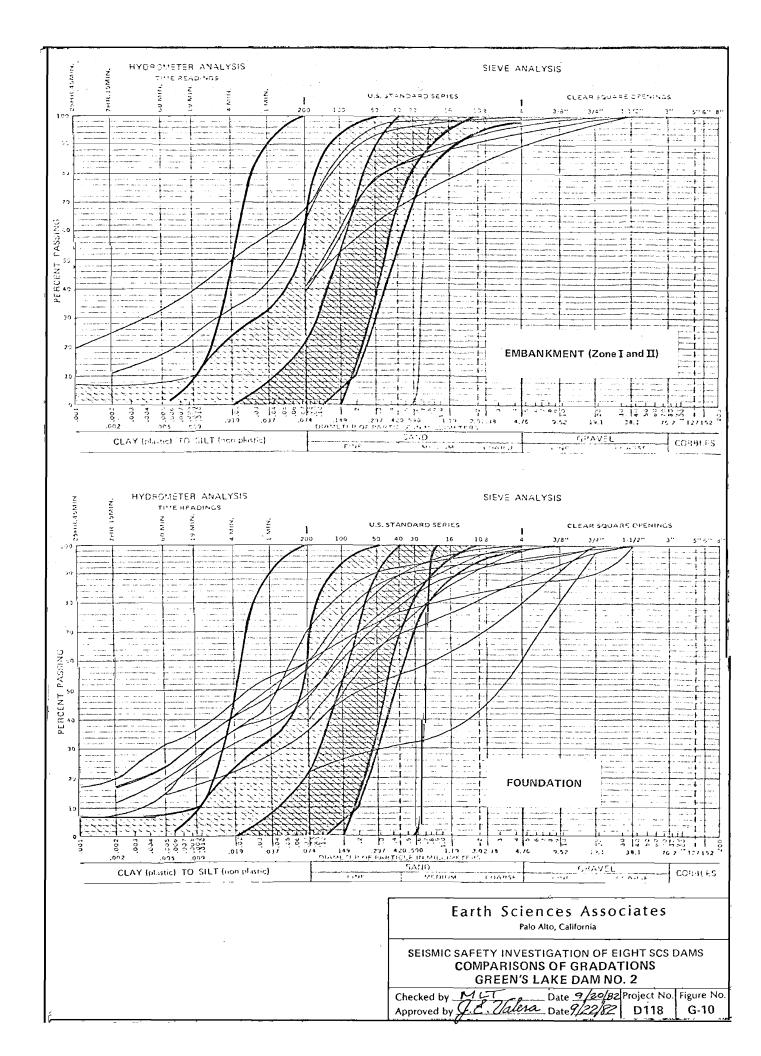
Earth Sciences Associates

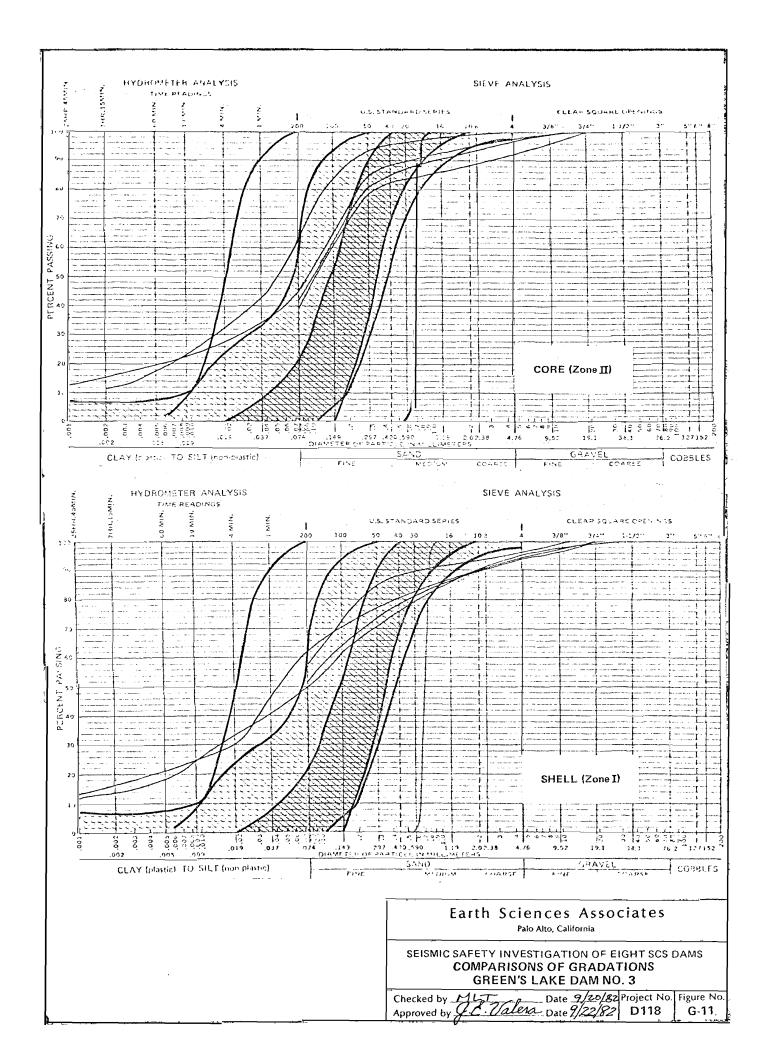
Palo Alto, California

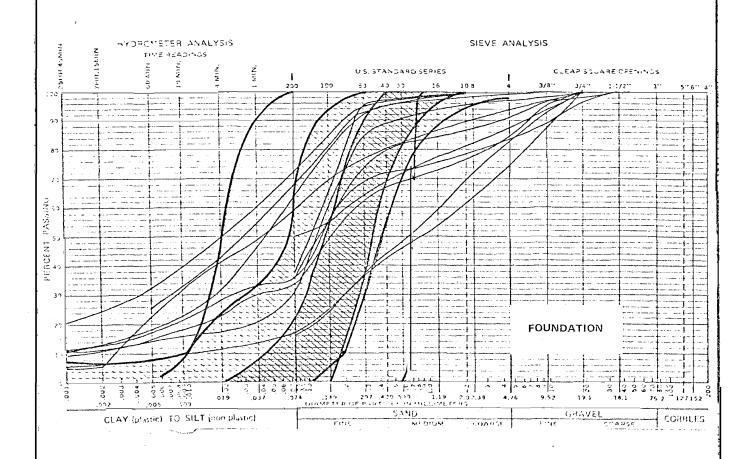
SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS REPRESENTATIVE EMBANKMENT CROSS SECTION IVINS DIVERSION DAM NO. 5

Checked by MUT Date5/27/82 Project No. Figure No. Approved by A Julian Date7/MAY 82 D118 G-8





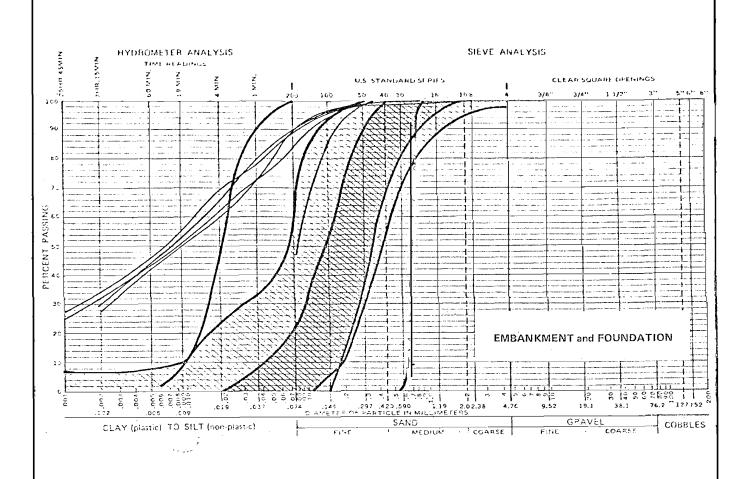




Palo Alto, California

SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
COMPARISONS OF GRADATIONS
GREEN'S LAKE DAM NO. 3

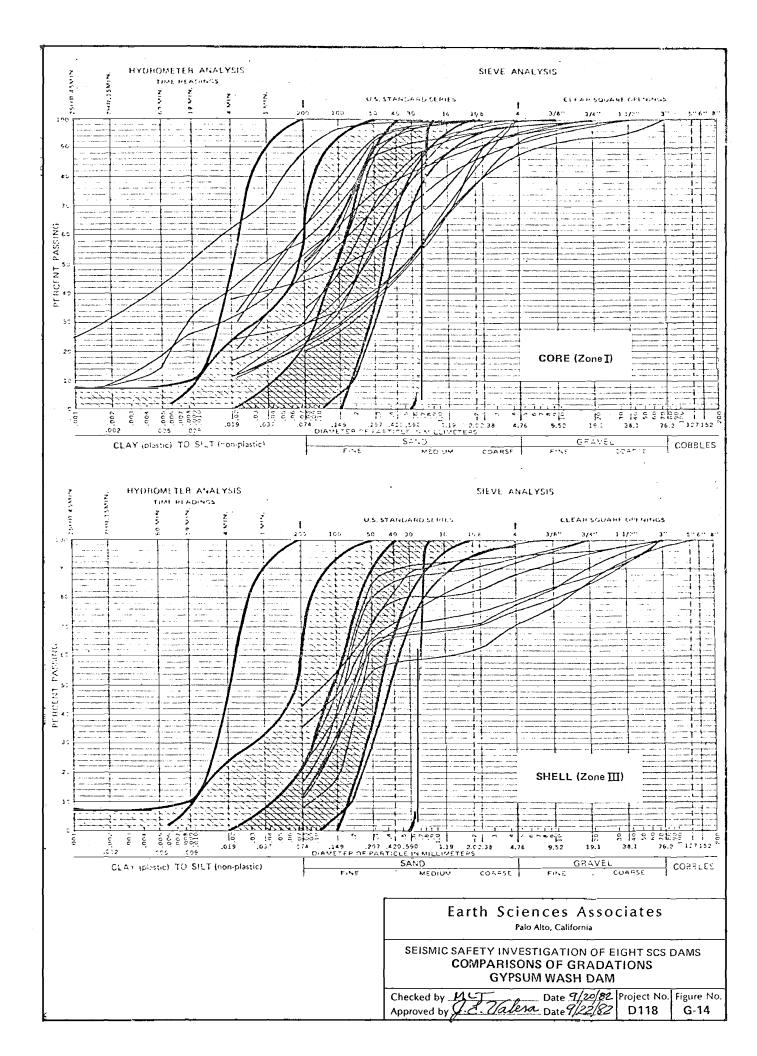
Checked by MLT Date 9/20/82 Project No. Figure No. Approved by C-C-Valera Date 9/22/82 D118 G-12

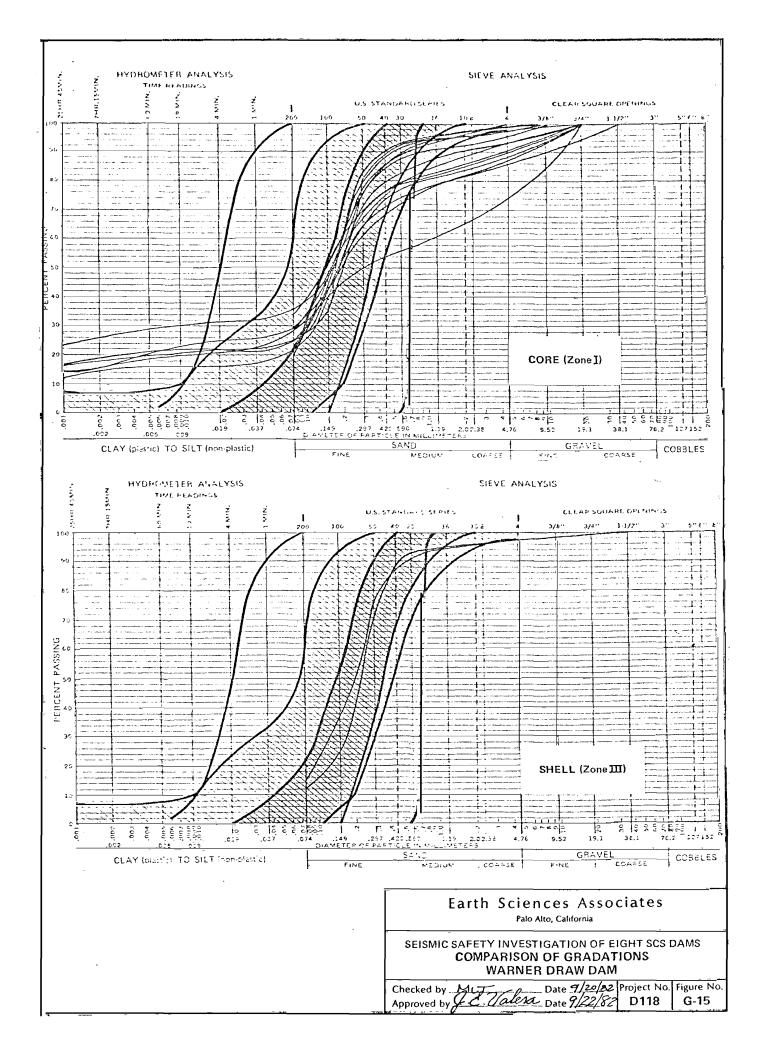


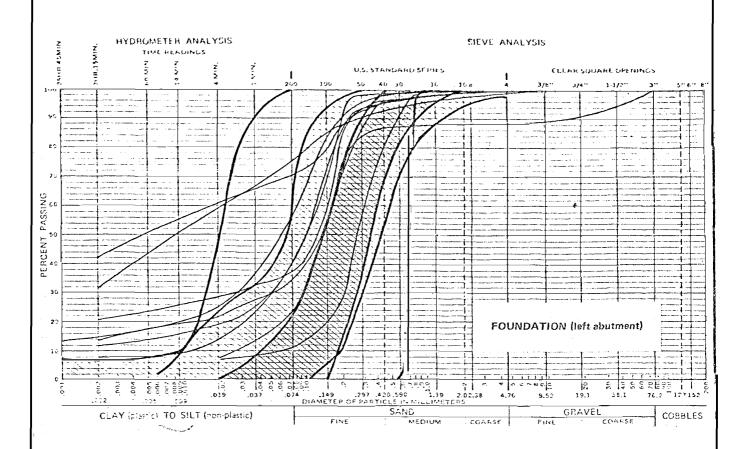
Palo Alto, California

SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
COMPARISONS OF GRADATIONS
GREEN'S LAKE DAM NO. 5

Checked by MCT Date 9/20/82 Project No. Figure No. Approved by C.C. Valena Date 9/22/82 D118 G-13



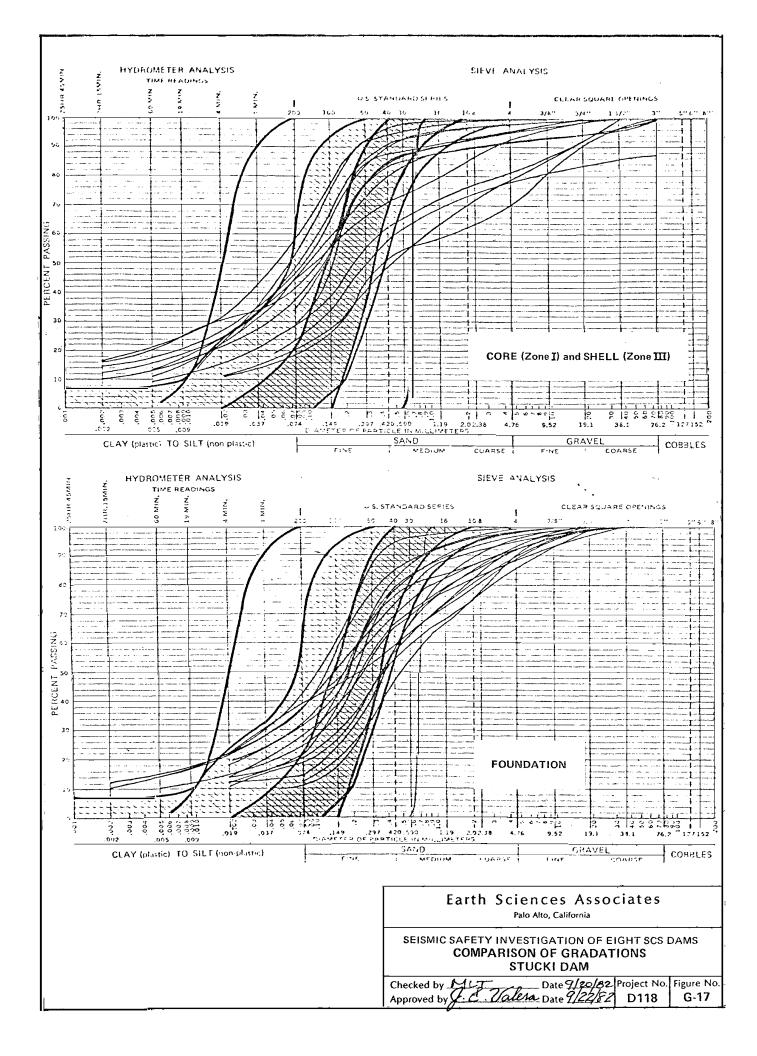


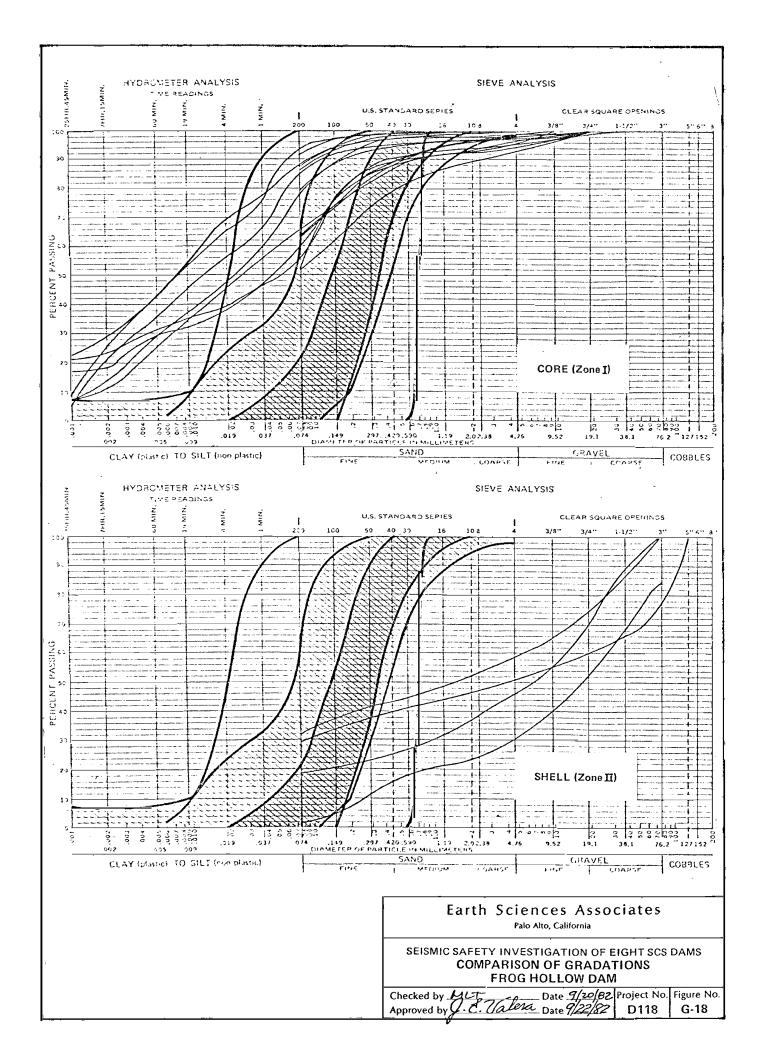


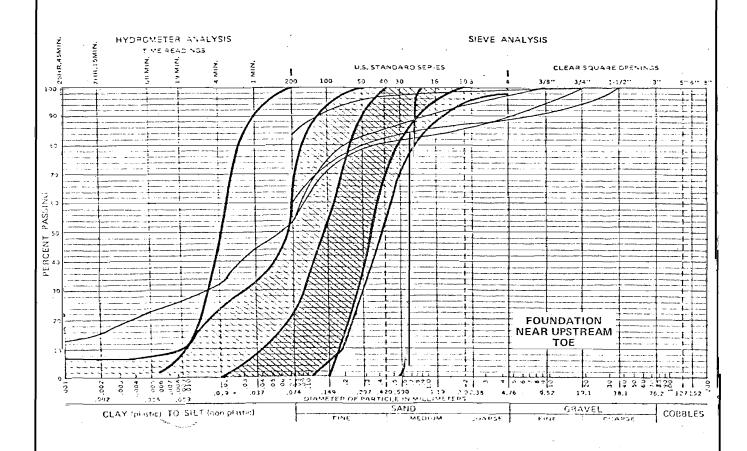
Palo Alto, California

SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
COMPARISON OF GRADATIONS
WARNER DRAW DAM

Checked by MLT Date 9/1982 Project No. Figure No. Approved by J.C. Walesa Date 9/2982 D118 G-16



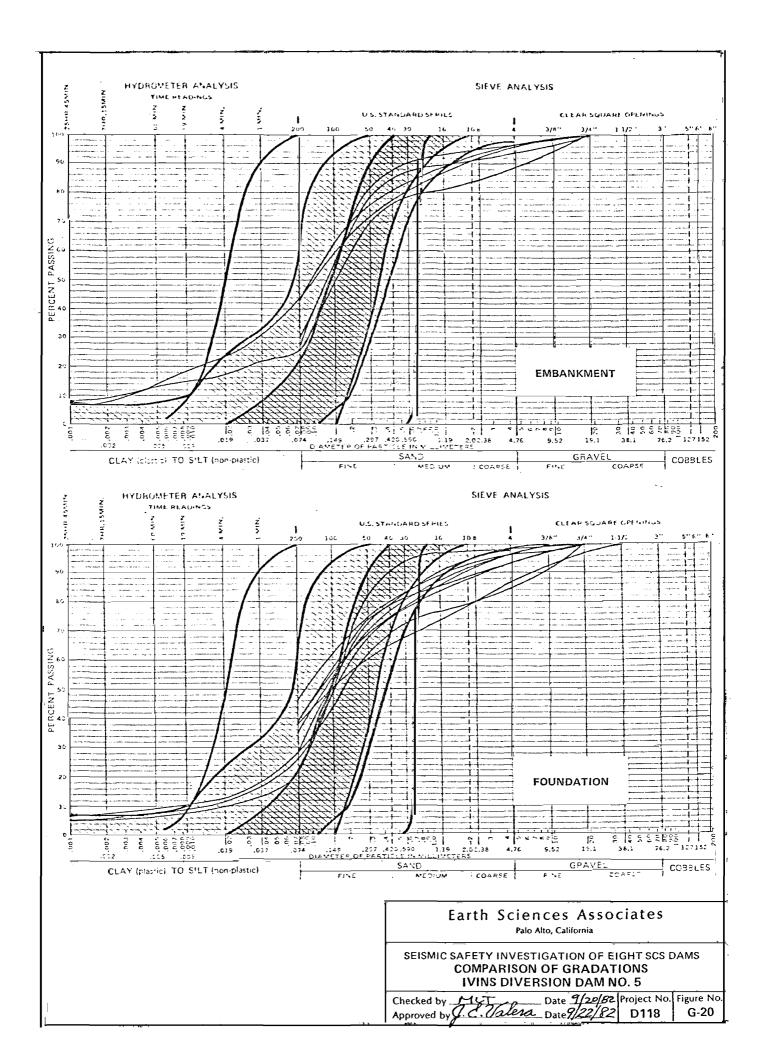




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SEISMIC SAFETY INVESTIGATION OF EIGHT SCS DAMS
COMPARISON OF GRADATIONS
FROG HOLLOW DAM

Checked by MLT Date 9/20/82 Project No. Figure No. Approved by J. C. Valera Date 9/22/82 D118 G-19



APPENDIX H

LISTING OF EARTHQUAKES OF $M \ge 3$ WITHIN 200 KM OF LAT. 30° 25' N, LONG 113° 15' W (1850-1981)

UTAH EARTHQUAKE CATALOG 1850 - June 1962 Compiled by Walter J. Arabasz and Mary E. McKee

(EXPLANATION)

This catalog chronologically lists all earthquakes in the Utah region, 36°45'N - 42°30'N and 108°45'W - 114°15'W, that have been identified since 1850, the year of publication of the first newspaper in Utah, through June 1962. The following data are listed for each event:

- Year (YR), date, and origin time (ORIG TIME) in Universal or Greenwhich Mean Time (GMT). "Local time" is 7 hours earlier than GMT (i.e., local time = GMT 7 hours) except for three time periods when it was 6 hrs earlier: (1) 02:00 Mar. 31, 1918 02:00 Oct. 27, 1918; (2) 02:00 Mar. 30, 1919 02:00 Oct. 26, 1919; and (3) 02:00 Feb. 9, 1942 02:00 Sept. 30, 1945. Origin time given in hours and minutes for non-instrumental locations, and in hours, minutes, and seconds for instrumental locations.
- 2. Earthquake location coordinates in degrees and minutes of north latitude (LAT-N) and west longitude (LONG-W). For non-instrumental locations, epicenter is assumed; in most cases, assigned coordinates correspond to location of town or city where felt effects were strongest. Epicentral accuracy ≅ ±25-50 km.
- 3. MAG, estimated Richter magnitude determined in one of four ways, as indicated by a suffix: (1) I implies estimate from maximum Modified Mercalli Intensity (INT) assuming the Gutenberg-Richter relation (Gutenberg and Richter, 1956, Bull. Seism. Soc. Am. 46, 105-145): MAG = 1 + 2/3 (INT); (2) M implies magnitude determined by Seismological Laboratory in Pasadena, (3) N implies magnitude estimated by University of Nevada (Reference 5, see below); and (4) X implies value arbitrarily assumed for event of unidentified size; X = 2.3 for non-instrumental locations, and 3.0 for instrumental locations.
- 4. INT, maximum Modified Mercalli Intensity. Unless otherwise noted, intensity is from Reference 1 (see below) for earthquakes through 1949, and from Reference 8 (see below) thereafter. Where sources disagree on maximum intensity, range is indicated as a comment and a maximum value has been interpreted. For events of unidentified size (X suffix in MAG column), intensity II arbitrarily assumed for non-instrumental locations—and no intensity assigned for instrumental locations.
- 5. Comments: Compilation of the 1850 1962 catalog has involved the careful checking and correlation of numerous sources--and extensive annotation. For convenience, several abbreviations and numbered references have been used, as outlined below. Earthquakes without comments generally are from Reference 1 for 1850-1949, and from either Reference 6 (instrumental) or Reference 8 (non-instrumental) for 1950-1962.

Abbreviations

LOC:	location	ASSGN:	assigned
INT:	intensity (Modified Mercalli)	A'SHOCK:	aftershock
PAS:	Pasadena, Seismological Laboratory	F'SHOCK:	foreshock
NEV:	University of Nevada, Reno	SALT LK:	Salt Lake
ID:	Idaho	MAG:	magnitude
UT:	Utah		

References and Footnotes

- (1) Williams, J. S. and M. L. Tapper (1953). Earthquake history of Utah, 1850-1949, Bull. Seism. Soc. Am. 43, 191-218.
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- (3) Townley, S. D. and M. W. Allen (1939). Descriptive catalog of earth-quakes of the Pacific Coast of the Unites States, 1769 to 1928, Bull. Seism. Soc. Am. 29 (1), Ch. 4, 5, 6.
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- (18) Stansbury, H. (1852). An Expedition to the Valley of the Great Salt Lake of Utah, Lippincott, Grambo and Co., 487 pp. (see pp. 149-150).
- Note: Coordinates listed in catalog for non-instrumental locations correspond arbitrarily to <u>center</u> of town or city where felt effects were strongest. Decimal-point accuracy is by no means implied for the earthquake epicenters.

UTAH EARTHQUAKE CATALOG: JULY 1962 - PRESENT (EXPLANATION)

Utah earthquakes for July 1, 1962 through present are listed in chronological order. The following data are listed for each event:

- 1. Year (YR), date and origin time in Universal Coordinated Time (UTC). Subtact seven hours to convert to Mountain Standard Time (MST).
- 2. Earthquake location coordinates in degrees and minutes of north latitude (LAT-N) and west longitude (LONG-W), and depth in kilometers. "*" indicates depth restricted to the initial trial depth due to poor depth resolution.
- 3. MAG, computed local magnitude for each earthquake. "W" indicates original Wood-Anderson records used. "M" indicates a magnitude less than 1.5.
- 4. NO, number of P, S and S-P readings used in solution.
- 5. GAP, largest azimuthal separation in degrees between recording stations used in the solution.
- 6. DMN, epicentral distance in kilometers to the closest station.
- 7. RMS, root-mean-square error in seconds of the travel-time residuals:

$$RMS = \left(\sum_{i} \left(W_{i} R_{i}^{2}\right) / NO\right)^{\frac{1}{2}}$$

where:

- R_i is the observed minus the computed arrival times of P, S, or S-P data at the i-th station,
- W_i is the relative weight given to the i-th station (0.0 for no weight through 1.0 for full weight) for the type of data (P, S, or S-P).
- NO as described above.

- 8. Q, quality class of the hypocenter.
 - a) July 1962 through September 1974.

Quality	NO	RMS	ERH
A B C D	6 6 6 Others	0.25 0.50 0.75	5.0 10.0 15.0

b) October 1974 through present--Q is the average of S and D defined as follows:

<u>S</u>	RMS	ERH	ERZ
A B C D	0.15 0.30 0.50 Others	1.0 2.5 5.0	2.0 5.0
\overline{D}	<u>NO</u>	GAP	DMN
A B C D	6 6 6 Others	90° 135° 180°	Depth or 5 km 2xDepth or 5 km 50 km

Where:

ERH and ERZ represent the largest horizontal and vertical deciation respectively in kilometers within the error ellipsoid.

NOTE: The values in these tables represent the minimum acceptable values for each quality class.

Earthquakes within 200km of 37 25N & 113 15W, Mas 3.0 and above

year	date	oris	time	lat-n	lons-w	denth	mas	km	82
1859	828	0	.00	37-50.52	112-49.63		3.7	60.2	38,3
1873	731	315			112-38,38		4.3	109.8	29.3
1876	1130	500	.00	37-40.97	113- 3,95		3.0	33.7	28.9
1876	1230	515	5,00	38-46,15	112- 5.02		3.0	181.7	34.3
1877	101	0	۰00	38-46.15	112- 5.02		3,7	181.7	34-3
1877	115	1200	۸۸	70.74.00	112-15.30		7 ^	155 0	74 1
1878	814	1////			112-34.80		3.0	155.9	34.1
1878	816	244			112-34,80		4.3 3.7	143.9	24.1
1881		215						143.9	24.1
	326				113-43.00		4,3	52.0	290.8
1881	804	430	.00	38-16,74	112-38,38		3.0	109.8	29.3
1885	905	335	.00	37- 2.84	112-31.34		3.9	74.5	122.4
	1026	510	,00	38-12.89	112-55,43		3.0	93.1	18.0
1885	1026	800	.00	38-25.80	113-17,40		3.0	112.5	358.2
1885	1026	900	,00	38-25.80	113-17,40		3.0	112.5	358.2
1885	1217	100	.00	38-10.32	112-16.27		3.7	120.3	45.8
1005	1217	320	.00	70-17 10	111-28,80		3.0	183.2	58.3
	1205				112-31.34		5.7	76.5	122.4
1891	1200	1355			113-34,41		5.0	44.8	219.8
1891	913	348			113- 3.95		3.0	33.7	28.9
1894	205	330			112-26,32		3.7	149.7	24.8
107-3	290	.,,,,,	**/*/	00.0000	112,2025		347	10747	7.4+0
1894		1500			112-26.32		3.0	169,7	24.8
	1110	900			112-38,39		3.7	109.8	39,3
19 1	1114	439	.00	38-46.15	112~ 5.02		7.0	181.7	34.3
19 2		0			112-39.38		3.0	109.8	29.3
19 2	731	700	.00	38-16.74	112-38.38		4.3	109.8	29.3
19 2	1117	1950	.00	37-23,58	113-31,20		5.3	24.0	263.7
	1205				113-31,20		5.0	24.0	264.2
		1630			112- 5,02		3.0	181.7	34.3
		2340			113-34.41		3.0	44.8	219.8
		1000			113-34,41		4.3		219.8
10.0	4 4 C	^	44	70 07 50	447 44		<i>=</i> ^	4 4 4 4	44.4
	415				11344		5.0	110,4	11.1
		1300			112- 8.98		5,0	170.5	34.5
1910		300			112- 8.98		5.0	170.5	34.5
		2112			111-30,00		5.7	186.1	123.1
1913	1020	1000	.00	37-49.36	112-26,02		3.7	85.0	58.0
1914	1214	530	.00	37-34,37	113-42,79		4.3	44.5	292.9
1914	1221	525	.00	37-34.37	113-42,79		3.0	44.5	292.9
1915	212	1950	.00	37-34.37	113-42,79		3.7	44.5	292.9
1915	213	830	.00	37-34.37	113-42.79		3.7	44.5	292.9
1915	1025	1713	,00	38-37.90	112-13.00		3.0	162.5	33.9

Earthquakes within 200km of 37 25N % 113 15W, Mas 3.0 and above

gear	date	oris	time	lat-n	โดกส-พ	denth	esm	ķm	az
1920	818	820	.00	38-16.74	112-38.38		3.0	109,8	29.3
1920	1126	0	,00	37- 4.38	113-34,41		1.3	44.8	219.8
1921	602	2130	.00	37-40.97	113- 3.95		3.0	33.7	28.9
1921	912	1018	<.00	38-45,15	112- 5.02		1.3	181.7	31.3
1921	912	1327	.00	38-46.15	112- 5.02		3.7	181.7	
1921	913	930	.00	38-46.15	112- 5.02		4.3	181.7	34.3
1921		1412			112- 8,98		6.3		
1921		230			112- 8,98		5.7		
		1532			112- 8.98		6.3		
		1227			112- 8,98		3.0	170.5	34.5
1991	1007	1415	. 00	70AA 07	112- 8.98		4.3	170.5	34.5
		1535			112- 8.98		3.7		
		945			112- 8.98		3.0		34.5
		1510			112- 8,98		3.7		
		1210			113-14.00		4.3	83.3	1.0
177.0	314	1210	100	30-10+00	113-14490		463	ପର୍ବଣ	1.0
1925	714	1347	,00	37-48.60	113-55.80		3.7	74.2	504+0
1926		1951			112-35+47		3.0		100.7
1926		1100			112-35,47		3.0		100.7
1926		520			112-35,47		3.0		100.7
		1515			112-35,47		3.0		100.7
1926	1004	2030	.00	37-19.04	112-35.47		3.0	59.4	100.7
	1023				112-35,47		3.0		100.7
	1112	340			112-35.47		3.0		100.7
	1116				112-35.47		3.0		100.7
	522	500			112- 5.02		3,0	181.7	34.3
	V	300							
1927	1123	930			112-35,47		3.0		100.7
1933	120	1310			112-49.53		5.0		33+3
1934	1225	1000	,00	37- 2,84	112-31.54		3.7	76.5	122.4
1935	1006	300	.00	37-55.47	111-25.61		3.7	170.4	70.7
1936	122	338	,00	36-18,00	113-30.00		4.3	125.9	190.2
1936	509	1025	.00	37-15.00	112-57.48		4.3	31.8	125.6
1936	902	2337	,00	38-30,00	112-23:00		3.0	142.3	32.4
1936	921	620	+00	3800	113-18.00		4.7	64.9	354.1
1937	218	415	.00	37-49.36	112-26:02		4.3	85.0	58+0
1937					112-26.02		3.7	85.0	58.0
1937	221	830	,00	37-49.36	112-24,02		3.0	85.0	58.0
1937		130			112-26.02		3.7	85.0	58.0
1938					113-42-00		3.0	29.0	206.7
	1228				11400			81.1	235.2
					115- ,00		3.0		234.4

Earthquakes within 200km of 37 25N % 113 154, Mas 3.0 and above

year	date	oris	time	lat-n	lona-⊎	derth	nas	k.m	5Z
1940	310	1801	54.00	3700	11500		3.0	162.1	253.4
1940	311		30.00		11500		3.0	162.1	253.4
1940	407				11500		3.0	162.1	253.4
1940	920	1221		36-30,00			3.0	186.1	236.9
1940	1026		24,00		113~ .00		3.0	68.4	18.8
								0017	1010
1941	322	1208	6,00	3600	114-34,00		3.0	198.2	217.5
1941	506		42.00	37-18.00	114-18.00		3.0	93,9	262.1
1942	328	1310	.00	38-10,32	112-16,27		3.7	120.3	45.8
1942	830	2208	.00	37-40.97	113- 3.95		5.0	33.7	28,9
1942	918	0	,00	37-40,97	113- 3,95		3.0	33.7	28.9
1942	918	200	.00	37-40.97	113- 3.95		3.0	33,7	28.9
1942	918	520	+00	37-40,97	113- 3,95		3.7	33.7	28.9
1942	926	1116	.00	37-40.97	113- 3.95		4.3	33.7	28.9
1942	926	1450	,00	37-40,97	113- 3,95		5.0	33.7	28.9
1942	928	1330	.00	37-15.00	112-57,48		3.7	31.8	125.6
1943	1.16	1150	18,00	37-40.97	113- 3,95		4.3	33.7	28.9
1943		2014	30.00	37-24.00	114- 6.00		3.0	75.3	268.6
1943	1103	930	,00	38-34,80	112-15.60		1.3	155.7	34.0
1943	1103	1130	.00	38-34.80	112-15-60		3.0	155.7	34.0
1943	1104	500	.00	38-34,80	112-15,60		3.0	155.7	34.0
1943		354			114-15.00		3.7	83.9	265.2
1943	1209	1604	,00	37-40.97	113- 3.95		3.7	33.7	28.9
1944	131	424	58.00	36-54.00	112-24,00		3.0	94.8	127.2
1944	308	754	51.00	37-36,00	114- 9,00		3.0	82.1	284.3
1944	320	931	.00	37-12.00	11400		3.0	70.7	250.1
1944		2345		37- 6,38	113-34,41		3.7	44.9	219.8
1944	1108	2030	6,00	38-48.00	112~54.00		3.0	156.6	11.3
1945	111	1156	.00	37-24,00	114~54.00		3.8	146.1	269.3
1945	1118	107	41.00	.00	11200		3.0	127.8	59.6
1945	1118				111-59,21		5.0	186.2	36.5
	1102			37- 6.38	113-34,41		4,7	44.8	219.8
1950	505	735			112-13.25		3.7	128.6	44.8
1951					112-32.40		3.7	78.2	126.3
1952	502	1015	,00	37-14-47	113~16,90		3.7	19,7	188,2
1952	524	415	15.00	36- 6.00	114-42.00		5.0	195.2	221.5
1952	524				114-42-00		3.7	195.2	221.5
1952	524				114-42.00		3.7	195.2	221.5
1952	525	1306	30.00	36- 6.00	114-42,00		3.7	195.2	221.5
1953	418	515	٠00	38-38-21	112- 7.40		3.7	167.7	36.1
1953	518	703	2,00	3600	114-30.00		3.0	192.8	215.4
								•	

Earthquakes within 200km of 37 25N & 113 15W, Mas 3.0 and above

sear	date	oris	time	lat-n	lons-w	derth	mas	ł m	az
1953	809	2200	2,00	37-15-00	114-30.00		3.0	112.3	260.5
	1022	300	,00	37-49,36	112-26,02		1.3	85.0	58.0
1,955	110	1007	.00	3700	114-30.00		4.3	120.2	247.4
1955					113-54,00		3.7	67.5	278.6
1959	227	2219	52.00	3800	112-30,00		5.0	92.6	45.6
1959					112-30.00		5.5	81.1	124-8
1959		1115			113-10.51		3.7	15-1	25.9
1959	917	620			112-13-80		3.7	145.5	38,1
1961					114-20.00		3.0	99.7	286.1
1962	215	712	42,90	36-54,00	112-24.00		1.5	94.8	127.2
1962	215				112-54.00		3.0	55.7	146.1
1962	217				111-36,00		3.0	195.6	47.8
1962					112- 6.00		3.0	120.3	57.4
1962					112- 5.31		3.2	124.3	55.5
1963	217	1734	20.64	37-48.10	113-54.21	7.0	3.7	71.8	305.5
1963	419	838	44,87	38- 1:10	112-31.50	7.0	3.7	92.1	43.7
1963	930	917	39,33	38- 5.90	111-13.05	7.0	4.3	194.4	67.1
1963	1113	617	30.12	38-17,72	112-39.39	7.0	3.2	110.5	28.2
1964	101	1643	5.82	37-33.22	112-43.85	7.0	3.1	48.4	71.7
1964	117	15	3 : 45	38-11-11	112-37.30	7.0	7.3	101.7	33.0
1964	824	151	55ء	38-46.33	112-13.93	7.0	3.1	175.0	30.7
1964	824	155	30,44	38-47,00	112-14.83	7.0	3.0	175.4	30.1
1964	921	36	23.17	38-47.76	112-12.75	7.0	3.0	178.1	20.7
1944	1129	931	34,35	38-58,05	112-13,81	7.0	3.1	193.9	27.4
1965	130	1348	53.21	37-32.24	113- 6.95	7.0	3.2	17.9	41.5
1945					112-39.00		3.0		
1965					112-26-49		3.0		46.5
1965					113-41.34		3.2	39.0	
1966					112-12.09		3.0	116.2	52.7
1944	520	1340	47,87	i≟-≃ci oo	111-51.25	7+0	4.1	138.2	85.8
1966					114- 9.07				
1966					111-11 18				
1966					117-15-01			38.6	
					112-15-79			134.6	
1966	923	13	57.11	37-30.50	113-17.47	7.0	Ĭ. 2	8,01	340.3
1966	1021	713	48,88	38-11.74	113- 9-45	7.0	4.2		
					112-13.01				30.5
					112- 9.39			1박구 7	
1967	1113	1648	53.77	37-34.97	113-20:32	7.3	3.5		
1948	320	1533	3.98	37-55,28	112-16,55	7 9	3.0	102.6	51.0

Earthquakes within 200km of 37 25N ≩ 113 15W, Mag 3.0 and above

year	date	oris	time	lat-n	lond-w	derth	៣១៩	k.n _i	87
1968	408	1608	14.65	37-22.40	114-17-12	, 9	3,3	91.8	267.0
1968	412				114-37,59		3.1		245.7
1968	924				112- 4.92	7.0	3.5	124.2	56.0
1949	410	837			112- 4.39		3.6	172.8	
1969	815	30			113-14,97	7.0	3.0	41.6	- 1
1970	418	1042	11,46	37-52,47	111-43.26	7,0	3.7	144.2	69+4
1970	523	2255	23.20	38- 3.47	112-28.13	7.0	3.9	99.0	44.1
1970	901	1152	17,81	37-26,47	113-47.17	7.0	3.0	47.5	273.3
1971	623	608	35.87	38-36.49	112-42.38	7.0	3.1	140.6	10.9
1971	1110	1114	10.32	37-41,60	113- 6.16	7.0	3.1	33.4	23.0
1971	1110	1124	34.89	37-41.87	113- 7.28	7.0	3.1	33.2	20.0
1971	1110	1338	14,19	37-42:14	113- 5.58	7.0	3,1	34.8	23.5
1971	1110	1410	23.01	37-48.03	113- 5.99	7.0	3.7	44.6	17.3
1971	1110	1443	58,51	37-48,25	113- 4.89	7.0	3.4	15.5	19.1
1971	1110	1608	37,68	37~42.01	113- 1.03	7.0	3,1	37,6	33.2
								,	
1971	1110	1941	33.89	37-48.37	113- 3.75	7,0	3.5	46.3	21.0
1972	103				112- 9,91		4.4	166.5	34.9
1972	427				113- 6.88		5.4	114.7	174.0
1972	427				113- 8.99		3.4	115,4	175.6
1972	514				113-30.10		3.3	118.4	190.9
		221. 3							
1972	602	315	48.21	38-40,27	112- 4.32	7.0	4.0	173.4	30+6
1972					113-50.77		3.0	53.6	269.4
1972					113-21.78		3.2	22,7	206.2
1972	1116	217	45,16	37~31.93	112-46-18	7.0	3.6	44.4	73.2
1973	122				112-57.90		3.0	35.5	134.6
1973	218	931	39,62	38- 5.87	113-10.57	7.0	3.3	75.9	4.9
1973					113- 8.74		3.3		6.7
1973		1620			113-31.37		3.1	04.0	337.9
		-			112-14.13		3.5	136.0	41.0
					113-28,30			72.7	
277.5		E-W-AL-V		2		,			
1976	813	1030	21.13	38-25.27	112-10.81	7.0	3,1	145,9	40.2
					112-50.60				
					112-29.38			95.5	
					112-28.73				
					112-31.32				
1978	1209	2349	8,00	38-38,84	112-31,08	5.5	3.3	151.0	25.2
1979					113-10.05		3.2	36.5	
1979					113-40.08				217.5
1979					113-49.01			90.5	213.8
1979					112-43.41			60.1	50.7
***	~ ~		* / / //	45.752.7	1.04.11	· • •	- ' -	****	10.7.1.6

Earthquakes within 200km of 37 25N % 113 15W, mag 3.0 and above

Rest	date	oris	time	lat-n	long-w	derth	esm	km	2 %
1980	428	2147	35.00	37-18.11	114- 5.00	7.0	3.4	74.9	260-2
1980	803				112-58,16		3.1	42.9	29+8
	1221				113- 4.93		3.3	15.2	78.0
	1227				113- 7.03		3.0	15.9	47.5
			•		113- 4.92	1.7	3.1	15.5	74.2
1981	116	1023	29,90	37-26,52	113- 5.51	3,5	3.4	14.3	78.6
1931	116	1450	45.61	37-26.34	113- 5.28	1.2	3.2	14.6	80.2
1981	201	221	47.37	37-34,47	113-15.56	٠0	3.8	17.5	357+3
1981	405	540	40.06	37-36.61	113-18.02	1.3	4.5	21.9	348.3
1981	529	309	14,98	37-27.13	111-16.11	7.0	3.0	175.4	88.7
1981	714				111-33,25	7,0	3,1	160.7	68.6
1981	808				112-48.51	.0	3,3	82.2	28 42
1912		2112			111-30.00		5.7	136.1	123.1
1936	122	339			113-30,00	. 0	1.3	125.9	190.2
1938	817	908	6.00	36-42.00	113-42.00	٠٥	0.8	89.0	206.7
	1217			36- ,00			3.5	170.9	203.1
	1121			36-30.00		-0	4.0	186.1	236.9
				36-30.00		,0	3.0	186.1	276.7
1940				36-30.00		•0	3.0	186.1	235.7
1941	522	1208	ሉ.00	36- 500	114-36,00	٠0	3.0	198.2	217.5
1945	107	2225	32.00	36-30.00	111-48.00	.0	4.9	164.4	128.2
1952	524	148	12,00	36- 6-00	114-42,00	.0	3.7	195.2	221.5
1952	524	952	52,00	36- 6.00	114-42.00	.0	3.7	195.2	221.5
1952	525	1306	30.00	36- 6:00	114-42.00	- 0	3.7	195.2	221.5
1953	518	703	2.00	3600	114-30.00	.0	3.0	192.8	215.4
1963					114-54.00		3.3	179.8	227.5
1963					114-42.00		4.4	195,2	221.5
1966	903				112-15,26			134.6	138.8
1966	121				113-54.00		3.7	146.9	203.2
1970	1124	1647	56,00	34-21:42	112-16-37	٤.0	3.0	146.3	143.5
1972	427	157	₹.1 <i>4</i>	36-23.31	113- 6.88	7.0	3,4	114.7	174.0
1972					113- 8.99			115.4	175.6
1972					113-30.10		3.3	118.4	190.9
1973					115-15.53		3.6	195.7	240.0
	1226				114-38,34		3.1	193.3	219.9
17/3	1240	918	10.00	30° 4'49	114730+34	5.0	ુલ.	573+3	4 1 7 + T
1979	815	211	30.27	36-41,36	113-49.01	7,0	3.2	90.5	213.8
1981	112	859	13.20	35-39.47	113-29.13	5.0	3.5	195,2	185.7
1934	330	1627	,00	37-42,00	115-18,00	٠0	4.9	183,8	279+8
1934	330	1851	.00	37-42.00	115-18.00	.0	4.0	184.8	279.0
1934	331	824	,00	37-42.00	115-18.00	, ſ)	4.0	183.8	278.8

1692	date	oris	time	lat-n	long-w	derth	mas	km	az
1934	415	1209	.00	3800	11500	٠0	5.0	167.4	292.8
1934	417	330	.00	38 00	11500	.0	4.0	147.4	297.8
1934	417	1446	.00	3800	11500	.0	4,0	167.4	292,8
1937	128	314	,00	37- 6,00	114-42.00	.0	4.0	133.3	254.7
1940	310		54.00		11500	٠0	3.0	162.1	253.4
1940	311	4	30,00	37- +00	11500	,0	3.0	162.1	253+4
1940	407	842	.00	37- ,00	11500	٠0	3.0	162.1	253.4
1941	504	311	42:00	37-18:00	114-18.00	. 0	3.0	93.9	262.1
1943	610	2225	30.00	37-18:00	11500	.0	3.5	155.6	265.2
1943	114	355	,00	37-21.00	114-15.00	, 0	1.0	88.9	285+2
1944	320	930	57.00	37-30.00	114-42.00	.0	4.0	128.6	274.1
1945	111	1156	,00	37-24.00	114-54.00	, ()	3.8	146.1	596.3
1953	809	2200	2.00	37-30.00	114-30.00	•0	4.5	111.0	274.48
1955	110	1007	28,00	37 +00	114-30:00	, 0	1.4	120.2	247.4
1961	926	2146	20.00	37-39.96	114-20.00	٥٠	3.0	99.7	286.1
1964	529	111	10,40	37-18.00	114-48-00	33.0	3.6	137,9	264.3
1964	821	2203	51.60	3700	115- 6.00	33.0	3,8	170.6	254.3
1965					115-12:00		3.7	173.6	276.7
1966	406				115-24,00		4.1	192.1	262.8
1966					111-18,00		3.7	93.0	268.9
1966	820	822	3,40	37-18.00	114-18.00	33.0	3,5	93,9	252 (1
1966		454	28,79	37-26,94	114-15.01	7.0	8.8	88. 6	272.3
1966	825	309	48.90	37-18.00	114-18.00	33.0	3.7	93.9	262+1
1966	827	1531	49,50	37-18:00	114-18-00	33.0	3.0	93.9	262+1
1967					115- 472		4.7	161.9	255.0
1947	' 509	42	24,10	37- 2,51	t15- +60	15.0	4.0	151.6	255.1
1947					115-12-00			176.5	258.5
	1211				115-12.00			174.5	232 . 1
1968					114-17.12		3.3	91.8	267.0
					111-37.59		3.1	134.2	245.7
1945	7 1008	3 1555	1.50	77-30.00	115-28.97	10.0	3.7	197.8	272.7
1969	7 1108	3 658	1.30	37- 8.74	1114-58.69	4.0	3.3	156.2	258.9
1970	1.00	2001	31.40) 37-25.70) [1:4-57,77	' ()	4.0	151+6	270.5
1970	208	155/	26.60	37-19-20) 115-34 17	, Q	7.5	190.9	266.8
1970	1103	3 1641	25.00	37-31.28	115-18-36	5 9	3,8	192.3	273.6
1971	L 1125	5 1827	7 25.00	37-47.70	114-57.00) ^	7 ع	155. 9	295.6
					115-16.97			190.9	278,5
1971	1208	1523	3 51.90	37-46.56	114-39,40	, , 0	3.5		
					5 114-25.50				271.2
					114-19.88				245.3
						-			

Soar	date	oris	time	lat-n	lons-u	derth	esm	km	82
1972	421	2248	34.80	38-16.79	114-57.60	٠0	3.0	178.4	302.5
1972	424	212	30,70	37-39,65	115- 3.24	٠.0	3.0	161.7	279.6
1972	927	1229	49,00	38-28.73	114-48.84	٠٥	3.8	181.1	310.6
1972	1028	1901	24,90	37-13.38	114-53.34	5.0	3.4	146.9	261.6
1972	1029	1135	2,50	36-59.81	114-37.86	•0	3,5	131.2	249.2
					115-13,43		3.2	175.5	265+1
					115-11,22	•0	3.4	172.9	263.2
					115- 4.44		3.7	162.3	276.5
					115-24.12	•0	3.5	195.2	283+3
1972	1117	950	8.80	37-53,63	115-16.79	· 0	3.6	186.8	286.5
1972	1117				115-14.27		3.8	182.6	286.1
	1123				115-22.13		3.8	194.1	285+6
1972	1129				114-47,22	.0	3.7	151.5	296.6
	1129				114-40.25	, 0	3.9	191.6	297.8
1972	1217	20	53.30	37-41.34	115-21.71	٠Û	3.0	189.0	279.2
1973	129	1438	50,20	37-57.18	114-54.88	, 0	3.1	161.2	291.7
1973	304	1112	51.40	38- 3.77	115-11.10	, ()	3.0	195.0	292.8
1973	527	239	29,80	37-10.97	115-13.50	.0	4.0	177.0	261,6
1973	610	557	16.10	37-41.15	114-55.20	.0	3.1	150.6	281.4
1973	730	918	42,10	37-36,65	115- 8.93	5.0	3.5	169.3	277.3
1973	914				115- 3.18		4.1	157.6	271.0
1973	1105	250	22,20	37-57,54	115- 6.59	.0	3.2	174.7	290.1
1974	405	27	31,80	37-41.27	114-50.22	.0	3.8	143.4	232.1
1971	411	1253	7,60	37-34,11	115-17-39	11.0	3,9	181.5	276.5
1974	826	805	41,70	37-19.50	115-11.93	٠0	3.6	172.9	266.6
1975				=	115-24.76				283.2
1976					115-24.21		4.5		276.8
1976					114-58,79		3.6		297.0
1977					114-58.31		3.3	154,0	262.4
1977	1113	1844	45.50	37-19-55	115-11.22	,0	3.5	171.9	266+6
					115-18.23				267.3
					114-26,94	٠٥	3,8		302.4
					114-36.77		3.5		299.5
					115-12-19				249.6
1979	812	1131	21.00	37-16.86	11572	.0	4.2	156.8	264.5
1979					115- 3,83		3.1	161.0	266+2
1979					115- 3,48		3.8	161-0	264.5
1979					114-52,13		3.4		269.4
1979					114-56.70				267.48
1979	813	2332	29,20	37- 8,27	115- 4,19	•0	3.0	164.3	259,1

Earthquakes within 200km of 37 25N % 113 15W, Mas 3.0 and above

acar	date	oris	time	lat-n	lons-9	desth	អខន	km	er
1979	814	303	16.20	37-14,88	115- 8.27	٠٥	3.9	168.4	263.6
					114-57.65	,0	3.7	151.9	266.0
1979	816	1549	58.90	37-18.84	115- 2.46	.0	3.1	159.1	265.9
1980	416	1124	36.30	37-10.70	115-27,18	, ()	3.3	197.1	242.3

APPENDIX I

LISTING OF EARTHQUAKE OF M>3 WITHIN 150 KM OF
CEDAR CITY AND ST. GEORGE (1850-1981)

Earthquakes within 150km of Cedar City site, Mas 3.0 and above

aesr	date	oris	time	1at−n	long-w	derth	៣៩៨	km	32
1859	828	0	.00	37-50.52	112-49.63		3.7	30.3	45.3
1873	731	315	.00	38-16.74	112-38,38		4.3	79.5	28.5
1876	1130	500			113- 3.95		3.0	3.7	8.0
1877	115	1200	.00	38-34.80	112-15.30		3.0	125.6	34.8
1878	814	0	00	38-36.00	112-34.80		4.3	113.9	22.2
1878	816	244	.00	38-36.00	112-34.80		3.7	113.7	22.2
1881	326	215	.00	37-35.00	113-48.00		4.3	64.7	263.4
1881	804	430	٠00	38-16.74	112-38.38		3.0	79.5	28.5
1885	905	335	.00	37- 2.84	112-31.34		3.0	82.47	144.0
1885	1026	610	.00	38-12-89	112-55.43		3.0	64.0	11.7
1885	1026	800	.00	38-25.80	113-17.40		3.0	88.7	347.5
1885	1026	900	.00	38-25.80	113-17.40		3.0	88.7	347.5
1885	1217	100	.00	38-10.32	112-16.27		2.7	91.2	50.5
1887	1205	1530	.00	37- 2.34	112-31,34		5,7	82.7	144.0
1891	420	1355	.00	37- 6.38	113-34.41		5.0	74.9	216.4
1891	913	348	+00	37-40.97	113- 3.95		3.0	3.7	8.0
1894	205	330	.00	38-48.30	112-26.32		3,7	139.7	23.4
1894	206	1500	.00	38-48.30	112-26.32		3.0	137.7	23.4
1899	1110	900	.00	38-16.74	112-38,38		3.7	79.5	28.5
19 2	601	0	.00	38-16.74	112-38,38		3.0	70.5	28.5
19 2	731	700	.00	38-16.74	112-38.38		4,3	79.5	28.5
19 2	1117	1950	٠00	37-23.58	113-31.20		4.3	48.3	234.3
19 2	1205	0	.00	37-23.68	113-31.20		5.0	48.7	234,4
19 3	1104	2340	+00	37- 6.38	113-34.41		3.0	74.9	216.4
19 3	1123	1000	.00	37- 6.38	113-34.41		4,3	74.9	216.4
19 8	415	0	.00	38-23.58	11344		5.0	82.7	3.9
1910	110	1300	.00	38-40,97	112- 8.98		5.0	140.3	35.2
1910	112	300	.00	38-40.97	112- 8.78		5.0	140.3	35.2
1913	1020	1000	.00	37-49.36	112-26.02		3.7	59.4	71.2
1914	1214	530	٠00	37-34.37	113-42,79		4.3	57.3	261.4
1914	1221	525	.00	37-34.37	113-42.79		3.0	57.3	261.4
1915	212	1950	.00	37-34.37	113-42.79		3.7	57.3	251.4
1915	213	830	.00	37-34.37	113-42.79		3.7	57.3	261.4
1915	1025	1713	.00	38-37.90	112-13.00		3.0	132.3	34.5
1920	818	820	.00	38-16.74	112-38.38		3.0	29.5	5612
1920	1126	0	.00	37- 6.38	113-34.41		4.3	74.9	216.4
1921	602	2130	.00	37-40.97	113- 3.95		3.0	3.7	8.0
1921		1412	.00	38-40.97	112- 8.98		4.3	140.3	35.2
	930	230	.00	38-40.97	112- 8,98		5.7	140.3	35.2
1921	1001	1532	.00	38-40.97	112- 8.99		5,3	140+3	75.2

Earthquakes within 150km of Cedar City site, Mas 3.0 and above

year	date	oris	time	lat−n	long-w	derth	mas	k.m	āZ
1921	1015	1227	.00	38-40.97	112- 8.98		3.0	140.3	35.2
1921	1027	1415	٠٥٥	38-40.97	112- 8.98		4.3	140.3	35.2
1921	1101	1535	.00	38-40.97	112- 8.98	•	3.7	140.3	35.2
1921	1220	945	.00	38-40.97	112- 8.98		3.0	140.3	35.2
1921	1220	1510	.00	38-40.97	112- 8,98		3,7	140.3	35.2
			¢						
1923	514	1210	.00	38-10.00	113-14-00		4.3	59.1	346.1
1925	714	1347	.00	37-48.60	113-55.80		3.7	77.7	283.2
1926	515	1951	.00	37-19.04	112-35.47		3.0	56.3	131.0
1926	605	1100	.00	37-19.04	112-35.47		3.0	56.3	131.0
1926	712	520	.00	37-19.04	112-35.47		3.0	56.3	131.0
	1001				112-35.47		3.0	56.3	131.0
	1004				112-35,47		0.5	56.3	131.0
	1023	0			112-35.47		3.0	56.3	131.0
	1112	340			112-35.47		3.0	56.3	131.0
1926	1116	1415	.00	37-19.04	112-35,47		3.0	55.3	131.0
1927	1123	930	۰00	37-19.04	112-35,47		3.0	56.3	131.0
1933	120	1310	.00	37-50.52	112-49.63		5.0	30.3	45.3
	1225				112-31.34		3.7	82.7	144.0
	1006	300		37-55.47			3.7	148.1	78,1
1936	509				112-57,48		4.3	45.5	167.2
1936		2337			112-23.00		3.0	112.0	32.6
1936	921	620			113-15.00		4.7	43.7	332.6
1937	218	415			112-26.02		4.3	59.4	71.2
1937	218	900			112-25.02		3,7	57.4	71.2
1937	221	830	.00	37-49.36	112-26.02		3.0	59.4	71.2
1937	226	130	.00	37-49+36	112-26.02		3.7	59.4	71.2
1938	S17	908	6.00	36-42.00	113-42.00		3.0	119.3	207.9
	1228				11400		3.0	109.4	228.8
	1026				11300		3.0	39.4	9.2
					114-18.00		3.0	115.4	250.3
									pu
1942		1310			112-16.27		3.7	91.2	50.5
1942		2208			113- 3.95		5.0	3,7	8.0
1942	918	0			113- 3.95		3.0	3.7	8.0
1942	918	200			113- 3.95		3.0	3.7	3.0
1942	918	520	.00	37-40.97	113- 3,95		3.7	3.7	8+0
1942	926	1116	.00	37-40.97	113- 3.95		4 - J	3.7	3.0
1942	926	1450			113- 3,95		5.0	3.7	8.0
1942	928	1330	.00	37-15.00	112-57.48		3.7	45.5	157.2
1943	116	1150	18.00	37-40.97	113- 3.95		4.3	3+7	8.0
1943	306	2014	30.00	37-24.00	114- 5.00		3.0	95.0	253.0

Earthquakes within 150km of Cedar City site, Mag 3.0 and above

1592	date	oris	time	lat-n	long-u	derth	mas	km	ar
1943	1103	930	.00	38-34.80	112-15,60		4.3	125.4	34.6
	1103				112-15.60		3.0	123.4	34+6
	1104	200			112-15.60		3.0		34.6
	1106				114-15.00		3.7		252.3
					113- 3.95		3.7		8.0
1775	1741	1000	* 100	37 70477	110 04/0		U+1	U + 7	0.0
1944	131	424	58.00	36-54.00	112-24.00		3.0	102.4	144.4
1944	308	754	51.00	37-36,00	114- 9.00		3.0	95,4	266.7
1944	320	931	۰00	37-12,00	114- +00		3.0	96.2	238.7
1944	503	2345	.00	37- 6.39	113-34.41		3.7	74.9	216,4
1944	1108	2030	6.00	39-48.00	112-54.00		3.0	128.5	٤.7
1945	1118	107	41.00	3800	11200		3.0	102.0	57.6
1949	1102	230	.00	37- 6.38	113-34.41		4.7	74.9	216.4
1950	505	735	.00	38-14.33	112-13,25		3,7	99.3	48.9
	305				112-32.40		3.7	94.2	
1952		1015			113-16.90		3,7	a9.0	200.3
1953	418	515	.00	38-38.21	112- 7.40		3.7	137.5	37.2
1953	809	2200	2,00	37-15.00	114-30.00		3.0	134.0	250.5
	1022				112-26.02		4.3	59.4	71.2
1955	110	1007		3700	114-30.00		4.3		
					113-54.00		3.7	95.4	230.2
			~	2. 2					
1959	227	2219	52,00	3800	112-30.00		5.0	63.6	52.3
1959	721	1739	29.00	3700	112-30.00		5.5	88.2	114.4
1959	727	1115	.00	37-32.36	113-10.51		3.7	15.3	216.7
1959	917	620	,00	38-26.93	112-13.80		3.7	115.4	39.8
1961					114-20.00		3.0	111.3	270.9
1962	215	712	42.90	36-54.00	112-24.00		4.5	102.4	144.4
1962	215	906	45.00	37- +00	112-54.00		3.0	73.7	168.1
1962	605	2229	45.00	400	112- 6.00		3.0	94.0	65.6
1962	819	1732	41,43	38- 3,08	112- 5.31	7.0	3.2	97.3	62.8
1963	217	1734	20.64	37-48.10	113-54.21	7.0	3.7	75.2	282.4
1017	410	070	ለለ ወግ	70 1 10	112-31.50	7 0	3.7	65.1	49.7
					112-31.30		3.2		
1964					112-43.85		3.1		
							3.3		
1964					112-37.30				
1964	824	151	. 55	35~46+33	112-13.93	7+0	3,1	114,6	30.6
1964	824	155	30.44	38-47.00	112-14-83	7.0	3.0	t 15.0	29,0
1964	921	36	23.17	38-47.76	112-12.75	7.0	3.0	147.8	30.46
1965	130	1348	53.21	37-32,24	113- 6.95	7.0	3.2	13+1	197.3
1965					112-59-00		3.0	10.4	48.6
1965					112-26,49		3.0	69+4	53.0

Earthquakes within 150km of Cedar City site, Mas 3.0 and above

year	date	oris	time	lat-n	lons-⊌	derth	សខន	km	28
1965	830	43	15.20	37-26.92	113-41.34	7.0	3,2	59.0	247.7
1966	520	1211	37.39	38- 3.03	112-12.09	7.0	3.0	88.6	59.9
1966	520	1340	47.87	37-58.99	111-51.25	7.0	4.1	113.4	71.0
1966	816	1802	32.85	37-27.81	114- 9.07	7.0	5.6	97.5	257.8
1966	819	1051	3,7.92	37-26.33	114-11.48	7.0	4.7	101.7	254.7
1966	824	454	28.79	37-26.94	114-15.01	7.0	3,8	106.5	257.9
1966	903				112~15.26		3.5	146.4	150.2
1966	923				113-17.47		3.2	25.0	231.0
1966	1021	713	48.98	38-11.74	113- 9.45	7.0	4,2	61.0	352.9
1967	722	2151	27.40	39-48.14	112-13.01	7.0	3.6	148.2	30.3
					112- 9.39		5.2	127.6	39.0
1967					113-20.32		3.5		252.5
1968					112-16,55	7.0	3.0	76.3	56+R
1938					114-17.12	٠9	$\Sigma_{+}\Sigma_{-}$	111.6	254.0
1968	924	210	49,63	38- 2.50	112- 4.92	7.0	3.5	97.4	63.5
1969	410	837	5.46	38-39,89	112- 4.39	7.0	3.6	142.5	37.8
1969	315				113-14.97		3.0		
1970	418				111-43.26	7.0	3.7	121.6	79,2
1970					112-23.13		3.9	69.7	49.5
1970					113-47.17	7.0	3.0	67.3	249.8
1971	623	608	35,87	38-36.49	112-42.38	7.0	3,1	111.1	16 9
1971	1110	1114	10.32	37-41.60	113- 6.16	7.0	3.1	5.5	330 . 4
1971	1110	1124	34.89	37-41.87	113- 7.28	7.0	3.1	6,0	320.5
1971	1110	1338	14.19	37-42.14	113- 5.58	7+0	3.1	5.1	342+0
1971	1110	1410	23.01	37-48.03	113- 5.99	7.0	3.7	15.9	351.5
					113- 4.89	7.0		17.1	357.1
					113- 1.03		3.1		
					113- 3.75			17.4	
					112- 9.91	7.0	4,4	136.2	35.7
1972	427	153	3.14	36-23.31	113- 6.88	7+0	3,4	140.0	181.6
1972	427	841	13.59	36-22.77	113- 8.99	7.0	3,4		
1972	514	238	40.66	36-22.14	113-30.10	7.0	3.3	147.2	195.1
1972					112- 4.32			143.3	37.7
1972	702	2007	1.15	37-20.15	113-50.77	7.0	3.0	76+9	243.0
1972	902	1530	35,47	37-14.01	113-21.78	7.0	3.2	52.9	209.2
1972	1116	217	45.16	37-31,93	112-46,18	2.0		29.7	116.1
1973					112-57.90		3.0		169.5
1973					113-10.57		3,3	50.5	349.5
1973					113- 8.74			53.3	
1973	520	1620	6.62	37-57.07	113-31.37	7.0	3.1	51.9	310.1

Earthquakes within 150km of Cedar City site, Mag 3.0 and above

1974 1104 902 26.60 38-20.45 112-14.13 7.0 3.5 106.2 43.8 1975 111 1820 24.95 38- 2.88 113-8.30 7.0 3.2 56.55 321.4 1976 813 1030 21.13 38-25.27 112-10.81 7.0 3.1 116.0 42.4 1978 224 1949 48.77 38-19.97 112-50.60 1.5 3.5 73.4 14.8 1978 830 1534 38.87 38- 1.76 112-29.38 7.0 3.0 56.3 50.6 1978 1116 818 57.33 38- 5.29 112-28.73 7.0 3.1 71.3 47.0 1978 1209 1459 48.36 38-39.36 112-31.32 4.6 3.3 121.6 23.3 1978 1209 2349 8.00 38-38.84 112-31.08 5.5 3.3 120.9 23.7 1978 1209 2349 8.00 38-38.84 112-31.08 5.5 3.3 120.9 23.7 1979 815 211 30.27 36-44.35 113-40.08 7.0 3.6 91.1 215.5 127 36.9 112 928 59.54 37-44.35 113-40.08 7.0 3.6 91.1 215.5 127 38.9 112 928 2147 35.00 37-18.11 114-5.00 7.0 3.2 13.0 319.4 1990 428 2147 35.00 37-18.11 114-5.00 7.0 3.4 97.5 24.6 1990 803 608 28.57 37-48.43 112-38.16 16.5 3.1 19.6 27.3 1980 1221 1825 9.12 37-26.71 113- 4.93 1.7 3.3 22.8 182.3 1980 1221 1825 9.12 37-26.71 113- 4.93 1.7 3.3 22.8 182.3 1980 1221 1825 9.12 37-26.71 113- 4.93 1.7 3.3 22.8 182.3 1981 116 1450 48.64 37-26.54 113-5.51 3.4 23.2 23.5 183.5 1981 126 1450 48.64 37-26.52 113-5.51 3.5 3.4 23.2 23.5 183.5 1981 201 221 47.37 37-34.47 113-15.56 .0 3.8 18.6 243.2 1981 808 620 17.09 38- 4.15 113-48.50 .0 3.0 114.2 19.2 24.4 1931 41.613 46.72 37-56.67 111-33.25 7.0 3.1 137-6 76.3 1981 808 620 17.09 38- 4.15 112-48.54 6.0 3.0 119.3 20.7 1966 903 753 18.55 36-30.27 112-15.26 7.0 3.0 119.4 29.9 147.3 20.7 1966 903 753 18.55 36-30.27 112-15.26 7.0 3.0 144.4 150.2 26.4 1972 427 841 13.59 36-22.77 113-8.99 7.0 3.4 141.2 182.8 1972 427 153 3.14 36-23.31 113-6.83 7.0 3.0 14.4 140.0 181.6 1972 427 841 13.59 36-22.77 113-8.99 7.0 3.4 141.2 182.8 1972 427 841 13.59 36-22.77 113-8.99 7.0 3.4 141.2 182.8 1972 427 841 13.59 36-22.77 113-8.99 7.0 3.4 141.2 182.8 1972 427 841 13.59 36-22.77 113-8.99 7.0 3.4 144.2 182.8 1972 427 841 13.59 36-22.77 113-8.99 7.0 3.4 144.2 182.8 1972 427 841 13.59 36-22.77 113-8.99 7.0 3.4 144.2 182.8 1972 427 841 13.59 36-22.77 113-8.99 7.0 3.0 114-30.0 1.0 4.0 144.8 263.4 1953 809 2200 2.00 37-30.00 114-30.00 .0	ñб	ar	date	oris	time	lat−n	loua-A	derth	กอร	k m	az
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1976 813 1030 21.13 38-25.27 112-10.81 7.0 3.1 116.0 42.4 1978 224 1949 48.77 38-19.79 112-50.60 1.5 3.5 73.4 14.8 1978 830 1534 38.87 38-1.76 112-28.73 7.0 3.0 36.3 50.6 1978 1116 818 57.33 38-5.29 112-28.73 7.0 3.1 71.3 47.0 1978 1209 1459 48.36 38-39.36 112-31.08 5.5 3.3 120.9 23.7 1979 112 228 59.54 37-44.33 113-10.05 .0 3.2 13.0 319.4 1979 815 211 30.27 36-44.36 113-49.01 7.0 3.2 130.8 213.2 1979 816 121 30.47 35.44 36.13 31.4 210.8 213.2 1979 815 211 30.47											
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1978 830 1534 38.87 38-1.76 112-29.38 7.0 3.0 66.3 50.6 1978 1116 818 57.33 38-39.36 112-31.32 4.6 3.3 121.6 23.3 1978 1209 2349 8.00 38-38.84 112-31.08 5.5 3.3 120.9 23.7 1979 112 928 59.54 37-44.33 113-10.05 .0 3.2 13.0 319.4 1979 815 211 30.27 36-44.35 113-49.01 7.0 3.2 13.0 319.4 1979 815 211 30.27 36-44.35 113-49.01 7.0 3.2 120.8 213.2 1979 816 1733 4.98 37-45.58 112-43.41 7.0 3.4 97.5 246.6 1980 933 608 28.57 37-46.58 112-43.41 7.0 3.4 97.5 246.6 1980 1221 182.5								1.5		73+4	14.8
1978 1209 1459 48.36 38-39.36 112-31.32 4.6 3.3 121.6 23.3 1978 1209 2349 8.00 38-38.84 112-31.08 5.5 3.3 120.9 23.7 1979 112 928 59.54 37-44.33 113-10.05 .0 3.2 13.0 319.4 1979 815 211 30.27 36-48.89 113-40.08 7.0 3.2 120.8 213.2 1979 816 1733 4.98 37-45.58 112-43.41 7.0 3.2 120.8 213.2 1980 428 2147 35.00 37-18.11 114-5.00 7.0 3.4 97.5 246.6 1980 903 608 28.57 37-48.43 112-58.16 16.5 3.1 17.6 22.3 1980 1221 1825 9.12 37-26.71 113-7 3.3 3.0 15.6 194.9 1980 1227 434									3.0	8+86	50.6
1978 1209 1459 48.36 38-39.36 112-31.32 4.6 3.3 121.6 23.3 1978 1209 2349 8.00 38-38.84 112-31.08 5.5 3.3 120.9 23.7 1979 112 928 59.54 37-44.33 113-10.05 .0 3.2 13.0 319.4 1979 815 211 30.27 36-48.89 113-40.08 7.0 3.2 120.8 213.2 1979 816 1733 4.98 37-45.58 112-43.41 7.0 3.2 120.8 213.2 1980 428 2147 35.00 37-18.11 114-5.00 7.0 3.4 97.5 246.6 1980 903 608 28.57 37-48.43 112-58.16 16.5 3.1 17.6 22.3 1980 1221 1825 9.12 37-26.71 113-7 3.3 3.0 15.6 194.9 1980 1227 434	19	78	1116	818	57.33	38~ 5.29	112-28.73	7.0	3.1	71.3	47.0
1978 1209 2349 8.00 38-38.84 112-31.08 5.5 3.3 120.9 23.7 1979 112 928 59.54 37-44.33 113-10.05 .0 3.2 13.0 319.4 1979 805 1910 21.35 36-58.89 113-40.08 7.0 3.6 91.1 215.5 1979 815 211 30.27 36-44.35 113-49.01 7.0 3.2 120.8 213.2 1979 816 1733 4.98 37-45.58 112-70.0 7.0 3.1 33.0 48.3 1980 428 2147 35.00 37-18.81 114-5.00 7.0 3.4 97.5 246.8 1980 1221 434 15.52 37-30.83 113-7.03 3.5 3.0 15.5 194.7 1980 1227 434 15.52 37-27.28 113-4.92 1.7 3.1 21.7 142.4 1981 116 1026											
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1981 116 1450 45.61 37-26.34 113-5.28 1.2 3.2 23.5 183.5 1981 201 221 47.37 37-34.47 113-15.56 .0 3.8 18.6 243.2 1981 405 540 40.06 37-36.61 113-18.02 1.3 4.5 20.7 257.6 1981 714 1613 46.72 37-56.67 111-33.25 7.0 3.1 137.6 76.3 1981 808 620 17.09 38-4.15 112-48.54 -0 2.3 52.0 26.4 1938 817 908 6.00 36-42.00 113-42.00 .0 3.0 119.3 207.9 1966 903 753 18.55 36-30.27 112-15.26 7.0 3.5 146.4 159.2 1972 427 1841 36-23.31 113-6.38 7.0 3.4 149.0 181.6 1972 427 841 13.59 36-22.77 113-8.99 7.0 3.4 149.0 181.6 1972	19	80	1229	712	52.50	37-27+28	113- 4.92	1.7	3.1		
1981 201 221 47.37 37-34.47 113-15.56 .0 3.8 18.6 243.2 1981 405 540 40.06 37-36.61 113-18.02 1.3 4.5 20.7 257.6 1981 714 1613 46.72 37-56.67 111-33.25 7.0 3.1 137.6 76.3 1981 808 620 17.09 39- 4.15 112-48.54 -0 3.3 52.0 26.4 1938 817 908 6.00 36-42.00 113-42.00 -0 5.0 119.3 207.9 1966 903 753 18.55 36-30.27 112-15.26 7.0 3.5 146.4 150.2 1972 427 153 3.14 36-23.31 113-6.38 7.0 3.4 149.0 181.6 1972 427 841 13.59 36-22.77 113-8.99 7.0 3.4 149.0 181.6 1972 514 238 40.63 36-22.77 113-8.99 7.0 3.4 141.2 182.8 <td>19</td> <td>81</td> <td>116</td> <td>1026</td> <td>29.90</td> <td>37-26,52</td> <td>113- 5.51</td> <td>3.5</td> <td></td> <td></td> <td>184.4</td>	19	81	116	1026	29.90	37-26,52	113- 5.51	3.5			184.4
1981 405 540 40.06 37-36.61 113-18.02 1.3 4.5 20.7 257.6 1981 714 1613 46.72 37-56.67 111-33.25 7.0 3.1 137.6 76.3 1981 808 620 17.09 38- 4.15 112-48.54 -0 3.3 52.0 26.4 1938 817 908 6.00 36-42.00 113-42.00 +0 3.0 119.3 207.9 1966 903 753 18.55 36-30.27 112-15.26 7.0 3.5 146.4 159.2 1972 427 153 3.14 36-23.31 113- 6.88 7.0 3.4 140.0 181.6 1972 427 841 13.59 36-22.77 113- 8.99 7.0 3.4 141.2 182.8 1972 514 238 40.66 36-22.14 113-30.10 7.0 3.3 147.2 195.1 1979 815 211 30.27 36-44.36 113-49.01 7.0 3.2 120.8 213.2 1941 506 311 42.00 37-18.00 114-18.00 0 3.0 115.4 250.3 1943 116 355 .00 37-21.00 114-18.00 0 4.0 109.4 252.3 1944 320 930 57.00 37-30.00 114-42.00 0 4.0 144.8 263.4 1953 809 2200 2.00 37-30.00 114-30.00 0 4.0 144.8 263.4 1953 809 2200 2.00 37-39.96 114-20.00 0 3.0 111.3 270.9 1966 818 928 56.90 37-24.00 114-18.00 33.0 3.7 115.4 250.3 1966 824 454 28.79 37-26.94 114-15.01 7.0 3.8 106.5 257.9 1966 825 309 48.90 37-18.00 114-18.00 33.0 3.7 115.4 250.3	15	81	116					1.2	3,2	23.5	
1981 714 1613 46.72 37-56.67 111-33.25 7.0 3.1 137.6 76.3 1981 808 620 17.09 38- 4.15 112-48.54 -0 3.3 52.0 26.4 1938 817 908 6.00 36-42.00 113-42.00 -0 3.0 119.3 207.9 1966 903 753 18.55 36-30.27 112-15.26 7.0 3.5 146.4 150.2 1972 427 153 3.14 36-23.31 113-6.38 7.0 3.4 140.0 181.6 1972 427 841 13.59 36-22.77 113-8.99 7.0 3.4 140.0 181.6 1972 514 238 40.66 36-22.14 113-30.10 7.0 3.3 147.2 195.1 1979 815 211 30.27 36-44.36 113-49.01 7.0 3.2 120.8 213.2 1941 506 311 42.00 37-18.00 114-18.00 0 4.0 109.4 252.3 </td <td>19</td> <td>31</td> <td>201</td> <td>221</td> <td>47.37</td> <td>37-34,47</td> <td>115-15.56</td> <td>٠,0</td> <td>8,2</td> <td>18.5</td> <td>243.2</td>	19	31	201	221	47.37	37-34,47	115-15.56	٠,0	8,2	18.5	243.2
1981 808 620 17.09 38- 4.15 112-48.54 -0 3.3 52.0 26.4 1938 817 908 6.00 36-42.00 113-42.00 .0 3.0 119.3 207.9 1966 903 753 18.55 36-30.27 112-15.26 7.0 3.5 146.4 150.2 1972 427 153 3.14 36-23.31 113-6.88 7.0 3.4 140.0 181.6 1972 427 841 13.59 36-22.77 113-8.99 7.0 3.4 141.2 182.8 1972 514 238 40.64 36-22.14 113-30.10 7.0 3.3 147.2 195.1 1979 815 211 30.27 36-44.36 113-49.01 7.0 3.2 120.8 213.2 1941 506 311 42.00 37-18.00 114-18.00 .0 3.0 115.4 250.3 1943 116 355 .00 37-21.00 114-15.00 .0 4.0 109.4 252.3	19	81	405	540	40.06	37-36.61	113-18.02	1.3		20.7	257.6
1938 817 908 6.00 36-42.00 113-42.00 .0 3.0 119.3 207.9 1966 903 753 18.55 36-30.27 112-15.26 7.0 3.5 146.4 150.2 1972 427 153 3.14 36-23.31 113-6.88 7.0 3.4 149.0 181.6 1972 427 841 13.59 36-22.77 113-8.99 7.0 3.4 141.2 182.8 1972 514 238 40.64 36-22.14 113-30.10 7.0 3.3 147.2 195.1 1979 815 211 30.27 36-44.36 113-49.01 7.0 3.2 120.8 213.2 1941 506 311 42.00 37-18.00 114-18.00 .0 3.0 115.4 250.3 1943 116 355 .00 37-21.00 114-15.00 .0 4.0 109.4 252.3 1944 320 930 57.00 37-30.00 114-30.00 .0 4.0 144.8 263.4	19	81	714	1613	46.72	37-56-67	111-33.25	7.0	3.1		76.3
1966 903 753 18.55 36-30.27 112-15.26 7.0 3.5 146.4 150.2 1972 427 153 3.14 36-23.31 113-6.38 7.0 3.4 140.0 181.6 1972 427 841 13.59 36-22.77 113-8.99 7.0 3.4 141.2 182.8 1972 514 238 40.63 36-22.14 113-30.10 7.0 3.3 147.2 195.1 1979 815 211 30.27 36-44.36 113-49.01 7.0 3.2 120.8 213.2 1941 506 311 42.00 37-18.00 114-18.00 .0 3.0 115.4 250.3 1943 116 355 .00 37-21.00 114-18.00 .0 4.0 109.4 252.3 1944 320 930 57.00 37-30.00 114-42.00 .0 4.0 144.8 263.4 1953 809 2200 2.00 37-30.00 114-30.00 .0 4.5 127.3 262.5	19	81	808	620			112-48.54	. ()	3.3	52.0	
1972 427 153 3.14 36-23.31 113- 6.88 7.0 3.4 140.0 181.6 1972 427 841 13.59 36-22.77 113- 8.99 7.0 3.4 141.2 182.8 1972 514 238 40.66 36-22.14 113-30.10 7.0 3.3 147.2 195.1 1979 815 211 30.27 36-44.36 113-49.01 7.0 3.2 120.8 213.2 1941 506 311 42.00 37-18.00 114-18.00 .0 3.0 115.4 250.3 1943 116 355 .00 37-21.00 114-15.00 .0 4.0 109.4 252.3 1944 320 930 57.00 37-30.00 114-42.00 .0 4.0 144.8 263.4 1953 809 2200 2.00 37-30.00 114-30.00 .0 4.5 127.3 262.5 1955 110 1007 28.00 3700 114-30.00 .0 4.4 145.7 240.3 1961 926 2146 20.00 37-39.96 114-20.00 .0 3.0 111.3 270.9 1966 818 928 56.90 37-24.00 114-18.00 33.0 3.7 112.1 255.7 1966 820 822 3.40 37-18.00 114-18.00 33.0 3.6 115.4 250.3 1966 824 454 28.79 37-26.94 114-15.01 7.0 3.8 106.5 257.9 1966 825 309 48.90 37-18.00 114-18.00 33.0 3.7 115.4 250.3	19	38	917	908	6.00	36-42,00	113-42.00	¥ 0			
1972 427 841 13.59 36-22.77 113-8.99 7.0 3.4 141.2 182.8 1972 514 238 40.66 36-22.14 113-30.10 7.0 3.3 147.2 195.1 1979 815 211 30.27 36-44.36 113-49.01 7.0 3.2 120.8 213.2 1941 506 311 42.00 37-18.00 114-18.00 .0 3.0 115.4 250-3 1943 116 355 .00 37-21.00 114-18.00 .0 4.0 109.4 252.3 1944 320 930 57.00 37-30.00 114-42.00 .0 4.0 144.8 263.4 1953 809 2200 2.00 37-30.00 114-30.00 .0 4.5 127.3 262.5 1955 110 1007 28.00 37- .00 114-30.00 .0 4.4 145.7 240.3 1961 926 2146 20.00 37-39.96 114-20.00 .0 3.0 111.3 270.9	19	66	903	753	18.55	36-30.27	112-15.26	7.0	3.5	146.4	150.2
1972 514 238 40.66 36-22.14 113-30.10 7.0 3.3 147.2 195.1 1979 815 211 30.27 36-44.36 113-49.01 7.0 3.2 120.8 213.2 1941 506 311 42.00 37-18.00 114-18.00 .0 3.0 115.4 250.3 1943 116 355 .00 37-21.00 114-15.00 .0 4.0 109.4 252.3 1944 320 930 57.00 37-30.00 114-42.00 .0 4.0 144.8 263.4 1953 809 2200 2.00 37-30.00 114-30.00 .0 4.5 127.3 262.5 1955 110 1007 28.00 37- .00 114-30.00 .0 4.4 145.7 240.3 1961 926 2145 20.00 37-39.96 114-20.00 .0 3.0 111.3 270.9 1966 818 928 56.90 37-24.00 114-18.00 33.0 3.7 112.1 255	19	72	427	153	3.14	36-23.31	113- 6.88	7.0	3.4	140+0	181.6
1979 815 211 30.27 36-44.36 113-49.01 7.0 3.2 120.8 213.2 1941 506 311 42.00 37-18.00 114-18.00 .0 3.0 115.4 250.3 1943 116 355 .00 37-21.00 114-15.00 .0 4.0 109.4 252.3 1944 320 930 57.00 37-30.00 114-42.00 .0 4.0 144.8 263.4 1953 809 2200 2.00 37-30.00 114-30.00 .0 4.5 127.3 262.5 1955 110 1007 28.00 37- .00 114-30.00 .0 4.4 145.7 240.3 1961 926 2145 20.00 37-39.96 114-20.00 .0 3.0 111.3 270.9 1966 818 928 56.90 37-24.00 114-18.00 33.0 3.7 112.1 255.7 1966 824 454 28.79 37-26.94 114-15.01 7.0 3.8 106.5 257	19	72	427	841	13.59	36-22.77	113- 8.99	7.0	3,4	141.2	182.8
1941 506 311 42.00 37-18.00 114-18.00 .0 3.0 115.4 250-3 1943 116 355 .00 37-21.00 114-15.00 .0 4.0 109.4 252.3 1944 320 930 57.00 37-30.00 114-42.00 .0 4.0 144.8 263.4 1953 809 2200 2.00 37-30.00 114-30.00 .0 4.5 127.3 262.5 1955 110 1007 28.00 37- .00 114-30.00 .0 4.4 145.7 240.3 1961 926 2145 20.00 37-39.96 114-20.00 .0 3.0 111.3 270.9 1966 818 928 56.90 37-24.00 114-18.00 33.0 3.7 112.1 255.7 1966 824 454 28.79 37-26.94 114-15.01 7.0 3.8 106.5 257.9 1966 825 309 48.90 37-18.00 114-18.00 33.0 3.7 115.4 25	19	72	514	238	40.66	36-22.14	113-30.10	7.0	3.3	147.2	195.1
1943 116 355 .00 37-21.00 114-15.00 .0 4.0 109.4 252.3 1944 320 930 57.00 37-30.00 114-42.00 .0 4.0 144.8 263.4 1953 809 2200 2.00 37-30.00 114-30.00 .0 4.5 127.3 262.5 1955 110 1007 28.00 3700 114-30.00 .0 4.4 145.7 240.3 1961 926 2146 20.00 37-39.96 114-20.00 .0 3.0 111.3 270.9 1966 818 928 56.90 37-24.00 114-18.00 33.0 3.7 112.1 255.7 1966 820 822 3.40 37-18.00 114-18.00 33.0 3.6 115.4 250.3 1966 824 454 28.79 37-26.94 114-15.01 7.0 3.8 106.5 257.9 1966 825 309 48.90 37-18.00 114-18.00 33.0 3.7 115.4 250.3	19	79	815	211	30,27	36-44.36	113-49.01	7.0	3.2	120.8	213.2
1944 320 930 57.00 37-30.00 114-42.00 .0 4.0 144.8 263.4 1953 809 2200 2.00 37-30.00 114-30.00 .0 4.5 127.3 262.5 1955 110 1007 28.00 37- .00 114-30.00 .0 4.4 145.7 240.3 1961 926 2145 20.00 37-39.96 114-20.00 .0 3.0 111.3 270.9 1966 818 928 56.90 37-24.00 114-18.00 33.0 3.7 112.1 255.7 1966 820 822 3.40 37-18.00 114-18.00 33.0 3.6 115.4 250.3 1966 825 309 48.90 37-18.00 114-18.00 33.0 3.7 115.4 250.3	19	41	506	311	42.00	37-18.00	114-18.00	٠0	3.0	115.4	250.3
1944 320 930 57.00 37-30.00 114-42.00 .0 4.0 144.8 263.4 1953 809 2200 2.00 37-30.00 114-30.00 .0 4.5 127.3 262.5 1955 110 1007 28.00 37- .00 114-30.00 .0 4.4 145.7 240.3 1961 926 2145 20.00 37-39.96 114-20.00 .0 3.0 111.3 270.9 1966 818 928 56.90 37-24.00 114-18.00 33.0 3.7 112.1 255.7 1966 820 822 3.40 37-18.00 114-18.00 33.0 3.6 115.4 250.3 1966 825 309 48.90 37-18.00 114-18.00 33.0 3.7 115.4 250.3	19	43	116	355	.00	37-21.00	114-15.00	,0	4,0	109.4	252.3
1953 809 2200 2.00 37-30.00 114-30.00 .0 4.5 127.3 262.5 1955 110 1007 28.00 37- .00 114-30.00 .0 4.4 145.7 240.3 1961 926 2145 20.00 37-39.96 114-20.00 .0 3.0 111.3 270.9 1966 818 928 56.90 37-24.00 114-18.00 33.0 3.7 112.1 255.7 1966 820 822 3.40 37-18.00 114-18.00 33.0 3.6 115.4 250.3 1966 824 454 28.79 37-26.94 114-15.01 7.0 3.8 106.5 257.9 1966 825 309 48.90 37-18.00 114-18.00 33.0 3.7 115.4 250.3											
1955 110 1007 28.00 3700 114-30.00 .0 4.4 145.7 240.3 1961 926 2146 20.00 37-39.96 114-20.00 .0 3.0 111.3 270.9 1966 818 928 56.90 37-24.00 114-18.00 33.0 3.7 112.1 255.7 1966 820 822 3.40 37-18.00 114-18.00 33.0 3.6 115.4 250.3 1966 824 454 28.79 37-26.94 114-15.01 7.0 3.8 106.5 257.9 1966 825 309 48.90 37-18.00 114-18.00 33.0 3.7 115.4 250.3											
1961 926 2145 20.00 37-39.96 114-20.00 .0 3.0 111.3 270.9 1966 818 928 56.90 37-24.00 114-18.00 33.0 3.7 112.1 255.7 1966 820 822 3.40 37-18.00 114-18.00 33.0 3.6 115.4 250.3 1966 824 454 28.79 37-26.94 114-15.01 7.0 3.8 106.5 257.9 1966 825 309 48.90 37-18.00 114-18.00 33.0 3.7 115.4 250.3											
1966 820 822 3.40 37-18.00 114-18.00 33.0 3.6 115.4 250.3 1966 824 454 28.79 37-26.94 114-15.01 7.0 3.8 106.5 257.9 1966 825 309 48.90 37-18.00 114-18.00 33.0 3.7 115.4 250.3											
1966 820 822 3.40 37-18.00 114-18.00 33.0 3.6 115.4 250.3 1966 824 454 28.79 37-26.94 114-15.01 7.0 3.8 106.5 257.9 1966 825 309 48.90 37-18.00 114-18.00 33.0 3.7 115.4 250.3	19	66	818	929	56.90	37-24.00	114-18.00	33+0	3.7	112-1	255.7
1966 824 454 28.79 37-26.94 114-15.01 7.0 3.8 106.5 257.9 1966 825 309 48.90 37-18.00 114-18.00 33.0 3.7 115.4 250.3											
1966 825 309 48,90 37-18,00 114-18,00 33,0 3,7 115,4 250,3											
											250.3
									3.6	115.4	250.3

Earthquakes within 150km of Cedar City site: Mas 3.0 and above

1592	date	oris	time	lat-n	lons-w	derth	ឃទដ	k.m	82
1968	408	1608	14.65	37-22.40	114-17,12	.9	3.3	111.5	254+0
1971	1209	47	11.20	37-26.15	114-25.50	.0	3.8	121.9	258.8
1971	1226	603	55.30	37-20.52	114-19.86	.0	4.2	116.5	252.9
1972	1129	553	24.60	3866	114-40.25	٠0	3.9	146.4	285.9
1977	1205	9	46.20	38- 1.25	114-26.94	٠0	3.8	128.0	288.8
1977	1216	1652	36.30	38- 1.74	114-36.77	٠0	3.5	142.0	287+2

Earthquakes within 150km of St. George site, Mag 3.0 and above

rsee	date	oris	time	lat-n	long-w	derth	esm	km	82	
1859	828	0	.00	37-50,52	112-49.63		3.7	104.9	33,6	
	1130	500			113- 3.95			78.9		
	326	215			113-49.00			65.0		
	905	335			112-31,34			85.5		
	1026	610			112-55.43			137.9		
	1010	010	*	22 12127	112 001 70			14, 1,	2474	
1387	1205	1530	.00	37- 2.84	112-31.34		5.7	85.5	90.6	
1891	420	1355	.00	37- 6.38	113-34.41		5.0	०,8	305.4	
1891	913	348	.00	37-40.97	113- 3.95		3.0	78.9	28.0	
19 2	1117	1950	۰00	37-23.58	113-31,20		6.3	37.7	355.0	
19 2	1205	0	٠00	37-23.68	113-31.20		5.0	37.8	355.1	
19 3	1104	2340	.00	37- 6.38	113-34.41		3.0	9+8	305.4	
		1000			113-34.41		4.3			
		1000			112-26.02		3.7			
		530			113-42,79		4.3	51.0		
		525			113-42,79		3.0	61.0		
1717	الم متد الم	020	100	3/ 37,3/	110 7,,,,,		J + V	01+0	W 1 / 1 W	
1915	212	1950	.00	37-34.37	113-42.79		3.7	&1.0	340.5	
1915	213	830	.00	37-34.37	113-42.79		3.7	51.0	340.5	
1920	1126	0	٠٥٥	37- 6.38	113-34.41		4.3	9.8	305.4	
1921	602	2130	.00	37-40.97	113- 3.95		3.0			
	514				113-14.00		4.3	125.3	10.1	
1925	23.4	1347	۸۸	77. AG /A	113-55.80		3.7	מ מס	77.4 7	
									334.7	
1926		1951			112-35.47		3.0	84.4	67.8	
1926		1100			112-35.47		3.0	84.4		
1926		520			112-35.47		3.0	84.4		
1926	1001	1515	•00	3/-19,04	112-35.47		0,6	84.4	44+8	
1926	1004	2030	۰00	37-19.04	112-35.47		3.0	34.4	69.8	
1926	1023	0	.00	37-19.04	112-35.47		3.0	84.4	69.8	
1926	1112	340	.00	37-19.04	112-35,47		3.0	34.4	69.8	
1926	1116	1415	.00	37-19.04	112-35.47		3.0	84.4	59.8	
1927	1123	930	.00	37-19.04	112-35.47		3+0	84.4	69.8	
1973	120	1310	.00	37-50.52	112-49.63		5.6	104.9	33.6	
		1000			112-31.34			85.5		
1936		338			113-30.00		4,3			
1936		1025			112-57.48			51.4		
1936	921	620	•00	20- +00	113-18.00		4.7	106.1	8*8	
1937	218	415	.00	37-49.36	112-26.02		4,3	126.0	47.5	
1937	218	900	+00	37-49+36	112-26.02		3.7	126.0	47.5	
1937	221	830			112-26.02		3.0	126.0	47.5	
1937	226	130	.00	37-49.36	112-26.02		3.7			
1938	817	908	6.00	35-42.00	113-42.00		3.0	43.9	206.1	

Earthquakes within 150km of St. Georse site, Mas 3.0 and above

Rest	date	oris	time	lat-n	lons-w	derth	mad	km	an .
1938	1228	437	36,00	3700	11400		3.0	46.4	262.4
				36-30.00			3.0	148.7	245.5
1940			54.00		11500		3.0	135.1	267.4
1940	311	6	30.00	37- +00	11500		3.0	135.1	267.4
1940	407	842	00	3700	11500	•	3.0	135.1	267.4
1940	920			36-30.00			3.0	148.7	245.5
1940	1026				113- +00		3.0	113.3	22,2
1941	506				114-18.00		3.0	77.5	290,6
1942		2208			113- 3.95		5.0	78.9	28.0
1942	918	0	.00	37-40.97	113- 3.95		3.0	78.9	28.0
1942	918	200			113- 3.95		3.0	78₊9	28.0
1942	918	520			113- 3.95		3.7	78.9	28.0
1942		1116			113- 3.95		4,3	78.9	28.0
1942		1450			113- 3.95		5.0	78.9	23.0
1942	928	1330	•00	37-15.00	112-57.48		<i>5</i> .7	51.4	δ 5.1
1943	116	1150	18.60	37-40.97	113- 3,75		4.3	78.9	28.0
1943					114- 6.00		3.0	66.8	305.0
	1106		-		114-15.00		3.7	75.5	295.7
					113- 3.95		3.7	78.9	28.0
1944	131	424	28:00	36-54,00	112-24.00		3.0	98.0	100.1
1944	308	754	51.00	37-34.00	114- 9.00		3.0	84.6	315.7
1944	320	931	.00	37-12+00	11400		0.5	48.7	289.3
1544	503	2345	.00	37- 6,38	113-34.41		3.7	9,3	305.4
1945	11i	1156	٠00	37-24.00	114-54.00		8,8	131.4	286+9
1949	1102	230	.00	37- 6.38	113-34-41		4.7	9.8	305.4
1951	305	2300	.00	3700	112-32.40		3.7	84.2	94.2
1952	502	1015	.00	37-14.47	113-16.90		3.7	27.3	40.7
1953	518	703	2,00	3600	114-30.00		3.0	148.3	217.9
1953	809	2200	2.00	37-15.00	114-30.00		Q * Q	92.9	283.5
1953	1022	300	۰00	37-49.36	112-26.02		4.3	126.0	47.5
								90.7	
1955	1120	1057	54.00	37- 6.00	113-54.00		3.7	37.4	277.7
1959								136.2	
1959	721	1739	29.00	37- +00	112-30+00		5.5	87.7	94.0
1959	727	1115	٠00	37-32,36	113-10.51		3.7	60 ₊3	26.9
1961					114-20.00				312.0
1962								98.0	
1962	215	906	45.00	37- +00	112-54.00		3.0	52.3	96+7
1963								90.8	335.9
1963	619	838	44.87	39- 1.10	112-31,50	7.0	3.7	136.4	38.4

Earthquakes within 150km of St. George site, Mas 3.0 and above

year	date	oris	time	let-n	lons-w	derth	ಹಕತ	km	az
1964	101	1643	5.82	37-33.22	112-43.85	7.0	3.1	86.7	50.3
1964	117	15	3.45	38-11.11	112-37.30	7.0	3.3	146.7	31.2
1965	130				113- 6.95	7.0	3.2	62.7	31.3
1965					112-59.00		3.0	85.3	31.3
1965					112-26.49		3.0	141.8	40.5
4010	070	4.77	45 04	77 0/ 00	3 4 17 4 4 77 4	-1 0	** .	. 7	
1965	830				113-41.34		3.2	47.3	337.3
1966					114- 9.07		5.6	74.6	307.4
1966					114-11.48	7.0	4.7	75.9	304.1
1966	824				114-15.01		3+8	80.9	302.7
1966	903	753	18.55	36-30.27	112-15.26	7.0	3.5	125.6	117.1
1966	923	13	57.11	37-30.50	113-17.47	7.0	3.2	53.1	18.7
1966	1021	713	48.88	38-11.74	113- 9.45	7.0	4.2	129.8	12.3
1967	1113	1648	53.77	37-34.97	113-20.32	7.0	3,5	60.0	12.3
1968					112-16:55	7.0	3.0	143.7	48.0
1968					114-17.12	.9	3.3	79.5	296.4
10/0	117	/ 4.0	7 7.5	7/ 65 15	11ል ንግ ድጠ	7 0	3.1	165.6	261.5
1968	412				114-37.59	7.0		102.9	
1969	815				113-14.97	7.0	3.0	84.3	14,2
1970					112-28-13	7.0	3.9	142,9	38.8
1970	901				113-47.17	7.0	3.0	50+6	3117.9
1971	1110	1114	10.32	3/-41.00	113- 6.16	7.0	2.1	79.5	25,5
1971	1110	1124	34.89	37-41.87	113- 7.28	7.0	3 + 1	78.2	24.2
1971	1110	1338	14.19	37-42-14	113- 5.58	7.0	3.1	79.7	25.7
1971	1110	1410	23.01	37-48.03	113- 5.99	7.0	3.7	89.4	22.3
1971	1110	1443	58.51	37-48.25	113- 4.89	7.0	3.9	90.4	23.2
1971	1110	1608	37.48	37-42,01	113- 1.03	7.0	3.1	82,7	30.0
1971	1110	1941	33.29	37-48.37	113- 3.75	7.0	3.5	91.3	24.1
1972	427	153			113- 6.88	7.0	3.4	81.0	156.0
1972	427				113 - 8,99	7.0	3.4	80.7	158.3
1972	514				113-30.10	7.0	3.3	76.1	181,2
1972					113-50.77				
1///-	, , , ,	LVVI	1+10	0) EV*10	110 00,,,	, , ,	.,,,		OI 193
1972	902	1530	35.47	37-14.01	113-21.78	7.0	3,2	22.5	28.4
1972	1116	217	45.15	37-31.93	112-46.18	7.0	3.6	∂2 ₊5	50.1
1973	122	1024	55.94	37-11,54	112-57.90	7.0	3.0	48.5	71.7
1973	218	931	39.62	38- 5.87	113-10.57	7.0	3.7	118.9	13.2
1973					113- 8.74		3,3	122.6	14.1
1973	520	1,520	A.42	37-57.07	113-31.37	7.0	3.1	99.5	358.0
1975					113-28.30		3.2	110.2	.5
1978					112-29.38		3.0	139.3	39.1
					112-28.73			145.0	37.7
1979					113-10.05		3.2	30.9	20.2
3///	2.14.	120	27437	U/ 77+UU	220 10100	**	~ + ~.	D.A. + \	i., '. J + st.

Earthquakes within 150km of St. George site, Mag 3.0 and above

1592	date	oris	time	lat-n	long-w	der th	mad	k m	82
1979	805	1910	21.35	36-58.89	113-40.08	7.0	3.6	18.3	243.5
1979	815				113-49+01		3.2		220.3
1979	816				112-43.41		3.1	103.2	40.7
1980	428	2147	35.00	37-18.11	114- 5.00		3.4	59.9	
1980	803				112-58.16		3.1	95.1	28.6
1980	1221	1825	7.12	37-26.71	113- 4,93	1.7	3.3	56.1	39.4
1980	1227	434	15.52	37-30.83	113- 7.03	3.5	3.0	60.4	32,5
1980	1229	712	52.50	37-27.28	113- 4.92	1.7	3.1	56.9	38.8
1981	116	1026	29.90	37-26.52	113- 5.51	3.5	3.4	55.2	39.0
1981	116	1450	45.61	37-26.34	113- 5.28	1.2	3.2	55.2	39.5
1981	201	221	47.37	37-34.47	113-15.56	٠,0	3.3	61.0	19.0
1981	405	540	40.06	37-36.61	113-18.02	1.3	4.5	63.7	14.7
1981	808	620	17.09	38- 4.15	112-48,54	٠0	3.3	127.4	27.9
1936	122	338	.00	36-18:00	113-30.00	÷ ()	4.3	83.8	181.0
1938	817	908	6.00	36-42,00	113-42,00	. 0	3.0	43.9	206+1
1938	1217	843	.00	3600	11400	٠Ú	3.5	125.9	301.6
1939	1121	2240	.00	36-30.00	11500	.0	4.0	148.7	245.5
1939	1121	2340	49.00	36-30.00	11500	.0	3.0	148.7	245.5
1940	920	1221	54.00	36-30,00	11500	•0	3.0	149.7	245.5
1953	518	703	2.00	36- +00	114-30.00	, ()	3.0	148.3	217+9
1966	903	753	18.55	36-30.27	112-15.26	7.0	3.5	125.6	119.1
1966	121	920	40.90	36-12.00	113-54.00	26.0	3.7	101.7	201.4
1970	1124	1647	56.00	36-21,42	112-16.37	6.0	3.0	133.0	175.6
1972	427	153	3.14	36-23:31	113- 6.88	7.0	3.4	31.0	150.0
1972	427	841	13.59	36-22.77	113- 8-99	7.0	F. E	80.7	158,3
1972	514	238	40.66	36-22.14	113-30,10	7.0	3.3	76.1	181.2
1973	1226	618	16.60	36- 4.86	114-38.34		3.1	149.6	223.7
1979	815	211	30.27	36-44.36	113-49.01	7.0	3.2	45.9	220.3
1937	128	314	.00	37- 6,00	114-42.00	, ()	4.0	108.3	272+6
1940	310	1801	54.00	37- +00	11500	٠0	3.0	135.1	267.4
1940					11500				267.4
1940	407	842	.00	3700	11500	.0	3.0		267.4
1941	506	311	42.00	37-18,00	114-18,00	• 0	3.0	77.5	290+6
1943	610	2225	30.00	37-18.00	11500	$\cdot g$	3.5	137.4	281.4
1943	116	355	.00	37-21.00	114-15,00	٠,0	4.0	75.5	295.7
1944	320	930	57.00	37-30.00	114-42.00	٠0	Δ,ζ)	118.7	
1945					114-54.00			131.4	
1953					114-30.00			102,8	
1955					114-30.00				266.1
1961	926	2146	20.00	37-39,96	114-20.00	٠٥.	3.0	101.3	312.0

Earthquakes within 150km of St. George site, Mag 3.0 and above

Rest	date	oris	time	lat-n	lons-w	derth	กอร	km	3.5
1964	529	111	10.40	37~18.00	114-48.00	33.0	3.6	120.1	283.1
1964					115- 6.00		3.8	144.0	267.6
1966	818				114-18,00		3.7	82.0	297.8
1966	820				114-18.00		3.6	77.5	290.6
1966	824				114-15.01		3,8	80.9	302.7
1966	825	309	48.90	37-19.00	114-18.00	33.0	3.7	77,5	290.6
1966	827	1531	49.50	37-18,00	114-18.00	33.0	3.6	77.5	290.6
1967	507	1801	35.70	37- 2.40	11572	15.0	4.7	136.0	267.3
1967	509				11560		4.0	135.8	269,4
1968	408				114-17.12		3.3	79.5	296.4
1968	412	640	7.05	36-55,12	114-37.59	7.0	3.1	102.9	261.5
1969	1108	658	1.30	37- 8,76	114-58.68	4.0	3.3	133.3	274.3
1970	102	2001	31.60	37-25.70	114-57.77	.0	4.0	137.7	287.5
1971	1209	47	11.20	37-26,15	114-25.50	٠٥	3.8	93.0	296.8
1971	1226	603	55.30	37-20.52	114-19,86	.0	4,2	81.7	292,9
1972	1029	1901	24.90	37-13.38	114-53.34	5.0	3 . 4	125.3	279.5
1972	1029	1135	2.50	36-59.81	114-37.86	٠0	3,5	102.3	266-4
1972	1129	553	24.60	38- +66	114-40,25	٠0	3.9	149.2	315.3
1973	610	557	16.10	37-41.15	114-55,20	٠0	3.1	145.2	298.8
1973	914	2219	4.50	37-26.57	115- 3.18	. 04	4,1	145.8	287+2
1974	405	27	31.80	37-41.27	114-50.22	, ψ	3.8	139.0	300.4
1977	721	1606	22.10	37-13.92	114-58.31	7.0	3.3	133.7	278+4
1977	1205	9	46.20	38- 1.25	114-26,94	٠0	3.8	137.0	321.5
1977	1216	1652	36.30	38- 1.74	114~36.77	٠0	3.5	147.1	317.3
1979	812	1131	21.00	37-16.86	11572	٠٥	4.2	138.1	280.5
1979	812	1253	3.90	37-19.20	115- 3.83	.0	3.1	143.4	281.8
1979	812	1546	39.50	37-16.62	115- 3.48	.0	3.8	142.0	280.0
1979	813	950	2.20	37-24.18	114-52.13	. 0	3.4	128.9	287.4
1979	813	1823			114-56.70	٠0	3.8	134.3	284.9
1979	813				115- 4.19	.0	3.0	141.3	273.7
1979	814	303	16.20	37-14,88	115- 8.27	.0	3.9	148.5	278,3
1979	816				114-57.65		3.7	134.5	202.7
1979	816				115- 2.46	.0	3.1	141.3	

APPENDIX J

LISTING OF UNIVERSITY OF UTAH EARTHQUAKE CATALOGUE
WITHIN A RECTANGULAR AREA ENCOMPASSING ST. GEORGE
AND CEDAR CITY (1850-1981)

SW Utah Earthquakes (36.75-38.00N,112.00-114.25W): 1850-June 1962

yr date orig	s time lat-n	long-w ma	es int	comments
1050 000	~~ ~~ ~~			
1859 828		112-49-63 3		
1878 1130 030	00 37-40 . 97	113- 3.95 3		PEHEDAL CHOCKE (1)
				SEVERAL SHOCKS (1)
				HEBRON, UT LOC FROM (16)
1885 905 033	35 37- 2.84	112-31+34 3	3.0 3	
1887 1205 153	30 37- 2.84	112-31.34 5	5.7 7	INT=6-7
	55 37- 6.38			ASSGN ST. GEORGE, UT (1)
	48 37-40.97			
	50 37-23,58	113-31.20 6	8 2.4	INT=7-8, FOUR SHOCKS (1)
19 2 1205	37-23.68	113-31.20 5	5.0 6	NUMEROUS A'SHOCKS (1)
19 3 1104 234	40 37- 6.38	113-34.41 3	3.0 3	F'SHOCK?
	00 37-6.38			
	00 37-49.36			
	30 37-34.37			(1)
1914 1221 052				LOC ASSUMED, A'SHOCK? (1)
		110 /2//		
1915 212 193	50 37-34+37	113-42.79	3.7 4	A'SHOCK? (1)
1915 213 083	30 37-34.37	113-42.79 3	3.7 4	A'SHOCK? (1)
1920 1126 006	00 37- 6.38	113-34,41	4.3 5	1NT=5-6
1921 602 213	30 37-40,97	113- 3.95 3	3.0 3	
1924 101 23	15 37-19+04	112-35.47	2.3 3	
1925 714 13-	17_10 L0	113-55.80 3	3.7 4	
1926 515 193		112-35,47		
1926 601 053		112-35.47		SWARM (1)
1926 605 110		112-35.47 3		
1926 622 034		112-35,47		
1720 022 03	40 3/-17.04	112-33+4/ 2	2+3 2	DWMRT (1)
1926 628 069				TWO SHOCKS, SWARM (1)
1926 712 053	20 37-19.04			SWARM (1)
1926 715	37-19.04	112-35.47 2	2.3 2	SWARM (1)
1926 1001 15	15 37-19.04	112-35.47	3.0 3	SWARM (1)
1926 1004 203	30 37-19.04	112-35,47	3.0 3	SWARM (1)
1926 1023 006	00 37-19.04	112-35,47 3	3.0 3	SWARM (1)
1926 1112 034		112-35.47		
1926 1116 14		112-35.47		
1927 1123 024				TWO SHOCKS, SWARM? (1)
1927 1123 093		112-35.47		SWARM? (1)
1071	ту д — — — — — — — — — — — — — — — — — —			443
1931 410 083		113- 3,95 2		(1)
1932 618 203		112-49-63 2		
1933 120 13		112-49+63 5		
1934 1225 100		112-31.34 3		
1935 1205 21	15 37-37,30	112- 5.20 2	2.3 2	

```
er date oris time lat-n
                             long-w mag int
                                                comments
                    37-15.00 112-57.48 4.3 5 INT=5-6, TWO SHOCKS (1)
1936 509 1025
                    38- .00 113-18.00 4.7 5 INT ASSUMED, NEV (5)
1936 921 0620
1937 218 0415
                    37-49.36 112-26.02 4.3 5 INT=4-5, SWARM (1)
1937 218 0630
                    37-49.36 112-26.02 2.3 2 SWARM (1)
1937 218 0900 -
                    37-49.36 112-26.02 3.7 4 SWARM (1)
                    37-49.36 112-26.02 3.0 3 SWARM (1)
1937 221 0830
1937 226 0130
                    37-49.36 112-26.02 3.7 4 SWARM (1)
1937 313 1140
                    37-49,36 112-26,02 2,3 2 SWARM (1)
1937 401 0441
                    37-49.36 112-26.02 2.3 2 SWARM (1)
1938 1228 0437 36.00 37- .00 114- .00 3.0 0 (6.8)
1940 1026 0124 24.00 38- .00 113- .00 3.0 0 (6.8)
1942 830 2208
                    37-40,97 113- 3,95 5,0 6 SWARM (1)
                    37-40.97 113- 3.95 3.0 3 SWARM (1)
1942 918 0000
1942 918 0200
                    37-40.97 113- 3.95 3.0 3 SWARM (1)
1942 918 0520
                    37-40.97 113- 3.95 3.7 4 SWARM (1)
1942 926 0700
                    37-40,97 113- 3,95 2,3 2 SWARM (1)
1942 926 1116
                    37-40.97 113- 3.95 4.3 5 SWARM (1)
                    37-40.97 113- 3.95 2.3 2 SWARM (1)
1942 926 1136
1942 926 1450
                    37-40.97 113- 3.95 5.0 6 SWARM (1)
                    37-15.00 112-57.48 3.7 4
1942 928 1330
1943 116 1150 18,00 37-40,97 113- 3,95 4,3 5 INT=4-5 (1,6,8)
1943 306 2014 30.00 37-24.00 114- 6.00 3.0 0 (6,8)
                    37-21.00 114-15.00 2.3 2 1ST OF THREE SHOCKS (6.8)
1943 1106 0345
                    37-21.00 114-15.00 3.7 4 (6.8)
1943 1106 0354
1943 1106 0356
                    37-21.00 114-15.00 2.3 2 (6.8)
                    37-40.97 113- 3.95 3.7 4 SWARM? (1)
1943 1209 1606
1943 1209 1821
                    37-40.97 113- 3.95 2.3 2 SWARM? (1)
                    37-40.97 113- 3.95 2.3 2 SWARM? (1)
1943 1210 0117
1943 1210 0730
                    37-40.97 113- 3.95 2.3 2 SWARM? (1)
1944 131 0424 58.00 36-54.00 112-24.00 3.0 0 (6.8)
1944 308 0754 51.00 37-36.00 114- 9.00 3.0 0 (6,8)
1944 320 0931 00.00 37-12.00 114- .00 3.0 0 (6.8)
1944 503 2345
                    37- 6.38 113-34.41 3.7 4
                    37-40.97 113- 3.95 2.3 2
1944 613 1048
1945 1118 0107 41.00 38- .00 112- .00 3.0 0 (6)
1949 1102 0230
                    37- 6.38 113-34.41 4.7 6 LOC ST.GEORGE, UT(1.5.6.8)
                    37- .00 112-32.40 3.7 4 LOC ASSUMED (8)
1951 305 2300
1952 502 1015
                    37-14.47 113-16.90 3.7 4 1ST OF TWO SHOCKS (8)
                    37-14,47 113-16,90 2,3 2 (8)
1952 502 1615
                    37-49.36 112-26.02 4.3 5
1953 1022 0300
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SW Utah Earthquakes (36,75-38,00N,112,00-114,25W): 1850-Jone 1962

яr	date	oris	time	lat-n	lons-w	derth	mas	no	설공원	dın	rms	a
63	217	1734	20.64	37-48.10	113-54.21	7.0*	3.7	5	326	216	.26	D
64					112-43.85				185		.63	
64	206				112-58.19				138		+66	
64	828				113- 8.46				135		.37	
65					112-50.86				102		.54	
65					113- 6.95		3.2		107		•55	C
65					112-26.88				309		, 33	D
65					112-59.00				137		•08	
65					113- 2.33				141		.44	
65	830	43	15.20	37-26,92	113-41.34	7.0*	3.2	6	166	192	.73	C
72	1204	1 ጋጋ ለ	40 OE	77 11	113-19.62	7.0*	2 0		161	כם	.36	ā
66					113-43.06				169		• 30 • 87	
66					112-22.95				107		+46	
					112-22.73				224		+40 +52	
66												
66	318	1312	40129	3/~3/+08	112- 9.91	7.0*	ಪ+ವ	/	115	117	+67	Ļ
66	526	2348	52,22	37-53.56	112- 9.51	7,0%	2.3	5	144	113	.78	D
66					112-18.30		2.1		203		.27	
66	716	943	18.93	37-52,97	112-36.50	7.Q*	2.2	6	110	97	+58	C
66	816	1802	32.85	37-27.81	114- 9.07	7.0%	5.64				،71	€
66					114-11.48		4.7W	6	126	171	,32	B
66	923	13	57.11	37-30.50	113-17.47	7.0*	3.2	7	215	162	• 56	C
67	517	658	35,45	37-51.06	112-18+13	7.0*	2.7	4	203	115	.17	Ü
67	1113	1648	53.77	37-34.97	113-20.32	7.0*	3,5		225		.41	B
88	115	303	26,86	37-25.62	113-26.73	7.0*	2.4	5	219	172	.39	Ţi
68	205	1417	27+40	37-26.17	113-57.58	7,0#	2.8	6	177	216	.63	C
40	718	100	07 EE	7749. 47	11757.00	7 64	2 5	A	202	227	.34	T)
68					113-56,20		2.5		207		-	
68					112-16.55				148 144		+45 ^0	
88 88					112-33,18 112-23,19				181		•63	
00	1008	1013	4/+61	3/~30+63	113-58.75	/ . ∪#	2.9	11	7.3	120	+4/	L
69	815	30	30.57	37-47.48	113-14.97	7.0*	3.0	8	153	20	1.00	Ţ!
69	1112	2011	40.41	37-46.06	112-25.90	7.0*	2.9	12	96	57	. 83	Ţi
69	1117	1637	22.51	37-31.68	113-23.31	7.0*	2.1	4	153	32	.08	Ţ)
70	524	154	34,08	37~52.19	112-28.33	7.0*	1.5	9	99	56	1.00	D
70	524	156	6.84	37-56.64	112-24.25	7.0*	2.7	6	109	65	+99	D
מכ	001	1150	17 01	77-24 47	113-47.17	7.0*	7 0	10	89	ΛŌ	1.01	n
					113-44.30							
70					113-44+30				108		.84	
70 71	1025				112-23.38				104 119		.42 1.24	
71					114- 6.93			9		104	+59	
11	247	104/	20+70	70.07	114. 0+13	7+47	ال جائد	7	72	104	+ J7	U

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er date oris time
                   lat-n
                            long-w
                                     derth mas no sar dmn rms o
71 1110 1114 10.32 37-41.60 113- 6.16 7.0* 3.1\(\mathbf{i}\) 13 74
                                                          3
                                                            .49 B
71 1110 1124 34.89 37-41.87 113- 7.28
                                      7.0* 3.1W 15 72
                                                          5
                                                             .44 B
71 1110 1338 14.19 37-42.14 113- 5.58 7.0* 3.1W 10 101
                                                             ,40 B
                                                          3
71 1110 1346 55.32 37-42.05 112-58.74 7.0* 2.6W 6 167
                                                          8
                                                            +48 B
71 1110 1410 23.01 37-48.03 113- 5.99 7.0* 3.7W 24 80
                                                        14
                                                             .63 C
71 1110 1443 58.51 37-48.25 113- 4.89 7.0* 3.4W 20 78
                                                        90
                                                             .51 C
71 1110 1454 31.50 37-20.66 113-20.73 7.0* 2.4
                                                             .08 D
                                                 4 324
                                                        44
71 1110 1608 37.68 37-42.01 113- 1.03
                                     7.0* 3.1W 13 77
                                                         5
                                                             .51 C
71 1110 1703 12.63 37-20.07 113- 4.54 7.0* 2.9
                                                 5 323
                                                             .56 D
                                                        37
71 1110 1907 34.51 37-42.28 113- 3.92 7.0* 2.94 7 107
                                                             +36 B
                                                          3
71 1110 1941 33.89 37-48.37 113- 3.75 7.0% 3.5W 17 81
                                                        14
                                                             .49 B
71 1115 1230 7.25 37-36.03 113-11.62 7.0% 2.1
                                                  4 313
                                                        13
                                                             .19 D
71 1130 337 43.07 37-37.17 113- 5.47 7.0% 2.7
                                                  6 319
                                                             .63 C
                                                         6
                                      7.0% 2.4
71 1208 554 22.83 37-41.83 113-51.81
                                                 5 286
                                                        70
                                                             .06 D
72 123 1129 19.37 37-49.67 113-10.44 7.0% 2.89 11 78
                                                        19
                                                             .47 B
72 317 2045 24.91 37-23.68 112-49.52 7.0* 2.2
                                                  6 284
                                                        37
                                                             .43 C
                                                             .27 B
72
   513 1340 25.66 37-25.87 113-19.40 7.0* 2.3
                                                 4 323
                                                        35
72 701 1050 19,75 37-27,45 112-18,83
                                     7.0* 2.4
                                                  5 253
                                                        70
                                                             .40 D
72
   701 2345 36.10 37- 5.29 113-29.51 7.0* 2.3W 8 124
                                                        20
                                                             .43 B
72 702 1434 20.71 37-18.47 113-49.85 7.0% 2.5W 10 107
                                                             +69 C
72
   702 2007 1.15 37-20.15 113-50.77 7.0* 3.0W 9 102
                                                             .85 D
                                                        42
72
   805 1115 42.00 37-58.08 112-27.33 7.0* 2.1
                                                  6 201
                                                        63
                                                             +24 A
72 902 1530 35.47 37-14.01 113-21.78 7.0* 3.2
                                                 7 324
                                                        55
                                                             .50 C
                                                 6 269
72 1006 828 44.48 37-42.03 113-15.74 7.0* 2.8
                                                        17
                                                             .38 C
72 1006 1211 8.65 37-35.91 113-17.84 7.0* 2.7
                                                 6 301
                                                        21
                                                             .36 B
72 1017 2334 57.56 37-41.12 112-55.80 7.0* 2.9
                                                  8 216
                                                        12
                                                             .56 €
72 1116 217 45.16 37-31.93 112-46.18 7.0* 3.6W 10 215
                                                        30
                                                             .49 B
72 1128 2053 24.50 37-34.32 113- 1.12 7.0* 2.2
                                                 7 114
                                                        12
                                                             .57 C
72 1130 2322 55,66 37-28,03 113-24,22 7,0* 2,4
                                                 6 150
                                                        25
                                                             •38 B
73 122 1024 55,94 37-11,54 112-57,90 7,0% 3,0
                                                  5 315
                                                        54
                                                             .34 D
73 122 1123 54.01 37-21.19 113-43.56 7.0% 1.6W 6 137
                                                        33
                                                             .33 B
73
    512 2339 50.60 37-22.94 114-12.56
                                      7.0* 1.2
                                                 9 106
                                                        75
                                                             .58 C
73
   514 101 9.84 37- 9.89 114- 7.18 7.0% 2.4 10 117
                                                        66
                                                             .69 C
73 520 1620 6.62 37-57.07 113-31.37 7.0% 3.1W 10 79
                                                         50
                                                             .74 C
73
    714 1054 1.28 37- 9.00 112-45.20 7.0* 2.3
                                                 6 102
                                                        16
                                                             ,91 D
73 1102 1644 5.47 37-58.30 112-19.89 7.0* 2.2
                                                  5 207
                                                        72
                                                             .50 D
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SW Utah Earthquakes (36.75-38.00N,112.00-114.25W): Oct 1974-Dec 1981

ar	date	oris	time	lat-n	long-w	derth	mad	no	452	dmn	rms a
74	1225	813	40.38	37-52.10	112-59.39	7.0*	2.7	4	165	22	.49 D
75	112				112-54-63	7.0*			161	38	.34 D
					113-19.00	7.0*			222	41	.44 D
76					112-14.47	7.0*			239	58	.36 D
76					112-51.79	7.0*			171	32	.51 D
76	808	346	55.77	37-28.72	112-21.12	7.0*	2.1	5	259	66	.21 D
76	1016	438	31.08	37-51.76	112-43.65	7,0*	2.1	7	194	36	.38 D
76	1025	2151	30.25	37-52.52	112-42.15	7.0%	2.6	4	202	39	.12 D
76	1222	1937	12.02	37-46,02	112-57.43	7,0*	1.7	5	157	14	.54 D
77	217	2321	9.10	37-34.30	112-37.15	7.0*	1.3	5	266	41	48 D
77	219				112-35.18	7.0*	2.1		267	44	.61 D
77	221				112-30.83	7.0≭			209	47	.55 D
77	324				112-51.91	7.0*			322	77	.05 D
7 7	423				112-25,65	7.0x	1.9		266	57	.34 D
77	502	2148	53,18	37-32.09	113-35,91	7.0*	1.5	7	317	49	,40 D
											
77					113-21.07	7.0*			274	26	.38 D
77					113-15.91	7.0*			248	24	.06 D
77					112-20.42	7.0*			203	65 66	.39 D
77					113-23.57	7.0*			306	29	.44 D
77	531	2128	15+14	3/~42,25	113-17.87	7.0%	1.4	6	282	20	.36 D
77	704	1.442	42.22	37-50.37	113-22,81	7.0*	1.7	5	265	32	.12 D
77	709				112-23.85	7.0%			109	53	.25 C
77	709				112-24.79	7.0*			108	55	.70 D
77	709				113- 4.76	7.0*			196	17	.47 D
77	718				112-27.47	7.0*			192	58	.25 D
• •		J.,	,,,,,,	5. 01,00	2.44	,,,,,	,	_			
77	1029	938	14.92	37-55.77	113-26:36	3.1	2.0	5	263	43	.06 D
77	1114	810	41.02	37-43.61	113- 9.11	1.5	1.2	6	247	9	.33 D
77	1114				113-10.55	1.5	1.2		261	10	.39 D
	1115	724			113-10.16	1.5	1.4		257	10	.48 D
77	1115	1738	18.11	37-43.12	113- 7.99	1.5	1.2	8	244	7	.44 D
77	1116	1902	56.93	37-45.41	113-13.59	1.5	.8	6	256	16	.15 D
					113- 9.33						.44 D
					113- 8.26				234		.28 D
					113-13-59						.56 D
					113-16.05				287	17	,62 B
.,		40 W W /	******	U/ 101U/	220 X0100	2 4 0 0	* * *	U	a s		,
77	1125	945	14.69	37-54.39	113- +64	7.0*	1.2	9	179	26	.22 €
					113-11.30				219	12	.45 D
					113- 9.49				238	11	.22 C
					113-11.39				259		.51 D
77	1206	1903	30.07	37-42+07	113-13.17	11.9	2.0	9	223	13	.79 D

ur date oris time lat-n lons-w deeth mad no dae dmn rms a 77 1227 1928 56.19 37-46.77 112-31.07 7.0* 2.6 11 101 49 .32 C 78 111 1859 16.76 37-41.01 113-16.80 7.0* 1.5 6 287 18 .44 D 6 304 78 112 2235 10.30 37-37.50 113-20.45 7.0* 1.1 24 .36 D 8 224 78 118 1748 54.46 37-46.52 113- 8.27 7.0* 1.3 12 +24 D 118 1856 50.00 37-39.28 113-12.97 7.0* 1.3 5 300 13 .51 D 78 118 2238 48.82 37-45.52 113- 8.74 1.2 2.2 9 232 11 +33 D 9 242 11 .31 D 78 121 2015 ·21 37-44.66 113- 9.62 4.6 1.1 714 8.96 37-43.65 113- 9.23 7.0* 1.7 10 216 9 .50 D 78 123 225 919 27.12 37-52.54 113- 6.41 7.0* .6 7 204 22 .23 D 78 228 1037 6,33 37- 7,39 113-22,02 7,0* 1,6 7 341 .50 D 18 402 1217 10.58 37-50.43 113- 4.98 7.0* 1.3 10 197 .32 D 78 402 1219 24.17 37-50.37 113- 3.88 7.0* .9 10 191 18 .37 D 6 327 102 78 522 750 51.40 37-35.17 112-32.27 7.0* .7 .12 D 347 6.76 37-43.19 113- 6.71 7.0* .7 7 231 .65 D 78 606 6 6 284 78 625 258 7.12 37-43.78 113-23.30 7.0* 1.3 28 .14 D 7 224 706 1016 12.23 37-52.05 112-30.61 7.0* 2.1 53 .31 D 78 707 1002 41,40 37-54,31 112-35,84 7,0* 1,8 5 163 48 .62 D 78 707 1244 30.02 37-51.77 112-33.72 7.0* 1.4 6 169 47 .36 C 78 1005 858 26.50 37-58.43 112-42.44 7.0* 1.3 4 330 57 .39 D 7 158 78 1011 856 54.99 37-33.03 112-54.38 3.5 1.7 19 .45 C .53 D 78 1025 2334 50.43 37-52.93 113-51.88 7.0* 1.4 8 271 73 78 1108 229 49.12 37-56.72 112-32.91 7.0* 1.0 7 308 51 .07 D 78 1111 2232 33.06 37-46.99 113- 1.48 11.5 1.2 10 174 12 .51 D 9 177 23 ,26 C 78 1120 900 28.88 37-13.40 112-53.42 7.0* 1.7 78 1201 609 16,77 37-50,44 113- ,99 12,5 6 178 18 .42 D 36 ,10 C 78 1203 2210 30.69 37-59.14 112-55.43 7.0* 1.1 6 165 5 266 37 ,27 D 78 1212 732 10,22 37- 6,89 113-13,66 7,0* 1,5 78 1214 2151 18.05 37-28.74 112-38.46 7.0* 1.8 6 178 53 .21 D 6 256 78 1218 749 7.49 37- 4.68 112-31.54 7.0* 1.7 27 .53 D +46 D 78 1218 1009 14,95 37-34,72 112-55,14 7,0* 2,8 6 284 17 8 284 53 .54 D 78 1218 1016 34,57 37-26,38 112-32,55 7,0* 2,1 5 221 58 .46 D 78 1220 559 34.11 37-26.24 112-25.88 7.0* .9 305 49.59 37-28.39 112-36.21 7.0* 1.7 8 183 46 .24 C 78 1221 107 836 12,50 37-40,47 113- 4,82 .8 1.6 8 208 1 .33 D 4 320 ,10 D 79 109 2237 6.32 37-46.85 112-18.17 16.0 64 • 0 9 217 .13 D 79 .0 3.2W 11 112 928 59.54 37-44.33 113-10.05 79 119 1807 47.77 37-45.20 112- 7.57 7.0* 2.0 5 220 68 .42 D 326 38.90 37-59.19 112-37.77 7.0* .8 8 191 51 .33 D 79 122 70 9 281 .36 B 943 47.87 37-20.64 113-52.35 7.0* 2.1 80 131 .37 D 7 278 17 79 203 543 56.93 37-34.39 112-55.02 7.0* 1.5

er date oris time lat-n long-w death mag no gas dmn rms a 208 2237 34.19 37-54.68 113-32.44 7.0* 1.8 6 282 49 .63 D 79 139 40.66 37-58.33 113- .38 7.0* 1.8 10 182 33 .52 D 79 322 230 10.74 37- 6.70 112-19.83 7.0* 1.8 45 .49 D 6 251 79 322 1446 30,20 37-57,26 112-36,67 7,0* 1,3 7 199 50 +39 D 611 1254 12.85 37-32.71 112-46.27 7.0* 1.6 .33 D 8 175 58 79 703 1904 49.39 37-48.04 112- 5.54 7.0* 2.0 10 140 63 .56 B 79 722 733 33.67 37-58.32 112-13.00 7.0* 1.1 9 312 43 .28 D 79 805 1910 21.35 36-58.89 113-40.08 7.0* 3.6W 8 291 75 .29 D 79 805 1932 35.49 36-48.32 113-55.11 7.0* 2.5 10 219 100 .84 D 79 806 1759 18.85 37-56.45 113-40.62 7.0% 1.4 6 260 93 .20 D 79 816 1733 4.98 37-45,58 112-43,41 7.0* 3.14 13 99 31 .25 C 79 816 1737 45.20 37-45.41 112-43.36 7.0* 1.8 12 100 31 .50 D 32 816 1743 13,91 37-45,50 112-43,16 7,0% 1,9 12 156 .23 C 816 1744 23.69 37-45.70 112-43.12 7.0* 2.2 79 15 101 32 .35 C .76 D 79 816 1854 18.82 37-42.86 112-22.48 7.0* 1.2 7 215 61821 1220 33.37 37-50.07 113- 3.73 3.5 1.4 10 190 .23 D 17 79 826 958 10.15 37-56.58 113-38.45 7.0* 1.2 10 158 58 .40 D 79 826 1941 27.20 37-16.72 112-48.28 7.0* 1.7 12 167 49 ,42 C 79 904 1822 31.60 36-55.18 113-33.58 7.0* 2.3 8 204 142 .46 D 79 916 46 6.23 37-34.01 112-18.66 7.0* 2.3 11 144 .37 D 79 920 840 27.39 37-44.28 112-26.48 7.0* 1.9 7 250 55 .58 D 79 1009 1235 23.96 37-38.61 112-56.96 7.5 1.3 10 252 11 .16 C 79 1112 816 17.58 37-48.11 112-51.58 20.1 1.1 23 8 193 .42 D 79 1118 2301 55.85 37- 7.21 113-24.70 7.0* 2.0 10 283 86 .44 D 79 1121 1650 5.08 36-54.41 112-51.12 7.0* 2.7 9 261 87 .69 D 79 1221 448 35.80 36-58.99 112-30.38 7.0* 2.2 7 187 28 .48 D 79 1224 537 53.18 37- 1.11 112-13.27 7.0% 1.2 9 175 53 .40 D 79 1226 138 20,89 37-31,31 113- 6,30 7,0* 1.7 17 9 332 .44 D 921 57.87 37-49.78 112-53.32 7.0* 2.3 14 134 .63 D 45 41.65 37-59.65 112-26.02 7.0* 1.4 80 112 7 303 42 .20 D 80 114 744 1.61 37-18.68 113-33.33 7.0% 2.1 4 339 59 .16 D 119 1950 47.54 37-51.30 112-53.31 7.0* 2.4 7 175 .40 C 80 25 80 315 1113 38.95 37-31.37 113- 9.52 10.3 1.9 8 334 18 .32 D 80 428 2147 35.00 37-18.11 114- 5.00 7.0* 3.4W 10 194 67 .39 D .82 D 80 429 1825 10.11 36-55.60 113-29.43 7.0% 2.7 9 224 60 80 603 1157 7,12 37-44,09 112-51,96 10,1 1,9 .27 C 8 140 19 .61 D 80 615 1034 58.09 37-39.62 113-16.28 16.1 2.4 5 273 18 80 731 433 13.25 36-50.48 112- 7.70 7.0* 2.3 6 215 49 .37 D 608 28.57 37-48.43 112-58.16 16.5 3.1W 8 152 .27 C 80 803 17 80 1221 1122 56.86 37-29.36 113- 4.84 9.2 2.1 11 154 20 .35 C

чr	date	oris	time	lat-n	lons-w	derth	mas	no	452	dmn	rms	Q
80	1221	1825	9,12	37-26.71	113- 4.93	1.8	3.3W	18	114	25	.41	C
	1227				113- 7.03	3.6	3.0W			18	+32	
					113- 5.68	2.8	2.8W			20	٠35	
	1228				113- 5.16	9,2	2.3		202	18	.32	
	1229				113- 4.92	1.8	3.1W			24	.50	
81	102	1720	22.76	37-36.52	113-17.36	1.1	2.0	11	252	20	.47	ď
81					113- 6.14	1.5	3.4W	25	113	25	+60	d
81					113- 6.49	2.2	2.5	19	150	16	٠40	C
81	116	1207	42.60	37-33.78	113- 6.61	6.2	2.1	14	198	13	.42	đ
81	116	1413	56,70	37-30.86	113- 5.30	• 1	2.2	19	151	18	.49	Ç
01	117	1.450	4E 09	77 0/ /0	117 5 74	4 /	3.2₩	7	-7	25	. 59	
81					113- 5.74 113- 5.49	1.6 1.5			7 151	23 17	+37 +45	
81 81					113- 5.01	2.8	2.4		150	16	.49	
81					113- 6.16	9,6	2.0		199	13	,54	
81	118				113 7.44	2.5	1.4		148	13	.53	
01	7.7.0	40	23+47	97. 99±00	110- /+44	4.4	7 4 7	si /	170	7.7	+00	u
81	118	556	22.53	37-34.41	113- 7.09	6.1	2,2	25	198	12	₊ 37	d
81	119				113- 5.56	3.6	2.1		153	20	. 44	C
81					113- 6,19	2.7	2.2		152	19	.45	
81	120				113- 5.78	1.9	2.3		152	18	+43	
81	120				113- 6.35	1.4	1.6		201	17	.52	ជ
81	120	1302	37.52	37-31.95	113- 7.04	1.8	1.5	16	200	16	.63	þ
81	121	218	11.49	37-30.05	113-13.43	7.0*	1.0	11	299	23	• 36	d
81	121	653	25,66	37-28.46	113- 7.51	6.5	1.8	15	209	22	.51	ď
81	122	837	17,48	37-58.95	112-58.83	7.0	1.9	9	112	35	.37	C
81	201	221	47.49	37-34.09	113-14.56	٠2	3.84	26	111	19	. 53	d
81	202	751	71 00	77-71 77	113-15.80	1.6	1.5	12	195	20	+49	ત
81	204				113-13,60	+7	2.6		148	21	+36	
81					113-17.13	3.6	2.8W			21	.64	
81					113-14.67	1.4	1.7		195	20	.49	
81					113-13.77					25	,44	
9.1	207	405	31+07	3/733+30	113-1/+40	7+2	1+4	7	202	23	: 44	u
81	217	1325	22.01	37-33.53	112-30.36	7.0*	1.5	9	216	50	.20	c
81	218	452	47.33	37-44.90	113-58.05	7.0x	1.4	5	172	45	.40	ď
81	226	1053	18,68	37-17.09	113-20.22	7.0×	1.4	13	226	49	.71	ď
81	227	1557	54.90	37-38.86	113-14.74	.0	1.5	11	201	15	.43	d
81	302	29	12.38	37-52,65	112-32.42				103	50	.47	
n i	711	1750	1 74	77_70 47	447.40.70	ግ ሊታ	1 A	17	101	777	nn.	بر
81 81					113-49.38 112-23.52				194 127		.87 .18	
81	314				112-23.32				178	55	.25	
81					112-30.43						+20	
81					112-23.81		1.7			59	.33	
OI	014	101	VA-11	5/ 7/+/0	117 20+01	7 + V P	1 + /	. ·	7.1.4	JI	+ 444	٠.1

Яľ	date	oris	time	1at−n	lon⊴-w	derth	mes	uo	ತಕ್ಕ	dmn	rms	G
81	314	1824	34.48	37-48.26	112-24.12	7,0*	2.3	23	121	72	.40	ส
81					112-22.68	7.0*			190	69	. 45	
81					112-25.03	7.0%			118	73	.29	
81	315				112-24.74	7.0*			257	72	.26	
81	316				112-29.43	7.0*		31		55	٠55	
	- ,			2, 23,,,	242 271 16				• •			-
81	316				112-30.43	7.0*	1.8		164	55	.58	ď
81	316	609	40.08	37-56+48	112-28.91	7.0*		20	165	56	.28	d
81	318				112-29.54	7.0×			164	55	.30	
81	318	638	11.07	37-56.81	112-29.73	7.0≭	1.7	22	164	57	+27	ď
81	318	1411	53.31	37-56.64	112-29.41	7.0*	1.5	18	114	55	.28	c
81	329	706	57.85	37-56.21	112-29.89	7.0*	1.6	12	165	56	.35	d
81	405			37-37.85	113-18.20	1.4	4.5W		144	20	, 34	
81	405			37~37,32	113-14.75	2.7	2.0		192	16	.42	
81	405	640		37-36.80	113-16.54	4.3	2.1		154	19	.41	
81	405	1052			113-16.45	2+6	1.9		154	19	.48	
81	406	2038	29.43	37-37.98	113-17.87	1.1	2.2	11	152	20	.51	d
81	407	2111	46.87	37-37.95	113-17.91	7.0*	1.4	11	199	85	.43	d
81	408	2325	13,46	37-38.39	113-17.90	7.0*	1.4	10	193	84	.36	ď
81	410	1343	27,92	37-36.40	113-15.75	5.1	1.4	10	201	18	+43	ರ
81	410	1345	21.29	37-37,00	113-16.35	3.1	1.4	14	200	18	. 49	ď
81	410	1352	3,29	37-37.34	113-15.75	4.9	•9	1.3	199	17	. 49	ส
81	411	737		37-36.09	113-16-14	2.3	1.8		155	19	.48	
81	413			37-37.29	113-14-64	4.9	1.7		152	16	,62	
81	413			37-37.23	113-14.70	5.4	1.5		200	16	.62	
81	413				113-15.27	4.8	1.1		202	18	.49	
Οï	417	JZ.4	42+01	27-30+00	113-13+27	7:0	1+1	12	202	10	177	u
81	414	1330			113-13.87	5.7	1.4		200	15	.40	
81		646			113-17.54	5.1	1.2		179	18	.50	
81			-		113-14.75	2.7	+9		210	17	, 48	
81					113-53.86	7.0%			294		.42	
81	423	254	55,35	37-37.41	113-17-18	1.6	1.7	9	196	19	•46	ď
81	423	626	46.23	37-39.75	113-18.42	7 .0 米	1.6	7	196	82	.52	ď
81	423	629	14.50	37-37.70	113-18-14		2.0		153	21	,64	
81	423				113-15,53	3.9	2.0	10	201	18	£43	
81	423				113-15.46	.6	1.6		196	17	٠57	ď
81	504				113-17.35	4.3	1.9		150	19	.57	
81	504	440	77.71	37 <u>-</u> 39 A2	113-17.72	1.5	1.9	17	186	19	، 52	r!
81					113-17-72	7.0*			220	42	+67	
81	611				113-30.07		1.7		189	21	.50	
81	627				113-11-77		1.9		97	15	.48	
81	705				113-18.70				179	43	.59	
O.Y	743	720	14,91	37-70+13	TT9-T0+\A	/ + V A	T + 1	17	1/7	42	+ 117	Ü

SW Utah Earthquakes (36.75-38.00N,112.00-114.25W): Oct 1974-Dec 1981

81 729 658 5.77 37-46.55 112-40.62 7.0* 1.6 9 183 35 .30 81 729 1652 53.11 37- 9.04 113-36.26 7.0* 1.8 12 201 72 .40 81 813 1840 55.00 37-39.85 112-29.29 7.0* 1.5 12 220 82 .28 81 827 1901 52.61 37-38.11 112-26.40 7.0* 2.1 13 134 87 .45 81 828 237 42.82 37-53.14 112-56.80 7.0* 1.6 12 179 43 .46 81 828 243 31.47 37-52.81 112-56.57 7.0* 1.4 12 181 44 .46 81 828 2119 9.00 37-51.96 112-55.91 7.0* 2.5 11 128 44 .39 81 828 2125 30.54 37-52.41 112-55.31 7.0* 1.5 12 127 45 .45 81 831 2316 5.54 36-46.38 112-53.95 7.0* 2.0 11 221 115 .53 81 929 2100 33.90 37-34.70 113-15.28 .4 1.6 12 204 28 .59 81 828 212 85 85 85 85 85 85 85 85 85 85 85 85 85	
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81 930 851 56.67 37-30.85 112-44.27 7.0* 1.8 10 215 67 .36	ď
81 1002 1956 27.43 37-39.90 112-58.68 23.7 2.0 8 156 7 .90	d
81 1004 2257 39.30 37-48.19 113- 3.60 .7 1.8 16 83 14 .65	ď
81 1016 1937 46.17 37-33.61 113-52.15 7.0* 2.2 15 166 44 .42	C
81 1031 1815 33.45 37-51.53 112-56.63 1.6 1.4 14 133 23 .93	d
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81 1116 432 49.33 36-55.03 112-31.39 7.0* 1.8 10 209 124 .47	d
81 1201 1339 11.94 37-19.51 113-19.42 7.0* 1.6 13 223 44 .55	đ
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APPENDIX K

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